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# Stratospheric observations of noctilucent clouds: a new approach in

2	studying large-scale mesospheric dynamics
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16	Abstract.
17	The experimental campaign Stratospheric Observations of Noctilucent Clouds (SONC)
18	was conducted on the night of 5-6 July 2018 with the aim of photographing noctilucent clouds
19	(NLC) and studying their large-scale spatial dynamics at scales of 10061450 km. An
20	automated high-resolution camera (equipped with a wide-angle lens) was lifted by a
21	stratospheric sounding balloon to 20.4 km altitude above the Moscow region in Russia
22	(~56°N; 41°E), taking several hundreds of NLC images during the flight that lasted 1.7
23	hours. The combination of a high-resolution camera and large geographic coverage (~1500
24	km) have provided a unique technique of NLC observations from the stratosphere, which is
25	impossible to currently achieve either from the ground or space. We have estimated that a
26	horizontal extension of the NLC field as seen from the balloon was about 1450 x 750 km
27	whereas it was about 800 x 550 km as seen from the ground. The NLC field was located in a
28	cold area of the mesopause (136-146 K), which is confirmed by satellite measurements. The
29	southmost edge of the NLC field was modulated by partial ice voids of 150-250 km in
30	diameter. A medium-scale gravity wave had a wavelength of 49.4±2.2 km with vertical
31	amplitude of 1.9±0.1 km. The final state of the NLC evolution was represented by thin
32	parallel gravity wave stripes. Balloon-borne observations provide new horizons in studies of
33	NLC at various distances from metres to thousands of km.





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35 Keywords: noctilucent clouds, mesospheric dynamics, balloon-borne stratospheric
36 observations, atmospheric gravity waves

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#### 1 Introduction

39 Night-shining clouds or noctilucent clouds (NLC) are the highest clouds in the Earthos 40 atmosphere observed at the summer mesopause between 80 and 90 km. NLC can be readily 41 seen from mid- and subpolar latitudes of both hemispheres. NLC are composed of water-ice 42 particles of 30ó100 nm in radius that scatter sunlight and thus NLCs are observed against the 43 dark twilight arc from May until September in the Northern Hemisphere and from November 44 to February in the Southern Hemisphere (Bronshten and Grishin, 1970; Gadsden and 45 Schröder, 1989; Liu et al., 2016). NLC are also observed from space and in this case they are usually called Polar Mesospheric Clouds (PMC) (Thomas, 1984). 46 47 NLC are almost always represented by a wave surface having a complex interplay 48 between small-scale turbulence processes of 10-1000 metres, atmospheric gravity waves 49 (GW) with wavelengths of 10-1000 km, planetary waves, solar thermal tides and lunar gravitational tides of about 10000 km (Witt, 1962; Fritts et al., 1993; Rapp et al., 2002; 50 51 Kirkwood and Stebel, 2003; Chandran et al., 2009; Dalin et al., 2010; Fiedler et al., 2011; 52 Taylor et al., 2011; Pertsev et al., 2015). Sometimes, distinguished non-linear mesospheric 53 phenomena like mesospheric walls or fronts appear at the mesopause which clearly separate 54 two volumes of the mesopause having cold and warm air masses with temperature difference of 20-25 K across a few km (Dubietis et al., 2011; Dalin et al., 2013). 55 56 NLC/PMC are systematically observed and studied from the ground (optical imagers, 57 lidars), as well as from space (AIM, Odin, SBUV instruments) (e.g., Karlsson and Gumbel, 58 2005; Dalin et al., 2008; Bailey et al., 2009; Fiedler et al., 2011; DeLand and Thomas, 2015); 59 there are also irregular (campaign-based) NLC observations conducted by using sounding 60 rockets and aircraft (Zadorozhny et al., 1993; Gumbel and Witt, 2001; Reimuller et al., 2011). 61 These techniques have advantages and disadvantages. In particular, ground-based 62 measurements provide a high horizontal resolution of ~20 m and high temporal resolution of 63 seconds (optical imagers) (Dalin et al., 2010; Baumgarten and Fritts, 2014) and high vertical 64 resolution of 50-150 metres using lidars (Baumgarten et al., 2009) but are limited to tropospheric weather conditions and restricted to a certain small region on the Earth& surface. 65



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66 Satellite measurements, on the other hand, provide global PMC coverage but have low spatial horizontal resolution (~5 km) as well as large spatial gaps of several hundreds of km between adjacent orbits at middle and subpolar latitudes. Thus, there is no perfect technique to observe 68 69 and study NLC/PMC so far. There is an obvious methodological gap in these techniques, 70 resulting in a gap of stratospheric altitudes (20-40 km), which are potentially available for comprehensive studies of NLC/PMC. This gap is due to lack of systematic stratospheric 72 balloon-borne experiments aiming at NLC/PMC observations. So far, there has been 73 conducted a single published experiment from a stratospheric balloon providing PMC 74 observations over Antarctica between 29 December 2012 and 9 January 2013 (Miller et al., 75 2015). The E and B Experiment (EBEX) was dedicated to another research field concerning 76 polarization in the cosmic microwave background (Reichborn-Kjennerud et al., 2010). At the same time, two star cameras of the EBEX experiment, having a narrow field of view of 4.1° x 78 2.7°, were able to register fine structures of PMC and turbulence dynamics, ranging from 79 several km down to 10 m. Another balloon-borne experiment (PMC-Turbo) was conducted 80 between 8 and 14 July 2018 over Sweden-Greenland-Canada territories in order to capture NLC with seven optical cameras and lidar (Fritts et al., 2019). The PMC-Turbo experiment 82 was launched about 2.5 days after the experiment described in the present paper. 83 In this paper, we report on scientific results of a new balloon-borne experiment dedicated 84 to studies of NLC large-scale dynamics at horizontal scales of more than 100 km (Dalin et al., 85 2019). Such experiment, conducted for the first time, opens new horizons for studies of large-86 scale dynamical features in combination with a high spatial resolution at the summer mesopause, currently unachievable for other techniques like ground-based and space measurements.

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## 2 Technique and method

The Stratospheric Observations of Noctilucent Clouds (SONC) experiment is a special balloon-borne experiment dedicated to studies of large-scale dynamical features in NLC. A high resolution high sensitive camera (Sony Alpha A7S), having a full frame 35 mm 12 megapixel sensor (4240 x 2832 pixels) and equipped with a wide-angle lens (field of view, FoV, is 109.7° x 81.6°), has been installed on a meteorological sounding balloon. This combination of a high resolution sensor and wide FoV yields spatial horizontal resolutions of ~30 m and ~3000 m, when looking at 83 km from 20 km at elevation angles of 90° and 0°, respectively. The horizontal coverage of a mesopause layer is over 2000 km, when viewing along the horizon at low elevation angles. The balloon was launched from the Moscow



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region, Russia (~56°N; 41°E), on the night of 5-6 July 2018. Since a gondola payload is constantly rotating and shaking during its flight, the NLC camera was installed on a special stabilized platform. The 3-axis motorized gimbal stabilized platform (Fig. 1) was designed and build by the Aerospace laboratory õStratonauticaö (http://stratonautica.ru), which has a wide experience in building such platforms and launching sounding balloons. The platform was designed to rotate in a 60° step in the azimuth angle in order to capture the whole hemisphere (360°) since NLC can appear in any direction as observed from mid-latitudes, including the southern part of the sky (Hultgren et al., 2011; Suzuki et al., 2016). The NLC camera took images every 6 s during the whole flight, obtained several thousands of images and several hundreds of images capturing NLC. Besides, automatic exposure bracketing was used to take four images in sequence with different exposures, allowing us to register various NLC brightness from very bright to very faint as well as faint stars, which are important information for the photogrammetric technique and georeference procedure of the images. The balloon was launched at 21:34 UT on 5 July 2018 and the total flight duration was about 1.7 hours. The ascent speed was around 5 m/s and the balloon reached its maximum altitude of 20.4 km where it burst; then the payload descended with a parachute and the payload was successfully recovered. A GPS receiver was installed onboard in order to obtain information on the balloon trajectory. The flight characteristics of the SONC balloon are shown in Fig. 2. A ground-based support consisting of three automated NLC cameras was established in the Moscow region in order to launch the balloon at the time of NLC appearance. Also, a number of amateur observers significantly contributed to the NLC observational programme before and during the flight. A launch window was preliminary chosen at the beginning of July based on long-term statistics of NLC observations conducted in the Moscow region since 1962 to present time. This statistics demonstrate that NLC appear at the beginning of July with about 65% occurrence probability on a clear night.

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#### 3 The observation

During the flight, the balloon-borne camera captured an extended NLC field with a number of interesting features discussed in section 4. One can note the following general characteristics of the NLC display:

a) NLC were observed between 20:30 and 23:15 UT on 5 July 2018.





b) NLC were located between 82.6 and 85.1 km. The NLC height was estimated by using synchronously taken images obtained from two ground-based cameras located in the Moscow region.

- c) NLC field extended along the horizon from NW to NE at low elevation angles from  $65^{\circ}$  to  $+11^{\circ}$  as seen from the balloon.
- d) NLC were modulated by atmospheric gravity waves of various scales having horizontal wavelength from 9 km to 50 km.
- e) NLC were traveling in a rather unusual direction from the south to north at the observed mean speed of ~43 m/s.
  - f) NLC were fading during the balloon ascent and they got very faint and less structured at the maximum balloon altitude of 20.4 km. The brightest and well-developed NLC were observed when the SONC balloon was between 6 and 13 km, that is why we analyze the most profound features of NLC images obtained at this height range.
  - Each analyzed NLC image was georeferenced using horizontal coordinates of referenced stars (at least 15 stars are needed). The technique of the NLC georeference, triangulation height estimation and error analysis can be found in Dalin et al. (2004, 2015).

## 4 Results and discussion

The projection of the NLC field on the surface along with the temperature map obtained with the Aura/MLS spectrometer is shown in Fig. 3. The description on the MLS temperature product and its validation can be found in Froidevaux et al. (2006) and Schwartz et al. (2008). One can see that the NLC field extended mostly from the west to east along an area filled with low temperatures of 136-146 K, and the NLC were located north of 58°N due to rapidly increasing temperature with decreasing latitude. That is why the NLC were observed at low elevation angles (far to the north as seen from the Moscow region) on this particular night.

Detailed analysis of five consecutive in time balloon-borne images (Figs. 4 and 5) has revealed the following features of the NLC display:

a) The horizontal extent of the NLC field from the western to eastern observable border was about 1450 km, and from the northern to southern border of about 750 km. Such distances are impossible to observe from the ground due to the Earth& curvature and limited area of the twilight arch. The central part of the NLC field, having extension of about 850 x 550 km, was seen from the ground but the western and eastern wings of the field as well as the northern edge were located below the local ground horizon, making it impossible to observe them. Thus, balloon-borne NLC observations have an obvious great advantage over





ground-based observations in terms of larger geographic coverage which is comparable to PMC observations made from space.

b) The southmost edge of the NLC field was modulated by partial circles (something like ice voids but with open southern border), which is shown by the red curves in Figures 4 and 5. The diameters of these partial ice voids are estimated to be in the range of 150-250 km. The mechanism of the formation of ice voids in NLC/PMC is not clear now, and it is an ongoing topic in atmospheric physics. One can mention three main mechanisms which are currently discussing in the literature. Trubnikov and Skuratova (1967) addressed a theory of cellular convection and demonstrated its principal possibility in the summer mesosphere in relation to NLC occurrences. The authors estimated convective cells to be in the range of 90-250 km in radius, that agrees well with sizes of partial ice voids obtained in the present study. However, there should be fulfill the main criterion for the convection to be developed, namely, the height gradient of the potential temperature should have negative values. We have carefully estimated the potential temperature gradient or the static stability (based on Aura/MLS temperature measurements) in the analyzed area and could not find any signatures of its negative values in the mesosphere and mesopause region. It means that in this particular case cellular convection could not be responsible for the observed partial ice voids in the NLC.

However, satellite measurements can easily miss a negative static stability at local scales due to poor horizontal resolution and local ice voids may be generated by a gravity wave breaking. Rusch et al. (2009) have hypothesized that ice voids could be caused by heating due to the passage of warm crests of a gravity wave. It is possible in the present case. However, we could not find any significant displacement of the partial ice voids (their boundaries) relative to the NLC field, i.e., the partial ice voids traveled with the same speed and direction as the entire NLC field did (~43 m/s from the south to north). One would expect an intrinsic phase speed and intrinsic direction of the movement of the partial ice voids if they were generated by a large-scale gravity wave of a wavelength of several hundreds of km. Thus, it is difficult to prove this hypothesis of the influence of a large-scale gravity wave on the formation of the observed partial ice voids.

Thurairajah et al. (2013b) have proposed another mechanism related to a shock wave generated by a meteorite, which expands and cools the air that in turn leads to the formation of large ice particles which fall out of an NLC field (analogously to hole-punch clouds due to the passage of an aircraft). However, we observe large-scale partial ice voids (150-250 km) in a broad area of the mesopause over 1000 km. It was hardly possible that any big meteorite



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could produce such large holes in such broad area, and we did not observe any meteor motion in our ground-based and balloon images.

Megner et al. (2018) have recently presented an interesting case study of a quasistationary ice void in NLC which did not follow the general wind, suggesting that it was formed by a localized warming at the summer mesopause. This is not the case in our case study, in which we have observed partial ice voids moving at the general wind speed in the same direction along with the entire NLC field.

In the present case study, the partial ice voids had irregular shape and sizes ranging from 150 to 250 km. Also, these partial voids moved along the wind, having the same speed and direction. Thus, it is difficult to connect these partial voids with regular wave disturbances. At the same time, as shown in Fig. 3, the southmost border of the NLC field was confined to the warm air mass located at sub-polar latitudes of ~58°N and lower. The mesopause temperature at this border was equal to ~147 K at 86 km altitude. The MLS data cannot reproduce the exact shape of this border due to low horizontal resolution (~15°) and temporal resolution of ~1.5 h. However, it is well known that tropospheric frontal systems have a meandering shape, sometimes with intrusions of warm and cold air masses as in case of the formation of a frontal wave cyclone (Ahrens, 1993; Stull, 2000). In our case the warm front at the mesopause and the NLC partial ice voids resemble a tropospheric frontal wave, in which there are intrusions of warm air masses, moving from midlatitudes, into the cold air mass located at sub-polar and polar latitudes (see Fig. 6). Therefore, we consider that the most probable source of these partial ice voids observed in the NLC in this particular case is the intrusion of warm air masses into the cold air mass with the NLC field, sublimating ice particles. A similar conclusion was proposed by Thurairajah et al. (2013a) who have analyzed a large ice void observed in PMC (using AIM/CIPS satellite images) and have concluded that õi warmer temperatures (warmer than the frost point temperature of ~144 K) at the location of the void may be related to increased tidal activity and transport of warm air from low latitudes. Ö Also, Bailey et al. (2009) and Thurairajah et al. (2013b) have demonstrated that southmost borders of PMC can be highly modulated by partial ice voids of several hundreds of km in diameter, and the authors have found the structural similarity between PMC images and those seen in tropospheric clouds.

c) Clear vertical modulation of the NLC layer is shown with the red arrow in Fig. 7. This is a unique view on a particular gravity wave seen at the local horizon of the balloon; that is why this wave modulation is viewed almost at the right angle to the line-of-sight. This geometry allows observing a thin layer of NLC modulated in altitude by propagating gravity





233 waves of small and medium scales. Such geometry is almost impossible to obtain from the 234 ground since NLC seen at the very horizon are usually masked by topography, tropospheric 235 clouds and, most importantly, by tropospheric aerosols, which are constantly present and 236 significantly absorb NLC brightness when looking at the very horizon. We have carefully 237 estimated parameters of this particular wave: its horizontal wavelength was equal to 49.4±2.2 238 km and its vertical amplitude was 1.9±0.1 km between the crest and trough. In this 239 calculation, the angle of 13.3° between the camera image plane and vertical plane at the NLC altitude was taken into account. Also note that since NLC are clearly seen both in the crest 240 241 and trough of the wave (ice particles did not completely sublimated in the wave trough), we 242 have estimated the wave amplitude both in the wave crest and trough. The amplitude 243 estimations are the same in the wave trough and crest (within the given uncertainty). All this 244 makes us confident in the estimation of the vertical amplitude of this particular wave. This is the most precise estimation of the amplitude of a gravity wave at the mesopause by using 245 NLC observations (Witt, 1962; Haurwitz and Fogle, 1969; Bronshten and Grishin, 1970; 246 247 Demissie et al., 2014). Since wave amplitude represents kinetic wave energy, this is an 248 important source of information for estimating the wave energy budget at the upper 249 atmosphere, and also can be used for future model studies to estimate a wave source in the 250 lower atmosphere (Fritts and Alexander, 2003; Demissie et al., 2014). 251 d) Small-scale billow-type gravity waves were estimated to have horizontal wavelengths 252 of 8-11 km (Fig. 7). Such small-scale gravity waves are well-known to be observed in NLC 253 layers (Witt, 1962; Dalin et al., 2010; Pautet et al., 2011; Baumgarten and Fritts, 2014; 254 Demissie et al., 2014), but we demonstrate this result in order to emphasize the ability to 255 resolve small-scale NLC structures by using a large FoV camera, having a high resolution 256 sensor, onboard a sounding balloon. 257 e) Figure 8 illustrates an NLC image taken from altitude of 20.3 km which is very close to the maximum reached altitude of 20.4 km. The NLC were rather faint by that time that is in 258 259 line with an idea of the intrusion of warm air masses from mid- to subpolar latitudes. These 260 large-scale warm air masses led to rapid sublimation of ice particles at large scales of about 261 1500 km. At the same time, on can see a very interesting feature to be considered. There were 262 several thin parallel gravity wave bands (stripes) with lengths of 50-200 km and widths of ~ 3-5 km in cross-section. The reasons of seeing such thin stripes are as follows: (a) The SONC 263 264 balloon was in the stratosphere, i.e., above the troposphere in which optically strong air turbulence is constantly present (b) The exposure time of this image was very short of 1/125 s. 265 266 All these made the image free from blurring (as minimum blurring as possible for moving





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267 NLC and balloon motion). This image demonstrates a final stage of the NLC evolution (NLC 268 disappeared in 20 min since the image was taken), and these thin stripes might represent a final morphological state of the NLC evolution. Further balloon-borne NLC observations of 269 270 very faint NLC are required to confirm this consideration. 271 272 **5 Conclusions** 273 The combination of high resolution images (~30 m) and large geographic coverage (over 274 1500 km) is a unique property intrinsic to stratospheric balloon-borne NLC observations, 275 which is impossible to achieve either from the ground or space. In general, a balloon-borne 276 NLC observation provides us with the following new opportunities in case of a long duration 277 flight of several days: 278 a) NLC imaginary can be obtained for 24 hours a day and during several days due to very 279 little Rayleigh atmospheric scattering in the visible subrange of the spectrum above 20 km 280 (Hughes, 1964); 281 b) Quantitative information on a wide range of waves (gravity and planetary waves, solar 282 tides), propagating through the summer mesopause can be obtained; 283 c) Neutral wind velocity at the mesopause and large-scale trajectory of NLC fields over 1500 284 km can be measured; 285 d) Quantitative information on long mesospheric fronts, solitons and other non-linear 286 processes can be obtained; 287 e) Quantitative information on small-scale turbulent structures (down to 1 m) can be 288 obtained in case of using a narrow field of view lens. 289 f) High resolution vertical NLC structure (wave modulation, double layers) can be retrieved 290 by observing NLC at the very horizon. Absence of any terrain obstacles and tropospheric 291 aerosol loading makes such stratospheric NLC observations unique. 292 g) Absence of optically strong tropospheric turbulence makes NLC images free from 293 atmospheric blurring that in turn results in well-defined fine structures of gravity waves 294 and turbulence in the mesopause region. 295 296 In the present study, we have estimated the following characteristics of the NLC field: 297 a) The horizontal extent of the NLC field as seen from the SONC balloon was about 298 1450 x 750 km whereas it was about 800 x 550 km as seen from the ground. This

emphasizes the great advantage of making large-scale balloon-borne observations over





10 300 medium-scale ground-based ones. 301 b) NLC field was traveling from the south to north at a mean velocity of 43 m/s; 302 c) The southmost edge of the NLC field was modulated by partial ice voids of 150-250 km in diameter, which were like generated by the intrusion of warm air masses 303 304 moving from mid- to sub-polar latitudes. The mesopause temperature at this edge was 305 equal to ~147 K, i.e., it was a threshold temperature separating the mesopause region 306 filled with NLC from the warm area without NLC. 307 d) A medium-scale wave had a wavelength of 49.4±2.2 km and vertical amplitude of 308 1.9±0.1 km. This is the most precise estimation of a gravity wave amplitude ever 309 made. 310 e) Small-scale billow-type gravity waves had wavelengths of 8-11 km. 311 f) The final morphology state of the NLC evolution was represented by thin parallel gravity wave stripes with lengths of 50-200 km and widths of ~3-5 km. 312 313 314 Data availability. The reader can access the SONC experiment images and balloon GPS 315 coordinates, used in the paper, via publically available project ftp server at the Swedish 316 Institute of Space Physics: <a href="ftp://ftp.irf.se/outgoing/pdalin/NLC/SONC">ftp://ftp.irf.se/outgoing/pdalin/NLC/SONC</a> experiment/ 317 318 Author contributions. PD wrote the paper, made calculations and plotted the figures. NP and 319 VP read and made suggestions appropriated for the paper. DE provided the raw balloon-borne images and balloon GPS coordinates. VR contributed to the image processing. All the authors 320 321 read and commented regarding the work and agreed with the content and submission of this 322 paper. 323 324 Competing interests. The authors declare that they have no conflict of interest. 325 326 Acknowledgments. The authors are grateful to Nikolay Gusev, Andrey Reshetnikov, Alexander 327 Dalin for their support of ground-based NLC observations during the SONC experiment. The 328 Aura/MLS data version 2.2 were obtained from the NASA Goddard Space Flight Center Data

and Information Services Center: https://mirador.gsfc.nasa.gov.



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472	Figure captions:
473	Figure 1. The 3-axis motorized gimbal stabilized platform, holding the NLC camera,
474	designed and build by the Aerospace laboratory õStratonauticaö. Photo by Denis Efremov.
475	
476	Figure. 2. (Left) the altitude of the SONC balloon as a function of time flight. (Right) the
477	vertical-horizontal trajectories of the SONC balloon: the red line is the upleg and the black
478	line is the downleg trajectories.
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480	<b>Figure 3</b> . The temperature map at the mesopause (86.1 km) as measured by the Aura/MLS
481	spectrometer on 5 July 2018. Nighttime measurements around the globe have been selected to
482	produce the map. Upon the temperature map, the outer borders of the NLC field are
483	overplotted: the red line is as seen from the SONC balloon, the black line is as seen from the
484	ground at the launch. The black dots mark the position of the balloon at 7.8 km at the ground
485	and ground-based observers.
486	
487	Figure 4. The NLC field as observed from the SONC balloon at 4092 m, 4947 m, 7836 m,
488	9077 m and 13928 m above the ground at 21:46 UT, 21:49 UT, 21:57 UT, 22:01 UT, 22:20
489	UT on 5 July 2018. The red curves indicate large areas free from NLC particles (partial ice
490	voids).
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492	Figure 5. Projection of the NLC fields (shown in Figure 4) as observed from the SONC
493	balloon on the surface. The red curves indicate large areas free from NLC particles (partial ice
494	voids).
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496	Figure 6. A schematic representation of the intrusion of warm air masses from mid- to sub-
497	polar latitudes, forming partial ice voids in the observed NLC. A general concept of this
498	scheme is analogous to the formation of a wave cyclone in the troposphere (see Figs. 8.18 and
499	8.19 in Ahrens, 1993).
500	
501	<b>Figure 7.</b> The SONC balloon image taken at 6222 m above the ground at 21:49 UT on 5 July
502	2018. The red arrow marks the vertical modulation of the NLC layer by a gravity wave of
503	medium scale. The green arrow indicates small-scale billow-type gravity waves.
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505	Figure 8. The SONC balloon image taken at 20.3 km above the ground at 22:48 UT on 5 July
506	2018 represents the final stage of NLC evolution on that night.





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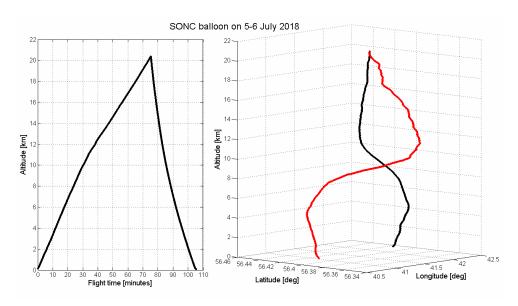


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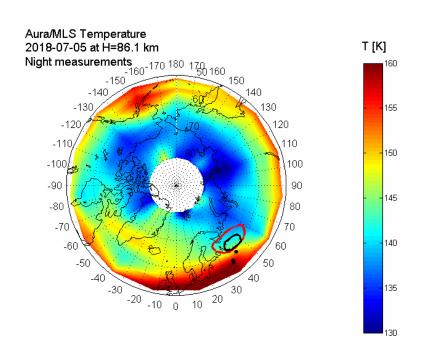


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**Figure. 2**. (Left) the altitude of the SONC balloon as a function of time flight. (Right) the vertical-horizontal trajectories of the SONC balloon: the red line is the upleg and the black line is the downleg trajectories.





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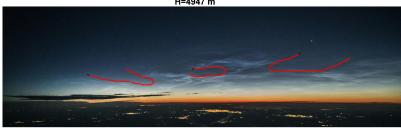
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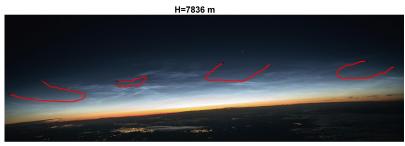




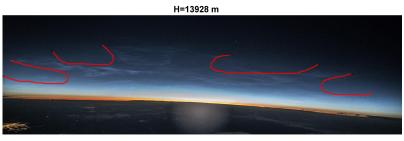












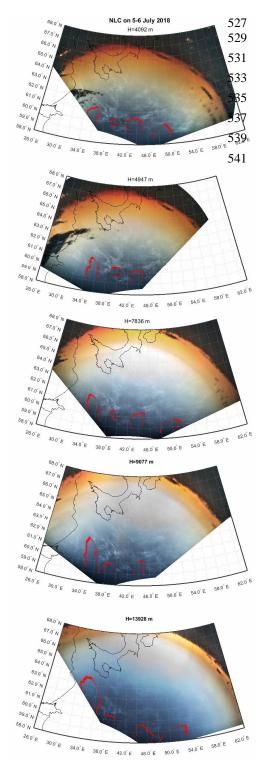
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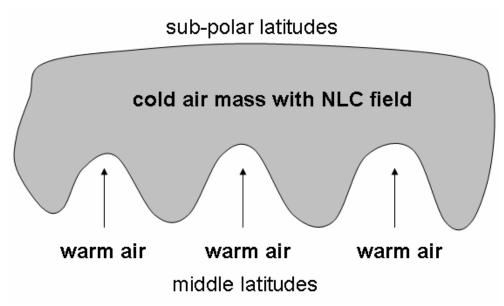
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Figure 8. The SONC balloon image taken at 20.3 km above the ground at 22:48 UT on 5 July

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