Responses to Editor and Reviewer

General Comments:

Dear Dr. Ana G. Elias.

First of all, thank you for consider our paper for revision. We also thank to the reviewers for the important insights presented during their revisions. We have done our best to address the concerns of both reviewers properly. Our point-by-point responses are listed as following and the necessary modifications were highlighted in the the end of this file.

Reviewer #1:

REVIEWER: "The variability in the amplitudes of the lunar semidiurnal tide is investigated using TEC maps over Brazil from January 2011 to December 2014. The authors find evidence of strong annual variation. Semiannual and intra-seasonal oscillations are found to be the second and third largest components, respectively. Among the short-period oscillations in the amplitude of the lunar tide, the most pronounced ones were concentrated between 7-11 days, which the authors ascribe to the normal mode westward propagating quasi 10 days planetary wave with horizontal wavenumber equal to 1. The presented results suggest a possible coupling process by modulation of the lunar semidiurnal tidal amplitudes that allows the propagation of the 7-11 days waves into the thermosphere-ionosphere system."

AUTHORS: We are grateful for the dedicated time in reading and suggesting improvements to our paper.

REVIEWER: "While the manuscript contains some interesting results, I cannot recommend publication in the present form for the following two reasons: (1) The language should be improved. (2) Additional observational and/or modeling work is needed to demonstrate that the \sim 9-day oscillation is indeed consistent with a westward quasi-10-day normal mode."

AUTHORS: The reviewer is right that the results are not conclusive at all. We have revised the manuscript excluding the non-conclusive results. We have tried to investigate the horizontal propagation of the wave observed in the amplitude of the lunar semidiurnal tide using data from stations separated by $\sim 3,600$ km, but the results were not conclusive as well. We have tried to use global TEC maps from CEDAR database to investigate the presence of 8.5 days directly in the TEC. As the amplitude of the Q8D oscillation is small compared to the other oscillations, the phases were not well resolved and we could not conclude about the propagation direction (eastward or westward). Maybe we need use specific receivers located in Brazil, Africa, India and Indonesia to be reach confident results, but we understand that this is out of the scope of this manuscript and we will not have enough time to do this kind of analysis within the schedule of publication of ANGEO.

Therefore, we kindly ask to the reviewer to check whether the modification implemented in the manuscript can address his/her concerns. In addition, we have included analysis of the wind data from FPIs. The results showed the presence of quasi 6 days oscillations which can be another possibility to explain the observed results. Even without a conclusion about the real coupling mechanism that are producing the Q8D oscillation in the amplitudes of the lunar semidiurnal tide in the TEC, we have to agree that the results are new and important to the scientific community and certainly it will start a relevant debate on this topic.

We have also revised the language according to the suggestions from the reviewer.

REVIEWER:

"Line 4 (and throughout): Intra-seasonal variability is usually referred to variations less than 90 days."

"Line 6: in special -> in particular"

"Lines 19-20: need reference"

"Line 22: Ozone - > ozone"

"Lines 26-27: these ranges are not consistent with other studies. Need reference."

"Lines 48-49: define what is meant with long and short period."

"Line 60: reference for the filter needed."

"Line 91: 30 days or 25 days as reported in the legend of Figure 4?"

"Line 131: maiximuma -> maxima."

"Line 161: thermosphere misspelled."

"Line 171: to notice - > to note."

"Line 201: preset -> present."

"Line 208: though -> through."

AUTHORS: All those minor points have been corrected according to the suggestions.

REVIEWER: "Lines 9-10 (see comment before): This sentence and result is highly speculative. More modeling or observational work should done to demonstrate a link to the 10- day normal mode."

AUTHORS: We agree with the reviewer and we have revised it.

REVIEWER: "Section 2: More details on how the lunar tide is calculated are needed."

AUTHORS: We have added a subsection to explain the determination of the lunar tide as suggested.

REVIEWER: "Line 77-78 and Figure 3: need to discuss the 70- to 80-day variations and 120 variations."

AUTHORS: We have added the description of these oscillations as mentioned.

REVIEWER: "Wavelet plots: need to include a confidence level."

AUTHORS: We have included it. Thank you for the suggestion!

REVIEWER: "Lines 133-134: the oscillation at 70-80 days is almost at large. Need to comment on this."

AUTHORS: Thank you. We have commented it.

REVIEWER: "Lines 144-148: this statement is highly speculative. Need additional modeling work or analysis of concurrent observations at different longitudinal locations."

AUTHORS: We have also revised this statement. Thank you for this contribution.

REVIEWER: "Lines 192-194: can the observed intra-seasonal variability be related to Madden-Julian Oscillation?"

AUTHORS: It is very difficult correlate these events without further analysis.

REVIEWER: "Lines 197-198: need to further elaborate this point."

AUTHORS: We have removed the analysis about the symmetry of these oscillations because it was discussed considering the magnetic latitudes instead of geographic ones. We are not sure how is the behavior of the symmetry of the planetary waves using magnetic coordinates, primarily in Brazil where there is a strong magnetic declination.

REVIEWER: "Lines 209-2010: analysis insufficient to support this statement."

AUTHORS: Yes, we have revised it.

Reviewer #2:

REVIEWER: "The authors examined seasonal and intraseasonal variability of the semidiurnal lunar tide in TEC over Brazil. The main finding is that the amplitude of the semidiurnal lunar tide in TEC often shows 7-11 day variations. The authors speculate that these variations are associated with the quasi-10-day wave in the middle atmosphere."

AUTHORS: We appreciate the revision and the contributions from the Reviewer # 2. We have done our best to address all of the concerns from the reviewer.

REVIEWER: "Although the results are interesting, I am not totally convinced that the authors were able to extract 7-11 day variations of the semidiurnal lunar tide. The authors used the technique of Paulino et al. (2017) to derive the amplitude of the semidiurnal lunar tide. This technique involves a 27 day window, which enables to distinguish between the semidiurnal lunar tide (12.42h) and semidiurnal solar tide (12.00h). The technique should largely eliminate variations with periods less than 27 days, even though the amplitude is calculated for each day. Thus, it is unclear whether the presented short- period variations are meaningful."

AUTHORS: We thank the reviewer for this important comment. The TEC in the tropical region is mainly produced by the absorption of the EUV and X-rays solar radiations. Thus the diurnal cycle is faraway dominant and should be removed. The determination of the lunar semidiurnal tides in TEC maps were done according to the Pedatella and Forbes (2010) methodology and only quiet days were considered ($K_p < 3$) in the analysis. Figure 1 (upper panel) shows a 29.5 days window of TEC from 27 July 2011 to 25 August 2011 measured at (15°S, 39°W).

After the elimination of geomagnetic influences, a Fourier analysis was performed

to extract the subharmonics of the solar day (diurnal, semidiurnal and terdiurnal oscillations). Effects of the solar rotation was removed using a 27-day window moving it forward one day at time to calculate the mean solar day centered in the window as can be seen in Figure 1 (middle panel). In addition, residual TEC was determined subtract the original TEC from the recovered one. Figure 1 (bottom panel) shows the residual TEC for this example, where the power of the diurnal cycle is reduced and other oscillations can be observed.



Figure 1: (Top panel) Original TEC from 27 July 2011 calculated at (15°S, 39°W). (Middle panel) recovered signal using sub-hamornics of the solar day within a 27 days window. (Bottom panel) Residual TEC.

In the relative residual data (residual TEC divided by the mean TEC), a least square analysis in a window of 29-day was applied using the following equation:

$$y(\tau) = \sum_{n=0}^{2} A_n \cos\left(n\tau + \phi_n\right) \tag{1}$$

where τ if the lunar time given by $\tau = t - \nu$, ν in the age of the Moon, which is set to be 0 at the New Moon. The solar time is represented by t, the amplitudes and phases of the lunar tide components are represented by A_n and ϕ_n , respectively.

We have included this explanation in the manuscript in order to clarify the methodology. It was requested by Reviewer #1 as well. Furthermore, the scope of the present work was investigate the day-to-day variability of the amplitude of the lunar semidiurnal tide, which was calculated for each day and change as well as shown in Pedatella and Forbes (2010) and Paulino et al. (2017). Additionally, the determination of the short-period oscillations were statistically significant as in the LS periodogram as in the wavelet analysis (the significance level were included in all plots).

REVIEWER:" The authors are advised to check the spectrum of the original TEC data (instead of the spectrum of the semidiurnal lunar tide) to confirm that a spectral peak exists at the semidiurnal lunar tide (12.42h) as well as the sideband frequency corresponding to the quasi-10-day wave modulation of the semidiurnal lunar tide."

AUTHORS: We thank to the reviewer for this comments. Regarding to the presence of the lunar tide in the TEC, Figure 1 of Paulino et al. (2017) shows very clear evidences in different periods.

On the other hand, we have followed the suggestion from Reviewer #2 to show the presence of the Q8D wave in the TEC. Figure 2(a) shows the data of the quiet day TEC maps for November 2013 (when the Q8D wave was strong in the amplitude of the semidiurnal lunar tide) at 8°S, 35° W (where there are confident GNSS receivers and the amplitude of the amplitude of the Q8D was strong). One can see that there is a strong day-to-day variability in the TEC. Figure 1(b) shows the Lomb-Scargle periodogram calculated using the data from Figure 1 (a). The diurnal cycle is very pronounced compared to the other oscillation. Figure 1(c) shows the same results of Figure 1(b) but ranged from 0 to 50 PSD units. One can see that the Q8D wave peak is above of the confidence level and it demonstrates what was suggested by the reviewer.



Figure 2: (a) TEC data calculated to November 2013 at (8°S and 35°W) (b). Lomb-Scargle periodgram for the data TEC data shown in panel (a). (c) Same as Figure (b), but for zoomed to the y-range from 0 to 50 PSD.

REVIEWER: "1. Equation (1) - This needs more explanations. What is "filter" on the left-hand side? How is it applied to the data?"

AUTHORS: Thank you for the comment. We have included a citation about the application of the filter as suggested by the Reviewer #1. To apply the filter, first we apply the FFT transform, then we multiply the "filter" by the FFT signal. Finally we apply the inverse FFT to recover the filtered signal.

REVIEWER: "Lomb-Scargle periodogram - Since the authors show wavelet spectra in Figures 3 and 5-8, Lomb-Scargle periodograms in Figures 2 and 4 do not seem necessary. I suggest to remove them." AUTHORS: The reviewer is right! Most of the aspects showed in Figure 2 and 4 can be seen in the wavelet charts. However, LS periodograms can give us a general idea about the periodicities using the whole period of analysis and comparisons between the latitudes are easily matched. Even so, if the reviewer thinks better to remove them, we can do it for the revised version.

REVIEWER: "Figure 9 - The antisymmetric mode such as the quasi-10-day wave has the phase structure that is antisymmetric about the equator, but not the amplitude structure. That is, when there is a strong quasi-10-day wave, we should expect the amplitude of the wave to be large in both northern and southern hemispheres but with the opposite phase. What is shown in Figure 9 is the anti-correlation of the amplitude between the northern and southern hemispheres, which does not necessarily support the involvement of the quasi-10-day wave."

AUTHORS: The reviewer is right! Figure 9 does not necessarily support the antisymmetry. Thank you for the comment. Besides, we are not sure about how is the symmetry of planetary waves regarding the magnetic coordinates. Therefore, we decide to remove these analysis. Thank you for this contribution.

Variability of the lunar semidiurnal tidal amplitudes in the ionosphere over Brazil

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Abstract. The variability in the amplitudes of the lunar semidiurnal tide was investigated using maps of Total Electron Content over Brazil from January 2011 to December 2014. Long period variability showed that the annual variation is always present in all investigated magnetic latitudes and it represents the main component of the temporal variability. Semiannual and intraseasonal terannual (~120 days) oscillations were the second and third components, respectively, but they presented

- 5 significant temporal and spatial variation variability without a well-defined pattern. Among the short period oscillations in the amplitude of the lunar tide, the most pronounced ones were concentrated between 7-11 days. These oscillations were stronger around the equinoxes, in special particular between September and November in almost all latitudes. In some years, as in 2013 and 2014, for instance, they appeared with large power spectral density in the winter hemisphere. There was also observed evidence of antisymetry in the amplitudes maxima and minima of the 7-11 days oscillation with respect to the magnetic
- 10 equator. These characteristics are compatible with normal mode westward propagating quasi 10 days planetary wave with horizontal wavenumber equal to 1. Besides, using data from a meteor radar located at low latitudes in Brazil for November 2013, when the amplitude of the 7-11 days oscillation was strong, it was possible to identify the presence of quasi 10 days oscillation in the both zonal and meridional component of the horizontal winds. These results suggest a possible coupling process by modulation of the lunar semidiurnal tidal amplitudes that allows the propagation of the 7-11 days waves into the
- 15 thermosphere-ionosphere system. These observed short period oscillations could be a result of a direct modulation of the lunar semidiurnal tide by planetary waves from the lower atmosphere or/and due to electrodynamics coupling of E and F region of the ionosphere.

1 Introduction

Planetary waves are produced by large-scale perturbations in the atmosphere that can have horizontal wavelength up to 40.000
20 40,000 km around the equator. Those waves can have periods which vary from a couple of days to weeks. Planetary waves are responsible for most of the temporal and spatial variation in the stratosphere and they also contribute substantially to the variability of the mesosphere and lower thermosphere (MLT). Basically, planetary waves have been classified in three types: (1) quasistationary midlatitude Rossby waves, (2) normal modes and (3) equatorial waves (Smith and Perlwitz, 2015).

Quasistationary Rossby waves are important to the midlatitude dynamics because they can largely influence the atmospheric

- 25 fields as-like wind and temperature and they are responsible for the distribution of the Ozone-ozone and other trace gases. Rossby normal modes, also known as free modes are predicted by the theory as oscillatory solutions of the Laplace's tidal equation without forcing. The Laplace's theory is constructed over an isothermal and non-damping atmosphere, thus, the real conditions of the real atmosphere can produces atmosphere can produce normal modes with some similarities to the theoretical ones. The class of planetary waves which occur near the equator and the most commonly observed in the MLT region are the
- 30 Kelvin waves, which are classified as low Kelvin waves (periods of $\sim 16-10-15$ days), fast Kelvin waves (periods of 6-7-6-10 days) and ultrafast Kelvin waves (periods of 3-4) days (Chen and Miyahara, 2012).

Dissipative processes in the atmosphere act significantly in the upward propagation of planetary waves in the atmosphere producing a pronounced damping above 100 km altitude. Among several mechanisms, the cooling by emission of heat and interaction with small-scale waves have been pointed out as the most important dissipative processes (e.g., ?, and references therein)

- 35 (e.g., Smith and Perlwitz, 2015, and references therein). However, large number of studies in the last decadeshave shown evidence for-, large number of studies have shown evidences of oscillations with periods of compatible with planetary-waves in the thermosphere (e.g., Forbes, 1996; Pancheva and Laštovička, 1998; Pancheva et al., 2002; Laštovička, 2006; Abdu et al., 2006, 2015; Jonal thermosphere-ionopshere (e.g., Forbes, 1996; Pancheva and Laštovička, 1998; Pancheva et al., 2002; Laštovička, 2006; Abdu et al., 2006, . Understanding how planetary waves can penetrate into the thermosphere-ionopshere system have raised up as one of the cur-
- 40 rent most important topic topics of research in the atmospheric layer's coupling. Recent works have given some insights in this topic (e.g., Forbes et al., 2014; Gasperini et al., 2017), but further observation observations and investigations are necessary in order to understanding this coupling.

The lunar semidiurnal tide with period of \sim 12.424 solar hours is the most important Moon's oscillation for the atmosphere in terms of amplitudes. Although the generation of the lunar semidiurnal tides comes from the lower levels of the atmosphere

- 45 due to the Moon's gravitational attraction and interaction with vertical motion of the oceans and solid Earth, it can propagate into the thermosphere with less influence of the dissipative process. As the source of the lunar tide is well known, variation associated to the sources are predictable. Then, it can be used as an important trace to observe changes in the atmosphere as it propagate propagates vertically. Furthermore, modulation of the lunar semidiurnal tidal amplitudes by planetary waves can be used to explain the presence of these waves in the thermosphere-ionosphere system, which is the main purpose of the present work. Additionally, variability of long period in the lunar semidiurnal tidal amplitudes is also investigated.
 - Data from a <u>network of</u> Global Navigation Satellite System (GNSS) <u>network receivers</u> over Brazil was used to calculate the amplitudes and phases of the lunar semidiurnal tide in the Total Electron Content (TEC) of the ionosphere from 2011 to 2014 (Paulino et al., 2017). In the present work, the temporal variability of the lunar semidiurnal tidal amplitudes was extensively investigated showing long and short (> 60 days) and short (< 25 days) period oscillations.

Analysis and results 2 55

Determination of the lunar tide in TEC maps 2.1

The determination of the lunar semidiurnal tides in TEC maps was done according to the Pedatella and Forbes (2010) methodology and only quiet days were considered ($K_p < 3$) in the analysis. After eliminate the geomagnetic influences, a Fourier analysis was performed to extract the subharmonics of the solar day (diurnal, semidiurnal and terdiurnal oscillations). Effects of the

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solar rotation was removed using a 27-day window moving it forward one day at time to calculate the mean solar day centered in the window. In addition, relative residual was determined dividing the residual variation of TEC by the TEC average. In the relative residual data, a least square analysis in a window of 29-day was applied using the following equation:

$$y(\tau) = \sum_{n=1}^{3} A_n \cos\left(n\tau + \phi_n\right) \tag{1}$$

where τ if the lunar time given by $\tau = t - \nu$, ν in the age of the Moon, which is set to be 0 at the New Moon. The solar time is represented by t, the amplitudes and phases of the lunar tide components are represented by A_n and ϕ_n , respectively. 65

2.2 Filtering

Further description of the methodology to calculate the TEC maps over Brazil was provided by Takahashi et al. (2016). TEC maps have also been used to calculate the amplitude and phases of the lunar semidiurnal tide from 2011 to 2014 over Brazil (Paulino et al., 2017). In the present study, the variability due to low and high frequencies in the lunar semidiurnal tide amplitudes observed in those TEC maps is investigated in details.

Figure 1 shows the filtering process in the amplitudes of the lunar semidiurnal tide calculated at 10° S (magnetic).



Figure 1. (a) Amplitudes of the lunar semidiurnal tide from 2011 to 2014 calculated at 10° S (magnetic latitude). (b) Low frequencies calculated using a Butterworth kernel filter from amplitudes of panel (a). (c) Same as Figure (b), but for high frequencies.

Figure 1(a) shows the raw amplitudes in TEC units from 2011 to 2014. One can see that there are low and high frequency oscillations in the amplitudes retrieved from the TEC. Figure 1(b) shows the filtered amplitudes considering periods greater

than 30 days and Figure 1(c) shows the high frequencies greater than 1/30 days⁻¹. This filtering process was done using 75 Butterworth Kernel low pass filter of order 1. Mathematically this filter can be written as:

$$filter = \frac{1}{\sqrt{1 - \left(\frac{\Omega}{\Omega_c}\right)^{2n}}}\tag{2}$$

where Ω is the frequency, Ω_c is the cutoff frequency and *n* is the order –(Roberts and Roberts, 1978). This filter is applied to the signal in the domain of the frequency and then, the filtered signal is recovered to the domain of the time.

2.3 Seasonal and intraseasonal terannual variability

80 Considering the filtered amplitudes for periods longer than 30 days, a spectral analysis was done and the results are shown in Figure 2. The Lomb-Scargle periodogram (Lomb, 1976; Scargle, 1982) was periodograms (Lomb, 1976; Scargle, 1982) were calculated using the filtered amplitudes from 2011 to 2014 for the magnetic latitudes 10° N, 0°, 10° S and 20° S.



Figure 2. Lomb-Scargle periodogram for (a) 10° N, (b) 0° , (c) 10° S and (d) 20° S. The horizontal dashed lines represent the confidence level of 99%, i.e., false alarm probability of 0.01.

Figure 2(a) shows strong the power density spectrum (PSD) associated to annual variation and semiannual variation variation the annual and semiannual variations. One can also observe that there is a third peak around 120 days, i.e., intraseasonal terannual

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variation. Similar patterns to what is observed at 10° N can also be observed at 10° S (Figure 2(c)) and 20° S (Figure 2(d)) as well. At 0° (Figure 2(b)), annual and semiannual variations were strongas well, but the intraseasonal, additionally the terannual variation was weak compared to the other latitudes. Comparing 10° N to 10° S, it is clear that there are more significant peak of oscillation in the South, indicating that the long oscillations are not symmetrical with respect to the magnetic equator.

In order to investigate when the periodicities shown in Figure 2 appear more frequently in the dataset, a wavelet analysis was

90 performed and the results are shown in Figure 3 with the respective magnetic latitudes. These wavelet charts were calculated based on the methodology of Torrence and Compo (1998).

Figure 3 shows that the annual variation is always present in the amplitude of the lunar tide. Figure 3(a) shows that the semiannual variation was present in the two first years and the intraseasonal terannual variation appear more pronounced in the beginning of 2013and 2014., which can be composed by oscillations from 80 to 120 days. In the beginning of 2014, the terannual variation appeared as well.

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Figure 3(b) (magnetic equator) shows that the semiannual variation were strong in the end of 2012 and beginning of 2013 with outspread of this peak to over 200 days. Figure 2(b) shows also this behavior in the Lomb-Scargle chart. One can also observe short oscillations of 70-80 days in the beginning of 2013 and 2014.

Figure 3(c) shows that the semiannual oscillation in the amplitude of the lunar tide becomes stronger than the annual in the 100 second half of 2013. It is important to observe that the intraseasonal terannual oscillation was present in the amplitude of the lunar tide from 2011 to March 2013 and became very strong at the end of this time range, compared to the other latitudes and times. Figure 3 (c) shows also oscillation with periods shorter than 100 days along the whole observed time.

Figure 3(d) shows that the intraseasonal terannual oscillation appeared in 2013 and 2014 and, at this magnetic latitude, the semiannual oscillation was weaker than the intraseasonal, terannual. Oscillations of 70-80 days have the same occurrence

105 observed at 10° N.



Figure 3. Wavelet analysis for (a) 10° N, (b) 0° , (c) 10° S and (d) 20° S. The heavy black lines in the contours represent the cone of influence. The light black lines show confidence levels of 95%.

2.4 Short period oscillations

The short period oscillations observed in the amplitudes of the lunar tide were the main focus of this work, primarily, because these oscillations also investigated in this work. These oscillation are important because they can be associated to the with planetary waves revealing important relevant aspects in the atmosphere-ionosphere coupling from below.

- Figure 4 shows the Lomb-Scargle periodogram for the same latitudes as used in Figure 2, but considering only periods shorter than 30.25 days. These periodograms were calculated using the high frequencies in the amplitudes of the lunar semidiurnal tide as exemplified in Figure 1(c). Figure 4(a) to 4(d), which represent the magnetic latitudes from 10° N to 20° S, show significant periodicities of 7-12 days from 7 to 12 days , when the time range from 2011 to 2014 is used. 2014. Hereafter, we are going to refer to these oscillations as quasi 8 days (Q8D) oscillation, although sometimes they can be either shorter or longer than 8
- 115 days. Similar assumption has been used by Ahlquist (1982). Please, note that other long periods were also observed above the significance levels, but they were not so pronounced like more sporadic than the Q8D oscillation as it is going to will be shown ahead.



Figure 4. Same as Figure 2, but considering only periods shorter than 25 days. Note that at 0° all periodicities were below of the confidence level since the amplitudes were smaller compared to the other latitudes.

In order to see further invertigate the temporal evolution of the short period oscillation in the amplitudes of the lunar semidiurnal tide, wavelet analysis was performed for each year of observation and the results are shown in Figures 5, 6, 7 and 8. One can observe that the dominant oscillation is the Q8D along the whole period of observation. Some particularities are also observed in each year, mainly regarding the epoch of the year in which the Q8D wave is stronger.

Figure 5 shows the wavelet results for the amplitudes of the lunar semidiurnal tide in 2011. The Q8D was stronger from September to November in almost all latitudes. There was a secondary peak of this oscillation from the middle February up to April, except at 10° S.

125 Figure 6 shows the results for 2012. Again the Q8D oscillation was the most important oscillation, but it appeared more frequently along the year, in special, out of the magnetic equator. One different aspect was the Q8D strong in February at 10° N and in May at 10° S.

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In 2013 (Figure 7), the strength of the Q8D oscillation was more concentrated in few months. At 10° N, the Q8D oscillation had more power spectral density in January and February. At 0° and 10° S, the oscillation appeared with more intensity from latter-later October to December. At 20° S there were two peaks of the Q8D oscillation in April and May.

Figure 8 shows the power spectral density contour for 2014. One can observed observe a regular behavior of the Q8D oscillation with two peaks around the equinox months.

An important result revealed from the observations is that the Q8D oscillation is always present in the equinox months. From September to November in almost all latitudes and years it was the dominant oscillation. One can also observe, that the Q8D oscillation appear strong during the winter in 2012 and 2013. 2013 for some magnetic latitudes.

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Figure 5. Same as Figure 3, but considering only periods shorter than 25 days during 2011.



Figure 6. Same as Figure 5 for 2012.



Figure 7. Same as Figure 5 for 2013.



Figure 8. Same as Figure 5 for 2014.

3 Discussion and summary

It is well known that the lunar semidiurnal tide has a predictable source. Thus, short and long variations observed in the amplitudes must reveal changes in the atmosphere where this tidal component is propagating. For instance, annual and semiannual variation in the amplitudes of the lunar tide <u>has_have</u> also been observed in the mesosphere and lower thermosphere (MLT)

140 region in the neutral wind (Paulino et al., 2015). Intraseasonal variation has Terannual variations have been observed and simulated in some atmospheric fields as well (e.g., Pedatella et al., 2012; Pedatella, 2014). However, more investigation is necessary to the better understanding the reason for those variability.

The results from Figures 2 and 3 show that the annual variation is always present in the amplitudes of the lunar semidiurnal tide in the TEC. At the magnetic equator, the PSD of the annual variation is comparable to the semiannual, for instance.

- 145 However, far from the equator, the annual variation is stronger. Furthermore, It it seems that the annual variation is out of phase at this latitude compared to the annual variation observed in MLT winds (Figure 2, bottom row of Paulino et al., 2015), i.e., the annual variation maximizes around January for all latitudes in the TEC and it maximizes around November in the MLT winds. It This reinforces that the lunar tide obey obeys the changes in the atmosphere and the observed variability is not due to changes in the sources, otherwise, we would expect variability almost in phase at different atmospheric levels.
- Although the semiannual oscillation raised up as the second peak in the Lomb-Scargle Periodogram (Figure 2) from 2011 to 2014, it appeared sporadically and with more intensity in lower latitudes. In contrast, the TEC observed in Brazil shows a semiannual variation and have maximum maxima around the equinox during both low and high solar activities (Jonah et al., 2015).

The intraseasonal terannual variation with period around 120 days, in average, was the third peak found in the amplitudes

- 155 of the semidiurnal tide, it. It was sporadic at 10° N, 0° and 20° S. At 10° S, it was present in almost the whole period of observation and it was stronger than the semiannual oscillation during the first two years of observations..., except at 20°S. Oscillations with 70-80 days period were also observed in all latitudes sporadically, mainly in the beginning of 2013 and 2014. In Paulino et al. (2017), one can it is possible to see that the intraseasonal terannual oscillation appears evident at magnetic latitudes out of the equator. Maybe It is probable that the combination of the annual variation with maximum in the austral
- 160 summer months and semidiurnal variation with maximums around the equinoxes (matching with the TEC maximums) is producing the intraseasonal-terannual variation in the amplitudes of the lunar semidiurnal tide.

Figure 1(a) shows roughly a short period oscillation in the amplitudes of the lunar tide which can be observed in all studied latitudes over Brazil. Sometimes, this short <u>period</u> oscillation is stronger and sometimes it is very tenuous. It This bahavior is quite interesting because it was also observed in Figure 3 of Paulino et al. (2017).

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- An interesting aspect reveled in this work was the periodicity close to 8 days. Based on the literature, there are two kind kinds of large scale waves with periods close to 8 days: (1) Fast Kelvin wave (e.g., Abdu et al., 2015, and references therein) and (2) Quasi 10 days planetary wave (e.g., Forbes and Zhang, 2015; Yamazaki and Matthias, 2019).

Fast Kelvin waves have been observed with period typically of 6-7-6-10 days. It is a kind of wave trapped in the equatorial region which has characteristics of for gravity waves, i.e., it obeys the dispersion relation of gravity waves. Fast Kelvin waves

- 170 are typically observed with large amplitude in the zonal wind <u>component</u> and insignificant amplitudes in the meridional one. As the present Q8D oscillation was observed close to the equator as well as at 20° magnetic latitude, which can be out of the tropic in the west part of Brazil, it is most unlikely that the Fast Kelvin wave is modulating the amplitudes of the lunar semidiurnal tidethis modulation can have contribution of other oscillations as well. Some observations have shown that the amplitudes of the Fast Kelvin wave dominate in altitudes below 90 km (e.g., Lieberman and Riggin, 1997). Dhanya et al. (2012) found out
- 175 periodicities close to 8 days in Equatorial Electro Jet and the Equatorial electrojet current and in the MLT winds and associated that oscillation to Fast Kelvin wave. Abdu et al. (2015) also observed Q8D oscillation in the vertical drift of F region and which modulated the spread-F development.

Figure 9 shows the zonal (solid line) and meridional (dashed line) thermospheric mean wind at 00:00 UT measured by two Fabry-Perot interferometers deployed at Cajazeiras (6.9°S, 38.5°W) and São João Cariri (7.4°S, 36.5°W) during November

180 2013. Further details about the operations of those equipments can be found in Makela et al. (2009). One can observe that there is an almost in phase oscillation of about one week in the wind field maybe suggesting the same origin of the observed Q8D oscillation in the amplitude of the lunar tide during this epoch.



Figure 9. Zonal (solid line) and Meridional (dashed line) means wind components measured two Fabry-Perot interferomenters over Cajazeiras and São João do Cariri measured at 00:00 UT during November 2013.

Figure 10 shows the Lomb-Scargle periodogram for those data. A strong oscillation of ~ 6 days can be observed and it is likely associated with the presence of fast Kelvin wave in the equatorial zone. One can also observe a peak of quasi 10 days in
185 both components, however, it was below the confidence level.



Figure 10. Lomb-Scargle periodogram for (a) zonal and (b) meridional wind during November 2013 in the thermosphere over the equatorial region. Horizontal dotted line represents a significancy level of 95%, i.e., false alarm probability of 0.05.

In the past years, the interest in studying the quasi 10 day wave (O10DW) has been recovered, primarily due to its association with polar Sudden Stratospheric Warmings (SSWs) (Yamazaki and Matthias, 2019; Mo and Zhang, 2020), Another motivation was the long term observation from satellites that allow to investigate seasonality, year to year variation and spatial (latitude x longitude x altitude) dependencies (Forbes and Zhang, 2015; John and Kumar, 2016).

- Although the present results concentrate into the oscillations around 8 days in the amplitudes of the semidiurnal lunar 190 tide, these oscillations have some characteristics similar to the Q10DWs as pointed out by Forbes and Zhang (2015). For instance, in some years, they have large amplitudes during the equinox months and during the winter months in both hemispheres. Another interesting characteristic was the anti-simetry as can be seen in Figure 9, i.e., the amplitudes of the Q8D oscillation is out of phase at the two hemispheres.
- 195 Wavelet phases using data from 10 September to 10 November 2011 for (a) 10° N and (b) 10° S.

Observation of wave/oscillation with periods 8-10 days have been made in the thermosphere-ionosphere-ionosphere. For example, Forbes (1996) used data of Medium Frequency radar and magnetometer to show the found Q10DW oscillations in the mesopause and low thermosphere region using data of Medium Frequency (MF) radar and a magnetometer. Pancheva and Laštovička (1998) observed fluctuations of 7-8 days from November to December 1994 during an international Campaign.

- 200 Abdu et al. (2006) studied variation in the Equatorial Electro Jet electrojet (EEJ) and current and in the MLT wind and showed the presence of 8-12 days oscillation around the equinox of 1999 in the equatorial region. Jacobi et al. (2007) observed 7-12 days waves in the TEC maps over Europe region. Jonah et al. (2015) have also observed 8-10 days oscillation in TEC over Brazil, primarily around the equinoxes. More recently, comprehensive studies on Q10DW using satellite data presented some important temporal and spatial characteristics of this wave below 110 km altitude (Forbes and Zhang, 2015; John and Kumar, 2016). Additionally, Yamazaki and Matthias (2019) and Mo and Zhang (2020) presented results associating the Q10W to
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Sudden Stratospheric Warming events.

It is important to notice note that the observations above showed periods varying from 8 to 10 days in both the mesosphere and thermosphere-ionosphere. Forbes and Zhang (2015) showed slight variation (0.4 days) in the period of Q10W (9.7 to 9.9 days) and concluded that the doppler shift produced by the horizontal wind can change the period of the wave. The present

- 210 study uses amplitudes of the lunar semidiurnal tide in lunar time, i.e., the lunar days is by about 0.036 days longer than solar day. Assuming that the observed oscillation has a period of 8.5 days in lunar time, it corresponds to 8.806 days in solar time. However, this is not enough to explain the discrepancy between the observed period of the Q10DW in the lower atmosphere and the present results.
- Investigations in other atmospheric fields have shown evidences of 8-10 oscillations with simultaneous periods. Figure 10 Figure 11 shows the horizontal wind at 93 km altitude over Cachoeira Paulista (22.7° S; 45.0° W) during November 2013 215 measured at 02:00 Universal Time. Further details about the meteor radar and its measurements wind measurements using meteor radar have been published elsewhere (e.g., Paulino et al., 2012). The temporal evolution of the winds shows some periodic oscillations.



Figure 11. Zonal (solid line) and Meridional (dashed line) wind components at 93 km height over Cachoeira Paulista measured at 02:00 UT during November 2013.

Figure 11–12 shows the Lomb-Scargle periodogram of the wind including all temporal measurements at 93 km. Quasi 10 220 days oscillation is shown in both zonal and meridional components of the horizontal wind. In the zonal component the peak of the oscillation was concentrated at 9.7 day, while in the meridional one, the peak was at ~9 days. The presence of this simultaneous oscillation in the MLT wind strongly suggests that the observed oscillation in the amplitude of the lunar tide might have origin in a coupling with the lower atmosphere. during November 2013 could have contribution of this oscillation, at least, out of the equator. But the coupling mechanism needs further investigations.



Figure 12. Lomb-Scargle periodogram for (a) zonal and (b) meridional wind during November 2013 at 93 km altitude <u>over Cachoeira</u> Paulista. Horizontal dotted line represents a significancy level of 95%, i.e., false alarm probability of 0.05.

225 The main results of this investigation can be summarized as follows:

- There is a strong temporal variability in the amplitudes of the lunar semidiurnal tides calculated in the TEC maps over Brazil from 2011 to 2014;
- Annual variation in the lunar semidiurnal tide is always present in all observed latitudes and it is dominant in lower latitudes;
- Semiannual and intraseasonal terannual (~120 days) were, respectively, the second and third most important long period oscillation observed in the amplitudes of the lunar tide. However, it was observed a temporal and spatial variability of these oscillations, which allow them to become dominant in a given time interval and latitude range;

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- It was The observed dominant oscillations between 8-11 days in the amplitudes of the lunar semidiurnal tide with maximums had periods between 8-11 days, with maxima around the equinoxes. In some years, as 2013 and 2014, the peaks occurred in the winter;
- Anti-symmetrical correlation between the Q8D oscillations calculated in conjugate latitudes from September to November 2011 were observed;
- Coincident measurements of the horizontal wind in low latitudes in the MLT during November 2013 showed show the presence quasi 10 days oscillations oscillation in MLT at low latitudes (23°S) and quasi 6 days oscillation in the equatorial thermosphere.

Based on the preset main results to present main results for the short period oscillations in the amplitudes of the lunar semidiurnal tide, i.e., the Q8D oscillation, it eould be verified that this oscillation has important characteristics of the westward propagating anti-symmetrical normal mode with s = 1 with period of quasi 10 days. However, the discrepancy associated to the low period of observation as compared to the Q10DW is an open question that must be further investigated.

- According is clear that there is discrepancy between the observed Q8D and other oscillations observed in the wind. Additionally, according to the theory of planetary waves, the dissipative process into the thermosphere does not allow direct propagation of these wave to high levels. Then, the explanation for observation of planetary waves in the thermosphere-ionosphere have been suggested, basically, based on two possibilities: (1) modulation of the tidal amplitudes, in special, the semidiurnal components which can propagate to higher altitudes into the thermosphere and/or (2) though through the theory of dynamo on the electro-
- 250 dynamics of the ionosphere. The present results suggest that the first possibility is more efficient to the propagation of waves with period of 8-11 days. maybe a combination of these two possibiliteis could be necessary in other to explain the observed modulation in the amplitude of the lunar semidiurnal tide. However, further investigations are necessary to understand this coupling mechanism.

Data availability. The TEC maps used in this work are available online in the EMBRACE website (http://www2.inpe.br/ climaespacial).
 Meteor winds can be requested to Dr. Lourivaldo Mota Lima (lourivaldo_mota@yahoo.com.br). Thermospheric winds for São João do Cariri and Cajazeiras can be dowloaded from Madrigal CEDAR data base.

Author contributions. ARP has written the manuscript and performed the analysis in the data base. FSA has worked in the Lomb-Scargle periodograms. IP has revised the manuscript and helped in some analysis of the data. CMW has provided the TEC maps. LML has provided the meteor winds. PPB is responsible for the meteor winds and has revised the manuscript. ISB has revised the manuscript.

260 Competing interests. The authors declare that they do not have competing interests

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