



A statistical study of the magnetic signatures of the unique Tonga volcanic explosion of 15 January 2022

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6 Abstract. For the first time, a statistical study has been conducted of the geomagnetic bay and quasi-periodic disturbances based on the datasets collected at 19 recording stations participating in INTERMAGNET Magnetic Observatories. In order 7 8 to identify the disturbances from the volcanic explosion, a preliminary analysis has been used of the state of space weather 9 during the catastrophic Tonga volcanic explosion of 15 January 2022. We summarize the main results as follows: The nonmonotony of the variations in the strength of all geomagnetic field components increased appreciably on the day of the 10 11 explosion as compared to the variations observed during the days used as a quiet time reference, while the eastward component of the geomagnetic field exhibited an up to 60-nT increase in variability. The duration and time delay of the bay 12 13 disturbances increased with distance from the volcano, while their amplitude decreased. The propagation speeds of the bay disturbances at various observatories were determined to be in the 700-1,000 m/s range. Six groups of time delays of quasi-14 sinusoidal disturbances have been identified in a simultaneous analysis for the first time; they correspond to the apparent 15 speeds of 4 km/s, 1.5 km/s, 1 km/s, as well as 500 m/s, 313 m/s, and 200 m/s. The time delay in each group increased with 16 17 distance away from the volcano. The agreement between theoretical estimates and the observational data testify to the 18 adequacy of the mechanism adopted for the generation of the disturbances.

19 1 Introduction

Five underwater Tonga volcanic explosions (20°54' S, 175°38' W) were observed to occur over the 04:00–05:00 UTC period 20 21 on 15 January 2022, with the second explosion at 04:15 UTC being the most powerful (Adushkin et al., 2022; Astafyeva et 22 al., 2022; Matoza et al., 2022a; Matoza et al., 2022b). The gas emissions reached 50-58-km altitude producing the highest 23 recorded eruption column, whereas the eruption columns of Krakatoa volcano on 26-27 August 1883 reached only 40-55 km (Chernogor, 2012; McNutt et al., 2015). The Tonga volcanic eruption thermal energy is estimated to be ~3.9 10¹⁸ J, and 24 the mean thermal power to be 9.1×10^{12} W (Chernogor, 2022a; Chernogor, 2022e; Chernogor, 2023a). The mass of the 25 26 erupted material attained 2.9 Gt and their volume 1.9×10^9 m³. The volcanic explosivity index (VEI) did not exceed 5.8, and 27 the explosive energy was estimated to be in the range from 4–18 Mt of TNT to 478 ± 191 Mt of TNT (Adushkin et al., 2022; 28 Astafyeva et al., 2022; Kulichkov et al., 2022).





29 The Tonga volcanic explosion was accompanied by essential disturbances in all components of the Earth 30 (lithosphere, ocean)-atmosphere-ionosphere-magnetosphere system (Chernogor, 2022a; Chernogor, 2022b; Chernogor, 31 2022c; Chernogor, 2022d; Chernogor, 2022e; Chernogor, 2023a; Chernogor, 2023b). More than 50 studies were concerned 32 with the effects caused by the volcanic explosion. Measurements were made of the earthquake of Richter magnitude 5.8 33 (Poli and Shapiro, 2022), of seismic wave propagation (Diaz et al., 2022; Matoza et al., 2022a; Matoza et al., 2022b; Poli 34 and Shapiro, 2022), of tsunamis (Carvajal et al., 2022; Imamura et al., 2022; Kubota et al., 2022; Ramírez-Herrera et al., 35 2022; Tanioka et al., 2022; Terry et al., 2022), of Lamb waves (Kubota et al., 2022; Kulichkov et al., 2022; Lin et al., 2022; 36 Matoza et al., 2022a; Matoza et al., 2022b; Otsuka et al., 2022), of atmospheric gravity, infrasound, and sound waves (Burt 37 et al., 2022; Chen et al., 2022; Chernogor and Shevelev, 2022; Lin et al., 2022; Matoza et al., 2022a; Matoza et al., 2022b; 38 Wright et al., 2022), as well as observations were made of volcanic signatures in the atmosphere and ionosphere (Aa et al., 39 2022a; Aa et al., 2022b; Ajith et al., 2022; Astafyeva et al., 2022; Chen et al., 2022; Chernogor et al., 2022; Harding et al., 40 2022; Hong et al., 2022; Lin et al., 2022; Muafiry et al., 2022; Rakesh et al., 2022; Shinbori et al., 2022; Sun et al., 2022a; 41 Sun et al., 2022b; Themens et al., 2022; Zhang et al., 2022a; Zhang et al., 2022b).

Theoretical studies of the chain of physical processes were performed by (Chernogor, 2012; Chernogor, 2022a;
Chernogor, 2022b; Chernogor, 2022c; Chernogor, 2022d; Chernogor, 2022e; Chernogor, 2023a; Chernogor, 2023b).

Sun et al. (2022b) have estimated disturbances in the electric current in the ionospheric *E* region caused by the Tonga volcanic explosion by making use of the data on geomagnetic field variations acquired by the global network of magnetometers. The *E*-region current density was estimated to be $J \approx 22-55$ mA/m, which changed the eastward components, *Y*, of the geomagnetic field by ~20–50 nT. The leading front of the disturbance traveled with a propagation speed of 740 m/s. Le et al. (2022) investigated the effect that the volcano had on the equatorial electrojet and revealed the reversal of the electrojet direction due to a strong eastward zonal wind.

50 The explosion was also accompanied by variations in the geomagnetic field (Adushkin et al., 2022; Chernogor, 2023c; Chernogor and Holub, 2023a, 2023b; Iyemori et al., 2022; Le et al., 2022; Schnepf et al., 2022; Soares et al., 2022; 51 52 Yamazaki et al., 2022). Adushkin et al. (2022) have described waves and disturbances in the atmospheric electric and 53 magnetic fields. The data collected at 14 stations in the global network of observatories, INTERMAGNET, which are located in the 2.790-6.225 Mm distance range from the volcano, have been used for investigating the magnetic effect. The 54 55 disturbances in the geomagnetic field have been deduced to occur on a global scale, and two groups of disturbance have been revealed. In the first group, the disturbances were virtually synchronously observed immediately after the explosion, whereas 56 57 in the second group, the magnetic disturbances appeared after the arrival of Lamb waves. Soares et al. (2022) described 58 quasi-periodic disturbances in the magnitude of the eastward component, Y, with amplitude of ~ 3 nT and an ~ 4 -min period 59 observed with onset time delay of 10 min at 835-km distance from the volcano. The geomagnetic variations at 3.8-mHz (period of $T \approx 4.4$ min) have been analyzed by (Iyemori et al., 2022; Yamazaki et al., 2022), who relate these variations to 60 61 the acoustic resonance. It is important to note that the oscillations at 3.8 mHz were observed simultaneously both in the 62 vicinity of the volcano (API station) and in the magnetically conjugate region (HON station). The amplitudes of these





virtually synchronous oscillations were observed to be 2 nT and 0.2 nT, respectively, while the time delay of the magnetic 63 64 effect did not exceed 6 min. However, analogous oscillations were not observed at distances, r, greater than 2.7 Mm. The study by Schnepf et al. (2022) is concerned with the investigation of geomagnetic variations in the 3-8-min period range 65 66 with amplitude of ~1 nT that were observed with a time delay of ~30 min (propagation speed of ~470 m/s). The authors relate these variations to the ionospheric wave, which was generated by the volcano, and explain the variations in the 13-93-67 and 5-100-min period ranges by the effects of tsunami and of atmospheric and ionospheric sources. Harding et al. (2022) 68 69 describe the multi-instrument studies of the magnetic effect of Tonga volcano. They utilized the data collected by 70 magnetometers at the ground and onboard the ICON and Swarm spacecraft to study the effect that the volcanic explosion 71 had on neutral winds and the ionospheric dynamo current system on a global scale. Despite significant progress made in 72 understanding the geomagnetic field disturbances related to the Tonga volcanic explosion, a further statistical and spectral 73 analyses of these variations is to advance understanding of this scientific issue.

The purpose of this paper is to present, for the first time, the inferences of the statistical and spectral analyses of the bay and quasi-periodic disturbances in the geomagnetic field that were observed to occur after the Tonga volcanic explosion on 15 January 2022. The data used for this research have been acquired at nineteen INTERMAGNET observatories closest to the volcano.

78 1 Information on Tonga volcano

Tonga volcano is located ~200 m below the oceanic surface. An intense volcanic eruption was recorded to occur from ~04:00 UTC to ~16:00 UTC on 15 January 2022 when the rates of eruption attained 67 kt/s or 44,000 m³/s. In total, the volcano was active for over 12 ± 1 h, whereas the energy of the blast wave was estimated to be 16–18 Mt TNT [Chernogor, 2022a; Chernogor, 2022e; Chernogor, 2023a]. Generally, Tonga volcano is among the five most powerful on record (Table 1).

Information	Krakatoa	St. Helen	El Chichón	Pinatubo	Tonga	
Data	26–27 August	19 May 1090	29 March and	15 June 1001	15 January 2022	
Date	1883	18 May 1980	3-4 April 1982	15 June 1991	15 January 2022	
Constant la settar	To to see to	USA, Skamania		DI '1'	Kingdom of	
Country, location	Indonesia	County	Mexico	Philippines	Tonga	
Geographic	6°06′ N,	46°12′ N,	17°22′ N,	15°7.8′ N,	20°54′ S,	
coordinates	105°25′ E	122°11′ W	93°14′ W	120°21′ E	175°38′ W	
Total eruptive mass	2.0×10^{13}	1.2×10^{12}	1.2×10^{12}	1.2×10^{13}	2.0×10^{12}	
(kg)	2.9×10^{13}	1.3×10^{12}	1.3×10^{12}	1.3×10^{13}	2.9×10^{12}	

84 Table 1. Basic information on volcanos.





Eruption column	40.55	10.25	30 32	33	50 58
height (km)	ght (km)		50-52	55	50-58
Mean mass flow rate (kg/s)	$5.5 imes 10^8$	2×10^7	1.5×10^{8}	8×10^8	6.7×10^{7}
VEI	6	5	5	6	5–6
Magnitude	6.5	5.1	5.1	6.1	5.5
Intensity	11.7	10.3	11.2	12	10.8
Notes	Current altitude 813 m, Vent 120 m	Altitude 2,549 m, Reduced by 400 m	Altitude 1,150 m, Vent 1,000 m, Crater depth 300 m	Altitude 1,486 m; before 1991, 1,745 m	Plinian underwater eruption at a 200 m depth

The volcanic magnitude can be estimated using the following formula of McNutt et al. (2015):

86
$$M = \log m - 7$$

85

where *m* is the erupted mass (in kg). Substituting $m = 2.9 \times 10^{12}$ kg yields $M \approx 5.5$, whereas the most powerful Krakatoa volcanic has a magnitude of $M \approx 6.5$ (Table 1). The mass eruption rate, \dot{m} , is characterized by the intensity, given by the relation (McNutt et al., 2015):

90 $I = \log \dot{m} + 3$.

91 Here \dot{m} is in kg/s. Given the averaged value of the mass eruption rate $\dot{m} \approx 6.7 \times 10^7$ kg/s, the intensity is $I \approx 10.8$, whereas 92 $I \approx 11.7$ for Krakatoa volcano (Table 1).

93 3 Analysis of the state of space weather

The state of space weather over the 12-18 January 2022 period is characterized by the data retrieved from the World Data 94 Center for Geomagnetism, Kyoto https://wdc.kugi.kyoto-u.ac.jp/ and from the Goddard Space Flight Center Space Physics 95 96 Data Facility https://omniweb.gsfc.nasa.gov/form/dx1.html. The sunspot number did not exceed ~100, while the daily 10.7 cm solar radio flux ($F_{10,7}$) was in the ~100–120 sfu range (1 sfu = 10⁻²² W /(Hz m²)). A substantial increase (by a factor of a 97 98 few times) in the solar wind parameters took place during the 14/15 January 2022 UTC night. The interplanetary magnetic 99 field B_z component showed a decrease from ±4 nT to -14 nT, the equatorial D_{st} index exhibited a decrease from 10 nT to -90 nT, while the auroral activity index A_p showed an increase from ~5 nT to 67 nT, and the 3-h range planetary K_p index from 100 101 ~1 to 5.7. Thus, a G2-moderate geomagnetic storm took place over the 14/15 January 2022 night. The recovery phase of the 102 storm proceeded over the 15–18 January 2022 period. It should be noted that the auroral electrojet (AE) index was observed

103 to be ~100 nT over 04:00-11:00 UTC period on 15 January 2022, i.e., geomagnetic conditions were quiet, whereas before





04:00 UTC and after ~11:00-12:00 UTC, the AE index exceeded 500 nT, which indicated a geomagnetic disturbance
(substorm). 13 January 2022, when the geomagnetic conditions were quietest, was chosen to be a quiet time reference. 17
January 2022 is also used, although partially, as a quiet time reference.

107 4 Instrumentation and techniques

The study is based on data from the INTERMAGNET magnetic observatory network, which were accessed through the https://www.intermagnet.org/. The list of the stations is presented in Table 2, and their locations around Tonga volcano are depicted in Figure 1. It is important to note that the stations are located around all cardinal points as seen from the volcano. We have analyzed the temporal variations in the northward, *X*, eastward, *Y*, and vertical, *Z*, components of the geomagnetic field acquired on 12, 13, 15, 16, 17, and 18 January 2022 with 1-min temporal resolution and the root-mean-square error not exceeding 1 nT.

114 The algorithm for finding the geomagnetic field response to Tonga volcanic explosion is as follows:

115 (1) Since the variations in the geomagnetic field may be caused by many powerful sources releasing significant 116 amounts of energy, any characteristic changes in the variations in the strength of the X, Y, and Z components that were 117 observed to occur after the volcanic explosion and could be associated with the explosion are highlighted at the first stage of 118 employing the algorithm. This condition is necessary but insufficient.

- (2) At the second stage, the variations analogous to those that occurred on quiet time days and were due to, forexample, diurnal variation, the solar terminator, etc., are filtered out.
- (3) Next, the possible time delays and apparent speeds are determined. The time delay should increase with distancefrom the volcano.

(4) If some apparent speeds at different stations are substantially close to each other, they are included in a particular statistic. The closeness of the apparent speeds in this particular statistic is considered a sufficient condition for this particular disturbance to be due to the volcanic explosion.

126 (5) The physical significance of the apparent speeds is an additional sufficient condition: these speeds must 127 correspond to the known speeds of waves of particular physical nature.

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8 (6) The results obtained are compared, if possible, with the results obtained for the volcanoes that exploded before.

It should be emphasized that the variations in the geomagnetic field components were generally more or less smooth on the days used as a quiet time reference, whereas they became non-monotonical after the Tonga volcanic explosion when aperiodic and quasi-periodic variations were observed to occur in the magnitude of the geomagnetic field components. The moving average process was first created by averaging, over 60-min intervals, the raw data X(t), Y(t), and Z(t) sampled at a 1-min time step to be subtracted from the temporal variations in the raw data to yield the $\Delta X(t)$ -, $\Delta Y(t)$ -, and $\Delta Z(t)$ component deviations, which were finally subjected to the Fourier and wavelet transforms (Chernogor, 2008).







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136 Figure 1: Map showing the sites of the recording stations. Star designates the volcano.

	Casaranhia	Magnetia		Distance	
	Geographic			from the	
IAGA [®] station code	latitude and	latitude and	Country	explosion	
	longitude	longitude longitude		(km)	
	13.8155° S	15.01° S	Western Company	940	
Apia (API)	171.7812° W	96.77° W	western Samoa	840	
Pamatai (Papeete)	17.5670° S	15.15° S	French	2 720	
(PPT)	149.5740° W	74.29° W	Polynesia	2,730	
	43.4740° S	46.56° S		2 700	
Eyrewell (EYR)	172.3930° E	106.28° W	New Zealand	2,790	
	35.3200° S	41.75° S	A . 11	2 00 6	
Canberra (CNB)	149.3600° E	132.81° W	Australia	3,806	
Charters Towers	20.0900° S	27.05° S	A . 11	2 000	
(CTA)	146.2640° E	138.47° W	Australia	3,990	
Macquarie Island	54.5000° S	59.32° S	Australia 4,34		

137 Table 2. Information on the INTERMAGNET magnetic observatories.





(MCQ)	158.9500° E	116.38° W			
	21.3200° N	21.65° N	United States of	5.004	
Honolulu (HON)	158.0000° W	88.98° W	America	5,024	
	23.7620° S	31.83° S	A (11	5 210	
Alice Springs (ASP)	133.8830° E	151.19° W	Australia	5,210	
	12.6900° S	20.96° S	A (1'	5 (0)	
Kakadu (KDU)	132.4700° E	153.66° W	Australia	5,602	
Isla de Pascua	27.1713° S	19.48° S	C1 '1'		
Mataveri (IPM)	109.4200° W	34.44° W	Chili	6,675	
	31.3560° S	40.34° S	A1:	C 997	
Gingin (GNG)	115.7150° E	170.60° W	Australia	0,887	
Learne outh (LDM)	22.2200° S	31.28° S	Assatualia	7 222	
Learmonth (LKM)	114.1000° E	172.67° W	Australia	7,255	
	36.2320° N	28.13° N	Ionon	7 950	
Kakioka (KAK)	140.1860° E	150.18° W	Japan	7,632	
Kanava (KNV)	31.4200° N	22.70° N	Ionon	9 125	
Kanoya (KINY)	130.8800° E	158.28° W	Japan	8,155	
Momembatan (MMD)	43.9100° N	36.09° N	Ionon	° 765	
Memambelsu (MMB)	144.1900° E	147.57° W	Japan	8,203	
Shumagin (SIIII)	55.3500° N	54.46° N	United States of	0 557	
Shumagin (SHU)	160.4600° W	100.96° W	America	8,337	
Dalat (DLT)	11.9400° N	2.60° N	Viotnom	0.068	
Dalat (DL1)	108.4800° E	178.89° W	vietiiaiii	9,008	
Cocos (Keeling)	12.1875° S	21.21° S	Australia	0.208	
Islands (CKI)	96.8336° E	168.97° E	Australia	9,508	
Gan International	0.6946° S	8.34° S	Maldivas	12 210	
Airport (GAN)	73.1537° E	145.40° E	ivialui ves	12,210	

138 ^aIAGA stands for International Association of Geomagnetism and Aeronomy





139 5 Analysis of temporal variations in geomagnetic field strengths

140 A preliminary analysis of the temporal dependences X(t), Y(t), Z(t) and of their time derivatives $\dot{X}(t)$, $\dot{Y}(t)$, $\dot{Z}(t)$ 141 determined that the character of the variations on 15 January 2022 was markedly different from that observed during the 142 quiet time reference periods when the variations were smoother and the values of the derivatives were noticeably smaller.

- 143 *API Station.* Geomagnetic bay disturbances were absent on 13 January 2022 (Figure 2), and the magnitude of 144 fluctuations did not exceed 1 nT. On 17 January 2022, used as a quiet time reference, synchronous geomagnetic bay 145 disturbances were absent (Figure 2). The magnitudes of fluctuations in all components exhibited insignificant variability 146 within the ± 1 -nT limits.
- 147 On 15 January 2022, the geomagnetic bay disturbances appeared with a time delay, τ , of ~16 min and lasted for ΔT_X 148 $\approx 120 \text{ min}, \Delta T_Y \approx 146 \text{ min}, \text{ and } \Delta T_Z \approx 130 \text{ min}$. They were observed to occur virtually synchronously in all three components 149 of the geomagnetic field (Figure 2). The peak deviations from the trend in the bay disturbances are estimated to be $\Delta X \approx 16$ 150 nT, $\Delta Y \approx 26$ nT, and $\Delta Z \approx -13$ nT. The more rapid fluctuations are superimposed upon these slow enough variations; they 151 appear with time delays of $\Delta t_0 \approx 6 \text{ min}, \Delta t_1 \approx 8.5 \text{ min}, \Delta t_2 \approx 14 \text{ min}, \Delta t_3 \approx 19 \text{ min}, \Delta t_4 \approx 33 \text{ min}, \Delta t_5 \approx 50 \text{ min}, \text{ and } \Delta t_6 \approx 75$ 152 min (Table 3).







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Figure 2: UTC variations of the northward, *X*, eastward, *Y*, and vertical, *Z*, components of the geomagnetic field at the API station during 15 January 2022, the day of the volcanic explosion, and during the days used as a quiet time reference. The vertical line marks the moment of the most powerful explosion. Arrows indicate possible moments of the onset of the magnetic field response.





	Δt_1	v'_1	Δt_2	v'_2	Δt_3	v'_3	Δt_4	v'_4	Δt_5	v'_5	Δt_6	v_6'
Station	(min)	(m/s)										
API	8.5	4,000	14	1,560	19	1,000	33	500	50	311	75	200
PPT	16	4,100	37	1,420	50	1,011	96	500	150	314	235	198
EYR	17	3,875	38	1,410	50	1,033	97	505	155	310	240	200
CNB	21	4,000	47	1,500	68	1,006	130	507	208	312	322	200
СТА	22	3,900	49	1,510	71	1,008	137	504	217	314	338	200
MCQ	23	4,030	53	1,510	77	1,007	150	500	237	313	368	200
HON	26	4,000	61	1,490	89	1,000	173	498	272	313	424	200
ASP	27	3,950	63	1,500	92	1,000	185	482	282	313	440	200
KDU	28	4,060	67	1,500	98	1,004	190	505	305	311	475	199
IPM	33	3,970	79	1,500	115	1,011	215	530	360	313	565	199
GNG	34	4,000	82	1,500	119	1,007	235	500	372	312	580	200
LRM	35	4,018	85	1,500	125	1,000	245	500	390	313	615	198
KAK	38	3,967	90	1,540	135	1,007	260	513	490	315	645	204
KNY	39	3,988	95	1,507	140	1,004	270	512	435	315	685	199
MMB	39.5	3,993	97	1,497	143	998	273	514	442	315	690	201
SHU	40	4,070	100	1,501	147	1,004	285	509	460	313	720	199
DLT	43	3,976	106	1,501	156	1,001	305	504	488	313	760	200
CKI	44	3,978	110	1,477	160	1,001	310	509	500	313	780	200
GAN	56	3,990	140	1,507	208	1,002	410	502	660	311	1,020	200

158 Table 3. Time delays and apparent speeds of disturbances in the geomagnetic field.





159 PPT Station. On 13 and 17 January 2022, used as a quiet time reference, the strength of the X-component showed 160 variations from about (-4)-(-5) nT to 2-5 nT (Figure 3) throughout the entire 03:00-17:00 UTC period, whereas the Ycomponent increased from (-6)-(-7) nT to 7-8 nT over the 03:00-05:00 UTC period and afterwards exhibited fluctuations 161 162 within the 2–3-nT limits, gradually decreasing from ~0 nT to (-15)–(-23) nT. On 13 January 2022, the Z-component exhibited undulating oscillation during the 03:00 to \sim 10:00 UTC period followed by a gradual decrease from \sim 0 to -10 nT, 163 whereas on 17 January 2022, it showed a broad maximum of 8 nT near ~04:00 UTC followed by a gradual decrease to a 164 165 minimum of -5.1 nT at ~10:30 UTC and later by oscillations with an amplitude of ~1 nT around the trend changing in the -5.1- to 0-nT range. 166

167 On the day of the volcanic explosion, the non-monotonousness in the magnitude of all components increased, the 168 fluctuations of the components also somewhat increased, while the trend substantially smoothed in all components. The 169 magnitude of the *X*-component increased from -10 nT to 20 nT, the value of the *Y*-component decreased from 15 nT to -20170 nT, and of the *Z*-component decreased from 10 nT to -5 nT. In addition, six groups of disturbances appeared with time 171 delays of $\Delta t_1 \approx 16 \text{ min}$, $\Delta t_2 \approx 37 \text{ min}$, $\Delta t_3 \approx 50 \text{ min}$, $\Delta t_4 \approx 96 \text{ min}$, $\Delta t_5 \approx 150 \text{ min}$, and $\Delta t_6 \approx 235 \text{ min}$ (see Figure 3). The 172 greatest disturbances (up to 10 nT) occurred after 14:00 UTC in the *X* component.

173 EYR Station. On 13 January 2022, the strength of the X-component increased, fluctuating, from -5 nT to 10 nT, and 174 then decreased from 10 nT to -3 nT (Figure 4). Since 06:00 UTC, the magnitude of the X-component fluctuated around the 2-nT value. The strength of the Y-component first showed a decrease from 12 nT to ~0 nT during the 03:30-05:00 UTC 175 period, then it fluctuated around ~ 0 nT, and afterwards its value was observed to decrease to -(5-10) nT. The magnitude of 176 the Z-component decreased from ~8 nT to 0 nT and fluctuated around 0 nT afterwards. On 17 January 2022, the magnitude 177 of the X-component was fluctuating around 0 nT, with excursions attaining 6-7 nT, while the magnitude of the Y-component 178 179 was decreasing, fluctuating, from ~ 20 nT to -10 nT. At the same time, the strength of the Z-component decreased from 7–8 180 nT to -10 nT, and then it increased, noticeably fluctuating, from -10 nT to 7 nT.

181 On the day the volcanic explosion occurred, the number of groups of disturbances was observed to attain six. The 182 most pronounced disturbances were negative geomagnetic bay disturbances with time delays of $\tau_X \approx 86 \text{ min}$, $\tau_Y \approx 51 \text{ min}$ and 183 $\tau_Z \approx 51 \text{ min}$. The drops in the *X*-, *Y*-, and *Z*-components attained -39 nT, -27 nT, and -22 nT, respectively. The drops in the 184 *X*-, *Y*-, and *Z*-component strengths were followed up by increases of ~38 nT, 30 nT, and 30 nT, respectively. The amplitudes 185 of other disturbances usually did not exceed a few nanoteslas, and they arrived with time delays of $\Delta t_1 \approx 15 \text{ min}$, $\Delta t_2 \approx 38 \text{ min}$, $\Delta t_3 \approx 50 \text{ min}$, $\Delta t_4 \approx 97 \text{ min}$, $\Delta t_5 \approx 155 \text{ min}$, and $\Delta t_6 \approx 240 \text{ min}$ (see Table 3).















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191 *CNB Station.* On 13 January 2022, the fluctuations in the magnitude of all components did not exceed 1–4 nT 192 (Figure 5). The strength of the *Y*-component showed a decrease from ~17 nT to -5 nT. The *Z*-component was observed to 193 drop from 4 nT to -7 nT over the 03:00–05:00 UTC period, and then its magnitude showed fluctuations around 1 nT. On 17 194 January 2022, the strength of the *X*-component gradually increased from -10 nT to 10 nT, fluctuating within the $\pm(5-7)$ -nT 195 limits, while the magnitude of the *Y*-component decreased from 27 nT to -10 nT. The strength of the *Z*-component increased 196 from -12 nT to 7 nT first, and then fluctuated within the $\pm(3-4)$ -nT limits.

During the course of the day the volcanic explosion occurred, the trend \overline{X} first increased from -10 nT to 10 nT, then it decreased from 10 nT to -10 nT, and once again increased from -10 nT to 30 nT, while the magnitude of fluctuations was observed to be $\pm(3-5)$ nT. The trend \overline{Y} first decreased from 16 nT to -40 nT, then it increased from -40 nT to 28 nT, and once again decreased from 28 nT to -30 nT. The trend \overline{Z} first decreased from -8 nT to -13 nT, then it increased from -13 nT to 14 nT, and once again decreased from 14 nT to -5 nT, which was followed by an increase in \overline{Z} from -5 nT to 15 nT. The variations with amplitude of a few nanotesla were superimposed on the slow trend in all components.

CTA Station. On the days used as a quiet time reference, the magnitudes of all components exhibited relatively 203 204 small fluctuations (Figure 6) except for the variations in the X-component on 17 January 2022 when its strength showed fluctuations within the ±5-nT limits while the trend \overline{X} increased from -10 nT to 5 nT. Instead, the trend \overline{Y} decreased from 205 25 nT to approximately -10 nT, and the trend \overline{Z} from 13 nT to -5 nT. On 13 January 2022, the trend \overline{X} decreased from 12 206 nT to -5 nT over the 03:00–05:00 UTC period and then remained at this level. The trend \overline{Y} decreased from 7 nT to -4 nT 207 from 03:00 UTC to 07:00UTC, then increased to 2 nT, and afterwards gradually decreased from 2 nT to -5 nT. The trend \overline{Z} 208 sharply decreased from 8 nT to -2 nT over the 03:00 UTC to 07:00 UTC period, and then it showed fluctuations around 0 209 210 nT. During 13 and 17 January 2022 used as a quiet time reference, synchronous geomagnetic bay disturbances were absent (Figure 6). 211

The magnitude of fluctuations in all components considerably increased on 15 January 2022. The trend \overline{X} increased from -15 nT to 20 nT. The *Y*-component, in addition to fluctuations, exhibited a deep drop from ~15 nT to -45 nT that occurred from 05:45 UTC to 08:30 UTC (see Figure 6). The *Z*-component also showed a drop from ~0 nT to -13 nT during the 05:20–07:15 UTC period, followed by a surge from -13 nT to ~23 nT. In addition, the disturbances appeared in all components with time delays of 22 min, 49 min, 71 min, 137 min, 217 min, and 338 min (see Figure 6).









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221 *MCQ Station.* On 13 January 2022, X(t), Y(t), and Z(t) components showed variations in the trend not exceeding 20 222 nT (Figure 7), as well as fluctuations within the ±3-nT range. On 17 January 2022, the strengths of all components showed 223 small variations up to 10:00 UTC, whereas the variations increased to 100–200 nT after 10:00 UTC.

On the day the volcanic explosion occurred, all components exhibited relatively small variations before 11:00 UTC, whereas they showed an increase of up to 300–400 nT after 11:00 UTC. Approximately from 06:00 UTC to 08:00–09:00 UTC, the strengths of the *X*-, *Y*-, and *Z*-component decreased by 80 nT, 40 nT, and 30 nT, respectively. Such a perturbation pertains to a bay disturbance. In addition, except for the bay disturbance, quasi-periodic disturbances occurred with strengths of 1–10 nT and $T \approx 5$ –10-min periods.

HON Station. On 13 January 2022, the *X* component exhibited weak fluctuations within the ± 1 nT limits from 00:00 UTC to 10:00 UTC (Figure 8). After 10:00 UTC, the level of variability noticeably increased. The strength of the *Y* component displayed a rise from -4 nT to 6 nT over the 00:00–04:00 UTC period, followed by a gradual decrease to 0 nT at 12:00, and the trend continued to decrease later. Throughout the 00:00–04:00 UTC period, the magnitude of the *Z* component showed first an increase from -3 nT to 1.3 nT, and then it decreased, fluctuating, from 1.3 nT to -2 nT. On 17 January 2022, all components showed insignificant (less than 1 nT) fluctuations from 00:00 UTC to 10:00 UTC. After 10:00 UTC, the magnitude of the fluctuations increased to ± 3 nT.

On 15 January 2022, the *X* component exhibited fluctuations within the \pm (5–7) nT limits, and the fluctuations in the strength of the *Y* component was noticeable enough (up to \pm (2–4) nT) as well. The *Z* component also showed a significant enhancement in fluctuations after the volcanic explosion, while all geomagnetic field components exhibited several groups of disturbance from 04:30 UTC to 12:00 UTC.

ASP Station. On 13 January 2022 used as a quiet time reference, the trend \overline{X} first decreased from ~15 nT to -(5-0) nT, and then it remained at a level of 0 nT (Figure 9). The trend \overline{Y} decreased from ~8 nT to -5 nT, whereas the trend \overline{Z} first sharply decreased from 13 nT to -3 nT, and then it exhibited variations between -3 nT and -1 nT. The strengths of fluctuations in all components usually did not exceed 1-2 nT. On 17 January 2022 used as a quiet time reference interval, the trend \overline{X} exhibited insignificant changes, and the strength of fluctuations did not exceed ±(3-5) nT. The trend \overline{Y} decreased from 27 nT to -21 nT, and its strength showed fluctuations attaining ±(8-10) nT. The trend \overline{Z} first sharply decreased from 19 nT to -5 nT, and then $\overline{Z} \approx -4$ nT; the magnitude of fluctuations did not exceed ±1 nT.

During the day the volcanic explosion occurred, all components of the geomagnetic field experienced geomagnetic bay disturbances that were superimposed on fluctuations with strengths of up to 4–5 nT. The trend \overline{X} exhibited a drop from 0 nT to -15 nT, whereas the trend \overline{Y} showed a considerably greater drop, from 12 nT to -40 nT, of almost 4 h temporal duration; a powerful surge from -40 nT to 20 nT of 5.5 h duration followed afterwards. The trend \overline{Z} first increased from -10 nT to -2 nT, then decreased from -2 nT to -10 nT, and increased from -10 nT to 22 nT afterwards; this surge in the trend \overline{Z} was followed by a decrease to 0–5 nT.





























258 Figure 9: Same as in Figure 2 but for the ASP station.





KDU Station. On 13 January 2022, the trend \overline{X} first sharply decreased from ~27 nT to -5 nT, and then it showed fluctuations around a strength of -5 nT (Figure 10). At the same time, the trend \overline{Y} gradually decreased from 2–3 nT to -4 nT, while the trend \overline{Z} first (up to 06:30 UTC) increased to 11 nT, then sharply decreased to -5 nT and did not change afterwards. During 17 January 2022, the behavior of the trends \overline{X} and \overline{Y} was qualitatively analogous to their behavior observed on 13 January 2022; however, fluctuations in the strength increased to 2–4 nT. The trend \overline{Z} , fluctuating, decreased from approximately 14 nT to -5 nT.

265 On the day the volcanic explosion occurred, the trend \overline{X} first decreased, fluctuating, from -7 nT to -15 nT, next it 266 increased from -15 nT to 10 nT before ~14:00 UTC, and then a drop was observed to occur to -10 nT over the 14:00-15:30 267 UTC period. The trend \overline{Y} first increased from -3 nT to 8 nT, next it decreased from 8 nT to -23 nT, then it increased from -268 23 nT to 20-22 nT, and finally it gradually decreased from 20-22 nT to -18 nT. The trend \overline{Z} first increased from ~3 nT to 9 269 nT, then it decreased from 9 nT to -8 nT, and once again increased from -8 nT to 13 nT. After this peak, the value of \overline{Z} was 270 observed to gradually decrease from 13 nT to -5 nT. Variations with amplitudes of a few nanotesla were superimposed on 271 the relatively smooth changes in all components.

IPM Station. On 13 January 2022 used as a quiet time reference, the magnitude of all components before 11:00–
12:00 UTC varied within the 5–7-nT limits (Figure 11). Quasi-periodic variations were virtually absent. During 17 January
2022 up to 12:00 UTC, the variations in *X*-, *Y*-, and *Z*-components did not exceed 3–5 nT.

On 15 January 2022, the day the volcanic explosion occurred, insignificant bay reductions of only (4–8) nT in the magnitudes of all components were observed to appear with time delays of 120–125 min and durations of 210–230 min, whereas quasi-periodic disturbances were virtually absent.

GNG Station. On 13 January 2022, used as a quiet time reference, the trend \overline{X} first sharply decreased from 20 nT 278 to -5 nT and then fluctuated around -5 nT (Figure 12). The trend \overline{Y} , fluctuating within the \pm (3–4)-nT limits, gradually 279 decreased from 12 nT to -5 nT. Over the 03:00–09:00 UTC period, the trend \overline{Z} substantially sharply decreased from 25 nT 280 to -8 nT, next it remained almost constant. On 17 January 2022, also used as a quiet time reference, the trend \overline{X} gradually 281 increased from -12 nT to 10 nT, while the strength of the X-component showed fluctuations within the $\pm(4-10)$ -nT limits. 282 The trend \overline{Y} first increased to 30 nT, and then gradually decreased from 30 nT to -20 nT showing fluctuations sometimes 283 attaining $\pm 5-10$ nT. The trend \overline{Z} increased to 38 nT by 05:30 UTC, then decreased to -17 nT by 12:00 UTC, and later 284 almost did not change; the amplitude showed fluctuations within the ± 4 -5-nT limits after 12:00 UTC. 285

Throughout the day the volcanic explosion occurred, all components showed variations qualitatively different from those observed over a quiet time period. Approximately since 06:00 UTC, all components reduced their strengths by 20–50 nT during 2–3 h. Next, their strengths increased by 15–40 nT over an almost 2-h interval. All components exhibited 5–9-nT variations superimposed on the slow changes.















293 Figure 11: Same as in Figure 2 but for the IPM Station.







295 Figure 12: Same as in Figure 2 but for the GNG Station.





LRM Station. On 13 January 2022, the trend \overline{X} nearly linearly decreased from 30 nT to -10 nT over the 03:30– 08:30 UTC period, and later it changed insignificantly, while the trend \overline{Y} showed variations not exceeding ~7–8 nT (Figure 13). The trend \overline{Z} first showed fluctuations about the 7-nT strength level, and later about the -5-nT level. On 17 January 2022, the trend \overline{X} showed fast fluctuations within the ±10-nT limits, while the trend itself first increased to 06:00 UTC, and then decreased from 14 nT to -10 nT before 10:00 UTC. The trend \overline{Y} increased from -15 nT to 25 nT, and then decreased nonmonotonically to -20 nT, while the amplitude of fluctuations attained ±(5–6) nT. The trend \overline{Z} decreased from 40 nT to -15 nT during the 04:00–10:00 UTC period and then fluctuated around -20 nT.

303 On 15 January 2022, the day the volcanic explosion occurred, the variations were observed to be substantially 304 different. After exhibiting insignificant fluctuations from 03:00 UTC to 06:00 UTC, the trends decreased by approximately 305 40 nT. After this, the trends were observed to increase by 20–50 nT over a 1–2-h interval. The amplitudes showed 306 fluctuations that did not exceed \pm (3–4) nT.

KAK Station. On 13 January 2022, used as a quiet time reference, the trend X increased from -20 nT to 5 nT over 307 the 03:00–08:00 UTC period, and then gradually decreased from 5 nT to 0 nT (Figure 14). The trend \overline{Y} decreased to -19 nT 308 by 05:00 UTC, after that it increased to 0 nT by 08:00 UTC and then remained essentially constant. The trend \overline{Z} increased 309 from -15 nT to 8 nT over the 03:00-06:00 UTC period, next it decreased from 7 nT to 0 nT from 07:00 UTC to 09:00 UTC, 310 and then almost did not change. The magnitude of fluctuations in all components did not exceed ±1 nT. On 17 January 2022 311 used as a quiet time reference, the trend \overline{X} increased from -10 nT to 5 nT, though a reduction in the dependence $\overline{X}(t)$ was 312 observed to occur from 08:00 UTC to 15:00 UTC, and $\overline{X} \approx 6$ nT from 15:00 UTC to 17:00 UTC. The trend \overline{Y} first 313 decreased to -17 nT, and then increased to 5 nT. The trend \overline{Z} increased to 2 nT by 06:40 UTC and dropped from 2 nT to -3314 nT over the 06:40–12:00 UTC period. Then $\overline{Z} \approx 2$ nT. The amplitude showed fluctuations attaining 2–3 nT in every 315 316 component.

317 During the course of the day the volcanic explosion occurred, the magnitude of fluctuations in every component increased noticeably. The trend \overline{X} , fluctuating, decreased from 6 nT to -7 nT from 06:00 UTC to 10:00 UTC. Next, it 318 increased from -7 nT to 10 nT over a 3-h interval, and finally the trend decreased to -13 nT. The trend \overline{Y} also first 319 increased from -27 nT to -7 nT, then it decreased by less than 10 nT over a 1.5-h interval, after which the trend \overline{Y} increased 320 to 17 nT at 13:30 UTC. A clear quasi-periodic perturbation with a period of $T \approx 55-60$ min and a strength of 4 nT was 321 recorded from approximately 11:00 UTC to 14:00 UTC. Other disturbances had amplitudes of 1-1.5 nT. The trend Z 322 showed fluctuations within the ±3-nT limits throughout the 03:00–07:00 UTC period, when the trend remained almost the 323 324 same, next, from 07:00 UTC to 12:00 UTC, the trend decreased from 3 nT to -5 nT, and finally it increased. The fluctuations 325 occurred with an amplitude of ~1 nT.







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327 Figure 13: Same as in Figure 2 but for the LRM Station.















KNY Station. During 13 January 2022, the trends \overline{X} and \overline{Z} increased from -21 nT to 5 nT from 03:00 UTC to 330 08:00 UTC (Figure 15), while the trend Y was observed to develop a deep drop from 4 nT to -25 nT. From 08:00 UTC to 331 17:00 UTC, the strength of fluctuations of the geomagnetic field was insignificant, ±1 nT. On 17 January 2022, a deep drop 332 occurred in the \overline{Y} and \overline{Z} trends. The magnitude of fluctuations in all components attained ±5 nT. On 15 January 2022, the 333 magnitudes of the trends in all components increased over the 03:00 UTC to 06:00 UTC period. The strength and time rate 334 of fluctuations also increased. Six groups of perturbation were observed to arrive with time delays of 39 min, 95 min, 140 335 min, 270 min, 435 min, and 685 min (see Figure 15); the amplitudes of the disturbances attained 4–5 nT. A pronounced Y-336 337 component oscillation with $T \approx 70$ -min period and an amplitude of 4-nT arrived with the time delay of 435 min.

338 *MMB Station.* On 13 January 2022, the trend of X(t) increased from -19 nT to 5 nT, and then fluctuated around 2–3 339 nT (Figure 16). The *Y* component showed a negative bay disturbance, with a strength reduction from -4 nT to -17 nT, which 340 persisted from 03:00 UTC to 08:00 UTC. Then, up to 17:00 UTC, an insignificant rise in this component strength was 341 observed to occur. The *Z*-component, instead, exhibited a positive bay disturbance over the 03:00 UTC to 09:00 UTC period. 342 On 17 January 2022, the strengths of the *X*- and *Y*-components showed variations attaining 10–15 nT, whereas the *Z*-343 component variations did not exceed 5–6 nT.

- On 15 January 2022, the bay reductions by 10 nT, 10 nT, and 3 nT in the strengths of the *X*, *Y*, and *Z* components, respectively, were observed to occur with time delays of 200–225 min and to persist for 210 min to 290 min. In addition, the amplitudes showed quasi-periodic disturbances with amplitudes of a few nanotesla and $T \approx 7-20$ min.
- 347 SHU Station. During 13 January 2022, the trend \overline{X} decreased from 3 nT to -2 nT over ~03:30 UTC to 17:00 UTC 348 period (Figure 17), while the trend \overline{Y} , instead, increased from -4 nT to 3 nT. The trend \overline{Z} decreased from 4 nT to -2 nT 349 over the 04:00 UTC to 10:00 UTC period and then gradually increased from -2 nT to 1 nT. All components showed 350 fluctuations with amplitudes not exceeding ~1 nT. On 17 January 2022, the trend \overline{X} was first found to fluctuate within the 351 0-5-nT limits; after 10:00 UTC, some surges and drops attained 10-20 nT, and their durations did not exceed 1 h, whereas 352 the time variations in the Y- and Z-components showed significant, up to 10-20 nT, fluctuations.
- On the day the volcanic explosion occurred, the trend \overline{X} first decreased from 5 nT to -17 nT over the 06:00 UTC 353 to 09:00 UTC period, then it increased from -17 nT to 15 nT. A drop from 15 nT to -15 nT in the X(t) dependence was 354 observed to occur from 15:00 UTC to 17:00 UTC, while a clearly observed oscillation of the strength with 4-nT amplitude 355 and $T \approx 50$ -min period persisted over the 11:00–13:30 UTC period. The trend \overline{Y} increased from -10 nT to 18 nT over the 356 357 time interval from 04:00 UTC to 09:40 UTC, next its decrease to -35 nT continued to 15:00 UTC; finally, a ~50 nT surge in the trend persisted for ~1.5 h. The trend \overline{Z} decreased from 20 nT to 0 nT over the 04:00 UTC to 11:00 UTC period. From 358 11:00 UTC to 16:00 UTC, the Z(t) showed a drop from 0 nT to -20 nT, while a clearly observed oscillation with a 6-7-nT 359 amplitude and an ~80-min period lasted from 12:00 UTC to 17:00 UTC. 360









362 Figure 15: Same as in Figure 2 but for the KNY Station.



















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368 *DLT Station.* On 13 January 2022, all components showed insignificant (a few nanotesla) fluctuations in their 369 strengths (Figure 18) over the time period beyond the data gap. 17 January 2022, the trend \overline{X} decreased from 36 nT to -17 370 nT from 04:00 UTC to 12:00 UTC, then it exhibited fluctuations within the 4–6-nT limits. The trend \overline{Y} decreased from 17 371 nT to -7 nT after 03:30 UTC. Next its strength showed fluctuations within the ±(2–3)-nT limits. The trend \overline{Z} remained 372 almost constant from 03:00 UTC to 06:00 UTC, then it increased from -25 nT to 10 nT, and after 09:00 UTC it again 373 remained almost constant.

Throughout the day the volcanic explosion occurred, all components showed considerable variations. The trend \overline{X} experienced a bay reduction from ~22 nT to -25 nT over the 06:00-11:00 UTC period. Another drop in the $\overline{X}(t)$ dependence was observed to occur from 13:00 UTC to 17:00 UTC. The \overline{Y} showed significant and long-lasting disturbances from 04:00 UTC to 17:00 UTC. The trend \overline{Z} increased from -9 nT to 4 nT over the 04:20-05:30 UTC period, next a drop in \overline{Z} from 4 nT to 2 nT occurred, which was followed by an increase from 2 nT to 17 nT. After that a steep fall in the trend to 0 nT was first observed, and then a slow decrease from 0 nT to -5 nT. After around 15:30 UTC, the trend once again showed noticeable variations (~5 nT).

CKI Station. On 13 January 2022 used as a guiet time reference, the trend \overline{X} showed an insignificant rise to ~26 381 nT before approximately 05:00 UTC, after which the trend experienced a sharp fall from 26 nT to -10 nT and later was 382 followed by insignificant fluctuations in its magnitude (Figure 19). The trend \overline{Y} first decreased from -3 nT to -13 nT over 383 the 04:00–06:00 UTC period, next, from 06:00 UTC to 10:00 UTC, it experienced a steep rise from -13 nT to 13 nT that was 384 changed by a gradual reduction in the trend to 0 nT at 17:00 UTC. The trend \overline{Z} fell from 5 nT to -3.5 nT over the 03:30 385 UTC to 06:30 UTC period, whereas it showed two considerable surges, from -2 nT to 2 nT and from -2 nT to 1 nT, over the 386 09:00-12:00 UTC and 12:00-17:00 UTC periods, respectively. On 17 January 2022, used as a quiet time reference, the trend 387 $\overline{X} \approx 20$ nT from 03:00 UTC to 06:00 UTC. From 06:00 UTC to 10:00 UTC, it was observed to steeply fall from 20 nT to – 388 15 nT. Noticeable surges (by 10 nT to 15 nT) were observed over the 12:00 UTC to 16:00 UTC and 16:00 UTC to 17:00 389 UTC periods. The trend \overline{Y} sharply increased from -40 nT to 10 nT from 03:00 UTC to 06:00 UTC, and then, fluctuating, 390 gradually decreased from 10 nT to -10 nT at 17:00 UTC. The trend \overline{Z} increased from 0 nT to 15 nT over the 03:00–05:50 391 UTC period, after this it sharply decreased to -5 nT over a 3-h interval. After 09:00 UTC, the trend \overline{Z} showed fluctuations 392 393 within the $\pm(2-3)$ -nT limits.























398 On the day the volcanic explosion occurred, all components showed a significant enhancement in their variations. The trend \overline{X} first increased to 05:30 UTC and then sharply decreased from 20 nT to -10 nT after 06:00 UTC. Noticeable 399 increases in \overline{X} from -10 nT to 0 nT occurred during the 10:00-12:00 UTC period. From 13:00 UTC to 17:00 UTC, the 400 trend exhibited a drop from 0 nT to -27 nT. The trend \overline{Y} increased from -5 nT to 10 nT from 04:00 UTC to 07:40 UTC. 401 Next, from 07:40 UTC to 12:30 UTC, \overline{Y} experienced a bay reduction from ~8–10 nT to –8 nT. After 12:30 UTC, the trend 402 \overline{Y} was observed to decrease to -14 nT at 17:00 UTC. The trend \overline{Z} sharply decreased from 15 nT to -13 nT during the 403 03:30–09:30 UTC period. From 09:30 UTC to 15:00 UTC, the dependence $\overline{Z}(t)$ exhibited a surge from –13 nT to 5 nT. Yet 404 another surge in \overline{Z} to 7 nT was observed to occur from 15:00 UTC to 17:00 UTC. 405

GAN Station. On 13 January 2022, the trend \overline{X} first increased from -12 nT to 21 nT over the 03:00 UTC to 08:00 406 UTC period, and then decreased from 21 nT to -14 nT during the 08:00-17:00 UTC period (Figure 20). The trend \overline{Y} first 407 increased from -1 nT to 5 nT from 03:00 UTC to 04:00 UTC, then sharply decreased from 5 nT to -25 nT from 04:00 UTC 408 to 08:45 UTC, next increased from -25 nT to 13 nT over the 09:00-12:00 UTC period, and finally gradually decreased from 409 13 nT to 8 nT at 17:00 UTC. The trend \overline{Z} exhibited two considerable surges, from -4 nT to 5 nT over the 03:00 UTC to 410 06:00 UTC period, and from -4 nT to 11 nT from 08:30 UTC to 14:00 UTC. On 17 January 2022, the trend \overline{X} first 411 increased from -10 nT to 40 nT from 03:00 UTC to 07:00 UTC, next decreased to -20 nT at 12:00 UTC, and then exhibited 412 fluctuations within the ± 5 -nT limits. The trend \overline{Y} showed short-term (~1–2 h) increases by up to 4–5 nT, in addition to a 413 powerful surge (from -20 nT to 26 nT) during the 03:00-11:30 UTC period. The trend \overline{Z} exhibited a powerful surge from -414 3 nT to 27 nT from 03:00 UTC to 10:00 UTC; relatively small undulations of up to 5-6 nT were observed to occur after 415 10:00 UTC. 416

On the day the volcanic explosion occurred, all components exhibited noticeably enhanced variability. The trend 417 X first increased from -7 nT to 28 nT over the 03:00 UTC to 06:00 UTC period and then reduced from 28 nT to -3 nT 418 from 06:00 UTC to 10:30 UTC; in addition to variations within the \pm 5-nT limits, a drop from 5 nT to -28 nT was observed 419 to occur over the 13:00–16:30 UTC period. The trend \overline{Y} also first increased from -20 nT to 7 nT over the 03:00–06:30 UTC 420 period, and then a deep drop occurred from 7-18 nT to -8 nT over the 06:30 UTC to 14:25 UTC period. Over the 14:25-421 17:00 UTC period, the trend \overline{Y} decreased from 18 nT to 2 nT. The trend \overline{Z} first increased to 22 nT before 06:00 UTC. 422 Next, a deep drop (from 22 nT to -20 nT) followed over approximately 7 h. And finally, the moderate (up to 10-15 nT) 423 variations in \overline{Z} were observed to occur. It should be noted that synchronous geomagnetic bay disturbances were observed to 424 425 occur uncertainly at this most distant recording station.









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428 **7** Statistical data analysis of the bay excursions in geomagnetic field strengths

Table 4 shows the basic parameters of the bay disturbances of the geomagnetic field, viz., the magnitudes ΔX , ΔY , ΔZ , the time delays τ , and the durations ΔT of the northward component, *X*, of the eastward component, *Y*, and of the vertical component, *Z*, at nineteen stations. As can be seen from Table 4, the values of ΔX , ΔY , and ΔZ were most often negative except for the data from the API and PPT stations, which were located at a distance, *r*, of 840 km and 2,730 km, respectively, away from the volcano. The time delay showed a tendency to increase with increasing distance from the volcano, and the duration of disturbances exhibited the same tendency as the time delay. Table 4 shows that the strength of disturbances exhibits a tendency to decrease with increasing *r*.

-	C (-);	ΔX	τ_X	ΔT_X	ΔY	τ_Y	ΔT_Y	ΔZ	τ_Z	ΔT_Z
	Station	(nT)	(min)	(min)	(nT)	(min)	(min)	(nT)	(min)	(min)
-	API	15	16	90	28	16	146	-13	16	130
	PPT	8	45	90	8	45	100			
	EYR	-40	50	120	-25	50	120	-15	50	120
	CNB	-20	100	120	-50	60	178	-15	60	120
	CTA	-18	105	130	-63	60	150	-30	60	150
	MCQ	-80	100	180	-50	100	150	-30	80	150
	HON	-10	75	180	-5	75	180	-2	75	180
	ASP	-15	100	230	-50	75	240	-15	75	240
	KDU	-10	110	210	-30	110	210	-15	110	200
	IPM	-10	110	220	-8	110	220	-2	100	220
	GNG	-15	120	270	-50	120	270	-30	120	270
	LRM	-10	140	270	-10	140	265	-5	140	265
	KAK	-10	180	270	-8	165	260	-8	165	300
	KNY	-10	115	270				-8	120	270
	MMB	-10	165	240	-8	165	240	-8	165	240
	SHU	-10	105	220	-10	150	200	-10	150	200
	DLT	-20	165	240	-8	165	240			
	CKI	-12	170	240	-15	175	250	-10	165	240
	GAN	-10	240	240	-8	240	240	-15	240	240

436 Table 4. Basic parameters of bay disturbances in the geomagnetic field.





Figure 21 presents scatter plots of time delay versus distance from the volcano, which reveal the following linear dependences:

440	$\tau_X = 17.17r + 9.3, \sigma \approx 22.1 \min, R^2 \approx 0.83,$	(1)
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- 441 $\tau_Y = 19.93r 10, \, \sigma \approx 12 \, \min, R^2 \approx 0.96,$ (2)
- 442 $\tau_Z = 19.63r 12, \sigma \approx 14.1 \text{ min}, R^2 \approx 0.94,$ (3)

where distance is in Mm, time delay is in min, σ is a root mean square error, R^2 is an adjusted coefficient of determination. The individual points are fit with the following straight lines (Figure 22):

- 445 $\Delta T_X = 18.44r + 86.5, \sigma \approx 36.3 \text{ min}, R^2 \approx 0.68,$ (4)
- 446 $\Delta T_Y = 14.29r + 115.6, \, \sigma \approx 34.5 \, \text{min}, \, R^2 \approx 0.60,$ (5)
- 447 $\Delta T_Z = 15.63r + 109.8, \, \sigma \approx 39.5 \, \text{min}, \, R^2 \approx 0.56.$ (6)
- 448 The relations (1) (3) and (4) (6) indicate that the time delay and the duration of disturbance indeed increase with distance
- 449 from the volcano. The formation of disturbance is close to root mean square deviations in time delays, i.e., to 12–22 min.







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Figure 21: Time delay of bay disturbance in the geomagnetic field vs distance, *r*, from the volcano and the estimated regression line superimposed on the scatter plot.







Figure 22: Duration of bay disturbance in the geomagnetic field vs distance from the volcano and the estimated regression line superimposed on the scatter plot.



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456 Histograms showing the distributions of all ΔX , ΔY , and ΔZ are presented in Figure 23. The most probable values of 457 these disturbances are seen to be as follows: for the northward component $\Delta X = -(9.0 \pm 5.1)$ nT, for the eastward component 458 $\Delta Y = -(10.5 \pm 5.6)$, and for the vertical component $\Delta Z = -(6.3 \pm 3.1)$ nT, and $-(25.0 \pm 5.0)$ nT.



460 Figure 23: Histogram showing the distribution of excursions in bay disturbances in the geomagnetic field.





461 8 Statistical data analysis of the quasi-periodic variations in geomagnetic field magnitudes

The time delays of a possible response of the magnetic field to the volcanic explosion and the apparent speeds for six groups of characteristic variations in the components of the geomagnetic field are presented in Table 3, which show that the variations in the eastward component *Y* are seen most clearly. Figure 24 presents a scatter plot of time delay versus distance from the volcano for all the data presented in Table 3, which reveal the following linear dependences:

466
$$\Delta t_1 = 4.157r + 5.1, \qquad \sigma = 0.32 \text{ min}, \qquad R^2 = 0.9995,$$
 (7)

467
$$\Delta t_2 = 11.14r + 4.6, \qquad \sigma = 0.55 \text{ min}, \qquad R^2 = 0.9998,$$
 (8)

468
$$\Delta t_3 = 16.66r + 4.6, \qquad \sigma = 0.47 \text{ min}, \qquad R^2 = 0.9999,$$
 (9)

469
$$\Delta t_4 = 33.13r + 4.6, \qquad \sigma = 1.60 \text{ min}, \qquad R^2 = 0.9998,$$
 (10)

470
$$\Delta t_5 = 53.11r + 6.1, \qquad \sigma = 9.98 \text{ min}, \qquad R^2 = 0.9969,$$
 (11)

471
$$\Delta t_6 = 82.97r + 7.7, \qquad \sigma = 2.61 \text{ min}, \qquad R^2 = 0.9999.$$
 (12)

472 If $r \rightarrow 0$, then $\Delta t_0 \approx 4.6-7.7$ min. Such a time interval is needed for the wave to reach ionospheric heights, or more precisely, 473 *E* region dynamo heights.

474 Use of relations in Eqs. (7) - (12) and the formula given by

$$475 \quad v = \left(\frac{d\Delta t}{dr}\right)^{-1}$$

476 yields the following average speeds: $v_1 \approx 4$ km/s, $v_2 \approx 1.5$ km/s, $v_3 \approx 1$ km/s, $v_4 \approx 503$ m/s, $v_5 \approx 314$ m/s, and $v_6 \approx 209$ m/s. 477 These values are close to the values inferred from the histograms in Figure 25.

478 The horizontal apparent speed of propagation of disturbances can be estimated from the following relation:

$$479 \quad v = \frac{r}{\Delta t - \Delta t_0}$$

480 where Δt_0 is the time taken for the blast wave to travel from the volcano to the *E* region dynamo.







Figure 24: Time delay of the onset of quasi-periodic disturbances in the geomagnetic field vs distance from the
volcano and the estimated regression line superimposed on the scatter plot.



484





Figure 25: Histogram showing the distribution of the apparent speeds of propagation of quasi-periodic disturbances
in the geomagnetic field.





487 9. Discussion

488 Bay disturbances of the geomagnetic field. During the day the Tonga volcanic explosion occurred, all three components usually exhibited geomagnetic bay disturbances whose absolute values were observed to be 10-60 nT, and the disturbances 489 490 themselves were more often seen to be negative. The eastward component Y experienced the largest disturbances, with the 491 average value of -53 nT, whereas disturbances in the X and Z components were observed to be, on average, -15 nT. The 492 smallest disturbance took place at the PPT station. The geomagnetic bay disturbances were virtually absent, or, more precisely, did not exceed -(2-8) nT also at the IPM station, which was located east of the volcano as well. This can be 493 explained by the location of these stations on the night side of the Earth where the electron and electric current densities 494 495 were approximately an order of magnitude smaller than in the sunlit ionosphere.

496 Disturbances were insignificant and unclear at the GAN station, the most distant station included in this study.

It should be stressed that the bay variations in the magnitudes of all geomagnetic field components did not exceed 5–10 nT during the days used as a quiet time reference. This observation supports the idea that the geomagnetic bay disturbances observed on 15 January 2022 were due to the volcanic explosion. However, this is a necessary but not sufficient condition for the volcanic explosion to be the cause of the effect.

A sufficient condition is a tendency for the time delay and duration of bay disturbance to grow with distance from the volcano, while a tendency for the disturbance strength was to decrease with distance from the volcano (see Figures 21 and 22).

The relations in Eqs. (1) – (3) suggest that, in the limit $r \rightarrow 0$, the minimum in the time delay, τ_{\min} , is determined by root mean square error in the approximation, which is close to 14–22 min for the *X*, *Y*, and *Z* components. The disturbance from the volcano takes such a time interval to travel from the volcano to the *E* region dynamo, $z \approx 90-150$ km altitude, and to generated magnetic disturbance. The relations in Eqs. (1) – (3) permit estimates of the average speeds of the disturbances to be made using the relation given by

$$509 \quad v = \left(\frac{d\tau}{dr}\right)^{-1}$$

510 Then, $v_X \approx 970 \pm 235$ m/s, $v_Y \approx 836 \pm 103$ m/s and $v_Z \approx 849 \pm 121$ m/s. These magnitudes of the speeds are close to the blast 511 wave speed [Chernogor, 2023b; Chernogor, 2023c].

512 It is important that the magnitudes of the speeds obtained are close to the speed of propagation of the disturbances 513 in the electron density, *N*, and in the total electron content [Chernogor, 2023a]. This means that the formation of the 514 ionospheric hole is the cause of the bay excursions in the geomagnetic field [Chernogor, 2023a].

Estimation of the magnitude of a bay disturbance in the geomagnetic field may be performed from the average daytime value of N in the E region dynamo of $(2-3) \times 10^{11}$ m⁻³ and a neutral wind speed of $w \approx 100$ m/s. Then, the electric current density in the ionosphere is given by





518 $j_0 = eNw \approx (3.2-4.8) \times 10^{-6} \text{ A/m}^2$

where *e* is the charge of an electron. The disturbance in *N* within the ionospheric hole is estimated to be 5–20%, which yields the perturbation in the ionospheric current of $\Delta j \approx (1.6-9.6) \times 10^{-7}$ A/m². The estimate of the disturbance in the magnetic field follows from Maxwell's curl equation, given by

```
522 \Delta B \approx \mu_0 \Delta j \Delta z
```

(13)

523 where μ_0 is magnetic permeability, $\Delta z \approx 50$ km is the thickness of the dynamo region. Substituting the numerical magnitudes 524 yields $\Delta B \approx 10-60$ nT, which is in excellent agreement with observations (~10-60 nT).

525 Thus, there is every reason to believe that the bay disturbances of the components of the geomagnetic field are 526 related to the generation of the ionospheric hole as a result of the explosion of Tonga volcano.

527 The effect of atmospheric acoustic resonance. The station nearest to Tonga volcano is the API Station. The Y 528 component exhibited the first perturbation over the 04:21–04:57 UTC period, i.e., the time delay was observed to be $\Delta t_0 \approx 6$ 529 min. The acoustic wave take such a time interval to travel to the ionospheric *E* region where dynamo electric fields are 530 generated and where the generation of this magnetic effect occur. The sound wave is reflected at an altitude of $z_r \approx \overline{v}_x \Delta t_0 \approx$

531 110–120 km (where $\overline{v}_x \approx 300-330$ m/s is an average speed of sound), i.e., in the *E* region dynamo. It is important that the 532 period of the disturbance $T_0 \approx 4-4.5$ min and its duration $\Delta T_0 \approx 32-36$ min. These values indicate that the magnetic effect 533 have been generated by the atmospheric acoustic resonance in the Earth–*E* region dynamo cavity, where the volcanic 534 explosion excited the vibrations.

535 Since the API station is located at a range of ~840 km from the volcano, the radius, r_L , of the footprint of the magnetic flux tube associated with the volcano is equal or greater than 1,000 km. This means that the magnetic effect from 536 537 the atmospheric acoustic resonance could be observed in the magnetically conjugate region. Indeed, oscillations with the same period, T_0 , duration ΔT_0 , and ~0.2-nT amplitude, were observed by [Iyemori et al., 2022; Yamazaki et al., 2022]. It is 538 important that the time delay was equal to $\Delta t_0 \approx 6$ min. This means that the disturbance from the API station was transferred 539 540 to the HON station along the magnetic flux tube ~10 Mm long at an Alfvén speed, v_A , of ~1 Mm/s for ~10 s, which much 541 shorter than Δt_0 . It should be noted that the HON station is located about 900 km from the center of the magnetic flux tube, and $r_L > 900$ km. 542

Quasi-periodic disturbances. Other disturbances with other time delays were superimposed on the disturbance due to acoustic resonance (see Table 3). In total, the number of such disturbances could be six. Table 3 shows that six groups of disturbances in the geomagnetic field also took place at other stations. It is important that the time delay increases with distance from the volcano. The premise of requiring the time delays of the magnetic disturbances due to the volcanic explosions to explain our observations is clearly supported by the INTERMAGNET Magnetometer Observatory data.

548 The values of the speeds were close to 4 km/s, 1.5 km/s, 1 km/s and 500 m/s, 313 m/s, and 200 m/s. All these 549 speeds have physical significance. The first and second group of speeds correspond to the speeds of the fast and slow MHD





waves [Sorokin and Fedorovich, 1982]. Approximately the same speeds were observed during powerful rocket launches 550 551 [Chernogor, 2009; Chernogor and Blaunstein, 2013]. The speed of $v_3 \approx 1$ km/s is characteristic of blast waves. This speed was revealed after the Tonga volcanic explosion by [Matoza et al., 2022a; Matoza et al., 2022b]. The speed v_4 pertains to 552 atmospheric gravity waves at ionospheric heights [Chen et al., 2022; Themens et al., 2022]. Lamb waves that are generated 553 554 by massive releases of energy (exceeding 10 Mt of TNT) propagate at a speed of $v_5 \approx 313$ m/s over the Earth's surface 555 virtually without damping and partially penetrate to ionospheric heights along their propagation paths [Chernogor, 2022a; Chernogor, 2022e; Chernogor, 2023a; Kubota et al., 2022; Lin et al., 2022; Zhang et al., 2022a]. The smallest speed of $v_6 \approx$ 556 200 m/s probably pertains to an average speed of tsunami, which was observed after the volcanic eruption and generated 557 558 ionospheric disturbances [Carvajal et al., 2022; Ramírez-Herrera et al., 2022; Tanioka et al., 2022; Terry et al., 2022].

559 *Estimation of the quasi-periodic effects.*

560 The amplitude of quasi-periodic disturbances usually showed variations not exceeding 1-3 nT. Such disturbances were 561 generated by quasi-periodic disturbances arising in the electric current density at *E* region dynamo heights from the action of 562 waves launched by the volcanic explosion.

563 The difference, w_m , in the drift velocities of ions and electrons, which are driven by the drag force of the neutral 564 atmosphere, causes the dynamo current density given by the relation

565 $j = eNw_m$.

566 The integrated in altitude current density is given by

567
$$J = \int_{\Delta z} j(z) dz$$

568 Then, the amplitude of the quasi-periodic disturbance in the geomagnetic field is given by the following expression:

569
$$\Delta B_a \approx \mu_0 J.$$

570 If $N \approx (2-3) \times 10^{11} \text{ m}^{-3}$ on the sunlit side of the Earth, and $w_m \approx 0.3-1.5 \text{ m/s}$, then $j \approx (1-7.2) \times 10^{-8} \text{ A/m}^2$, $J \approx (4.8-36) \times 10^{-5}$ 571 ⁴ A/m, and $\Delta B_a \approx 0.6-4.5 \text{ nT}$. These estimates are seen to be close to magnitudes observed (~1-3 nT).

Thus, the disturbances in the geomagnetic field described above were observed on 15 January 2022 and were absent during the days used as a quiet time reference. Consequently, they were most probably due to the volcanic eruption. These disturbances were transported by the waves of various physical nature, viz., the fast and slow MHD waves, blast waves, atmospheric gravity waves, Lamb waves, and ionospheric waves that arises from the tsunami.

576 8 Conclusions

577 Analysis of the data acquired at nineteen INTERMAGNET magnetic observatories revealed the following.





(1) During the day of the Tonga volcanic explosion, the variations in the magnitude of all components of the geomagnetic field varied less monotonically than during the days used as a quiet time reference. The strength of fluctuations also enhanced. All these factors indicated that the volcanic explosion led to the registered magnetic effect.

581 (2) The geomagnetic bay disturbances in all components of the geomagnetic field were observed to occur with a time delay increasing with distance from the volcano from a few tens of minutes to 100-200 min. The magnitude of the 582 effect changed from ~ 10 nT to ~ 60 nT. The eastward component (Y) exhibited the greatest variations. The time delay and 583 584 duration of the disturbances increased with distance from the volcano, but amplitudes of the disturbances, instead, decreased. The speed of propagation of the bay disturbances was close to the speed of the blast waves, approximately 700–1,000 m/s. 585 Geomagnetic bay disturbances were weakly expressed or were virtually absent on the night side of the planet. The premise 586 587 that the geomagnetic bay disturbances are closely related to the volcanic blast wave-induced formation of the ionospheric 588 hole has been validated.

(3) The quasi-periodic disturbances in the geomagnetic field arrived at the magnetic observatories with different time delays. Six main groups of disturbances were identified. It is important that the time delay increases with distance from the volcano in each group. The apparent speeds of propagation of the disturbances in each group have been estimated, and the values of these speeds are as follows: 4 km/s, 1.5 km/s, 1 km/s and 500 m/s, 313 m/s, and 200 m/s. The first two speeds pertain to the fast and slow MHD waves, the third to the blast wave, the fourth to the atmospheric gravity wave, the fifth to the Lamb wave, and the six speed pertain to the tsunami.

595 (4) The magnetic effect due to the atmospheric acoustic resonance in the Earth – *E* region dynamo cavity where 596 vibrations were excited by the volcanic explosion was observed at API Station, the nearest to Tonga volcano. The period of 597 the disturbance was estimated to be $T_0 \approx 4$ -4.5 min, the amplitude to be 2 nT, and its duration to be $\Delta T_0 \approx 32$ -36 min. 598 Similar effect was observed in the magnetically conjugate region at the HON station; however, its amplitude was an order of 599 magnitude smaller.

600 (5) Estimates of the bay and quasi-periodic disturbances are in good agreement with the parameters of disturbances601 inferred from INTERMAGNET data.

602 Competing interests

603 The contact author has declared that none of the authors has any competing interests.

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614 Data Availability Statement.

615 The data sets discussed in this paper are freely accessible on the Internet at <u>https://www.intermagnet.org</u>.

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