https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

Annales Geophysicae Discussions

Supernova effects on middle and upper atmospheric nitric oxide and stratospheric ozone

[1]David E. Siskind, [2]McArthur Jones Jr., [2,3]Jeffrey W. Reep
[1]Computational Physics Inc., Springfield, VA, USA
[2]Space Science Division, Naval Research Laboratory, Washington, DC, USA
[3] now at Institute for Astronomy, University of Hawaii at Mānoa, Pukalani, HI, USA

Corresponding authors:
David Siskind (dsiskind@cpi.com)
McArthur Jones Jr. (mcarthur.jones16.civ@us.navy.mil)

Abstract

We provide a quantitative test of the recent suggestion (Brunton et al., 2023) that supernovae could significantly disrupt planetary ozone layers through a multi-month flux of soft X-rays that produce ozone-destroying odd nitrogen (e.g. NO and NO_2). Since soft X-rays do not directly penetrate down to the ozone layer, this effect would be indirect and require downward transport of NOx from the mesosphere. Mirroring previous studies of the indirect effects of energetic particle precipitation (EPP-IE), we call this the X-ray Indirect Effect (Xray-IE). We use the NCAR Thermosphere-Ionosphere-Mesosphere-Electrodynamics General Circulation Model (TIME-GCM) to simulate the production of NO and its transport into the stratosphere. We model the soft X-ray flux as if it were a multi-month long solar flare and use our previously developed solar flare model to simulate the soft X-ray enhancement. Our results yield significant enhancement in stratospheric odd nitrogen, most dramatically in the Southern Hemisphere. The most global effects are seen in the upper stratosphere at pressure surfaces between 1-3 hPa (about 42-48 km) consistent with previous observations of the EPP-IE. We then use a detailed stratospheric photochemistry model to quantify the effects of this NOx enhancement on ozone. Widespread ozone reductions of 8-15% are indicated; however, because these are limited to the upper edges of the ozone layer, the effects on the ozone column are limited to 1-2%. We thus conclude that the effects of a multi-month X-ray event on biologically damaging UV radiation at the surface is also likely to be small.

1. Introduction

As discussed by Airapetian et al., (2019) and summarized by Garcia-Sage (2023), the explosion of new discoveries of exoplanets and the search for life in the universe as led to increased recent interest in how space weather can influence the climate and habitability of the earth and possible life-bearing exoplanets. As the above articles discuss (see also Kahler and Ling, 2023), these extreme space weather events can include solar/stellar flares, coronal mass ejections, solar/stellar energetic particles (SEPs) and/or cosmic rays. There is, however, a parallel line of inquiry that has long considered the effects of supernovae on planetary biospheres (Gehrels., et al., 2003). As we will discuss, there is significant conceptual overlap in the specific mechanisms, which is a motivation for our present study.

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

Recently Brunton et al (2023) have proposed a new mechanism by which supernovae could threaten the existence of planetary biospheres. The classical mechanisms have traditionally invoked ozone depletion either due to gamma ray emission which would occur promptly (within 100 days) with the event, or from cosmic ray fluxes which could be emitted over a period on the order of 10-100 years (Gehrels, et al., 2003). Brunton et al., (2023) suggest a third mechanism from enhanced X ray emissions that might result from interactions between the supernova blast wave and the local interstellar medium. They present observed light curves showing X-ray emissions occurring over periods ranging from 6 months to several years after the initial eruption. They suggest that these emissions might represent a heretofore unexplored mechanism for planetary ozone destruction.

An important consideration for understanding the effect of enhanced X-rays on the ozone layer, which Brunton et al (2023) discuss, is the fact that X-rays with energies less than 10-20 keV are absorbed in the mesosphere, above the ozone layer. While Brunton et al., recognize that there may be X-ray emission from a supernova with greater energies, much of their data is limited to these softer X-rays. As a result, they suggest that the effect of X-rays would be more indirect and they quote some aeronomic studies (Solomon et al., 1982; Randall et al., 2006) of how perturbations to nitric oxide in the mesosphere and lower thermosphere could be transported down to the middle atmosphere where they would catalytically lead to ozone loss. Conventionally this coupling mechanism is due the production of nitric oxide (NO) in the auroral zones near 100 km altitude by energetic electron impact on N_2 followed by descent through the mesosphere into the stratosphere under the cover of polar night which limits the dissociation of the enhanced NO by UV sunlight. Randall et al., (2007) labeled this as the Energetic Particle Precipitation Indirect Effect (EPP-IE). Here, motivated by Brunton et al.'s hypothesis, we consider an analogous indirect effect on stratospheric odd nitrogen and ozone from continual soft X-ray influx, which we dub the "X-ray IE".

Brunton et al. (2023) provide estimates for the total amount of X-ray energy that might threaten planetary ozone layers and compared them to the integrated energy emitted by a multi-year solar flare. Specifically, they argue that a so-called Carrington flare (X45, i.e. 4.5 x 10^{-3} W m^{-2}), near the upper limit of flare energy release by the Sun (see e.g. Cliver et al 2022), would have to persist for 2.8 years to provide the requisite energy. Using this analogy, we will use an existing solar flare model (Siskind et al., 2022) and consider the consequences of previously considered solar flares extending for over a year. We will show how the X-ray IE can lead to a significant influx of nitric oxide entering the stratosphere and quantitatively model to what extent this influx could reduce ozone abundances. Ultimately, we conclude that due to the specifics of how NO is transported in the middle atmosphere, while significant effects are probable, the global destruction of the Earth's ozone layer is less likely.

The general outline of the paper is as follows. In Section 2, we introduce the solar flares that form the basis of our study, look at the initial response of lower thermospheric NO and compare our calculations with previously published observations of the nitric oxide response to solar flare. In Section 3 we document the descent of this flare-produced NO down through the mesosphere using a three-dimensional model of chemistry and transport of the middle and upper atmosphere

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

(the NCAR Thermosphere Ionosphere Mesosphere Electrodynamics General Circulation Model (TIME-GCM)). To validate the X-ray IE we will put it into context of our calculated EPP-IE which can be compared with the extensive literature on that topic. Finally, in Section 4, we perform photochemical modeling of the sensitivity of stratospheric ozone to the various enhancements in middle atmospheric nitric oxide suggested by the TIME-GCM. One limitation that we will discuss is that the 30 km bottom boundary of the TIME-GCM is right at the peak of the ozone layer. Thus our photochemical simulations are required to be able to extrapolate down to encompass the entire ozone column.

2. Solar Flare and thermospheric NO modeling

2.1. Solar Flare modeling

Our approach follows the suggestion of Brunton et al. (2023), namely, to model the multi-month soft X-ray flux as if it were a solar flare that lasted for months rather than the 30-60 minutes which is typical (cf. Rodgers et al., 2010; Table 3; also Reep et al., 2023). The advantage of this approach is that it allows us to utilize existing flare spectra (Siskind et al., 2022). These spectra were developed with the NRLFLARE model, a physical model of solar flare irradiance, which uses a series of flaring loop simulations to reconstruct the soft X-ray light curves of both GOES/XRS channels, and from those loop simulations, synthesizes full spectra from approximately 0.01 to 200 nm (Reep et al 2020; Reep et al 2022). The ratio of the two GOES/XRS channels is commonly used as a proxy for temperature, which the model uses to derive heating rates to drive those simulations (see e.g. Garcia 1994). The loop simulations are run with the open-source radiative hydrodynamics code HYDRAD (Bradshaw & Cargill 2013; Reep et al 2019, https://github.com/rice-solar-physics/HYDRAD), which solves the Navier-Stokes equations for plasma constrained to travel along a magnetic flux tube. The full model and spectral synthesis are described in detail in Reep et al 2022.

NRLFLARE was designed to reproduce X-ray spectra from solar flares, so it is important to discuss the differences and similarities to supernova X-ray spectra. In both cases, the spectra in soft X-rays (around 1 to 20 keV or so) are dominated by optically thin thermal bremsstrahlung emission with a power law shape, with notable line emissions from hot ions such as Fe XXV (a prominent line at 6.7 keV appears in spectra of both). There are two important differences. First, the elemental abundances are not the same, which will cause the relative strength of the emission (particularly line emission) to differ. Second, solar flares are expected to be in collisional equilibrium, while supernova remnants have low enough densities that the collisional timescale is long, so they are typically not in equilibrium. See the reviews by Vink (2012) for X-ray emissions in supernovae and Fletcher et al. (2011) for solar flares (Sections 6 of both reviews). For our purposes, the exact spectral shape is less important than the total energy input driving the atmospheric response.

One of the main subjects of the Siskind et al., (2022) paper was the September 10, 2017 X8.3 flare and a spectrum at flare peak was presented in that paper. We will use that as our primary case. Table 1 summarizes key aspects of that flare that are relevant for this paper. First, it is

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

Annales Geophysicae Discussions

important to note that in 2020, NOAA removed a 0.7 recalibration that had historically been applied to GOES 13-15 data (cf. https://ngdc.noaa.gov/stp/satellite/goes/doc/GOES_XRS_readme.pdf; also Reep et al., 2022)

Thus, the true X-ray irradiance for older flares is 1/0.7 brighter. This means that the 2017 flare, originally labeled as 8.3 in Siskind et al. 2022 and earlier works (Qian et al., 2019; Redman et al., 2018) should be re-classified as X11.8. Table 1 shows the calculated peak energy by the NRLFLARE model as being about 12% greater than observed by GOES, thus effectively making this flare an X13.3 event. We will thus use the label "X13" to describe this event as we discuss our atmospheric simulations.

Table 1

Event	**NOAA class**	**Calculated 0.1-0.8 X-class with NRLFLARE**	**Calculated energy flux, 0.1-1.0 nm (W/m^2)**	**Calculated energy flux, 1-2 nm (W/m^2)**	**Integrated energy ≥ 1 keV after 1 year (kJ/m^2)**
Sept 10, 2017	11.8	13.3	1.55e-3	.0017	64.4
Oct 28, 2003	X25	X27	.004	.007	171.4

Table 1 also shows the integrated energy in several energy bins. The division into 0.1-1.0 nm and 1-2 nm bins is to compare with the calculations of Rodgers et al., (2010), discussed below. The final column extrapolates our flare duration to a year. In particular, it shows that if we assumed the X13 flare persisted for an entire year, it would deliver 64.4 kJ/m^2 to the atmosphere. This is less than the 400 kJ/m2 that Brunton et al., (2023) use as a critical threshold for ecologically destructive Xray energy input. We will therefore also consider the energy input from a spectrum calculated for the October 28, 2003, the so-called Halloween event. The effects of this flare on thermospheric nitric oxide were first discussed by Rodgers et al., 2010 and we will compare our calculations to theirs. Again, due to the NOAA recalibration, this flare, which was originally classified as X18, should really be classified as X25. As seen in Table 1, our calculated energy at flare peak was about 8% higher than measured by GOES and thus we label this as an X27. If this flare were to persist at peak level for a year, Table 1 indicates it would deliver about 171 kJ/m^2 to the atmosphere. As shown by Brunton et al. (2023, their Figure 3), it is not uncommon for supernova X-ray events to persist for over a year. Table 1 shows that if our calculated X27 event were to last about 2.3 years it would deliver about 400 kJ/m^2 which is the energy input postulated by Brunton et al. (2023) as being biospherically destructive. Unfortunately, the problem with the X27 simulation is that when this spectrum was input continuously into the atmospheric model (TIME-GCM, discussed below), the model crashed after 8 days of the simulation. Thus, in our discussion of ozone chemistry effects, we will discuss extrapolations based upon comparisons of the nitric oxide response from the first 8 days of each simulation.

Figure 1 compares the spectra from our X13 and X27 calculations at their respective peak minutes. The figure shows the calculated spectrum at the native spectral resolution of

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

NRLFLARE (0.5 Å) and then integrated in 1 nm bins so that it can be compared to that derived by Rodgers et al (2010, see their Figure 3). Like Rodgers et al. (2010) our spectrum shows a significant increase in the flare spectrum from 1-2 nm relative to the shorter wavelengths less than 1 nm. As discussed by Siskind et al., (2022) this seems consistent with Orbiting Solar Observatory (OSO) data presented by Neupert et al (1967), although this spectral region is not well covered with modern spectra. Comparing our results in detail with Rodgers et al., suggests that our 0.1 – 1 nm result (.004 W/m^2) agrees well. Our 1-2 nm integrated energy is about 20% lower than theirs at flare peak. For the purposes of this paper, this difference is not significant; when we compare our calculated nitric oxide variation to Rodgers et al (2010), we can account for this difference by using integrated energy as the independent variable to normalize both our calculations. This will be discussed further in Section 4.

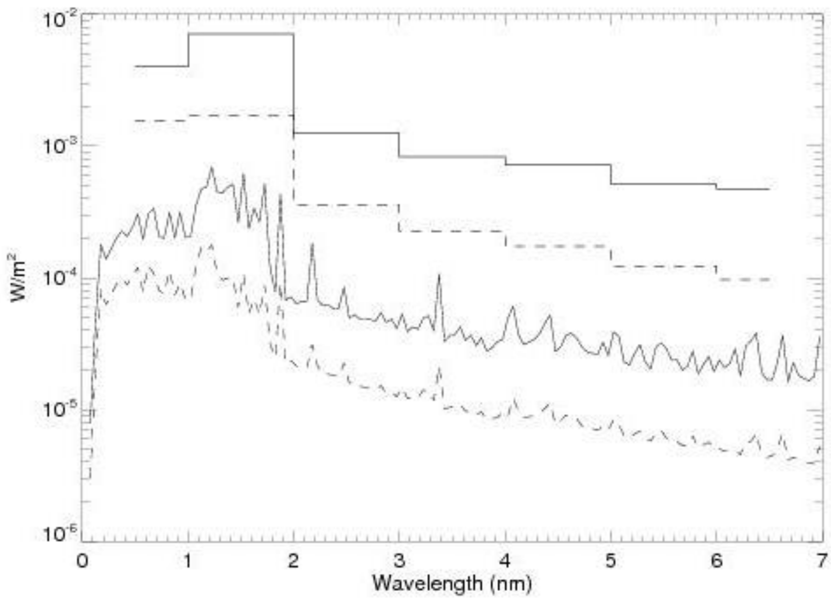

Figure 1. Calculated spectra for the peak of the X27 event of October 28, 2003 (solid lines) and the X13 event of Sept 10 2017. The rationale for the classifications is discussed in the text. The bottom two curves are at 0.5 A resolution. The histogram format for the top two curves is the integrated energy over 1 nm bins.

2.2. Atmospheric modeling with the TIME-GCM

The solar spectra shown in Figure 1 were used as inputs into the photoelectron ionization model presented by Siskind et al (2022) and incorporated into the NCAR TIME-GCM. The NCAR TIME-GCM is a hydrostatic general circulation of the middle and upper atmosphere that solves the continuity, electrodynamic, energy, and momentum equations from first principles on a

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

regular longitude and latitude and log pressure grid in the vertical (Roble and Ridley, 1994). The model resolution is 2.5° x 2.5° (longitude x latitude) and 4 grid points per vertical scale height extending from 12 to 4.6 x 10^{-6} hPa (or roughly 30 to 450-600 km depending on solar activity).

The photoelectron ionization model presented by Siskind et al (2022) defines 12 new wavelength bins for the soft X ray energy range to give better spectral resolution (and hence better altitude resolution of energy deposition) than the original NCAR spectral model presented by Solomon and Qian (2005). Note, there is a typographical error in Table 3 of Siskind et al., (2022), bin #7 for the O_2 cross section. It should read 1.5E-20, not 1.5E-21. It is correctly implemented in the model.

One difference in how we used the TIME-GCM from the short term (< 1 day) simulations of Siskind et al (2022) concerns the dynamics of the mesosphere. In the standard version of the TIME-GCM (i.e., the model setup used in Siskind et al., (2022)) climatological background horizontal winds, temperatures, and geopotential are used at the model lower boundary in combination with monthly mean diurnal and semidiurnal tides from the Global Scale Wave Model (GSWM; Zhang et al., 2010a,b). However, this standard model configuration does not properly simulate the downward transport of NOx from the mesosphere into the stratosphere. In order to do so, we constrained TIME-GCM upper stratospheric and mesospheric horizontal winds and temperatures between the model lower boundary (~30 km) and ~75 km with Modern Era Retrospective-analysis for Research and Applications - version 2 (MERRA-2, Gelaro et al., 2017) using four dimensional tendency nudging (originally termed 4D data assimilation by Stauffer and Seaman, (1990, 1994)). This nudging procedure is described in great detail by Jones et al. (2018), and involves adding an additional acceleration and energy tendency term to the conversation equations that is proportional to the modeled and MERRA-2 horizonal wind and temperature differences up to ~75 km.

In previous studies (e.g., Jones et al., 2020; 2023), TIME-GCM was constrained using a high-altitude version of the Navy Global Environmental Model (NAVGEM-HA, Eckermann et al., 2018; McCormack et al., 2017), which provides dynamical fields up to ~97 km. Note the MERRA-2 reanalysis product used herein does not extend as high as NAVGEM-HA, and therefore, we had to make a small modification to equation 5 of Jones et al. (2018). This equation describes the vertical weighting distribution of nudging, which in part controls the strength of the additional tendency term. The vertical weighting distribution used here takes the same functional form as equation (5) of Jones et al. (2018), but the z_{max} variable (representative of the TIME-GCM log-pressure level where the model becomes unconstrained) is equal to -10.5 or ~75 km. For reference, a vertical weighting factor of 0.5 occurs roughly at 55 km (or 0.2 hPa), above (below) which the nudging term is more weighted toward TIME-GCM (MERRA-2) dynamical fields.

2.3 Initial thermospheric response to multi-month solar flare

As discussed above, we model the effects of supernova induced soft X ray event as if it were a multi-month solar flare. Specifically, for the X13 event, we performed a simulation which continues through the end of 2017 and then covers a complete additional year. In the analyses discussed below, we present the results of the X13 and X27 simulations with a baseline run that

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

only includes the EPP-IE effect. The difference between the X13 or X27 and baseline runs serve to quantify the possible response of the middle and upper atmosphere to a multi-month soft X-ray event. We also note that for TIME-GCM simulations performed herein geomagnetic activity was held constant with Kp ≅ 3 in order to exclusively highlight flare impacts.

Figure 2 shows the initial response at low latitudes (averaged from 30S-30N), plotted every two hours, as a function of longitude for the first day. The solid line is 1600 UT which was just at flare onset (the peak of the Sept 10, 2017 flare was around 1606 UT). The four dashed lines are for 1700, 1900, 2100 and 2300 UT and show how the NO increases both in the thermosphere (panel (a)) and in the mesosphere (panel (b)) immediately after flare onset. Note how the longitudinal response progresses westward for the equatorial plots, tracking the sub-solar point. This is consistent with our implicit assumption that the supernova will be aligned with the ecliptic plane. While perhaps not always true (the galactic plane is tilted 60^{o} with respect to the ecliptic plane (cf. https://en.wikipedia.org/wiki/Astronomical_coordinate_system), any supernova will nonetheless rise and set like the sun, and the peak effects will, like with a solar flare, be concentrated at the sub-stellar longitude. Thus we conclude that our approach of using an extended solar flare event as a means of simulating a supernova soft X-ray event is acceptable.

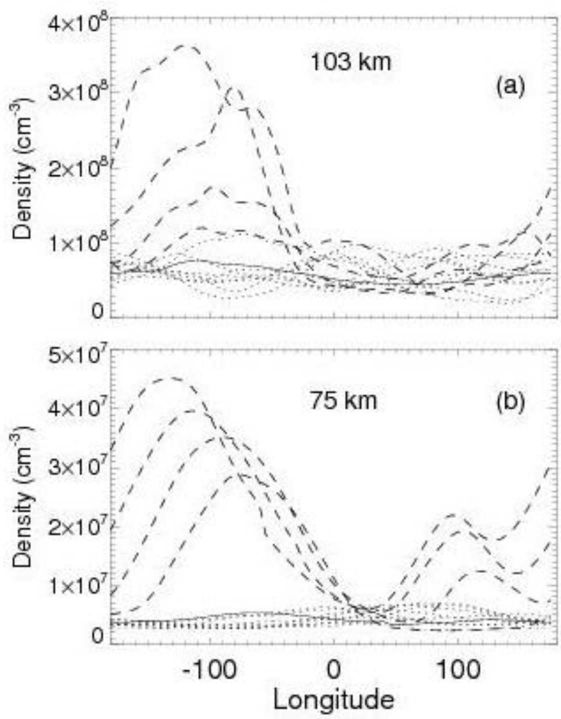

Figure 2. Initial response of thermospheric (panel **(a)**) and mesospheric (panel **(b)**) nitric oxide density to the onset of the extended flare. The solid line in each panel is for 1600 UT, which roughly corresponds to the onset of the flare. The dotted lines are for times prior to that. The dashed curves which progressively

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

increase and phase to the left according to the sub-solar point are for hours 1700, 1900, 2100 and 2300 UT.

Figure 3 shows daily averaged profiles for the first 10 days for the event, both at low and at high latitudes. The effects are largest at the equator, but are still significant at 59S, and extend well down into the mesosphere. Note the changes appear to level off after several days, suggesting that the initial response is saturating. Indeed we find that all the thermospheric response occurs in the first 10-14 days. The middle atmosphere response includes both this initial effect and then later, seasonal effects as NO is transported down from the upper mesosphere/lower thermosphere.

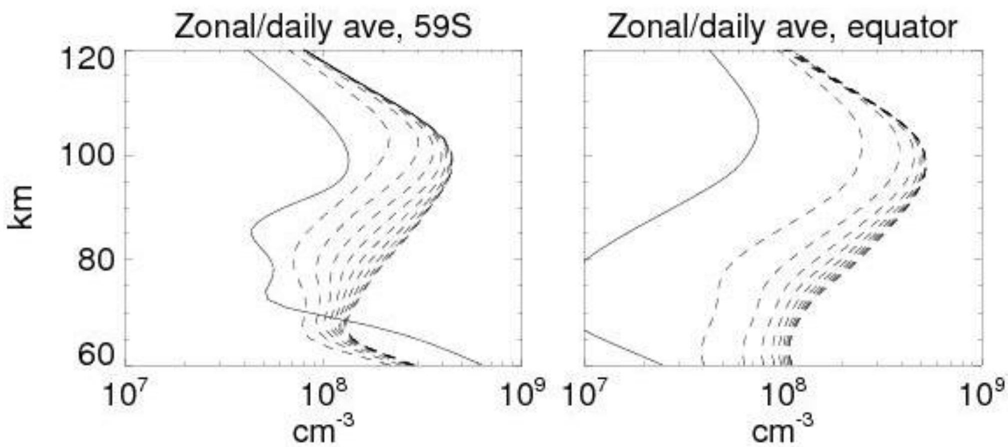

Figure 3. Profiles of the first 10 days of the nitric oxide profile at two latitudes. The individual days are not labeled, but the day-to-day increase in NO density is monotonic with time. The solid lines are pre-flare.

3. Seasonal Variation of the Xray-IE in the middle atmosphere

Figure 4 compares the seasonal variation of the TIME-GCM NOx (defined as NO + NO_2) from our extended flare calculation with our baseline run that only includes the EPP-IE. It thus shows the seasonal variation of how the Xray-IE leads to NOx buildup in the middle atmosphere beyond that caused by energetic electron precipitation. To understand this, we first focus on our baseline EPP-IE simulation and how it compares with the recent simulations of the EPP-IE from Pettit et al., (2021), specifically their Figures 9-10 which they compared with Michaelson Interferometer for Passive Sounding (MIPAS) data in the Southern Hemisphere. Ultimately, we will conclude that the Xray-IE shows similar behavior to the EPP-IE simulation, except with a larger magnitude and for a more prolonged seasonal duration. Thus to highlight the longer impact, we show the entire year whereas Pettit et al. (2021) just showed April-October.

In comparing with Pettit's results, we see that our baseline simulation underestimates the descent of the MIPAS NOx data at the higher latitudes. The MIPAS data show the 16 ppbv contour

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

descending to below 35 km for the month of August, whereas our simulation (panel a) has this contour remaining above 40 km for the late austral winter period. There are likely two reasons for this. First, is likely the simple fact that TIME-GCM has a bottom boundary at 30 km and thus the descent will decay as this boundary is approached. Indeed, analyses of data from both the Halogen Occultation Experiment (HALOE) on board the Upper Atmospheric Research Satellite (UARS)and Polar Orbiting Aerosol Measurement (POAM) data have shown that enhanced NOx can routinely be detected below 30 km in the Southern Hemisphere (Siskind et al., 2000; Randall et al., 2007). Second, our model does not have the medium energy electron ionization that Pettit et al (2021) discuss. They show that models without this component of energetic electrons underestimate the descent into the mid-stratosphere.

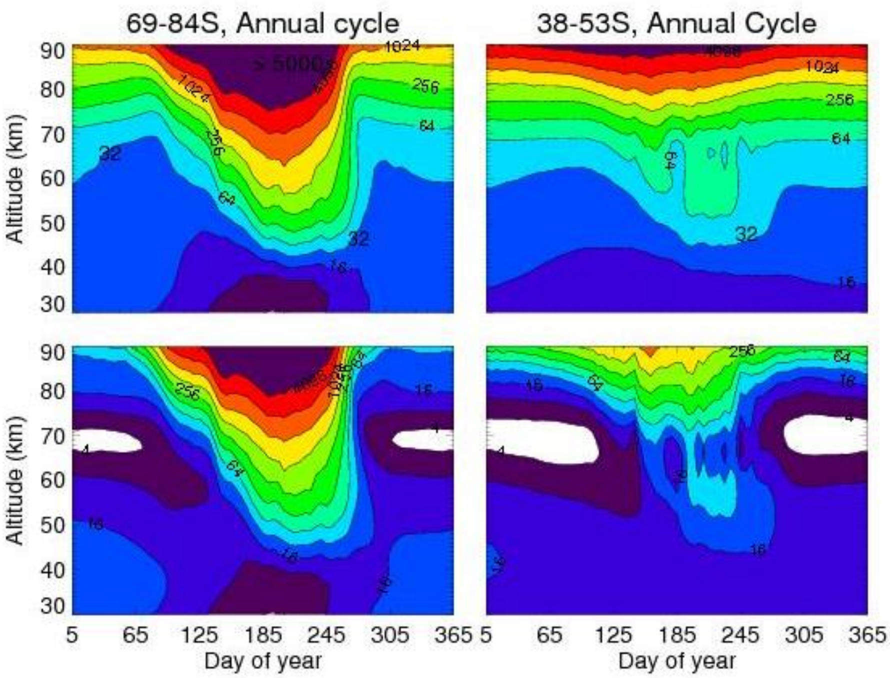

Figure 4. Annual cycle of NOx descent into the upper stratosphere from TIME-GCM for two latitude bands. The bottom row is for a baseline simulation that only includes the EPP-IE. The top row additionally includes the Xray-IE from the X13 simulation presented in Figures 1-3. The year shown is 2018 thus representing the period about 4-12 months after flare onset on Sept, 10, 2017. The values on the contour labels are in units of ppbv. The white colored regions in the baseline run are for mixing ratios < 4 ppbv.

On the other hand, our baseline simulation does much better at mid-latitudes (38-53S in the figure). It shows the 16 ppbv contour dipping down to 45 km for a couple of months. This is quite similar to the MIPAS data shown by Pettit et al., (2021) and is consistent with Funke et al (2005) and Arnone and Hauchecorne (2011) who pointed out that there are two components to the descent of upper atmospheric NOx into the stratosphere. One component is directly into the

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

stratospheric polar vortex and descends down into the mid-stratosphere; as we note above, our model cannot capture this. However, there is a second component that is dispersed into middle latitudes in the upper stratosphere. It appears that our model does capture this and it could be argued that from a global biospheric perspective, this second component is more important since a greater region of the globe is affected.

Regarding our Xray-IE simulation, dramatic effects are clearly seen in the mesosphere, both mid and high latitudes. The mesospheric minima near 70 km are completely filled in and mixing ratios of over 32 ppbv, up to near 100 ppbv, are seen for most of the year. However, for considerations of impacts on ozone, we focus more on the stratospheric effects. Here, at first glance, for the higher latitudes, the IE-Xray effect appears somewhat muted. We see no difference in the maximum value of NOx descending below 50 km between our baseline and constant X13 simulation. However, the IE-Xray effect is somewhat more prolonged in its NOx enhancement. The baseline simulation shows the 16 ppbv contour curving sharply upward around Day 270. Thus NOx values near 50 km decrease abruptly and this is similar to what is seen in Pettit et al.'s MIPAS data. However, the X13 simulation shows the upper stratospheric NOx values remaining between 16-32 ppbv for the entire austral spring.

At mid-latitudes, the effect of the continual soft X-ray flux is more pronounced. Whereas the baseline simulation shows 16 ppbv descending to about 45 km, the flare simulation has about double that. Like the high latitude case, after approximately Day 270, the baseline case NOx values fall below 16 ppbv, in agreement with the MIPAS data. By contrast, in the X13 simulation we see NOx values of 32-64 ppbv descending to 45-50 km and the entire upper stratosphere remains flooded with enhanced NOx values greater than 16 ppbv for the whole year.

Figure 5 also compares our baseline (EPP-IE only) simulation with that including the Xray-IE, this time for two pressure surfaces as a function of latitude and time: one near the stratopause (the indicated pressure roughly corresponds to altitudes of 45-48 km) and one lower down towards the middle stratosphere (approximately 38-40 km). The figure shows how the NOx from the flare/supernova spreads over the Southern Hemisphere. It is useful to first look at our baseline case; it clearly shows that the EPP-IE effect is mainly in late winter/early spring in the Southern Hemisphere and covers the latitudes from -80 to about -20 or -30. Note, there is no evidence for this seen at 3.0 hPa whereas in actuality, there should still be a spring time enhancement in the highest latitudes as we discussed above. When we compare this with the top row in the figure, the effects of the soft X-rays are very apparent. The late winter/spring enhancement at 1.1 hPa is about twice as large and there is now seen an enhancement at 3.0 hPa whereby values of NOx of 10-12 ppbv at Southern mid-latitudes are now replaced by values of 14-16 ppbv. Importantly, there is no evidence for significant enhancements in the Northern Hemisphere although there does seem to be a general global increase in NOx of about 2 ppbv- about 20% above the baseline values. This lack of significant NH enhancement is consistent with observations of the EPP-IE which show generally weaker effects in the NH relative to the SH (Funke et al., 2014). This is generally believed to be due to the weaker descent in the NH and the greater horizontal mixing due to mesospheric planetary waves (Siskind et al., 1997), although NH enhancements are seen in specific years with very strong dynamical perturbations (cf. Funke

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

Annales Geophysicae Discussions

et al., 2017). In the present case, while we will consider the effects on stratospheric ozone below, it does suggests a limit as to how biospherically destructive the soft X-ray event could be since the effects are likely to be much more muted in the NH

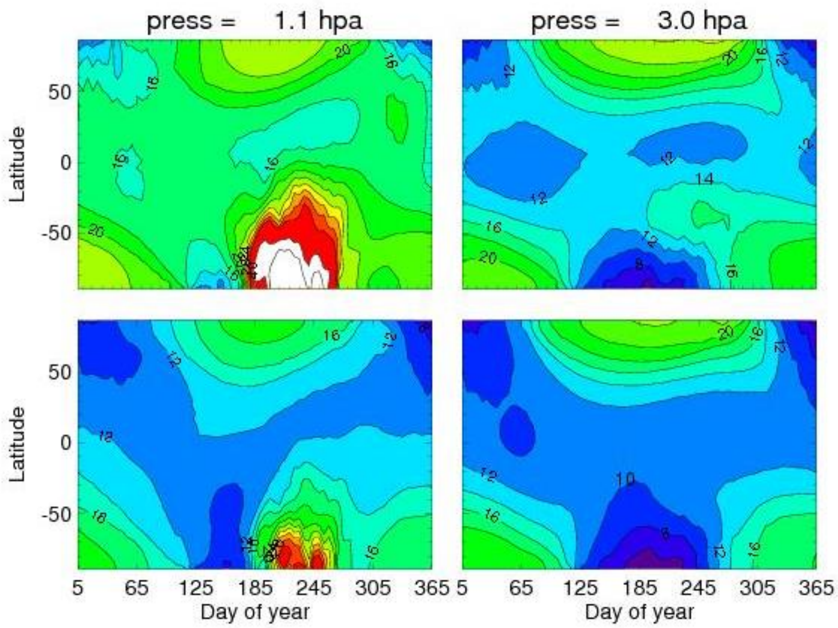

Figure 5. NOx (ppbv) vs latitude and day of year. The period of time is the same as shown in Figure 4. The bottom row is for the baseline case without enhanced soft X-rays; the top row includes the continuous X13 flux. The red regions are NOx values greater than 28 ppbv; the white regions are NOx values greater than 40 ppbv.

One final consideration in looking at the annual cycles in the upper stratosphere mesosphere in Figures 4 and 5 is that there appears to be no evidence for any continual buildup of NOx. The NOx at the end of 2018 is not much different than at the beginning. This is consistent with Figure 3 in that the day-to-day NO increase in the thermosphere decreases such that after 10 days the NO profile showed little change. This will be important when we try to extrapolate from our X13 simulation to stronger events.

Figure 6 shows the global change in ozone for the X13 simulation compared with our baseline EPP-IE only case for four pressure surfaces ranging from 0.68 to 3.0 hPa. The values are less than 1.0 globally for the entire year which means lower ozone for the X13 simulation. However, there is a clear maximum in the reduction for the late winter/early spring period in the SH, consistent with the global distribution of the enhanced NOx shown in Figure 5. Note that the fractional reduction is larger at the lowest pressures. Normally, at these altitudes in the lower

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

mesosphere, ozone loss is dominated by the HOx catalytic cycle (Brasseur and Solomon, 2005). However, with NOx enhancements on the order of 100 ppbv, the NOx catalytic cycle can dominate up to higher altitudes (lower pressures) than is conventional. At the same time, since the bulk of the ozone density is in the stratosphere, the effect of a 3-4% reduction at 3.0 hPa is of greater impact than a 10% reduction at 0.68 hPa.

The results show here clearly suggest a potentially global effect on the ozone, albeit limited to a couple of months when the SH NOx enhancement has spread to the equator. The effect is not large- about 5% locally in the upper stratosphere and thus unlikely to be biospherically significant. However, there are important caveats to this statement that we will explore in the subsequent section. First, as we noted above, our input Xray energy is much smaller than the supernova soft Xray events postulated by Brunton et al (2023). Second, the TIME-GCM is limited by a bottom boundary at 30 km. About half of the stratospheric ozone column lies below this altitude and must be considered before drawing any conclusions. We consider both these issues in the sections below.

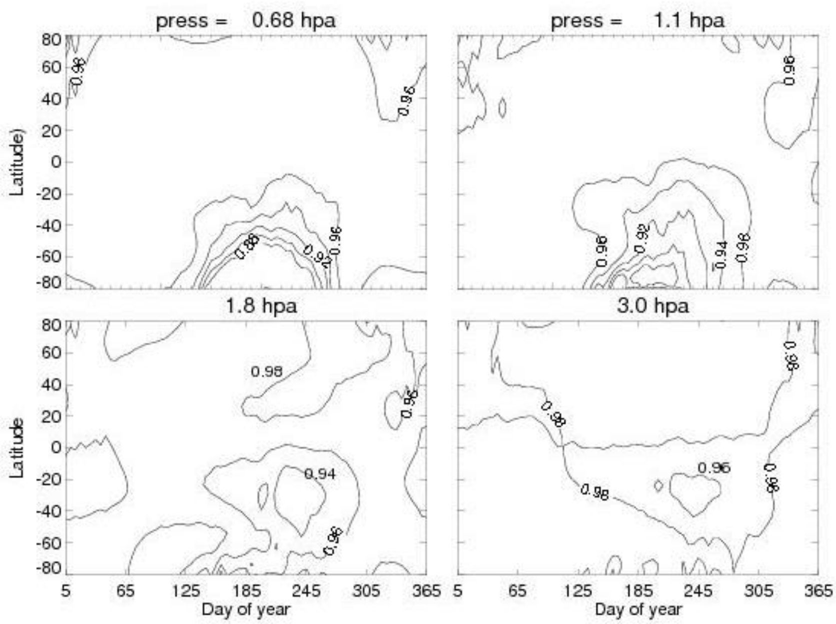

Figure 6. Annual variation of the ratio of ozone from the X13 simulation compared with the baseline simulation at the 4 indicated pressure surfaces

4. Extrapolation to higher Xray fluxes and impact on stratospheric ozone

To extrapolate our NO/flare response, we first seek to compare our results with observations of the NO response to solar flares. The only quantitative analysis of the response of nitric oxide to a

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

solar flare that we are aware of is that by Rodgers et al. (2010) using data from the Student Nitric Oxide Explorer (SNOE). SNOE was particularly well suited to study the NO response to a solar flare because it was in a sun-synchronous orbit with an equator crossing time in the late morning when the sun was relatively high in the sky. Rodgers et al. calculated the NO column change observed by SNOE and plotted it versus the integrated soft Xray input energy derived from a catalog of 11 flares.

Figure 7 compares the TIME-GCM results to Rodgers. The figure shows the integrated energy from the four strongest X-class flares observed by SNOE with the largest being the so-called Halloween event of October 28, 2003. As noted above, this event, labeled as X18 in Rodgers et al.'s Table 3, is now recalibrated to be X25, and in our simulation with NRLFLARE it is a bit higher at X27. Also shown are the TIME-GCM calculated hourly column NO from the local equatorial sub-solar longitude for each of the first 24 hours of our model simulations for the X13 and X27 events.

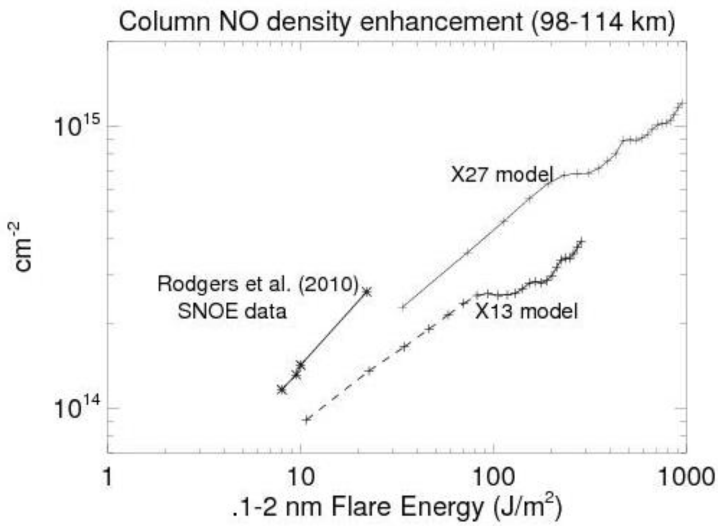

Figure 7. Calculated TIME-GCM NO column density enhancement from the X13 and X27 simulations compared with the observed NO increases reported by Rodgers et al. (2010) for the 4 strongest flares listed in their Table 3. The plus symbols on the model curves represent output for every hour. The first points shown for each of the model account for the number of minutes after each integral hour that the flare peaked. Thus the X13 flare peak was at 16.1 UT (cf. Table 1 of Siskind et al., 2022) and thus the first point shown for the X13 model represents 54 minutes of photon flux. Like Rodgers et al. (2010), we subtracted the pre-flare NO column in the model before calculating the enhancements shown.

In general, the figure shows a quasi-linear relationship between column NO and the integrated energy for both SNOE and the two model simulations. It appears that the rate of energy input is important for the NO increase. Thus after two model hours, the X13 simulation accrues the same energy input as the 27 minute long October 28, 2003 flare and yet the NO column response is

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

well below the observations. The column NO for the X13 simulation takes over 4X the energy input of the observed flare to reach the same enhancement as observed by SNOE. The column NO for our X27 simulation, which is designed to simulate the October 28, 2003 flare comes closer and matches the SNOE data just after the first hour of the model simulation (actually 51 minutes since the flare peak was at 9 minutes past 11 UT and model output was only saved hourly). However, since the actual October 28 flare only lasted 27 minutes, it means that the TIME-GCM is calculating a smaller NO column for the same energy input than was recorded by SNOE. Rodgers et al. (2010) reported an observed column enhancement of 2.6E14 cm^{-2} for solar X ray input of 22.4 J/m^2 where, reading from the graph, the TIME-GCM requires closer to 40 J/m^2 before reaching this level of NO enhancement.

After 24 hours, Figure 7 shows that the X27 simulation produces about a factor of 3 more NO than the X13 simulation. Figure 8 shows the daily averaged, zonal mean column NO for both models extended out to the full 8 days of the X27 simulation before the model crashed. Similar to Figure 3, it shows that both models level out after several days. The ratio of the two column densities equilibrates to a slightly smaller value than seen in Figure 7, about a factor of 2.6. The fact that the column densities level out can offer a useful guide for extrapolating our middle atmosphere NOx enhancements even without completing a full year with the X27 simulation. It suggests that the reasonable enhancements might lie in the range of a factor of 2-3 over the X13 simulation. We will explore this below.

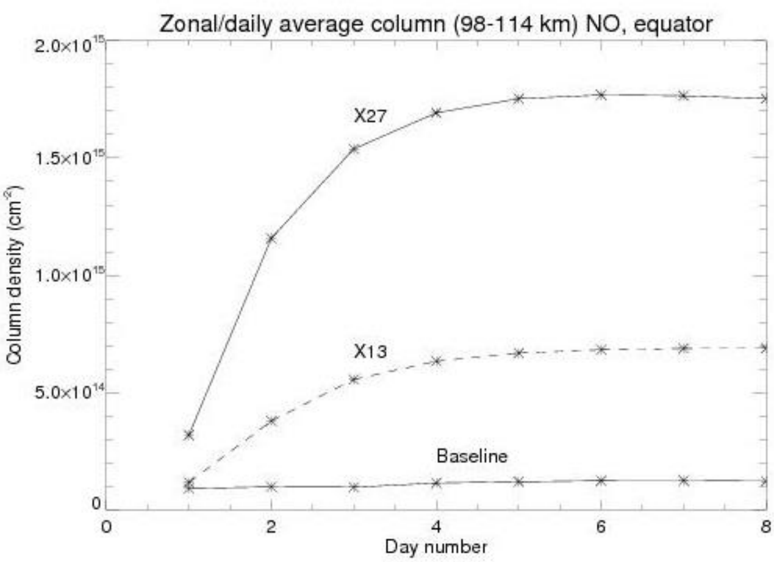

Figure 8. Daily and zonally averaged equatorial column densities for the X27 (solid line with stars) and X13 (dashed line with stars) TIME-GCM simulations. A baseline case run for the conditions of September 2017, but with no flare/supernova and which remains at approximately 1 x 10^{14} cm^{-2} is also shown.

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

To evaluate in detail how ozone may be reduced for the X27 simulation, we will use the CHEM1D photochemical box model. This model has previously been used to model satellite observations of mesospheric OH (Siskind et al., 2013) and validate ground based measurements of ClO (Nedoluha et al., 2020). It is important to first evaluate the model's ability to calculate stratospheric ozone since, as is most recently discussed by Diouf et al. (2024), chemical models of upper stratospheric and lower mesospheric ozone historically fall short of fully reproducing observations.

Figure 9 shows a comparison of CHEM1D and TIME-GCM ozone with two observations from September 2nd, (Day of year 245) 2018 at a latitude of 38-40S. This period and location was selected because it corresponds to the time and location of the most significant upper stratospheric ozone depletions indicated by the TIME-GCM in Figure 6. The observations are from the 9.6 μm measurement of the Sounding of the Atmosphere with Broadband Emission Radiometry (SABER) instrument on board the NASA TIMED satellite and the Microwave Limb Sounder (MLS) from the NASA Aura satellite. SABER and MLS data have long been the standards for measuring middle atmospheric ozone globally. Figure 9 shows, first, that TIME-GCM is ill suited for model-data comparisons of stratospheric ozone. This is perhaps not a surprise- the model was designed to study middle atmospheric dynamics and transport and its coupling to the upper atmosphere (Roble et al., 1994). For example, TIME-GCM does not include all the active chlorine and nitrogen species that are required for a comprehensive model of stratospheric ozone. Thus for chlorine, TIME-GCM has Cl and ClO, but not HOCl. For nitrogen, TIME-GCM only has NO and NO_2, but not HNO_3 or N_2O_5. By contrast, CHEM1D does include these species. The comparison with CHEM1D very closely matches that seen by Siskind et al. (2013), who used CHEM1D for mesospheric ozone and hydroxyl and Diouf et al. (2024), who used the model of Bertaux et al. (2020) and compared with MLS ozone and SABER $O_2(^1\Delta)$ 1.27 μm emission. In all cases, the model falls short of completely reproducing the observations. Both Siskind et al. (2013) and Diouf et al. (2024), having exhausted all possibilities for reaction rate changes and possible temperature inputs, invoked the possibility of an additional source of ozone from vibrationally excited oxygen as hypothesized by Slanger et al., (1988) and Price et al., (1993). The purpose here is not to answer this long standing question; rather, Figure 9 shows that CHEM1D does as well as could be expected given our understanding of middle atmospheric ozone photochemistry. Our purpose here is to perform sensitivity studies for varying amounts of NOx, guided by our TIME-GCM simulations. Figure 9 shows that CHEM1D is adequate for this task. We should additionally note that as one moves towards higher pressures greater than 5 hPa, the chemical lifetime of ozone becomes longer such that it is no longer under pure chemical control but also dynamical influences. Thus, the apparent improved agreement with the observations near 10 hPa should not be over-interpreted.

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

Annales Geophysicae Discussions Open Access EGU

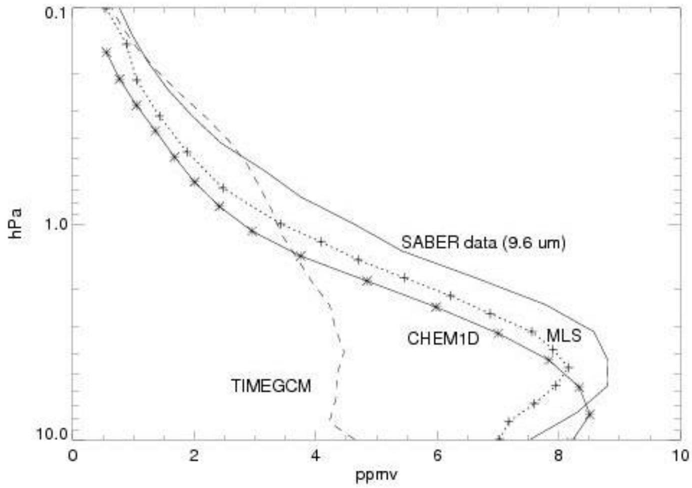

Figure 9. Comparison of the TIME-GCM (long dashes) and CHEM1D (solid line with stars) models with SABER (solid line) and MLS (dotted line with plus symbols) observations of ozone. The location is 38-40S and the time of year is September 2nd, 2018. CHEM1D used temperature and pressure and NOx abundances from the TIME-GCM as input. The approximate altitude range corresponding to the y-axis is about 30-62 km.

We now show the fractional ozone depletions, as a function of pressure, from the enhanced NOx due to a multi-month solar flare. Figure 10 presents the calculated ozone loss ratios (panel a) for two models of CHEM1D that use enhanced NOx compared with the baseline simulation presented in Figure 9. The location and time of year is the same as in Figure 9. The NOx enhancements (panel b) are taken from the X13 simulation shown in the previous figures plus an extrapolated enhancement (the greater of the curves in Figure 10) based upon the short term response shown in Figure 8. Figure 10 also shows the vertical profile of the TIME-GCM ozone change taken from Figure 6.

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

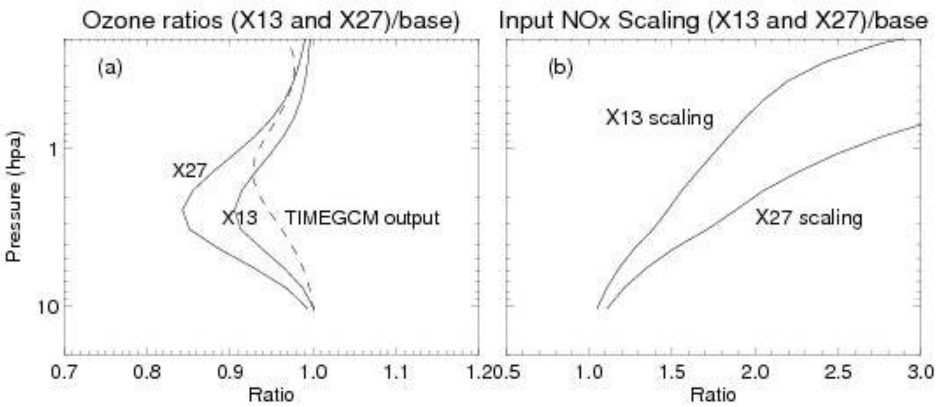

Figure 10. (a) Ratios of calculated ozone from CHEM1D compared with a baseline (no flare) case for September 2, at a latitude of 39S. The two solid lines use NOx input according to the scaling ratios shown in panel **(b)** The X13 scaling is based upon the NOx shown in Figures 3-5. The X27 scaling is a hypothesized extrapolation based upon Figure 8 and discussed in the text. Also shown as the dashed line in panel (a) is the ozone ratio from the TIME-GCM as per the surface contour plots shown in Figure 6.

Figure 10 shows that for the X13 case, we could expected ozone depletions of up to 8% in the upper stratosphere. For the more significant X27 case (i.e. for a more intense supernova X-ray event), we might see ozone reductions of up to 15-18% in the upper stratosphere. Figure 10 also shows the vertical profile of the TIME-GCM ozone reduction. It does not exactly match the profiles from CHEM1D in terms of shape and altitude of peak reduction, but it is very close to the X13 CHEM1D simulation in terms of giving a peak loss of 6-7% in the upper stratosphere. The TIME-GCM result is useful because it allows our detailed CHEM1D calculations to be placed in the global context shown in Figure 6.

Based upon Figure 10 and Figure 6, we can conclude that a supernova X-ray event could cause widespread ozone loss in the 10-20% range in the upper stratosphere for late winter/early spring in the Southern Hemisphere. While this would likely be easily observable with suitable instrumentation, it is less likely to have a dramatic biospheric effect. This is because most of the stratospheric ozone is found at altitudes from 20-35 km (5 hPa-50 hPa pressure levels). The losses shown in Figure 10 are only the upper edge of that layer. This is shown in Figure 11, which shows the actual ozone mixing ratios (panel (a)) and ozone density profiles (panel (b)) which correspond to the scaling ratios shown in Figure 10. In the case where the model output is shown as ozone densities, the curves are almost indistinguishable. The change in the total column ozone, which is most relevant for surface UV exposure, is 1% for the X13 simulation and 2% for the X27 extrapolation.

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

Annales Geophysicae Discussions

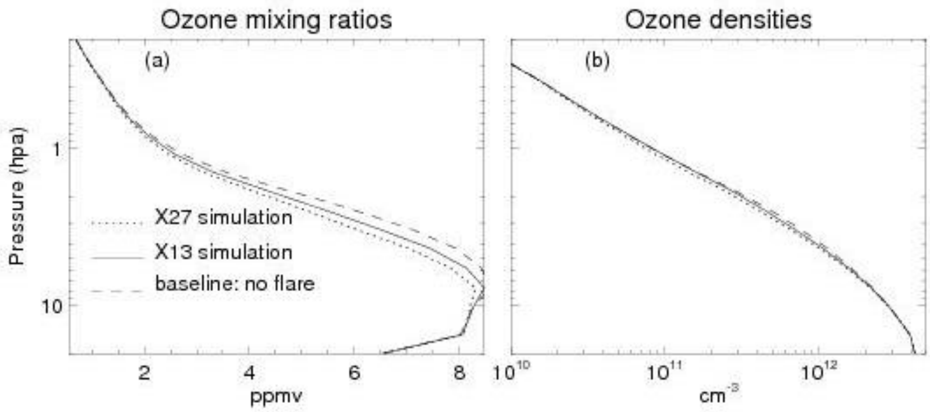

Figure 11. Absolute ozone abundances corresponding to the ratios presented in Figure 10. The three simulations are labeled in panel **(a)**. They are identically shown in density units in Panel **(b)** but are almost indistinguishable because the 8-15% reductions are very hard to see on a graph that covers over two orders of magnitude.

5. Discussion and conclusions

Our results clearly suggest the strong possibility of globally widespread ozone loss in the upper stratosphere, at least for a period of a couple of months in the Southern Hemisphere. However, at the same time, we conclude that this is unlikely to have a global biospheric impact because the depletion is limited to the upper edges of the ozone layer. This limitation is derived from our simulations showing that, like the EPP-IE, the Xray-IE does not penetrate below 35-40 km on a global basis. At polar latitudes, our results allow us to speculate that a supernova could greatly exacerbate the ozone hole. Or even, for atmospheres without anthropogenic chlorine, create an ozone hole. Indeed, it has already been noted that the EPP-IE has been confused with an expansion of the ozone hole due to volcanic aerosols (cf. Siskind et al., 2000 and discussion therein). However, since the hole is generally confined to the polar vortex, the effects of the Antarctic ozone hole have not caused widespread global ecological destruction although regional effects may be occurring (Robinson et al., 2024). There are likely other more subtle hypothesized effects of the enhanced NOx that we do not address. For example, we do see moderate NOx enhancements throughout the Northern Hemisphere and it has been suggested that EPP-IE in the Northern Hemisphere has effects on stratospheric and possibly tropospheric meteorology (Seppala et al., 2009). Our work here cannot rule this out for the Xray-IE.

Certainly, our results come with large uncertainties that would be useful to address. Perhaps the biggest is that the TIME-GCM, with a bottom boundary above the peak of the ozone layer, is not designed to study stratospheric chemistry. Moreover, the 30 km bottom boundary prevents us from studying descent of NOx enriched air down to the lower altitudes where the EPP-IE has been observed in the SH polar vortex (Randall et al., 2007). Thus our comments about the ozone hole are necessarily speculative. In addition, our simulation of the NO produced during solar flares appears to be less than observed by SNOE. This might mean that the NO response to a flare would be greater than we suggest, perhaps by as much as a factor of 2. Here it would be very helpful if there were another dataset that could corroborate the NO response reported by Rodgers et al., (2010). As we noted above, the local time of the sun-synchronous SNOE orbit was ideal for observing solar flares. By contrast, more recent NO observations which are summarized in Table 1 and Figure 3 of Emmert et al., (2022) are less well suited. Emmert et al. (2022) show that, for example, the Atmospheric Chemistry Experiment (ACE) and the Solar Occultation for Ice Experiment (SOFIE) on the NASA/AIM satellite used the technique of solar occultation which by definition means sunrise or sunset. This type of observation is not well suited to observing the effect from a flare which would be less noticeable at local sunset or sunrise. Likewise the ODIN satellite which measured NO with the Sub-millimeter radiometer (SMR) was in a dawn-dusk synchronous orbit. Based upon Emmert et al., (2022) it appears that only MIPAS on the ENVISAT satellite was in a proper daytime orbit to see flares. An examination of the MIPAS data might be an interesting test of some of our SNOE-based results.

Ultimately, however, even if we did underestimate the NO production by a factor of 2 or even 3, the effects on the ozone column are likely not catastrophic because they will be limited to above 35-40 km. We point to the simulations of Thomas et al., (2007) of a possible solar proton event that may have accompanied the 1859 Carrington flare event. Solar protons penetrate much deeper into the stratosphere than soft X rays and thus the effect on NOx is more direct rather than indirect as simulated here. Indeed, they obtained much larger NOx increases down to 30 km and localized ozone losses near 35-40 km of greater than 30%. Despite this greater increase in NOx and greater ozone loss, their calculated perturbation to the ozone column was less than 15% because the bulk of the ozone density between 20-30 km remained unaffected from the proton flux. By contrast, other phenomena linked to supernovae, such as cosmic rays, are known to be absorbed by the atmosphere near the peak of the ozone layer in the 20-30 km altitude range (Melott et al., 2017) and, in our assessment, those are likelier candidates for causing global ozone destruction that would greatly enhance the flux of destructive UV radiation to the surface. However, we should conclude by noting that the destructiveness of both the gamma ray and cosmic ray mechanisms have also been recently called into question (Christoudias et al., 2024). Our calculations here are therefore consistent with Christoudias et al, (2024) in showing how the earth's atmosphere can shield its biosphere.

Code and Data Availability. The TIME-GCM code is available by contacting the National Center for Atmospheric Research. The model output produced herein is reproducible from the TIME-GCM model source code following the discussions and implementations of the nudging schemes and lower boundary conditions described thoroughly in Sections 2.4 and in Jones Jr. et al. (2018) and Jones Jr. et al. (2020). Daily NCAR TGCMs outputs in netCDF format from this study are archived on the DoD HPCMP long-term storage system. MERRA-2 middle atmospheric horizontal winds and temperatures used for constraining TIME-GCM dynamics are available at https://disc.gsfc.nasa.gov/datasets?project=MERRA-2. The SABER and MLS data used in Figure 9 were respectively obtained from https://saber.gats-

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

Annales Geophysicae Discussions

inc.com/ and https://mls.jpl.nasa.gov/eos-aura-mls/data.php. Other model output such as CHEM1D and specific supernova output from TIMEGCM are now in the process of getting approval at NRL for eventual public release and will be made available once this paper is accepted for publication.

Author Contributions. DES conceived the study, performed the analysis of the TIMEGCM output, conducted the CHEM1D analysis and led the writing. MJJr. configured the TIMEGCM, both to be nudged by MERRA and to input the NRLFLARE spectra, performed the simulations and wrote Section 2.2. JWR is the developer of NRLFLARE; he provided the soft Xray spectra used by the TIMEGCM and wrote Section 2.1.

Competing Interests The contact author has declared that none of the authors has any competing interests.

Acknowledgements. This work was supported by the Office of Naval Research. We also acknowledge the NASA Living with a Star program for supporting development of NRLFLARE and the development of the supernovae soft X-ray spectra. Computational resources for this work were provided by the U.S. Department of Defense (DoD) High Performance Computing Modernization Program (HPCMP).

Financial support. This research has been supported by the Office of Naval Research (6.1 funding).

References

Airapetian, V. S., Barnes, Cohen et al., Impact of space weather on climate and habitability of terrestrial-type exoplanets (2020), *Int'l Journal of Astrobiology,* https://doi.org/10.1017/S1473550419000132

Arnone E. and Hauchecorne A., (2011) Stratospheric NOy species measured by MIPAS and GOMOS Onboard ENVISAT 2002-2010: Influence of plasma processes onto the observed distribution and variability, *Space Sci Rev.*, 168, 315-332, DOI 10.1007/s11214-011-9861-1

Bertaux, J. L., Hauchecorne, A., Lefevre, F., Breon, F. M., Blanot, L., Jouglet, D., et al. (2020). The use of the 1.27 µm O_2 absorption band for greenhouse gas monitoring from space and application to MicroCarb. *Atmospheric Measurement Techniques*, *13*(6), 3329–3374. https://doi.org/10.5194/amt-13-3329-2020

Bradshaw, S. J. and Cargill, P. J., "The Influence of Numerical Resolution on Coronal Density in Hydrodynamic Models of Impulsive Heating", The Astrophysical Journal, vol. 770, no. 1, IOP, 2013. doi:10.1088/0004-637X/770/1/12.

Brasseur G. and S. Solomon, 2005, *Aeronomy of the Middle Atmosphere,* D. Reidel Publishing Co.

Brunton, I. R., O'Mahoney, C., Fields, B. D., Melott, A. L., and Thomas, B. C., (2023) X-Ray-luminous Supernovae: Threats to terrestrial biospheres, *Astrophys. J.*, 947:42, https://doi.org/10.3847/1538-4357/acc728

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

Cliver, E. W., Schrijver, C. J., Shibata, K., and Usoskin, I. G., "Extreme solar events", Living Reviews in Solar Physics, vol. 19, no. 1, 2022. doi:10.1007/s41116-022-00033-8.

Christoudias, T., Kirkby, J., Stolzenburg, D., Pozzer, A., Sommer, E., Brassuer, G. P., Kulmala, M., Lelieveld, J., (2024) Earth's atmosphere protects the biosphere from nearby supernovae, Nature Communications: Earth and Environment, https://doi.org/10.1038/s43247-024-01490-9

Diouf M. M. N., Lefevre, F., Hauchecorne, A., and Bertaux, J.L., (2024) Three-Dimensional Modeling of the $O_2(^1\Delta)$ Dayglow: Dependence on Ozone and Temperatures, *J. Geophys. Res.*, https://doi.org/10.1029/2023JD040159

Eckermann, S. D., Ma, J., Hoppel, K.W., Kuhl, D. D., Allen, D. R., Doyle, J. A., et al. (2018). High-altitude (0–100 km) global atmospheric reanalysis system: Description and application to the 2014 austral winter of the deep propagating gravity wave experiment (DEEPWAVE). Monthly Weather Review, 146(8), 2639–2666. https://doi.org/10.1175/MWR-D-17-0386.1

Emmert, J. T., Jones, M. Jr., Siskind D. E., Drob D. P., Picone, J. M., Stevens M. H., Bailey S. M., Bender, S., Bernath, P. F., Funke, B., Hervig, M. E., and Perot, K. (2022) NRLMSIS 2.1: An empirical model of NO incorporated into MSIS, *J. Geophys. Res.*, *127*, e2022JA030896. https://doi.org/10.1029/2022JA030896

Fletcher, L., et al. "An Observational Overview of Solar Flares", Space Science Reviews, vol. 159, no. 1–4, pp. 19–106, 2011. doi:10.1007/s11214-010-9701-8.

Funke, B., Lopez-Puertas, M., Gil-Lopez, S., von Clarmann, T., Stiller, G. P., Fischer, H., Kellman, S., (2005) Downward transport of upper atmospheric NOx into the polar stratosphere and lower mesosphere during the Antarctic 2003 and Arctic 2002/2003 winters, *J. Geophys Res.* ,D240308, 0148-0227/05/2005JD006463

Funke, B., Lopez-Puertas M., Stiller, G. P. and von Clarmann, T., (2014) Mesospheric and stratospheric NOy produced by energetic particle precipitation during 2002–2012, *J. Geophys. Res*, 119, 4429, 10.1002/2013JD021404

Funke B., Ball, W., Bender, S., Gardini, A., Harvey, V. L., Lambert, A., et al. (2017) HEPPA-II model–measurement intercomparison project: EPP indirect effects during the dynamically perturbed NH winter 2008–2009, *Atmos. Chem. Phys*, 17, 3573-3604, doi:10.5194/acp-17-3573-2017.

Garcia, H. A., "Temperature and Emission Measure from Goes Soft X-Ray Measurements", Solar Physics, vol. 154, no. 2, pp. 275–308, 1994. doi:10.1007/BF00681100.

Garcia-Sage, K, Farrish A.O, Airapetian, V. S., (2023), Star-exoplanet interactions, A growing interdisciplinary field in heliophysics, *Frontiers in Astronomy and Space Science*, vol. 10, doi: 10.3389/fspas.2023.1064076

Gehrels, N., Laird, C. M. Jackman, C. H., Cannizzo, J. K., Mattson, B., J., and Chen, W., (2003) Ozone depletion from nearby supernovae, *The Astrophysical Journal.*, 585:1169-1176.

Gelaro, R. et al. (2017), The modern-era retrospective analysis for research and applications, version 2 (MERRA-2), J. Clim., 30(14), 5419–5454, doi:10.1175/JCLI-D-16-0758.1.

Jones, M., D. P. Drob, D. E. Siskind, J. P. McCormack, A. Maute, S. E. McDonald, and K. F. Dymond (2018), Evaluating Different Techniques for Constraining Lower Atmospheric Variability in an Upper Atmosphere General Circulation Model: A Case Study During the 2010 Sudden Stratospheric Warming, J. Adv. Model. Earth Syst., 10(12), 3076–3102, doi:10.1029/2018MS001440.

Jones, M. Jr., Goncharenko, L. P., McDonald, S. E., Zawdie, K. A., Tate, J., Gasperini, F., et al. (2023). Understanding nighttime ionospheric depletions associated with sudden stratospheric warmings in the American sector. Journal of Geophysical Research: Space Physics, 128, e2022JA031236. https://doi. org/10.1029/2022JA031236.

Jones, M. Jr., Siskind, D. E., Drob, D. P., McCormack, J. P., Emmert, J. T., Dhadly, M. S., et al. (2020). Coupling from the middle atmosphere to the exobase: Dynamical disturbance effects on light chemical species. Journal of Geophysical Research: Space Physics, 125, e2020JA028331. https://doi.org/ 10.1029/2020JA028331

Kahler, S. W and Ling, A. G, (2023) Solar Stellar Connection: X-ray flares to energetic (> 10 MeV) particle events, *Astrophy. J.*, 956, https://doi.org/10.3847/1538-4357/acf1ff

McCormack, J., Hoppel, K., Kuhl, D., deWit, R., Stober, G., Espy, P., et al. (2017). Comparison of mesospheric winds from a high-altitude meteorological analysis system and meteor radar observations during the boreal winters of 2009–2010 and 2012–2013. Journal of Atmospheric and Solar-Terrestrial Physics, 154, 132–166. https://doi.org/10.1016/j.jastp.2016.12.007

Melott, A. L., Thomas, B. C., Kahcelriess, M., Semikoz, D. V., and Overholt A. C., (2017) A supernova at 50 pc: Effects on the Earth's atmosphere and biota, Astrophys. J., 840, https://doi.org/10.3847/1538-4357/aa6c57

Nedoluha G. N., Gomez, R. N, Boyd, I., Neal H., Parrish A., Connor B., Mooney T., Siskind, D. E., Sagawa, H., and Santee, M., (2020) Initial Results and Diurnal Variations Measured by a New Microwave Stratospheric ClO Instrument at Mauna Kea, *Journal Geophysical Research.*, 125, https://doi.org/10.1029/2020JD033097

Neupert, W. M., Gates, W., Swartz, M., & Young, R. (1967). Observation of the solar flare X-ray emission line spectrum of iron from 1.3 to 20 A. *The Astrophysical Journal*, *149*, L79–L83. https://doi.org/10.1086/180061

Pettit, J. M., Randall, C. E., Peck, E. D., Harvey, V. L., (2021) A new MEPED-based precipitating electron data set, *J. Geophys. Res.*, https://doi.org/10.1029/2021JA02966

Price, J. M., Mack, J. A., Rogaski, C. A. and Wodtke, A. M. (1993) Vibrational-state-specific self-relaxation rate constant. Measurements of highly vibrationally excited O2 (v=19-28), *Chemical Physics., 175(1)*,83-98, https://doi.org/10.10.1016/0301-0104(93)80230-7.

Qian L., Wang, W., Burns, A. G., Chamberlin, P. Coster, A., Zhang S-R., Solomon, S., (2019) Solar Flare and Geomagnetic Storm Effects on the Thermosphere and Ionosphere During 6–11 September 2017, *J. Geophys. Res.*, 124, 1298, https://doi.org/10.1029/2018JA026175

Randall, C. E., Harvey, V. L., Singleton, C. S. et al., (2006) Enhanced NOx in 2006 linked to strong upper stratospheric Arctic vortex, *Geophys. Res. Lett.*, 33, L18811, doi:10.1029/2006GL027160.

Randall, C. E., Harvey, V. L., Singleton, C. S. et al., (2007) Energetic particle precipitation effects on the Southern Hemisphere stratosphere in 1992-2005 *J. Geophys. Res.*, 112., D08308, doi:10.1029/2006JD007696,

Redmon, R. J., Seaton, D. B., Steenburgh, J. R., He, J., & Rodriguez, J. V. (2018). September 2017's geoeffective space weather and impacts on Caribbean radio communications during hurricane response. *Space Weather*, 16(9), 1190–1201. https://doi.org/10.1029/2018SW001897

Reep, J. W., and Airapetian, V., (2023) Understanding the duration of solar and stellar flares at various wavelengths, *Astrophys. J.*, vol. 98, DOI 10.3847/1538-4357/acf45a

Reep, J. W., Bradshaw, S. J., Crump, N. A., and Warren, H. P., "Efficient Calculation of Non-local Thermodynamic Equilibrium Effects in Multithreaded Hydrodynamic Simulations of Solar Flares", The Astrophysical Journal, vol. 871, no. 1, IOP, 2019. doi:10.3847/1538-4357/aaf580.

Reep, J. W., Siskind, D. E., and Warren, H. P., "Solar Flare Irradiance: Observations and Physical Modeling", The Astrophysical Journal, vol. 927, no. 1, IOP, 2022. doi:10.3847/1538-4357/ac4784.

Reep, J. W., Warren, H. P., Moore, C. S., Suarez, C., and Hayes, L. A., "Simulating Solar Flare Irradiance with Multithreaded Models of Flare Arcades", The Astrophysical Journal, vol. 895, no. 1, IOP, 2020. doi:10.3847/1538-4357/ab89a0.

Robinson, S. A, Revell, L. E., Mackenzie, R., Ossala, (2024) Extended ozone depletion and reduced snow and ice cover- Consequences for Antarctic biota, Global Change Biology, https://doi.org/10.1111/gcb.17283

Roble, R. G., and Ridley, E. C. (1994) A thermosphere-ionosphere-mesosphere-electrodynamics-General Circulation Model (TIME-GCM): Equinox solar cycle minimum simulations (30-500 km), *Geophys. Res. Lett.* 417-420.

Rodgers, E. M., Bailey, S. M., Warren, H. P., Woods, T. N., and Eparvier, F. G., (2010), Nitric oxide density enhancements due to solar flares, *Adv. Sp. Res.*, 45,28-38.

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

Annales Geophysicae Discussions

Seppala, A., Randall, C. E., Clilverd, M. A., Rozanov, E., and Rodger, C. J., (2009) Geomagnetic activity and polar surface air temperature variability, *J. Geophys. Res*., 114, https://doi.org/10.1029/2008JA014029

Siskind, D. E., Bacmeister, J. T., Summers M. E., and Russell J.M. III, (1997) Two-dimensional model calculations of nitric oxide transport in the middle atmosphere and comparison with Halogen Occultation Experiment data, Journal of Geophysical Research, 102/D3, 3527-3545.f

Siskind, D. E., Nedoluha, G. N, Randall, C. E., Fromm, M. and Russell, J. M. III (2000) An assessment of Southern Hemisphere stratospheric NOx enhancements due to transport from the upper atmosphere, *Geophys. Res. Lett*., 329-332.

Siskind, D. E., Stevens, M. H., Englert, C. R. and Mlynczak M. G., (2013) Comparison of a photochemical model with observations of mesospheric hydroxyl and ozone, *J. Geophys. Res*., 118, 195–207, doi:10.1029/2012JD017971

Siskind, D. E., Jones, M., Jr., Reep., J. W., Drob, D. P. Samaddar, S., Bailey, S. M., Zhang, S-R., (2022) Tests of a new solar flare model against D and E region ionospheric data., *Sp. Wea.*, 2, e2021SW003012, https://doi.org/10.1029/2021SW003012.

Slanger, T. G., L.E. Jusinski, G. Black and G. E. Gadd (1988), Vibrationally excited O_2, *Science*, 241, 945.

Solomon, S. C., and Qian L., (2005) Solar extreme-ultraviolet irradiance for general circulation models, *Journal of Geophysical Research*, 110, A10306, https://doi.org/10.1029/2005JA011160.

Solomon S., Roble, R. G., and Crutzen, P. J., (1982) Photochemical Coupling Between the Thermosphere and the Lower Atmosphere, 1. Odd nitrogen from 50 to 120 km, *J. Geophys. Res*., 87, C9, 7206-7220.

Stauffer, D. R., & Seaman, N. L. (1990). Use of four-dimensional data assimilation in a limited-area mesoscale model. Part I: Experiments with synoptic-scale data. Monthly Weather Review, 118(6), 1250–1277. https://doi.org/10.1175/1520-0493(1990)118⟨1250:UOFDDA⟩2.0.CO;2

Stauffer, D. R., & Seaman, N. L. (1994). Multiscale four-dimensional data assimilation. Journal of Applied Meteorology, 33(3), 416–434. https://doi.org/10.1175/15200450(1994)033⟨0416:MFDDA⟩2.0.CO;2

Thomas B.C., Jackman, C. H., Melott, A.L. (2007) Modeling atmospheric effects of the September 1859 solar flare, *Geophys. Res. Lett*., 34, L06810, doi:10.1029/2006GL029174.

Vink, J., "Supernova remnants: the X-ray perspective", Astronomy and Astrophysics Review, vol. 20, Springer, 2012. doi:10.1007/s00159-011-0049-1.

Zhang, X., Forbes, J. M., & Hagan, M. E. (2010a). Longitudinal variation of tides in theMLT region: 1. Tides driven by tropospheric net radiative heating. Journal of Geophysical Research, 115, A06316. https://doi.org/10.1029/2009JA014897

https://doi.org/10.5194/angeo-2024-15
Preprint. Discussion started: 11 September 2024

Annales Geophysicae Discussions

Zhang, X., Forbes, J. M., & Hagan, M. E. (2010b). Longitudinal variation of tides in the MLT region: 2. Relative effects of solar radiative and latent heating. Journal of Geophysical Research, 115, A06317. https://doi.org/10.1029/2009JA014898