On the set of deterministic phenomena preceding the earthquake June 25, 2021 with a magnitude of 5.4 near the city of Yayladere (Turkey)

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Abstract. The article examines the 5.4 magnitude earthquake that occurred on June 25, 2021 in the vicinity of Yayladere (Turkey). The analysis of the geomagnetic measurements carried out revealed a set of deterministic processes that preceded this event and are interpreted as its precursors. An estimate is made of the average time between the interval of existence of such phenomena and the moment of the earthquake under consideration. As a result, close values of the corresponding averaged time intervals were obtained for the statistics of all three components of the geomagnetic field considered in the article. The proposed technique can be used to predict seismic processes in various regions of the world in a near real-time mode.

1 Introduction

During the last years a number of scientists published the results of researches, indicative of possibility of registration of harbingers of strong earthquakes in the distance of more than 5000 km, and in some cases more than 10 000 km (Hasanov & Keramova 2006; Lyubushin 2008; Sobolev et al. 2008; Khain and Khalilov 2008).

Philosophy of short-term forecasting of earthquakes hasn’t undergone essential changes during the whole history of its presence. The basis of all technologies of short-term forecasting the earthquakes is to create the network of stations, which register the changes of geophysical, geochemical, hydro-geological and other parameters of geological medium before strong earthquakes near potential sources of possible earthquakes.

This article discusses deterministic processes that were detected in the last few hours before an event with a magnitude of M = 5.4 with an epicenter of 39.196° N, 40.165° E, which occurred at 18:28:37 (UTC) on 25.06.2021, 8 kilometers from the city of Yayladere (Turkey). The object of the study was the statistics of magnetic field measurements carried out at the test site of the Radio Astronomy and Geodynamics Department of the Crimean Astrophysical Observatory in Katsiveli (Crimea) during the period 00:00:00 on 24.06.2021 – 23:59:59 on 25.06.2021. In order to search for precursors of an impending earthquake, the technique proposed in (Volvach et al. 2022a,b,c,d; Volvach et al. 2023; Kogan et al. 2021; Kogan 2015)¹ and based on the
change in the properties of the probability density of any random process when even a very small term appears in its composition, the nature of which is weakly related to a set of commonly occurring phenomena, was applied. As will be shown below, the proposed approach made it possible to detect a number of phenomena that indicate a high probability of an imminent seismic event of significant magnitude. Thus, statistical methods are applicable not only within the framework of the Gutenberg-Richter theory (Gutenberg & Richter 1956; Amitrano 2012; Sanchez & Vega-Jorquera 2018), which makes it possible to assess the possibility of earthquakes of a given magnitude on medium-term, on the order of months and years, of time intervals, but also in relation to small, up to several hours, intervals before the onset of impending seismic events, which is close to the current time regime.

2 Mathematical apparatus used in solving the objective

The article introduces the assumption used in (Volvach et al. 2022a,b,c; Kogan et al. 2021; Kogan 2015) that any physical fields $x(t)$ measured in seismically active regions can be written in the form

$$x(t) = x_1(t) + x_2(t).$$  \hspace{1cm} (1)

Here $x_1(t)$ is the background noise associated with the influence of a set of ordinary phenomena and processes. Whereas the term $x_2(t)$ is solely due to the impact of an impending seismic event. At the same time, the hypothesis about the weak statistical dependence of these two random variables (RV) is introduced into consideration.

As indicated in the Introduction, below we will consider the properties of magnetic field statistics. In this case, the entire measurement period is divided into implementation segments of one minute duration, and each such segment is assigned 60 measuring counts (one per second). As in (Volvach et al. 2022), to each implementation segment is matched value of the statistical functional

$$L(n) = \frac{A}{M} \sum_{l=n-(M-1)}^{n} L_l, \quad L_l = \sum_{m=0}^{N-1} (-1)^m P_{m,l}. \hspace{1cm} (2)$$

Here, $M = 100$, the factor $A = 1000$ is necessary to obtain a convenient for analysis range of values $L(n)$, $l$ is the number of the implementation segment, the value $n$ corresponds to the moment of observation, that is, the completion time of the implementation segment with this number, and $P_{m,l}$ is the probability of the measured value of the function $f[x(t)]$, corresponding to the indicated segment number $l$, falls into the $m$-th interval of the range of values of the form

$$f[x(t)]_{\min} + mh \leq f[x(t)] < f[x(t)]_{\min} + (m + 1)h, \quad 0 \leq m \leq N - 1 \hspace{1cm} (3)$$

In (3) $f[X] = \sin(X)$ (sec$^2$) and $f[x(t)]_{\min} \geq -1$ is the smallest value $f[x(t)]$ for a given implementation segment. (The $x(t)$ values were measured and accepted in calculations in nT units.) In addition, in (3) the sampling interval is set equal to $h = 0.1$, therefore, in (2) and (3) $N = 20$. According to (Volvach et al. 2022a,b,c; Kogan et al. 2021; Kogan 2015), functional (2) essentially depends on the level of entropy of the values of the random process, the value of which significantly increases when an independent stochastic term $x_2(t)$ appears in (1).
3 Data and analysis of the properties of the functional $L(n)$ constructed from the measurements of the magnetic field

The International Real-time Magnetic Observatory Network - the global network of observatories, monitoring the Earth's magnetic field and the data are available as digital data files (www.intermagnet.org, ftp.seismo.nrcan.gc.ca).

Fig. 1 shows the dependence $L(n)$, constructed from measurements of the H-component of the magnetic field on June 24 and 25, 2021, with a magnetometer at the above-mentioned test site of the CrAO in Katsiveli. Hereinafter, the horizontal axis corresponds to UTC time in minutes starting from $n = M = 100$ min after the first moment of the first day of the considered 48-hour time interval (see (2)). In all figures, the solid vertical red straight line (drawn at $n = 2459$ min) denotes the moment of the earthquake under consideration. Let us introduce a number of definitions, illustrating them with examples from Fig. 1.

1. We call the main extrema the set of points of maxima and minima of the dependence $L(n)$, which on the intervals of both increasing and decreasing corresponds to the variation $\delta L$ of the values of this functional, satisfying the condition

$$\delta L \geq 0.2 \Delta L_{\text{max}}. \quad (4)$$

In (4), $\Delta L_{\text{max}}$ is the difference between the largest maximum and the smallest minimum of functional (2) from the start to the moment of the considered earthquake. At the same time, an additional condition is introduced that any point considered as the main extremum must be horizontally removed from any other point of the $L(n)$ curve by at least 35 minutes. Here and below, unless the opposite condition is set, small-scale variations in the values of $L(n)$ are not taken into account. This term will denote fluctuations that are much smaller in modulus than the right side of (4) and, at the same time, are small in amplitude and duration compared to the variation of the functional $L(n)$ in the region of the corresponding extremum. Thus, local trends are rather large quasi-rectilinear segments of the $L(n)$ dependence, the boundary points of which are extrema, horizontally separated by a given minimum distance from any other points of the curve $L(n)$. In addition, within the limits of each such section, there should be a quasi-constant rate of change of this dependence. In Fig. 1, the main extrema, in particular, are the points $h, a, b, c, d, f, e, g, p$.

2. We will call the sections of the dependence $L(n)$, which have the properties of quasi-monotonicity and quasi-constant rate of change (also neglecting small-scale variations), as local trends. We assume that they correspond to the $\delta L$ variation satisfying (4). In this case, the geometric deviation of the points of the local trend from the segment of the straight line connecting the points of its beginning and end should not exceed 15% of the modulus of the difference between the values of the dependence $L(n)$ at these points. We emphasize that for the main extremums defined above, the corresponding intervals of increase and decrease can also be considered as local trends. Examples of local trends in Fig. 1 are, for example, segments $a - b, i - j$ and $e - u$ of a given curve.

3. Let us denote as global extrema those points of maxima and minima of the dependence $L(n)$, which correspond to the largest or, correspondingly, the smallest values of it over an interval of at least 24 hours. Moreover, they should be located near the middle of such a time interval. Such a point is, for example, point $q$, see Fig. 1.

4. Let’s call the set of the main and global extrema, as well as the boundary points of trends, the guide points.
5. We will call a channel a collection of two straight lines, each of which is drawn through two guide points, provided that the difference \( \Delta \alpha \) of the directions of these straight lines satisfies the constraint

\[
\Delta \alpha \leq 1.5^\circ.
\]  

(5)

Here and below, we assume that the angle \( \Delta \alpha \) corresponds to the condition of equality of the geometric dimensions of the length units of the vertical and horizontal axes. At the same time, in this article, in all the figures, the unit of the abscissa axis is slightly greater than the unit of the ordinate axis. The angle \( \Delta \bar{\alpha} \) corresponding to these figures between the directions of the "shores" of any of the channels is related to \( \Delta \alpha \) by the ratio

\[
\Delta \alpha \leq 0.8 \Delta \bar{\alpha}.
\]  

(6)

The duration of the channel's existence, that is, the time interval between the guide points through which its boundaries pass, must be at least 150 minutes.

6. We will call a sliding boundary (SB) any straight line that is drawn through two guide points and is tested by the curve \( L(n) \) at two more points, one of which is also a guide (we will call it additional). The term "testing" means either a simple intersection of the curve (2) and a given straight line, or its passage from the guiding point at such a close distance that the deviation coefficient \( \delta \) satisfies the condition

\[
\delta \leq 1.6\%.
\]  

(7)

where \( \delta = \frac{\Delta S}{L_g(n)} \cdot 100\% \). Here \( \Delta S \) is the minimum vertical distance on the plane \( \{n, L(n)\} \) of the Cartesian variable between the straight line under consideration (both the SB and the channel boundary) and the given guide point, and \( L_g(n) \) is the value of the functional \( L(n) \) in this point.
An additional point can be located either between the two guides through which the SB line passes, or after them, but until the fourth test. In Fig. 1, the sliding border is straight line 1 drawn through the global extremum q and the guide point j. Point m is additional. Her area is highlighted with an orange ellipse. In this case, the coefficient $\delta \approx 0.95\%$ satisfies (7).

7. We also define that if in the region of small-scale fluctuations there are more than three approximations of the curve $L(n)$ to the corresponding straight line when fulfilling (5), then any such segment of this curve will be assumed to correspond to one test.

Note that in Fig. 1, the variation value for the $w - t$ segment of the $L(n)$ curve is less than the right-hand side of (4) and therefore does not satisfy the definition of the local trend. It is also obvious that w is not the main or global extremum. In addition, at this point and its vicinity, the intersection of the curve under consideration and straight 1 does not occur. Therefore, despite the fulfillment of condition (7) for w (as applied to straight 1), the point w is not included in the number of testing points.

Thus, straight line 1 is both the SB and the bank of the channel at the same time. Hereinafter, those channel boundaries that are not sliding boundaries will be denoted by dashed straight lines, and the SB themselves will be denoted by solid straight lines.

With regard to channel boundaries, an additional point, if any, can also be located either between those two guide points through which this boundary passes, or after them. In this case, for one channel boundary, there can be no more than three consecutive tests. If this rule is violated, any segment of the corresponding straight line containing the indicated three sequentially tested points is considered the channel boundary. (These three points include also those two guide points through which the given channel boundary is drawn.)

We now turn directly to the search for the precursors of the earthquake under consideration. In Fig. 1, three vertical dashed straight lines in the right part of it are drawn from the points of the fifth testing of the curve $L(n)$ of the bank lines of channels 1–2, 3–4, and 5–6. (Here and below, similar vertical dashed straight lines mark the moments of time corresponding to the appearance of the "graphic precursors" considered below). In this case, the lower boundary of the first of these channels is assumed to be the segment $q - j$ of straight line 1 containing an additional point $m$. The fact of the existence of such channels can be interpreted as the emergence of a set of deterministic processes immediately before an earthquake.

The time intervals from the indicated points of the fifth testing to the moment of the earthquake are $T_{H,1} = 296 \text{ min}$, $T_{H,2} = 278 \text{ min}$ and $T_{H,3} = 197 \text{ min}$. As will be shown below, the phenomena of the fifth testing of the channel boundaries precede the considered seismic event and when carrying out a similar statistical analysis for the E- and Z- components of the measured geomagnetic field. Therefore, we will consider the moment of the indicated fifth testing as the time of realization of one of the types of "graphic precursors" of an impending earthquake.

In addition, as will be confirmed below, in the study of all three components of the magnetic field, the point of the fourth testing of the sliding boundary from the side of the curve $L(n)$ is also a recurring precursor of an earthquake. This effect takes place in Fig. 1 at point $v$ in relation to sliding border in the form of straight line 1. (At this point, the indicated fourth testing
of this SB takes place, also marked with a vertical dotted straight line, see Fig. 1.) The appearance of SB can also be called the emergence of a deterministic phenomenon a few hours before the impending earthquake.

This condition, as well as all other restrictions specified in the article (see, in particular, (4) - (8)) are introduced in order to minimize the number of linear structures under consideration in the range from several units to about eight to ten. In this case, the conditions for analyzing the degree of concentration of the considered precursors over limited time intervals are significantly simplified. As follows from the results of this article, the beginning of the considered earthquake is preceded by a sharp increase in the concentration of these precursors over the last few hours.

Further, when analyzing the process of “final preparation” of a seismic event, as well as when determining the average duration of the interval from the moment of registration of a precursor to the onset of an earthquake, we will take into account only those precursors that are separated from the time of its onset by no more than an interval

\[ T_{\text{max}} = 720 \text{ min} \]  \hspace{1cm} (8)

Therefore, in this case, the time interval from point \( \nu \) in Fig. 1 and two similar intervals from the first two precursors on Fig. 3 until the moment of the seismic event is not taken into account.

In Fig. 2 shows the dependence \( L(n) \) for the E-component of the magnetic field (data from the magnetometer in Katsiveli).

Here, channels 1–2, 3–4 and 5–6 are drawn, for which the angles \( \Delta \alpha < \Delta \bar{\alpha} \ll 1^\circ \) (see (5) and (6)). Segments \( a - b, i - j, c - d, e - f, m - l \) and \( l - h \) (their boundaries contain all but \( q \), the guide points through which all the indicated straight are drawn) satisfy the definition of local trends. For points \( q \) and \( e \), the definitions of the main and, accordingly, global extremum are fulfilled.

In the region of small-scale fluctuations, marked with an orange ellipse, there are more than three approximations of the curve \( L(n) \) to straight line 5 when (5) is fulfilled. According to definition VII introduced above, we consider that such a segment of this curve corresponds to one test. Taking this into account, we find that the specified straight is a sliding boundary, the fourth testing of which takes place at point \( g \), see Fig. 2.

By analogy with Fig. 1, we assume that the points of the fifth testing of the channel boundaries of the curve (2) or the fourth testing of the line of the specified SB correspond to the time of realization of the "graphic precursors" of an impending seismic event. Therefore, also by analogy with Fig. 1, we postpone the intervals \( T_{E,1} = 434 \text{ min}, T_{E,2} = 251 \text{ min}, T_{E,3} = 231 \text{ min}, \) and \( T_{E,4} = 181 \text{ min} \) from the moments of the appearance of such precursors to the time of the beginning of this earthquake.

As in the previous case, in Fig. 2 these intervals correspond to green horizontal lines.

In Fig. 3 shows the dependence \( L(n) \) for the Z-component of the magnetic field (data from the magnetometer in Katsiveli).

Here, channels 1–2 are drawn (here \( \Delta \bar{\alpha} \approx 0.8^\circ \), as well as 2–3, 4–5 and 8–9, for which the angles \( \Delta \bar{\alpha} \ll 1^\circ \) (see (5) and (6)).

Note that straights 1, 2 and 3 are almost parallel. Sliding boundaries 4, 6 and 7 are drawn through the first and third (counting from the left) test points; areas of additional points are marked with ellipses. Here, the guide point \( a \), through which this line passes, is the beginning of the local trend \( a - b \). For SB 7, the region of small fluctuations within the corresponding ellipse is associated with one test. Taking into account (8) for SB 6 and 7, the intervals from the time of the fourth testing to the moment
of the earthquake are not considered. Intervals $T_{Z,1} = 596 \, \text{min}$, $T_{Z,2} = 588 \, \text{min}$, $T_{Z,3} = 396 \, \text{min}$, $T_{Z,4} = 350 \, \text{min}$ and $T_{Z,5} = 144 \, \text{min}$ depicted by green horizontal lines.

For ease of comparison in Table 1 shows all the obtained values of the intervals $T_H$, $T_E$ and $T_Z$ (first 5 lines), as well as their average values of the form $\langle \ldots \rangle_{720}$ and $\langle \ldots \rangle_{360}$ over the last 720 and, respectively, 360 minutes before the earthquake.

4 Discussion of the results

The results presented in this paper rely on data collected at magnetic observatories. We used the INTERMAGNET stations - the global network of observatories, monitoring the Earth's magnetic field (www.intermagnet.org).

As a result of analyzing the statistics of magnetic field measurements carried out in the last two days before the earthquake, which occurred at 18:28:37 (UTC) on June 25, 2021, 8 kilometers from the city of Yayladere (Turkey), we come to the following conclusions.

External forcing influences geomagnetic data. The analysis on the geomagnetic observations consider the solar forcing. In Fig. 4 shows the K-index of solar radiation changes during the period 00:00:00 on 24.06.2021 – 23:59:59 on 26.06.2021. As can be seen from the chart, the K-index was in the quiet (green) zone.

1. In the last hours before a given seismic event, there is an effect of concentration of recurring deterministic phenomena identified in the previous section of the article and interpreted as a kind of "graphic precursors" of an impending seismic event. In this case, for all three components of the geomagnetic field, a kind of "ladder of precursors" appears, see Fig. 1 – 3.

2. The most common precursor is the phenomenon of fivefold dependency testing (2) of the boundaries of the channels defined above. These channels are formed by a pair of almost parallel straight lines (see (5)), which limit the values of the $L(n)$ curve for at least 150 minutes.

3. Sliding boundaries in the form of straight lines, each of which is tested at least four times by the dependence in question, is also an essential type of "graphical precursor". In this case, the first three tests of this kind must correspond to topologically selected points of the curve (2).

4. As follows from the analysis of the data in Table 1, in the last 6 hours before the event, the values of the intervals $T_H$, $T_E$ and $T_Z$ are comparable. In particular, this manifests itself in the relative closeness of the average values $\langle T_H \rangle_{360}$, $\langle T_E \rangle_{360}$ and $\langle T_Z \rangle_{360}$, which correspond to averaging over the last 360 minutes before the onset of this earthquake. As the authors believe, this fact is an argument in favor of the reliability of the proposed method.

5. The discovered effects of the existence of deterministic phenomena in the form of channels and sliding boundaries are quite reliable (outside the scope of this work, they were tested on a large array of data on various earthquakes and corresponding measurements of fields of a very different nature), but at the same time they are difficult to explain. It is possible that the existence of such straight linear boundaries of the values of functional (2) is associated with an almost constant (on a 48-hour scale) speed of movement of lithospheric plates; this assumption is completely hypothetical. In addition, significant difficulties in explaining the discussed phenomena can be associated with poor theoretical knowledge of the area of probability theory.


corresponding to the phenomena under consideration, which can be attributed to "anti-Gaussian" random processes associated with the properties of the sum of a large number of strongly dependent random variables.

6. The study of the influence of cracks on various physical properties of rocks is one of the main tasks of earthquake prediction. The conductivity of natural media is due to the transfer of electric charges by the through current of electrons, ions, holes. In the epicenter of an earthquake, in addition to an increase in the number and length of elementary cracks, their closure and the appearance of a main crack, there is also a change in the electrical resistance of rocks. When the soil level and density change, the specific electrical conductivity of rocks changes by several orders of magnitude compared to the initial value, which leads to a change in the characteristics of the magnetic field.

With the help of the INTERMAGNET international network, the Earth's magnetic field is monitored, which makes it possible to create a technique that can be used to predict seismic processes in various regions of the world in close to real time.

7. Note that Volvach et al. 2022a,b,c,d; Volvach et al. 2023a, using an approach similar to that used in this article, precursors for 37 earthquakes are considered. For 26 of them (with the exception of aftershocks that are not far removed in time from the previous event), almost all of the considered precursors are recorded before their “own” earthquake at a time interval of no more than 20–25% of the total considered measurement period $\Delta t_{\text{glob}}$. Outside a given period of time, the number of such phenomena is much less than within it. Moreover, in the absolute majority of the cases under consideration, this interval is on the order of 5–10% or less of $\Delta t_{\text{glob}}$. The same effect also takes place in relation to the present article, which indicates the stability of the results of the method used. This fact means a high probability that the statistical effects under consideration are related precisely to the process of the “final preparation” of the approaching earthquake.

8. As follows from this article, as well as Volvach et al. 2022a,b; Volvach et al. 2023a, if at least 3 considered precursors occur within an interval of 300 minutes or more, according to the measurement statistics of at least one of the magnetic field components, the probability that an earthquake will occur within an interval of about an hour after the registration of the last specified precursor is very high. We emphasize that this conclusion is obviously preliminary, subject to verification in the course of further studies and relating to the case of a total duration of the analyzed measurement period of the order of two days.

9. Note that the relative closeness of the mean values of the considered time intervals corresponding to different components of the magnetic field (see Table 1) is an argument in favor of the stability of the applied approach. Which in turn makes it possible to speak about the objectivity of the results obtained.

We also point out that the different time of occurrence of precursors in the analysis of the statistics of different components of the magnetic field can be explained by different amplitudes of the random process $x_2(t)$ for each of the components of this field. Such spatial anisotropy can be associated, in particular, with the existence of distinguished directions for the development of cracking in the zone of preparation of a seismic event. The consequence of this difference in amplitudes is a different degree of averaging of fluctuations in the probability density of the background noise for different components of the magnetic field. This effect manifests itself both in the noncoinciding form of the $L(n)$ curves corresponding to different components of the magnetic field, and in different times and in the number of detected precursors. In this case, the indicated proximity of the
average values of the indicated time intervals corresponding to various components of the magnetic field can be explained by a synchronous change in the properties of the statistics of all components of the magnetic field, the moments of the beginning and/or completion of the next stage of the "final preparation" of the approaching event. Thus, such a difference in the time of registration of individual precursors under consideration for different field components, in combination with a small relative difference in the corresponding average values, is quite consistent with the logic of the calculations.

10. We also note that, according to the statistical meaning of the functional $L(n)$, its maxima and minima with high probability correspond to the smallest and, accordingly, the largest level of variations of the random process $x_2(t)$, independent or weakly dependent on background noise, as applied to a local sequence from $M$ segments of realization, see (2), which precede the given moment of time. The article introduces the hypothesis that such an independent or quasi-independent process is determined by a set of phenomena associated with the fracture of lithospheric plates in the area of preparation for an approaching earthquake. (It is taken into account that there are no significant geomagnetic disturbances associated with solar events during the period under study.) Therefore, if this hypothesis is true, the maxima and minima of the dependence $L(n)$ correspond to the minimum and, accordingly, the maximum values of the level of compression of lithospheric plates in the region of the future hypocenter.

Therefore, the appearance of a channel means the appearance of boundaries determined by deterministic or quasi-deterministic functions of time, within which the process of oscillations of the corresponding seismic pressure occurs. Similarly, the formation of a sliding boundary means the existence of such a limit, limiting from above or below the values of the specified pressure. It is in this sense that the term about the occurrence of deterministic phenomena used in this article should be considered.

We also note that, taking into account the definition and properties of the functional $L(n)$, the linear form of the boundaries of the indicated graphical objects, whether they be channels or sliding boundaries, can correspond to the hyperbolic form of quasi-deterministic time dependences that determine the boundaries of seismic pressure fluctuations during the “final preparation” period approaching earthquake. The validity of this assertion will be investigated in future works.

Taking into account the foregoing, let us additionally explain the physical meaning of the chosen criteria for registering the precursors of an approaching earthquake. The fifth test of the curve $L(n)$ of the boundaries of any of the channels means either a double “internal” oscillation within the boundaries of this channel (with the possibility of going beyond its limits at the point of the fifth test, see channel 5–6 in Fig. 1). This means a double approach of the seismic pressure level to some time-varying critical level, which is determined by a quasi-deterministic dependence, with the possibility of its “breakthrough” in the end. Or such a five-time test means from four to five consecutive “attempts to break through” (either only from above, or only from below) both boundaries of a given channel (see channel 1–2 in Fig. 1), also with the possibility of going beyond its limits during the last test. The physical meaning of such behavior of the $L(n)$ curve is reduced to five successive approaches of the seismic squeezing level to a certain critical level, which is also determined by a quasi-deterministic function of time. As follows from Volvach et al. 2022a, in both cases, such phenomena with a fairly high probability take place at the “final preparation” stage of the forthcoming earthquake.
The physical meaning of the criterion for registering precursors after the fourth test of the $L(n)$ curve of the sliding boundary line is similar. It comes down to identifying at least four consecutive facts of reaching a certain time-varying critical level, which is determined by a quasi-deterministic dependence, with the possibility of its “pushing through” as a result. Also according to Volvach et al. 2022a, such effects with a fairly high probability occur in a fairly short, on average, about several hours, time interval before the onset of the upcoming earthquake.

5 Conclusions

In the article, based on the analysis of the results of measurements of the geomagnetic field, the deterministic phenomena associated with the process of "final preparation" of the earthquake with a magnitude of 5.4, which occurred on June 25, 2021, near the Turkish city of Yayladere, were considered. Based on the results obtained, "graphic precursors" were identified in the form of repeating phenomena of the emergence of channels and sliding boundaries, the properties of which were investigated in this work. The authors consider the most promising direction of further research to determine, within the framework of the methodology proposed in the article, not only the time, but also the coordinates of the future epicenter, and, possibly, the magnitude of the expected earthquake.

Data availability statement

The data underlying this paper are available in the paper and through the electronic resources: www.intermagnet.org, ftp.seismo.nrcan.gc.ca.

Acknowledgments

The results presented in this paper rely on data collected at magnetic observatories. We thank the national institutes that support them and INTERMAGNET for promoting high standards of magnetic observatory practice (www.intermagnet.org).

References


Table 1. Values of intervals from the time of realization of «graphic precursors» to the moment of the beginning of the earthquake, as well as their averaging.

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<th>$I$</th>
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<th>$T_{E,i}, \text{ min}$</th>
<th>$T_{Z,i}, \text{ min}$</th>
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Figure 1. Dependence $L(n)$ for the H-component of the magnetic field from measurements at the Katsiveli.
Figure 2. Dependence $L(n)$ for the E-component of the magnetic field from measurements at the Katsiveli.
Figure 3. Dependence $L(n)$ for the $Z$-component of the magnetic field from measurements at the Katsiveli.
Figure 4. H, E and Z-component of the geomagnetic field according to the magnetometer SIM (curves 1, 2 and 3) and Kp index (curve 4).