

Interactive comment on “Impact of Gravity wave drag on the thermospheric circulation: Implementation of a nonlinear gravity wave parameterization in a whole atmosphere model” by Yasunobu Miyoshi and Erdal Yiğit

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Received and published: 25 July 2019

Thank you for your constructive and helpful comments. According to your comments, we have revised manuscript.

(1) It would be helpful to the reader if the Introduction briefly described the approach of GWD parameterization and the challenges presented by extending these parameterizations from the middle atmosphere to the thermosphere, where different physical processes apply (e.g., page 4 lines 1-2). It may help to first cite a basic reference

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describing the difference between linear and nonlinear approaches (e.g., Fritts and Alexander, Rev. Geophysics, 2003 may be one such reference) to GWD parameterization and then cite and briefly describe in more detail why the thermospheric environment requires modified or additional physical terms. The introduction should also cite other related studies involving GCM simulations of the thermosphere at high resolution (e.g., Liu et al., GRL 2014 <https://doi.org/10.1002/2014GL062468>) or with parameterized GWD (Becker, JAS, 2017, <https://doi.org/10.1175/JAS-D-16-0194.1>, or England et al. JASTP 2006 <https://doi.org/10.1016/j.jastp.2005.05.006>), and describe how the present study fits in with previous work. These are just some examples, the authors no doubt are more aware of recent work in this field than I am, but the point is that there is already a fair amount of research in this area that should be noted in the manuscript.

The introduction now cites some of the relevant references suggested by the reviewer. Also, we have discussed in the beginning of introduction GW parameterizations in GCMs in the context of the upper atmosphere physics above ~105 km (turbopause). Overall, the rationale for the extension of GW parameterization to the thermosphere is discussed in a number of previous publications (e.g., Yiğit et al., 2008; Yiğit and Medvedev. (2013)

(2) The experimental design needs more description. Specifically, a. what is the model time step? b. How long was the perpetual June simulation carried out? Is this just 30 days of simulation? Figure 1 caption states only 30 days of results were averaged (June 1-30). c. Is there any spin up to the model to reach a steady state? This is important to know; is the SW2 signal steady or varying strongly with time throughout the simulation? d. Can the authors explain why is there no experiment where the “linear” middle atmosphere scheme is applied above 100 km? It is not surprising that a simulation with parameterized GWD artificially cut off at 100 km will produce a different thermospheric state than a simulation with parameterized GWD extending throughout the entire model domain. In its present form, the study isn’t really telling us

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that the YiÅ§it scheme is needed for a GCM in the thermosphere, it's just telling us that more drag on the winds gives a different SW2. This is rather superñAcial. To be more substantial, it would help to know the beneñA of the nonlinear YiÅ§it scheme? What dissipation terms are importantããã(nonlinear interaction, radiative damping, diffusion, ion drag, etc?). Running a simulation with the linear scheme in the thermosphere could answer this.

The descriptions of numerical simulation in the previous manuscript were incorrect. Numerical simulation started on 1 June, and we conducted 2-year numerical integration with seasonal variation. The GCM was nudged by Meteorological reanalysis data (JRA55) up to 40 km height to simulate realistic temporal variations in the lower atmosphere. The time step of the GCM is 30 s, and the data are sampled every 1 h during the numerical simulation. The data from 1 June to 30 June in the second year are analyzed in this study. The GCM has realistic temporal variations, so that the SW2 has significant day-to-day and seasonal variations.

The rationale for the extension of GW parameterizations into the thermosphere has been discussed in a number of research (YiÅ§it et al., 2008, 2009; YiÅ§it and Medvedev., 2013) and review papers (Yi?it and Medvedev, 2015; YiÅ§it et al., 2016). In fact, now the extended introduction summarizes the need for the whole atmosphere scheme, i.e., YiÅ§it et al. (2008) scheme, for GCMs extending in to the thermosphere. There are a number of differences between the conventional linear schemes and the nonlinear scheme of YiÅ§it et al. (2008). Linear schemes all use intermittency factors as they unrealistically overestimate GW drag, while the YiÅ§it scheme does not contain any intermittency or artificial factors. Second, while linear schemes assumes that waves propagates independently, as if other waves are not present, the YiÅ§it scheme accounts for the mutual interaction between different GW harmonics in the entire GW spectrum. An illustrative intercomparison of the linear approach with the nonlinear whole atmosphere approach of our GW scheme has already been performed in the initial work of YiÅ§it et al. (2008)..

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(3)Figure 1 plots zonal mean zonal wind, but not temperature. Since Figure 2 shows the SW2 in temperature, and since the authors claim the GW thermal effects are very important (page2line6), it makes sense to add temperature to Fig1 to give a complete description of the differences in the background thermospheric state between EXP1 and EXP2.

The following sentences are inserted in the manuscript (section 3.1)

Figures 2a and 2b show the height–latitude distribution of the zonal mean meridional wind obtained by EXP1 and EXP2, respectively. In both experiments, southward flow from the summer pole to winter pole is dominant at 50–100 km height, whereas northward flow appears between 100 and 120 km height. These flows are stronger in EXP2 than in EXP1, which is explained by the enhanced GW drag in EXP2 as shown later. Above 130 km height, southward flow is dominant in both experiments. The magnitude of the southward wind between 130 and 250 km height is weaker in EXP2 than that in EXP1 except for southward of 30° S (Figure 3c). This weaker meridional wind in EXP2 is caused by the meridional component of the GW drag. On the other hand, the difference of meridional wind between EXP1 and EXP2 is small above 250 km height (less than 10 %). Figure 3a and 3b shows the height-latitude distribution of the zonal mean temperature obtained by EXP1 and EXP2, respectively. At 80–100 km height, cooling and warming occur at 30–90° N and at 60–90° S, respectively (Figure 3c). This cooling and warming effects are caused by the enhanced southward wind (meridional circulation) at 80–100km height in EXP2. Namely, the cooling (warming) at 30–90° N (60–90° S) is due to enhanced upward (downward) wind. It is noteworthy that cooling prevails above 100 km height. In particular, cooling at high latitudes in the NH exceeds 60 K. This cooling is caused by the GW thermal effect. This indicates that GW induced cooling also affects thermal structure in the upper thermosphere. Our results support the conclusion of the GCM work by YiÅ§it and Medvedev (2012), who showed for the first time that GWs cool the thermosphere during low solar activity conditions.

(4)Effects of GWD on SW2 are discussed, but not DW1 or other tidal modes, partic-

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ularly nonmigrating tides. Why were these modes not considered? Earlier, Miyoshi et al (2014) used this exact same model at slightly higher resolution (1 degree latitude/longitude) and found that the diurnal tide is important above 200 km. Since these tides are not acting in isolation to one another, it makes sense to describe the impact of GW drag on both DW1 and SW2 at least. Please also mention in section 3.2 how the tidal amplitudes are obtained from the model.

The impact of GWD parameterization on DW1 was studied in detail by YiÅšit and Medvedev (2017). The impact on DW1 obtained in this study is similar to that presented by YiÅšit and Medvedev. Therefore, we have not focused on DW1 tide in this study. Effects of GWD on TW3 are newly added. The following sentences are inserted in section 3.3.

Figure 6a shows the height–latitude distribution of the temperature component of the migrating terdiurnal tide (TW3) amplitude in June obtained by EXP1. The amplitude peak is located at 15° N latitude, and secondary peak appears at 25–30° S. The maxima are 23 K at 15° N and 130 km height, and 18 K at 17.5° S and 165 km height. Figure 6b shows the temperature component of the TW3 amplitude obtained by EXP2, and Figure 6c shows the amplitude difference between EXP1 and EXP2. The latitudinal structure of the TW3 in EXP2 is quite similar to that in EXP1. However, the amplitude is weaker in EXP2 than in EXP1 by about 20–40%. The amplitude difference is significant in the 120–220 km height range. The TW3 is also attenuated by the GWD. Forbes et al. (2008) indicated that the TW3 amplitude at 110 km height is between 5 and 8 K. However, there are only a few studies concerning the satellite observation of TW3 amplitude in the 120–220 km height range. A detailed comparison of the TW3 amplitude between the simulation and observation is a subject of a future study.

(5) The description of the agreement between EXP2 SW2 and observations (Section 3.2, not 1.2 as in the manuscript) is not entirely convincing. SABER estimates are quoted as 15–20 K but it's not clear if this is for June conditions, over what years, or what kind of uncertainty is associated with this number. Does 15–20 K mean that is the

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typical range of values? Do the SW2 amplitudes from EXP1 and EXP2 vary widely over the simulation period? It would be most helpful if Figure 2 plotted SABER SW2 results as a function of latitude and altitude for June to provide a comprehensive comparison. The authors should also do the same for DW1 – that is, compare DW1 amplitude from EXP1, EXP2 and SABER.

I think my description of the observed SW2 tide in the original manuscript is unclear. The following sentences are inserted in section 3.2.

Pancheva (2011, IAGA book) studied climatology (6-year mean from 2002 to 2007) of SW2 temperature tide using SABER observation. Figure 2.3 in Pancheva (2011) indicates that the SW2 in June at 110 km height has peaks at 20–30° N and 20° S. The maxima at 15–30° N and at 15–25° S are 25–28 K and 15–20 K, respectively. The peak values of the monthly mean SW2 amplitude in EXP2 at 110 km height are 26 K at 20° N and 21 K at 15° S. The SW2 amplitude obtained in EXP2 is consistent with the SABER observation.

The SW2 also has significant day-to-day variations. For example, the SW2 amplitude at 20° N in EXP1 (EXP2) ranges from 27 (22) to 37 (31) K, and the standard deviation of day-to-day variations in the SW2 amplitude at 20° N in EXP1 and EXP2 are 2.8 K and 2.9 K, respectively. Similar day-to-day variations in the SW2 amplitude are found below 100 km height. This indicates that day-to-day variations in the SW2 amplitude are primarily generated in the lower atmosphere and propagate into the lower thermosphere.

Please also note the supplement to this comment:

<https://www.ann-geophys-discuss.net/angeo-2019-36/angeo-2019-36-AC1-supplement.pdf>

Interactive comment on Ann. Geophys. Discuss., <https://doi.org/10.5194/angeo-2019-36>, 2019.

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