Investigating the Role of Gravity Waves on Mesosphere-Lower-Thermosphere (MLT) Inversion Layers at Low Latitudes (3-15^0)

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Abstract

The Mesosphere and Lower Thermosphere (MLT) transitional region, encompassing a height range of 60-100 km, is a distinct and highly turbulent zone within the Earth's atmosphere. This region is particularly significant due to dynamic processes like gravity waves, which contribute to the formation of the Mesospheric Inversion Layer (MIL). Investigating these phenomena is crucial for understanding the dynamics of the middle and upper atmosphere, especially regarding stability and energy transfer. Inversion layers significantly influence the stability of the atmosphere in this region, playing a crucial role in atmospheric dynamics. Inversion layers are associated with energy transfer processes, vital for understanding the overall dynamics of the atmosphere. Despite extensive study on an inversion, the formation mechanisms of mesospheric inversions remain poorly understood. This article explores the upper and lower inversion phenomena and their causative mechanisms. It uses long-term SABER observations from 2005 to 2020 over the latitude range of 3-15° N and longitude range of 33-48° E. The results show that the upper inversion occurs more frequently, with a frequency below 40%, compared to the lower inversion, which occurs below 20%. The upper inversion occurs within the height range of 78-91 km, with an inversion amplitude of approximately 20-80 K and a thickness of around 3-12 km. In contrast, the lower inversion is confined to the height range of 70-80 km, with an inversion amplitude of about 10-60 K and a thickness of around 4-10 km. Moreover, the gravity wave indicator potential energy shows high energy (below 100 J/kg) in the upper mesosphere region (85-90 km) compared to the lower mesosphere region (70-75 km) with less than 50 J/kg. Considering gravity waves, the Brunt-Väisälä frequency (N^2) stability criteria indicate instability in the upper mesosphere region with very low values compared to the lower mesosphere region. This suggests that the high amount of gravity wave potential energy is a consequence of the higher instability in the upper inversion compared to the lower inversion.

Keywords: Mesosphere and Lower Thermosphere (MLT), Upper and Lower Inversions, Perturbed Temperature, Causative Gravity Waves, Potential Energy, Brunt-Väisälä Frequency, Atmospheric Instability.

Introduction

The Mesosphere and Lower Thermosphere (MLT) region serves as a transitional zone for wave processes from the lower and upper atmospheres, including tidal, planetary, and gravity waves. Gravity waves (GWs) originating from the lower atmosphere are known to propagate into the mesosphere. There, they break and dissipate their energy and momentum, affecting the thermal structure, global atmospheric circulation, and variability of the mesosphere. This process also influences the formation of mesospheric inversion layers (MILs), which are associated with increased temperature variability in the mesosphere. MILs indicate wave saturation when the lapse rate falls below the dry adiabatic lapse rate (Sica et al., 2007). Temperature inversions in the mesosphere have been widely observed and studied using various techniques, including lidar, radar, rocket sondes, and satellites, across different geographic locations. Sivakandan et al. (2014) utilized TIMED/SABER kinetic temperature data to examine the occurrence and characteristics of mesospheric inversions over the equatorial Indian region (0 to 10° N and 70 to 90° E) for the years 2002 and 2008. However, they did not explore the causative factors. This study aims to investigate the causes of these inversions, focusing specifically on the role of atmospheric gravity waves.

Gravity waves and mesospheric inversion layers (MILs) are interconnected phenomena within the Earth's atmosphere, particularly in the mesosphere and lower thermosphere. MILs are layers within the mesosphere where the temperature profile shows an inversion. This means the temperature increases with altitude, contrary to the typical decrease. These inversion layers often form due to dynamic processes, including the breaking and dissipation of gravity waves. As gravity waves propagate upwards, they can grow in amplitude because the atmospheric density decreases with altitude. When these waves reach a critical amplitude, they become unstable and break. This breaking process releases energy and momentum into the surrounding air, causing localized heating. The energy dissipation from breaking gravity waves causes localized heating. This heating can create or enhance mesospheric inversion layers by increasing the temperature at certain altitudes. The breaking of gravity waves can also generate turbulence, which further influences the structure and stability of inversion layers. This process also contributes to momentum and energy deposition.

The momentum and energy deposition of gravity waves (GWs) are believed to be key drivers of large-scale atmospheric circulation, the coupling of different atmospheric layers, and inversion phenomena (Fritts and Alexander, 2003; Lindzen, 1981; Smith, 2012). Additionally, researchers are exploring the influence of gravity wave-breaking on mesosphere dynamics to understand its impact on inversion phenomena, particularly in mid- and high-latitude regions (Gan et al., 2012; Walterscheid and Hickey, 2009; Collins et al., 2011; Szewczyk et al., 2013). Both observational and modeling studies have investigated GWs as a cause of these inversions (Fritts, 2018; Collins et al., 2014; Sridharan et al., 2008; Ramesh and Sridharan, 2012; Ramesh et al., 2013, 2014, 2017). Despite extensive exploration, our understanding of the impact of gravity waves on mesosphere inversions, particularly in terms of temperature variability, remains incomplete, especially in mid- and high-latitude regions (Singh and Pallamraju, 2018; Fritts et al., 2018). As a result, the study of inversion phenomena and their underlying causes remains a key area of focus in mesosphere dynamics research.

Research on the temporal and spatial variability of the mesosphere inversion phenomenon, particularly about atmospheric waves and gravity wave activity, is notably lacking in low latitudes. To address this gap, our study investigates the mesosphere inversion phenomenon and its association with gravity wave activity and stability criteria. We use Brunt-Vaisala frequency (N^2) over the low latitudinal band (3°–15° N) with long-term SABER observations from 2005 to 2020. The study is organized as follows: Section 2 details the data and methodology used to analyze the mesosphere inversion phenomenon and their causative gravity waves via the potential energy. Section 3 presents the results, and Section 4 concludes the findings.

2. Observation and Data analysis

2.1 SABER Observation

The TIMED/SABER satellite, launched on December 7, 2001, operates in an elliptical orbit at approximately 625 km altitude with a 74° inclination relative to the equator. Since its launch, SABER has been a crucial tool for atmospheric research, providing extensive data on the middle atmosphere. SABER is a limb-viewing radiometer working in the infrared region (1.27–17 microns) and can measure radiative emissions across a wide range of altitudes. It offers nearly global coverage and continuous 24-hour data over 60 days. The instrument completes 15 orbits daily, each taking about 97 minutes, and collects around 1400 data profiles per day, with each profile taking 58 seconds. SABER's high-resolution temperature profiles are essential for studying the dynamics and wave processes in the mesosphere. It provides temperature measurements with an accuracy of 1 to 2 K between 15 and 60 km. The accuracy decreases to 5 K below 85 km and increases to 6.7 K to 10 K near 100 km. This data has been instrumental in understanding the thermal structure and dynamical processes in the mesospheric region, as highlighted by various studies (Garcia et al., 2008; Gan et al., 2012, 2014; Bizuneh et al., 2022; Lingerew et al., 2023; Rezac et al., 2015; Meriwether and Gerrard, 2004; Fechine et al., 2008; Dou et al., 2009; France et al., 2015).

We utilized SABER vertical temperature profiles taken within the 60–100 km altitude range. These profiles cover the period from 2005 to 2020, spanning latitudes from 3°N to 15°N and longitudes from 33°E to 48°E. Figure 1 shows the monthly mean of SABER temperature data for the mesosphere and lower thermosphere. The data aim to illustrate the MLT temperature variability, which helps us identify the inversion layers (MIL). The monthly mean temperatures in the MLT region show a maximum of 200-240 K at altitudes of 60-70 km. Then it decreases to around 160-180 K at 95-100 km throughout the entire period. While the temperature patterns in the 70-90 km altitude range suggest an inversion, these inversions are not visible.

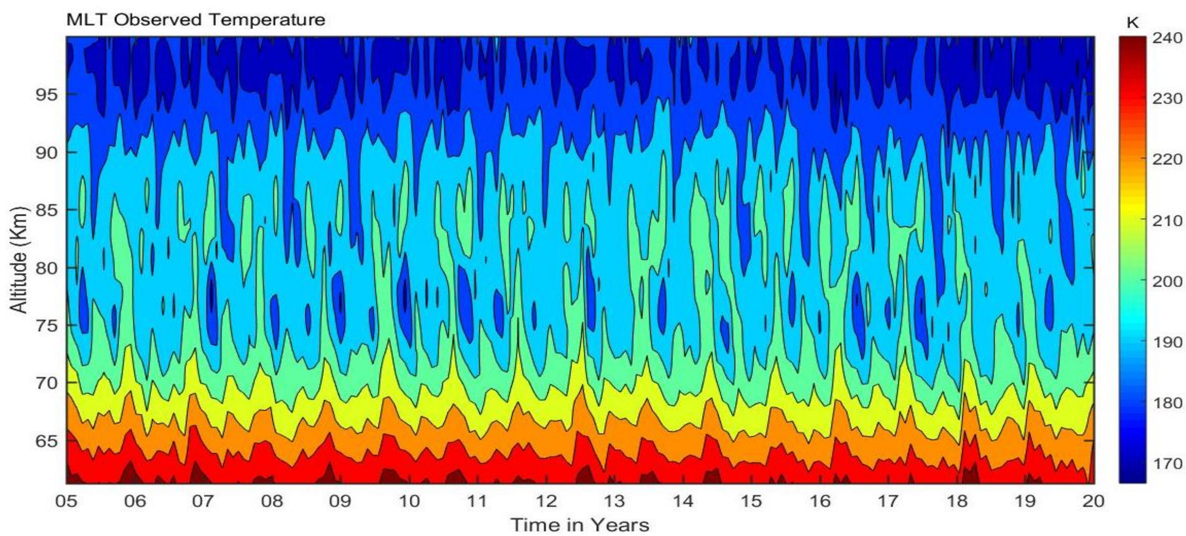

Figure 1. The monthly mean of MLT temperature variability in the height range of 60-100 km during 2005-2020 over the low-latitude.

2.2 Analysis Technique

The Earth's middle atmosphere typically has a negative temperature gradient, but some reports have shown positive temperature gradients in the mesosphere (Meriwether and Gardner, 2000; Gan et al., 2012). This phenomenon is called the “mesospheric inversion layer (MIL)”. MILs are identified using a procedure detailed by Leblanc and Hauchecorne (1997) and Fechine et al.(2008). Mesospheric inversions are characterized by their thickness (the altitude difference between maximum warming and cooling) and amplitude (the temperature difference between these points) (Meriwether and Gardner, 2000). Here are the identification criteria for these inversions:

1. The bottom level of the lower inversion is above 70 km, and the top level is below 80 km. For the upper inversion, the bottom level is above 80 km, and the top level is below 92 km.
2. The amplitude is considered larger than 5 K.
3. The thickness is greater than or equal to 3 km.

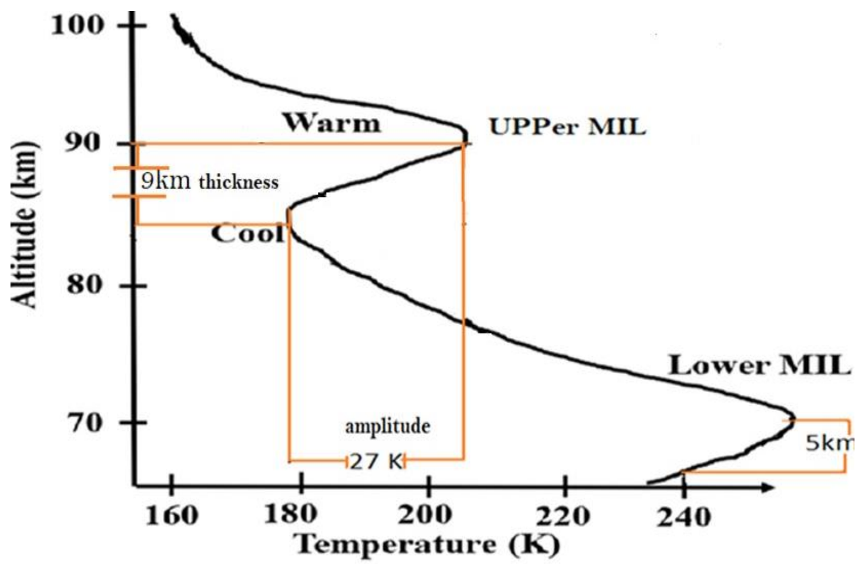

Figure 2. Schematic of upper and lower mesospheric inversion layers shown in the temperature profile for the MLT regions (Adapted from Meriwether and Gerrard, 2004).

Figure 2 illustrates this concept, showing a positive temperature difference between the top and bottom levels of the inversion. This method has been utilized in numerous previous studies investigating mesospheric inversions (Leblanc et al., 1998; Meriwether and Gardner, 2000; Duck et al., 2001; Duck and Greene, 2004; Cutler et al., 2001; Siva Kumar et al., 2001; Ratnam et al., 2003; Gan et al., 2012). The frequency occurrence rate of mesospheric inversion layers (MILs) is derived during the period 2005–2020 in the upper and lower MLT regions. This rate is calculated by dividing the number of inversion days in each month by the total number of days in that month over the 16-year observation period (2005–2020).

Mesospheric temperature inversions are related to instabilities in atmospheric dynamics. To identify the causative, short-period atmospheric gravity waves, a high-pass filter with a one-hour interval cutoff frequency is applied using the Brunt-Väisälä frequency (N2). Another important concept to estimate the Brunt-Vaisala frequency is the potential temperature (θ). It represents the air parcel's temperature when it is displaced adiabatically to a standard pressure level, p_0, from the current pressure level, p. This is based on the first law of thermodynamics.

$$\frac{dT}{T} = \frac{R}{c_p}\frac{dp}{p} \Rightarrow \quad \int_T^0 \frac{dT}{T} = \int_p^{p_0} \frac{R}{c_p}\frac{dp}{p} \qquad \text{(1) it yields}$$

$$\theta = T\left(\frac{p_0}{p}\right)^{R/c_p} \qquad (2)$$

Therefore, the motion of a vertical atmospheric air parcel can be described by (Liu, 2011; Vadas and Fritts, 2005) as follows in equation (2.3). This equation calculates the Brunt-Vaisala frequency of the parcel due to the buoyant and gravitational forces acting on it.

$$\frac{d^2s}{dt^2} = -g\,\frac{\rho-\rho_0}{\rho}\ sina \qquad (3)$$

Based on the hydrostatic equation, $\rho = \rho_0$, and $p = p_0 \Rightarrow \frac{\partial p}{\partial z} = \frac{\partial p_0}{\partial z} = -g\rho_0$ (4) and the ideal gas law, $\rho = p/RT = p_0/RT$ gives the parcels motion of an equation:

$$\frac{d^2s}{dt^2} = -\frac{g}{\rho}\left(\frac{d\rho}{dp}\frac{\partial p_0}{\partial z} - \frac{\partial \rho_0}{\partial z}\right) z \qquad (5)$$

Following the same approach using the hydrostatic equation (4) and adiabatic equation (6)

$$dln\rho = \frac{dlnp}{\gamma}\ , \gamma = c_p/c_v \qquad \text{(6) yields}$$

$$\frac{d^2s}{dt^2} = -\frac{g}{\rho}\left(\frac{\rho}{\gamma p_0}\frac{\partial p_0}{\partial z} - \frac{\partial \rho_0}{\partial z}\right) z = g\left(\frac{\partial ln\rho_0}{\partial z} - \frac{1}{\gamma}\frac{\partial lnp_0}{\partial z}\right) z \qquad (7)$$

For the ideal gas law of p = ρRT, the natural logarithm is taken for altitude, z on both sides, yielding

$$\frac{\partial ln\rho}{\partial z} = \frac{\partial lnp}{\partial z} - \frac{\partial lnT}{\partial z} \qquad (8)$$

Then after, the potential temperature (θ) of the parcel is calculated as follows based on the equation (2):

$$\frac{\partial ln\theta}{\partial z} = \frac{\partial lnT}{\partial z} - \frac{R}{c_p}\frac{\partial lnp}{\partial z} = \frac{1}{T}\left(\frac{\partial T}{\partial z} + \frac{g}{c_p}\right) = \left(1 - \frac{R}{c_p}\right)\frac{\partial lnp}{\partial z} - \frac{\partial ln\rho}{\partial z} \qquad \text{(9) to derive the}$$

Parcels acceleration based on equations (7) to become:

$$\frac{d^2s}{dt^2} = -g\,\frac{\partial ln\theta_0}{\partial z} z\ sina = -g\,\frac{\partial ln\theta_0}{\partial z} ds.sin^2a \qquad (10)$$

Whereas by introducing the frequency, N, with $N^2 = g\,\frac{\partial ln\theta_0}{\partial z}$

The Brunt-Vaisala frequency, N^2 is calculated based on the following mathematical formulation used to characterize atmospheric stability.

$$\mathrm{N}^2(z) = \frac{\mathrm{g}(z)}{T_0(\mathrm{z})}\left(\frac{\partial T_0(z)}{\partial z} + \Gamma_d\right) \quad (11)$$

Where g is the acceleration due to gravity, N is the Vaisala frequency, T_0 is the background temperature (estimated based on third-order polynomial fitting), $\Gamma_d = {}^{\mathrm{g}}/_{c_p}$ is the adiabatic lapse rate, and $c_p = 1004\, J\, K^{-1}\, kg^{-1}$ is the specific heat capacity of the atmosphere at constant pressure. When the Vaisala frequency, N^2, is positive, the atmosphere is stable. When N^2 is negative, the atmosphere is unstable. In this case, the atmospheric lapse rate, $\Gamma = -\frac{\partial T}{\partial z}$, is larger than the adiabatic lapse rate, ${}^{\mathrm{g}}/_{c_p} \approx 9.5\, K\, km^{-1}$. To estimate the Brunt-Vaisala frequency, a third-order polynomial fit of the least squares has been applied to the SABER observed temperature (T) profile to estimate the background temperature (T0), following the procedure of Leblanc and Hauchecorne (1997). After estimating the perturbed temperature (Tp) from equation (12), it is identified by subtracting the background temperature from the observed temperature data (T).

$$\mathrm{T_p} = \mathrm{T} - \mathrm{T_0} \quad (12)$$

After estimating the perturbed temperature (Tp), a high-pass band filter is applied. This filter removes low-frequency components associated with planetary and tidal waves, retaining the high-frequency components related to short-period gravity waves (John and Kumar, 2012). This process isolates the influence of gravity waves and accurately calculates their potential energy. The high-pass filter is based on known frequency ranges, typically below the period one-hour).

$$\mathrm{E_p}(\mathrm{z}) = \frac{1}{2}\left(\frac{\mathrm{g}(\mathrm{z})}{\mathrm{N}(\mathrm{z})}\right)^2 \left(\frac{\mathrm{T_p}'(\mathrm{z})}{\mathrm{T_0}(\mathrm{z})}\right)^2 \quad (13)$$

The potential energy of the waves, a function of altitude (z), is used to determine the impact of atmospheric gravity waves on atmospheric inversions.

3. Results and discussion

3.1 Identification and Characteristics of the Lower and Upper MLT Inversion

Daily SABER temperature profiles, covering altitudes of 60–100 km from 2005 to 2020, are shown in the contour plots of Figure 3. Figures 3 (a and b) depict the upper mesosphere, while Figures 3 (c and d) show the lower mesosphere. The horizontal panels of Figures 3 (a) and 3 (c) show observed temperatures ranging from approximately 180–220 K, before accounting for inversion layers. The horizontal panels of Figures 3 (b) and 3 (d) show inversion day temperatures, ranging from 180–225 K. These inversion day temperatures are higher than those shown in Figures 3 (a) and 3 (c). This indicates that maximum temperatures occur on inversion days in both the upper and lower MLT regions.

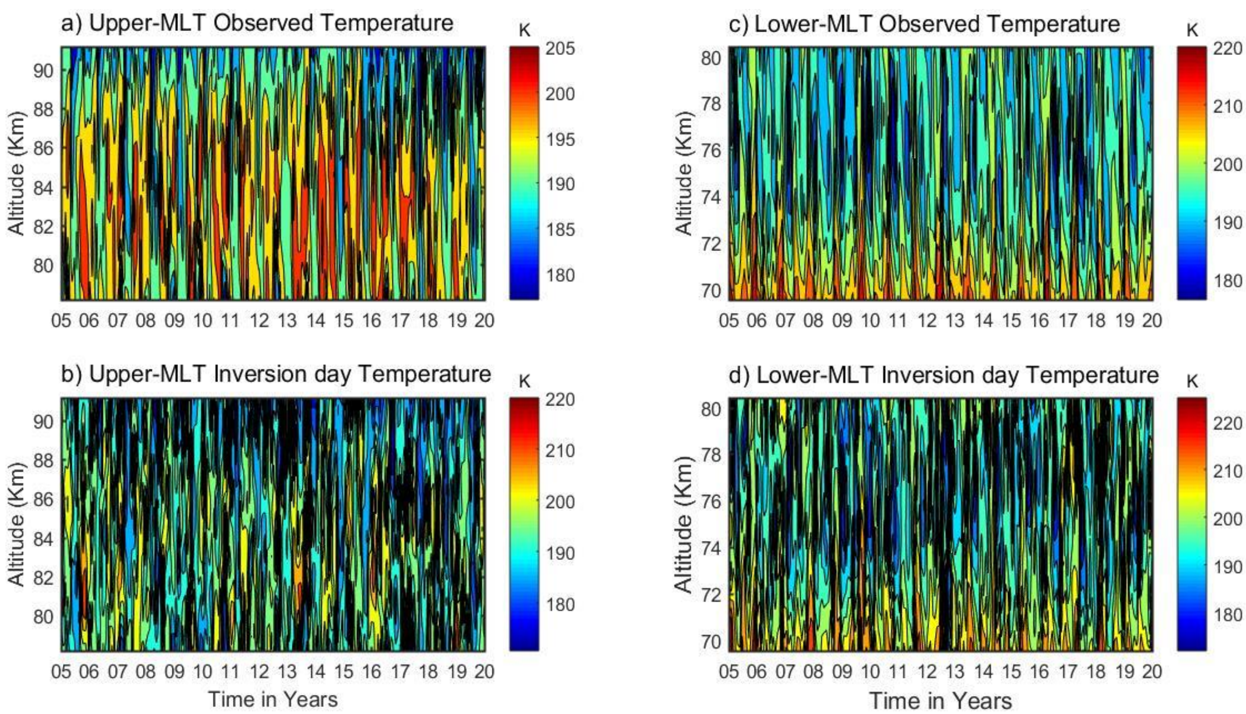

Figure 3. The upper and lower mesosphere observed temperatures in the first horizontal panel at (a and c) with their inversions in the second horizontal panel at (b and d).

The upper left panel of Figure 3(a) shows the observed temperature in the upper mesosphere. It ranges from approximately 180–205 K at altitudes of around 80–90 km. The upper right panel of Figure 3(c) shows the lower mesosphere, with temperatures ranging from about 180–220 K at altitudes of approximately 70–80 km. In contrast, the lower left panel of Figure 3(b) shows an upper-mesosphere inversion day temperature. It ranges from 180–220 K at an altitude of approximately 80–90 km. The lower right panel of Figure 3(d) shows a lower-mesosphere inversion day temperature. It ranges from 180–225 K at an altitude of approximately 70–80 km. These inversion day temperatures in Figures 3(b) and 3(d) suggest a temperature gradient shifting from negative to positive. This could be due to factors such as atmospheric gravity waves, chemical reactions, or solar radiation. Our temperature observations for the lower MLT region on an inversion day, within the altitudinal range of 70–80 km, align with those reported by Sivakumar et al. (2001). They identified inversion day temperature variability in the altitudinal range of 73–79 km. Additionally, Sivakandan et al. (2014) examined mesospheric inversions in the 60–105 km altitude range over low-latitude regions. Their findings closely match our results.

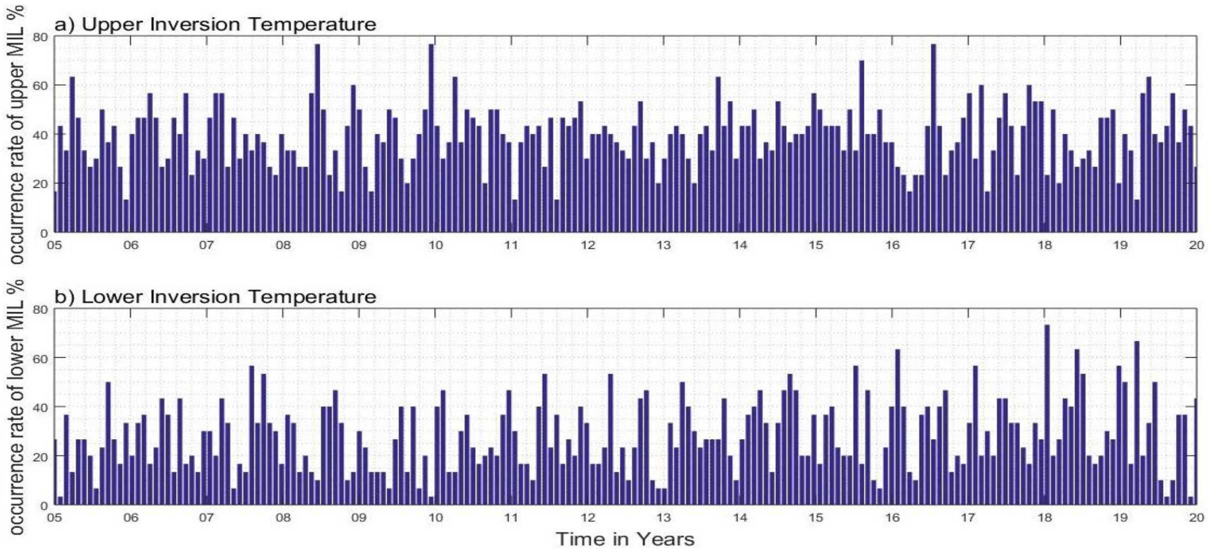

Figure 4. The frequency occurrence rate (percentage) of the (a) upper and (b) lower inversion temperatures during 2005-2020 over low latitudes.

Figure 4 shows the frequency occurrence rate (%) of mesospheric inversion layers (MILs) in histograms. Figure 4(a) shows the occurrence rate for upper MILs, while Figure 4(b) shows the rate for lower MILs. The mean frequency occurrence rate of upper inversions is approximately below 40%. Peak rates range from 60% to 78%, notably in the years 2008, 2010, and mid-2016. In contrast, the mean occurrence rate for lower inversions is below 20%. Overall, the occurrence rate for upper inversions is relatively higher compared to lower inversions. This may be related to atmospheric wave activities, mainly gravity waves. Hauchecorne et al. (1987) and France et al. (2015) discuss the effects of gravity waves on inversion variability in the upper and lower mesosphere. Based on these findings, Figure 5 examines the characteristics of inversion layers, including their amplitude and thickness. Figure 5 illustrates the characteristics of mesospheric temperature variability on an inversion day. It focuses on base height, amplitude, and thickness, before examining the effects of gravity waves on an inversion. Histograms show the frequency of amplitude, thickness, and base height for inversion day MLT temperature variability. These histograms feature best-fit Gaussian distribution curves, represented by red lines. The observed distributions align with Gaussian curves, indicating that the number of mesospheric inversion layers (MILs) follows a normal distribution. This suggests that the attributes are real-valued random variables. The left column has three rows showing histograms of (a) amplitude, (b) thickness, and (c) base height of the inversion day temperature variability for the upper MLT.

These histograms also include their statistical values, mean, and standard deviations (SD). The corresponding right column has three rows representing (d) amplitude, (e) thickness, and (f) base height of the inversion day temperature variability for the lower MLT region.

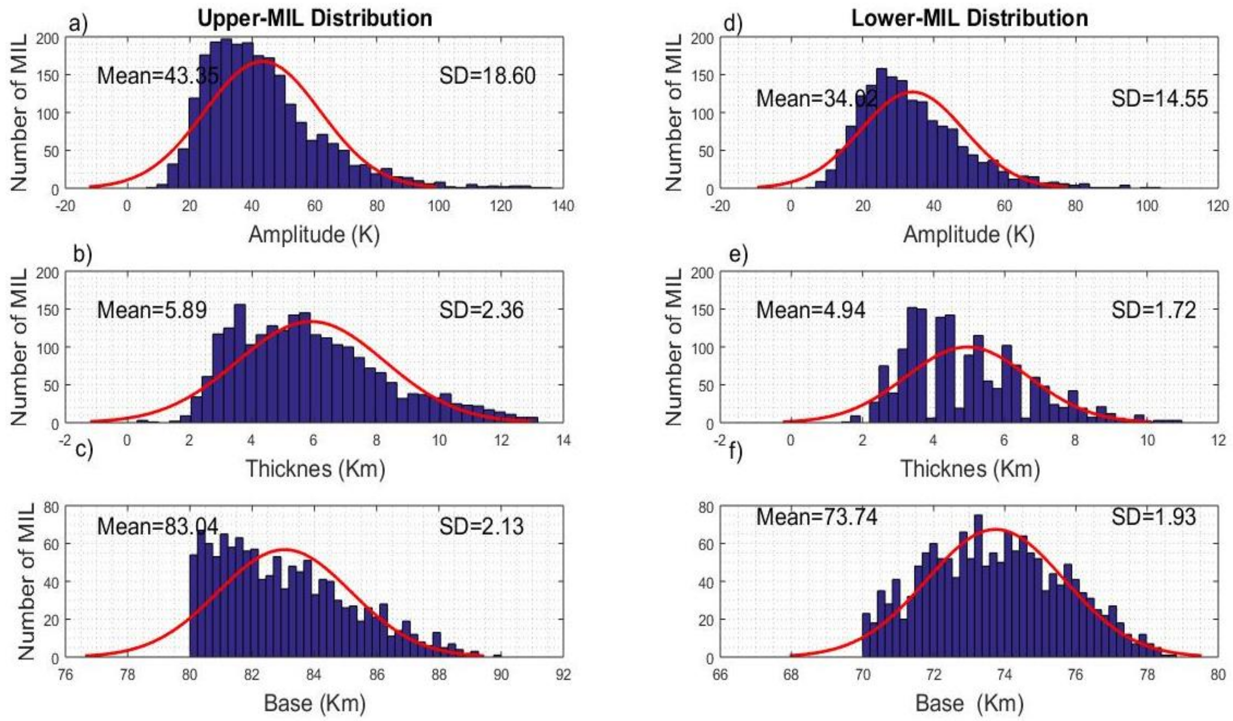

Figure 5. The histograms depict the occurrence of MLT inversion day temperature variability. The first vertical panel shows the distribution of (a) amplitude, (b) thickness, and (c) base height for the upper inversion day. The second vertical panel presents the corresponding distribution for the lower inversion, including (d) amplitude, (e) thickness, and (f) base height.

The amplitude of upper inversion day temperature variability in Figure 5(a) ranges between 20 and 80 K, with a peak value of 38 K. This follows a Gaussian distribution with a large standard deviation (SD) of 18.6. The thickness of the inversion layer for upper MILs, shown in Figure 5(b), ranges from 3 to 9 K. The most probable value is 5.5 K, with a low SD of 2.3. The base height of the upper MIL in Figure 5(c) spans from ~80 to 90 km, with a peak value of around 83 km. This indicates a large number of upper MLT inversions, with an SD of 2.13. The highest number of upper inversions between 2005 and 2020 is observed at 82 km. This may be attributed to gravity wave breaking and energy dissipation, influenced by waves generated from lower atmospheric regions and solar flux impacts. The lower inversion amplitude, depicted in Figure 5(d), ranges between 10 and 60 K, with a peak value of 25 K and a standard deviation (SD) of 14.5. The thickness of the lower inversion, shown in Figure 5(e), ranges from 3 to 8 km, with the most probable value at 3.8 km and a low SD of 1.72. The base height of the lower inversion in

Figure 5(f) ranges from 70 to 80 km, with a peak value of around 74 km and a lower SD of 1.93. Previous investigations by Sivakandan et al. (2014) from the Indian sector reported amplitudes ranging from 14–39 K in 2002 and 15–42 K in 2008. The thicknesses ranged between 2.7–7.5 km in 2002 and 2.8–7.3 km in 2008, under the influence of solar flux. These findings align well with the present study, indicating no significant variation in characterizing mesospheric inversion based on amplitude and thickness in the low-latitude region within the altitude range of 60 to 90 km.

3.2 Latitudinal Variations of Mesospheric Inversion Layers (MILs)

This section examines the spatiotemporal (latitudinal-time) variability of upper and lower mesosphere inversion phenomena. Contour plots of time vs. latitude in Figures 6 and 7, respectively, characterize this variability based on amplitude, thickness, and base height over the low-latitude band (3-15°) during 2005–2020. The Upper MILs phenomenon is observed around 80–90 km. The maximum amplitude, in the range of 90–120 K, occurs over latitude bands (5-12°) during 2005, 2007, mid-2011, 2013, 2015, 2016, mid-2019, and 2020 (Figure 6(a)). The second horizontal panel in Figure 6(b) shows the inversion thickness, with a maximum range of ~(8–12 km) across the entire latitudinal region (3-15° N). Figure 6(c) shows the relative maximum inversion base height, around ~(84-88 km), in the latitudinal range between 4° and 14° N during 2006, 2008, 2010, 2012, 2016, and 2018.

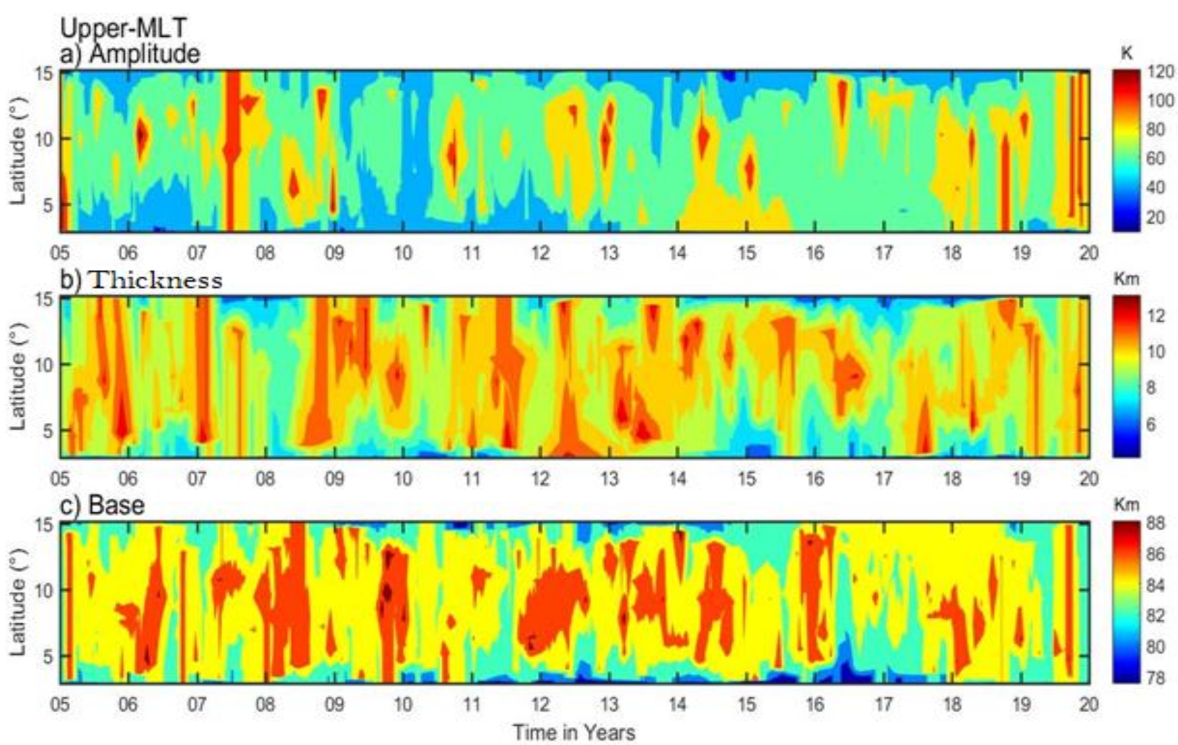

Figure 6. The daily upper inversions (~80-90 km) of (a) amplitude, (b) thickness, and (c) base height during 2005-2020 over latitudinal variation.

Contour plots in Figure 7(a, b, and c) show the latitudinal variations of the lower inversion (MILs) phenomenon, based on amplitude, thickness, and base height, respectively. These plots cover an altitudinal range of ~70-80 km. The lower inversion amplitude generally ranges from ~30-60 k across all latitudinal bands. However, it reaches a maximum range of ~(80-100 k) during 2013, 2015, 2016, and 2019 in various latitudinal regions between 5 and 140 N. Figure 7(b) shows an inversion thickness of 5-7 km across the entire latitude band, except for a maximum thickness of 8-10 km. Figure 7(c) shows a base height of 76-80 km for most latitudes and periods. However, there are exceptions: 2008, 2014, and mid-year 2018 exhibit a maximum base height.

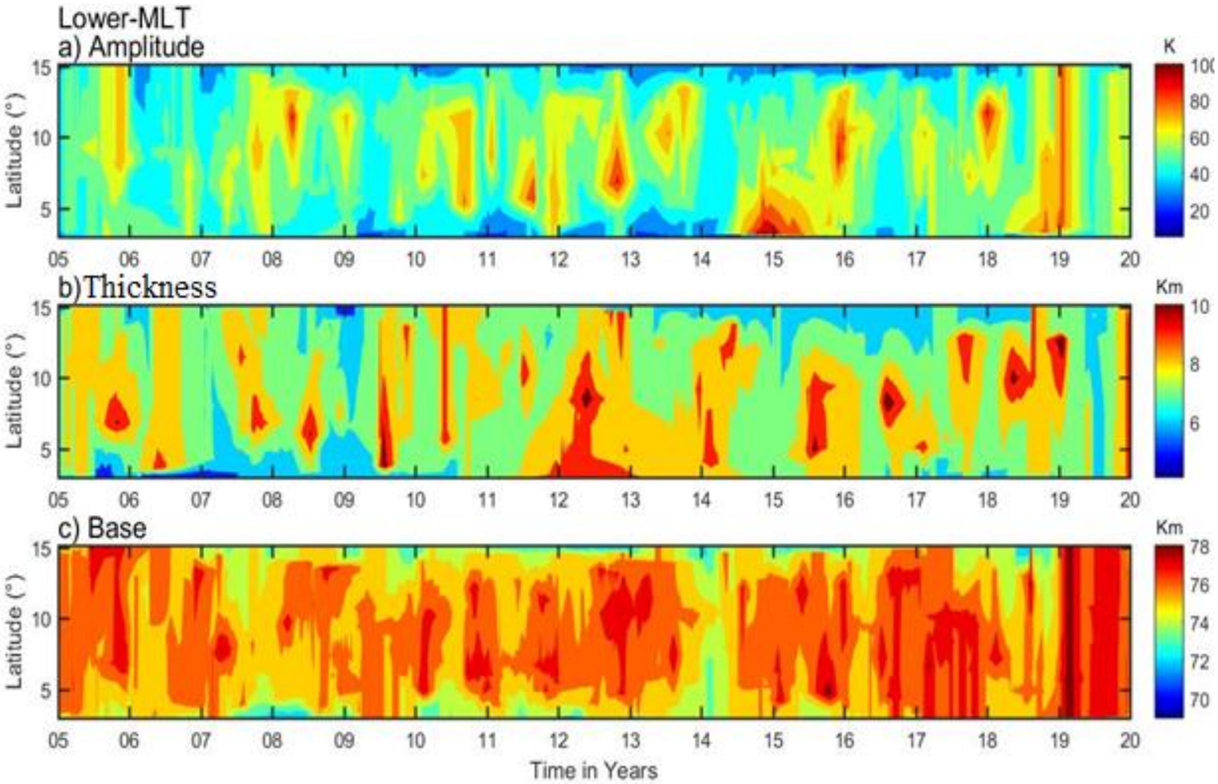

Figure 7. Same as Figure 5, but for the lower mesosphere inversions (~70- 80 km).

Figures 6 and 7 demonstrate that the upper inversion exhibits higher amplitude and thickness compared to the lower inversion, suggesting a highly dynamic phenomenon in the upper mesosphere region. Satellite measurements, particularly those from TIMED/SABER, provide valuable insights into latitudinal variations in MILs. These observations confirm that MILs also occur at low latitudes (Gan et al., 2012).

3.3 Separations of the Perturbed Temperature in the Mesosphere Region

The perturbed temperature (T_p) of the upper and lower MLT inversions during the period of 2005-2020 can further be used to calculate their derived potential energy of gravity waves and the Brunt-Väisälä frequencies (N^2). First, the upper-temperature inversion profiles are identified in the MLT region during the entire observational period of 2005-2020, as displayed in the contour plot of Figure 8(a). The observed temperature ranges from ~170-220 K with minimal variability. A 3rd-order polynomial fit is applied to calculate the background temperature (T_0), as shown in the contour plot of Figure 8(b). This background temperature exhibits periodic variability over an altitude of around ~82-87 km, ranging from ~195-197 K. The perturbed temperature profiles (Tp), calculated as the difference between the observed inversion temperature (T) and the corresponding background temperature profiles (T_o), range from -25 to +25 K, as shown in Figure 8(c).

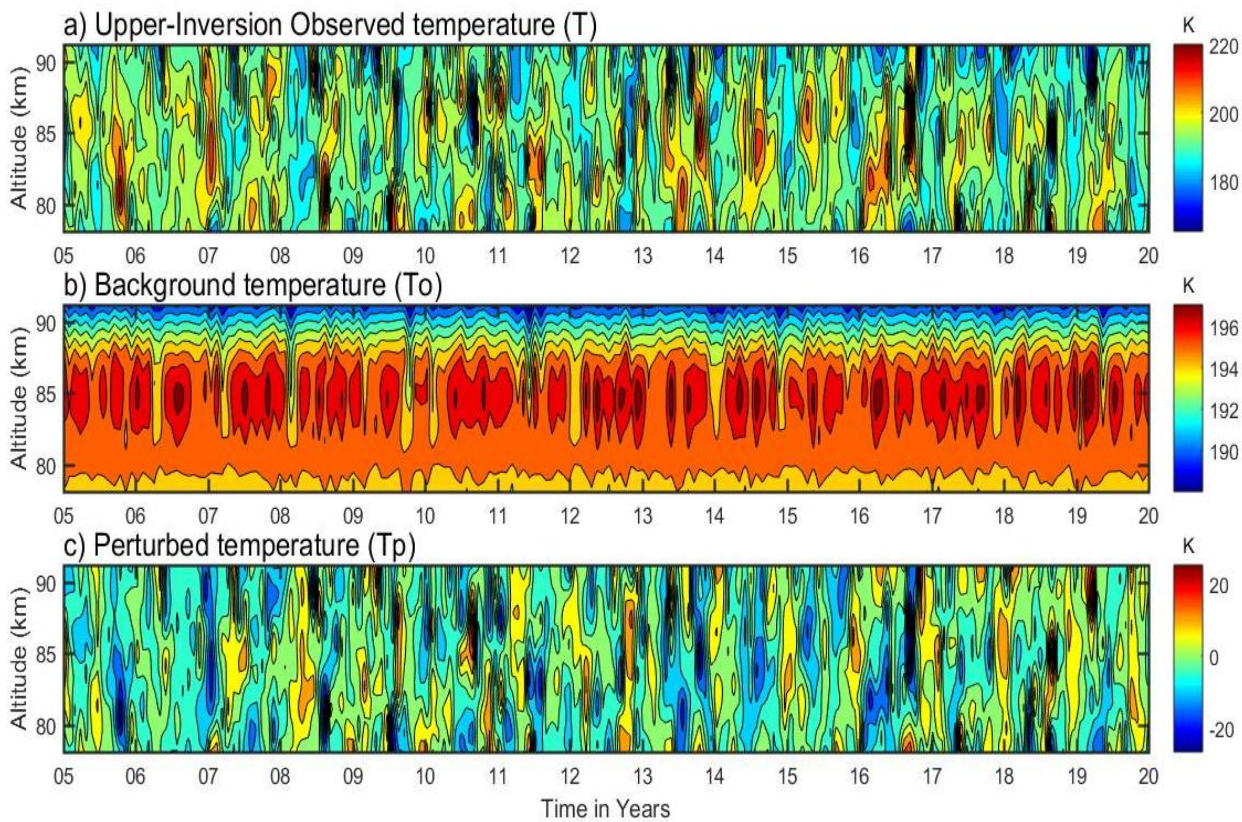

Figure 8. The upper mesosphere temperatures in the vertical panel are: (a) inversion day observed temperature; (b) background temperature; and (c) perturbed temperature in the upper mesosphere region.

A similar procedure is used to calculate the perturbed temperature (T_p), observed temperature, and background temperature in the lower mesosphere region from 2005 to 2020. Their corresponding contours are displayed in Figure 9(a-c). In Figure 9(a), the observed temperature of the lower inversion ranges from ~170-220 K. The background temperature of the lower inversion ranges from ~ 195-210 K, with maximum values of ~200-210 K at a height of ~70-72 Km, as shown in Figure 9(b). The perturbed temperature in Figure 9(c) ranges from -25 to 20 K. Notably, the upper mesosphere perturbed temperature is at its maximum compared to the lower mesosphere region, possibly due to a highly dynamic phenomenon.

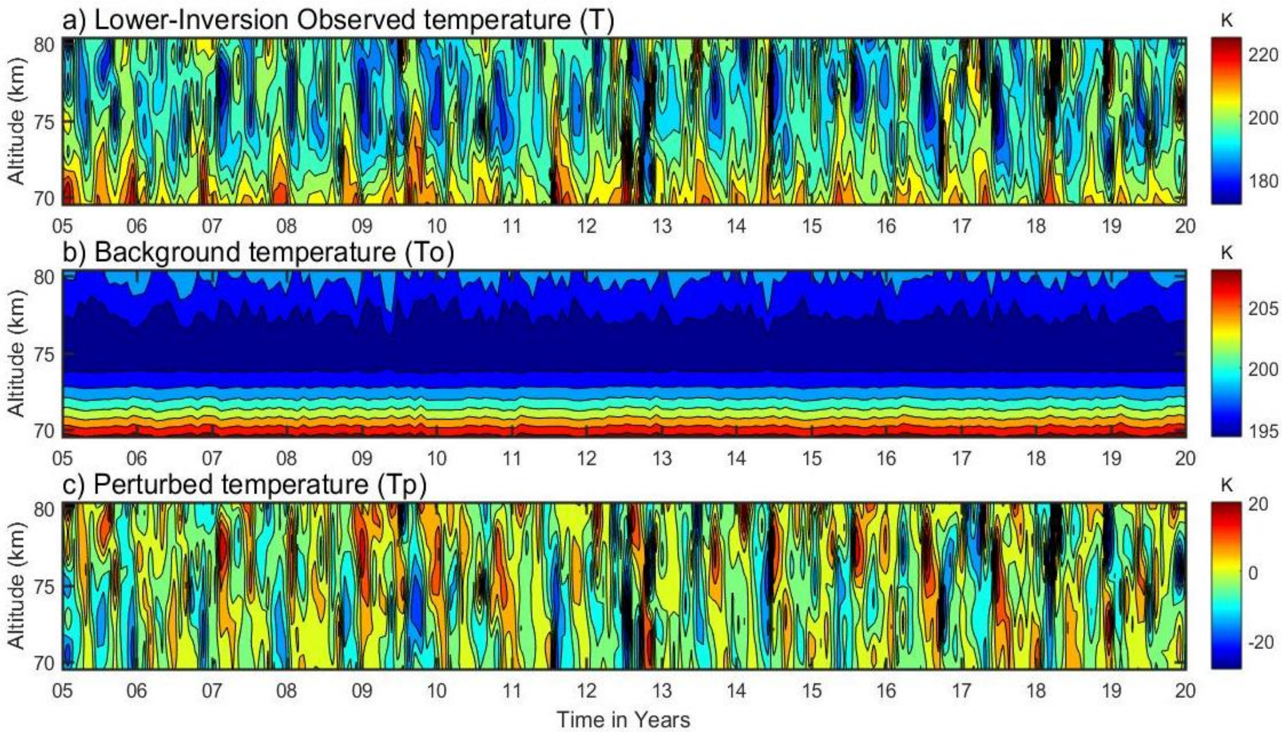

Figure 9. Same as Figure 7, but for the lower mesosphere atmospheric region.

3.4 Effects of Gravity Waves on Mesosphere Inversions and Associated Instability

Gravity waves form when air parcels oscillate due to the restoring force of gravity after being transported vertically. Several factors contribute to these waves, including airflow over mountains, convection, and wind shear. Propagating vertically, the waves break and dissipate, releasing energy and momentum into the surrounding atmosphere. This process, frequently responsible for the formation of inversion layers, is further investigated by identifying gravity wave potential energy (Ep) and its impact on inversion layers at selected MLT regions. This approach assumes that gravity wave activity is represented by potential energy, as described by numerous authors (Tsuda et al., 2000; Wang and Geller, 2003; Liu et al., 2014; Thurairajah et al., 2014). The gravity wave contribution is quantified by calculating the potential energy and evaluating instability through the Brunt-Väisälä frequency (N^2), derived from perturbed temperature (Tp') data spanning 2005–2020. The analysis focused on altitudes of 90, 85, 75, and

70 km by applying a high-pass filter with a one-hour interval to the Tp' data (see Figure 10 a-d). In the upper mesosphere (90 and 85 km), the filter reduces the amplitude of wave oscillations from approximately ±20 K to ±10 K, as shown by the blue curve in Figure 10a and b, compared to the red curve. Similarly, in the lower mesosphere (75 and 70 km) at (Figure 10 c & d), the amplitude decreases from ~(-20 to 20 K) to ~(-8 to 8 K) by filtering out higher amplitudes. By removing the impact of long-period wave oscillations, such as tidal and planetary wave contributions, the filter effectively isolates the gravity waves (Gw) on the MLT inversions.

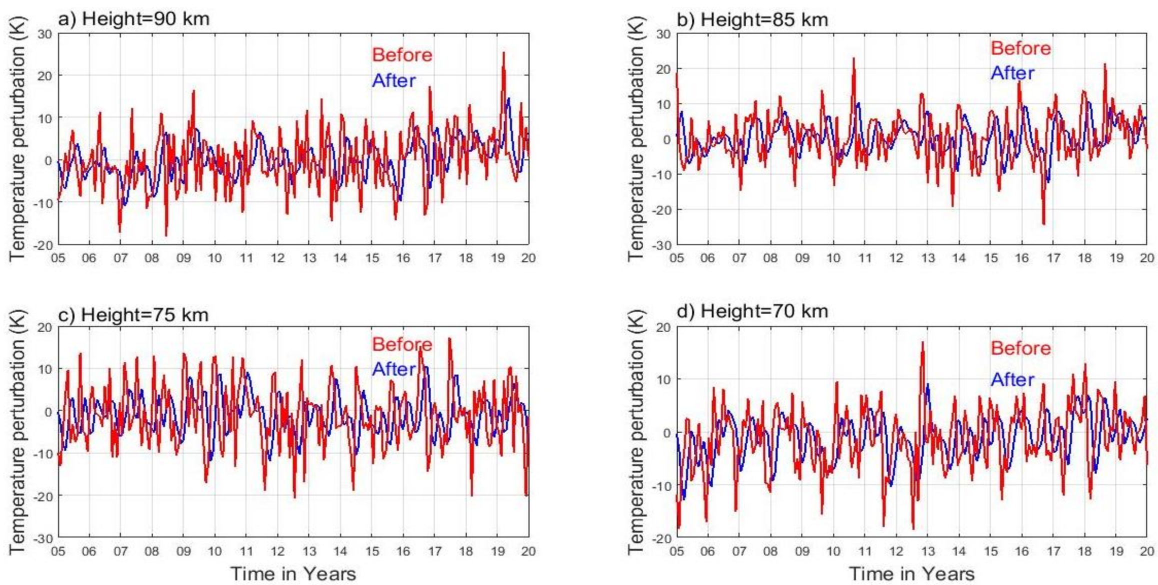

Figure 10. Perturbed temperature profiles before (red color) and after (blue color) applying the high-pass filter for the upper (90 and 85 km) and lower (75 and 70 km) regions.

In the MLT atmospheric region, gravity wave breaking typically dissipates their potential and kinetic energy, leading to increased turbulence and mixing. This energy transfer can alter thermal patterns and impact the overall dynamics of the upper atmosphere. As illustrated, gravity wave propagation and dissipation are major forces in the MLT region (Lindzen, 1981; Holton, 1983), influencing middle and upper atmospheric inversions. This has a substantial impact on the MLT's thermal structure, particularly the increase in temperature variability with elevation, known as inversion. Holton et al. (2003) and Holton and Hakim (2013) has demonstrated an interaction between the potential energy of gravity waves and inversions. Interestingly, in this investigation, the dissipation of gravity waves can lead to the mesospheric inversion layers (MILs). The study we conducted has clearly shown that the occurrence of an inversion is maximum at the upper MLT region relative to the lower MLT region from Figure 4. In a comparable manner, Figure 11 depicts the highest potential energy of gravity waves in the upper

MLT regions, demonstrating the interactions between inversion and gravity wave potential energy.

Figure 11 (a-d) demonstrates that the spatiotemporal variability of gravity wave potential energy, showing over the upper MLT at (90 and 85 km) and the lower MLT at (75 and 70 km). Figure 11(a) of upper MLT inversions at 90 km shows maximum gravity wave potential energies, ranging from ~70 to 90 J/kg, over the longitudinal regions of 45-47° E, as well as at 43° E, and 44° E during 2011, 2017, and 2019. In contrast, potential energies around ~10 to 60 J/kg are present across the entire longitudinal region from 33-48° E. Figure 11(b) shows maximum potential energies of ~70 to 100 J/kg over the longitudinal regions of 34°, 44°, and 46° E during 2014, 2016, and 2018 at 85 km. Minimum potential energies between 20 and 70 J/kg appear over the longitude regions from 33-48° E. Figures 11(c) and 11(d) depict the gravity wave potential energy in the lower MLT regions at 75 and 70 km, respectively. At 75 km, Figure 11(c) shows a relative maximum potential energy of 40-50 J/kg over the longitudinal regions of 46°, 42°, 40°, 37°, 36°, and 38° E during 2011, 2012, 2017, 2013–2015, 2018, and 2020. Similarly, Figure 11(d) illustrates gravity wave potential energy ranging from 2-30 J/kg at 70 km across the longitudinal region of 33-48° E. Maximum potential energy of 25-30 J/kg is observed in certain longitudinal regions over time.

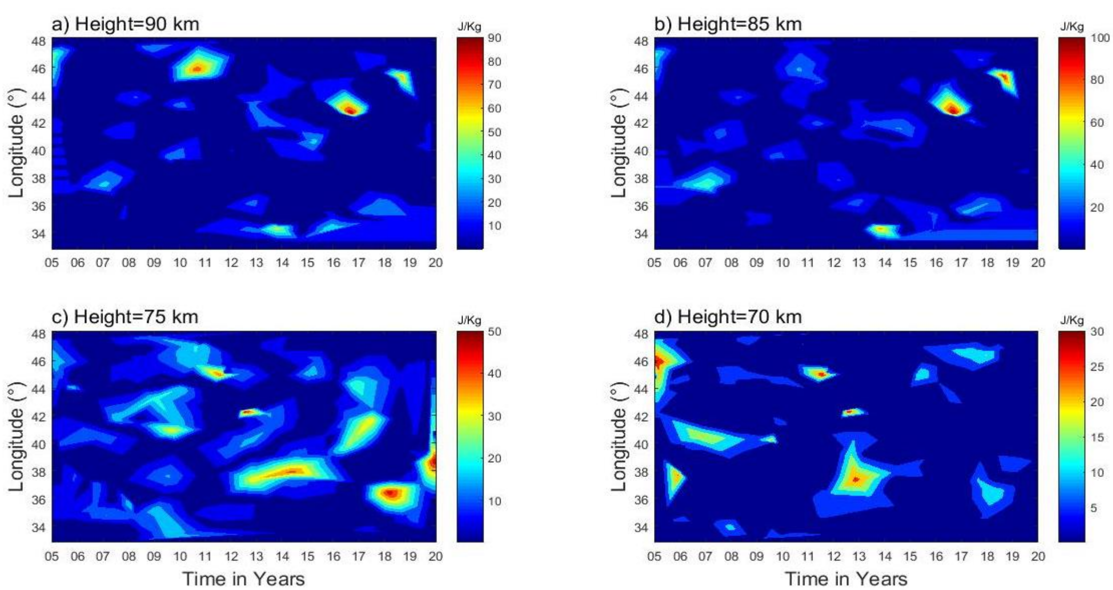

Figure 11. Gravity wave potential energy for the upper (90 and 85 km) and lower (75 and 70 km) MLT regions.

Gravity waves generated in different sizes, with smaller waves being the main derivers of instability and turbulence in the MLT (mesosphere and lower thermosphere) region (Liu and

Meriwether, 2004; Szewczyk et al., 2013). Hauchecorne et al. (1987) proposed a model where a series of breaking gravity waves leads to the formation of MILs through the gradual accumulation of heat, which contributes to instability. Conducting mesospheric inversion layer (MIL) phenomena is crucial for understanding MLT atmospheric dynamics, especially when it comes to stability and energy transfer. To investigate the role of gravity waves in MLT instability, the Brunt-Väisälä frequency was used. Contour plots in Figure 12 (a-d) show the spatiotemporal variability of the Brunt-Väisälä frequency, with Figures 12(a and b) representing the upper mesosphere (90 and 85 km) and Figures 12(c and d) representing the lower mesosphere (75 and 70 km). The Brunt-Väisälä frequency (N^2) shows that the upper MLT region is more unstable (~0.027 at 90 km and ~0.029 at 85 km) relative to that of the lower MLT region (~0.033 at 75 km and ~0.035 at 70 km).

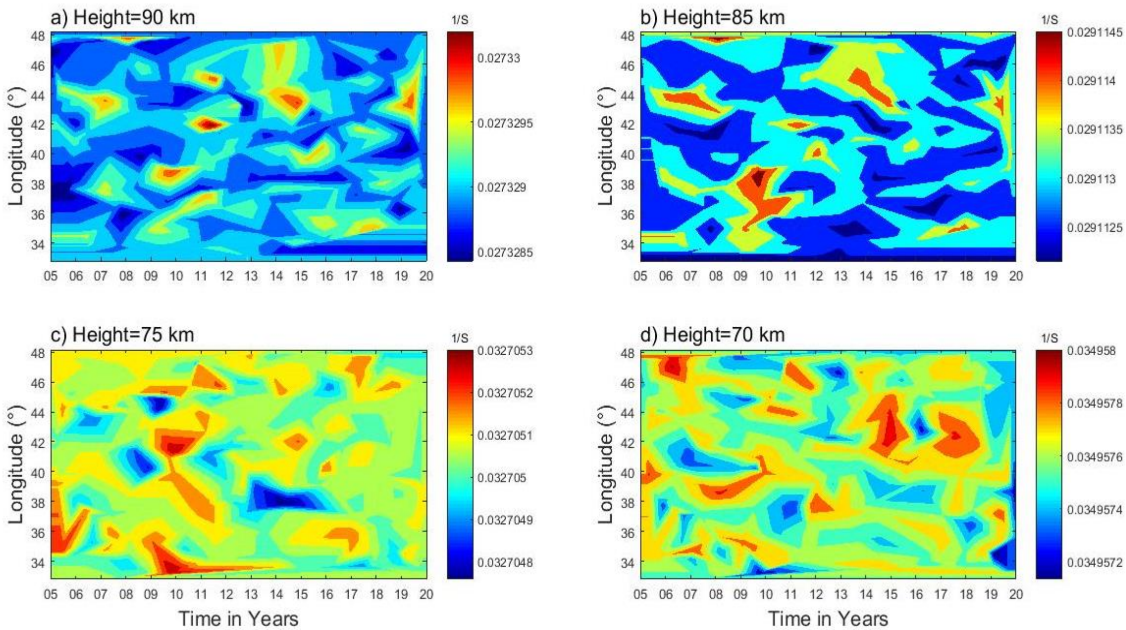

Figure 12. Brunt-Vaisala frequency (N^2) for the upper (90 and 85 km) and lower (75 and 70 km) MLT regions.

4 Summary

In this article, 16 years of SABER MLT temperature profiles are utilized to investigate the MIL phenomenon and its causative mechanism through gravity wave potential energy (P_E) and instability criteria of Brunt-Väisälä frequency (N^2) over low latitude bands. The following conclusions are drawn from the observations in this article:

- ✓ The upper mesosphere inversion frequency occurs more often than the lower mesosphere inversion.

- ✓ Analysis of the MIL characteristic features reveals the most probable values for the upper inversion: amplitude of 38 k, thickness of 5.5 km, and base height of 78 km. The lower inversion has an amplitude of 25 K, a thickness of 3.8 km, and a base height of 73 km.
- ✓ The upper mesosphere region has higher gravity wave potential energy compared to the lower mesosphere region.
- ✓ The high potential energy in the upper mesosphere region is likely due to the deposition of energy and momentum by gravity wave breaking. This could influence the dynamics of the inversion phenomenon.
- ✓ The Brunt-Väisälä frequency (N^2) indicates that the upper mesosphere region is less stable than the lower mesosphere region. This lower stability contributes to the high potential energy in the upper mesosphere, which leads to larger inversion phenomena.
- ✓ Atmospheric processes vary significantly from region to region, with altitude, and over time.

D***ata availability.*** The SABER data are freely available via the link at http://saber.gats-inc.com/index.php.

Author contribution. Chalachew Lingerew: data curation, investigation, software, visualization, writing the original draft, and writing review. U. Jaya Prakash Raju; supervision, and editing.

Competing interest. The authors declare that they have no conflict of interest relevant to this study.

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