Dear Editor,
Please consider our revised manuscript according to the comments of the Reviewer 3. We thank the Reviewer 3 for her/his useful comments, which led to improvements of our manuscript. In the revised version of the manuscript we have carefully addressed the comments of the Reviewer 3 which are highlighted in green. Our detailed replies are provided below.

Submitted on 20 Nov 2019
Anonymous Referee #3

Suggestions for revision or reasons for rejection (will be published if the paper is accepted for final publication)

This revised paper describes NLC observations from a stratospheric balloon carried out during a campaign in the Moscow area during the 2018 NLC season. The observations are novel and have the potential to provide new and interesting insight into NLC and the processes affecting NLCs. The paper does not really present spectacularly new results, but is still an interesting contribution to the field and should be published subject to some minor revisions, in my opinion. The paper is overall well written and easy to follow.

Specific comments:

Line 183 following: the vertical gradient of potential temperature was determined from MLS temperature measurements. The vertical resolution of MLS temperature profiles in the upper mesosphere is with 12 – 14 km quite low. Given this poor vertical resolution, how reliable can estimates of the potential temperature gradient at NLC altitude be? I'm not convinced one can draw robust conclusions.

We use the newest version of the Aura/MLS temperature data (ver. 4.23). According to the data quality document (available at https://mls.jpl.nasa.gov/data/v4-2_data_quality_document.pdf), the vertical temperature resolution is 6 km at 0.01 hPa and 8–10 km at 0.001 hPa. Since we consider the temperature data in the mesosphere and up to the mesopause level (0.0046 hPa), the actual vertical temperature resolution is about 6–8 km (not 12–14 km). Since we have considered the temperature data covering a large geographic area (the NLC area observed from the balloon) and we could not find any negative values of the potential temperature gradient, we consider this conclusion to be robust.

Line 241 following: How can you distinguish between an altitude variation and a variation of the horizontal distance between the cloud and the balloon. In other words, how can you distinguish between a vertical and a horizontal wave pattern? I don't think a unique altitude determination is possible based on the ballon images only.

First. We consider relative variations of a wave motion, i.e., ±variations relative to some undisturbed level. We use snapshots with a short exposure time of 1/8 sec. The horizontal speed of the balloon was about 16 m/s at the time of observing the wave with vertical disturbances. Thus, the horizontal motion of the balloon was equal to 2 metres during a snapshot which yields negligibly small uncertainties for estimating relative vertical variations of the wave motion. The

position of the balloon was known and was taken into account in the calculations.

Second. For this peculiar wave (vertical wave pattern) we assume that the observed wave disturbances occur in the vertical plane only. If a significant horizontal wave pattern were present at that time we would get different vertical amplitudes for the wave crest and trough due to different changes in the elevation angles in the wave crest and trough. But since we have estimated the same wave amplitudes both for the wave crest and trough it means that the wave motion was indeed in the vertical plane.

Thus a unique altitude determination based on the balloon images is possible. We have added this information on lines 247-250 in the revised manuscript.

Line 258: “kinetic wave energy“
Strictly speaking, the „E“ in the equation provided is not an energy, i.e., its unit is not Joule.
Yes, it is correct. This is the wave energy per unit mass. We have corrected it on lines 261-262 in the revised manuscript.

Line 531: „at the mesopause (86.1 km)“
Is this geometric or geopotential altitude?
This is geometric altitude. We have corrected it in the caption of Figure 3.

Typos, grammar etc.:
Line 67: “are limited to” -> “are limited by”
It has been corrected.

Line 174: “those centers” -> “whose centers” ?
It has been corrected.

Line 525: “build” -> “built”
It has been corrected.

Line 529: “trajectories” -> “trajectory”
It has been corrected.

Stratospheric observations of noctilucent clouds: a new approach in studying middle- and large-scale mesospheric dynamics

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Abstract.

The experimental campaign Stratospheric Observations of Noctilucent Clouds (SONC) was conducted on the night of 5-6 July 2018 with the aim of photographing noctilucent clouds (NLCs) and studying their large-scale spatial dynamics at scales of 100–1450 km. An automated high-resolution camera (equipped with a wide-angle lens) was lifted by a stratospheric sounding balloon to 20.4 km altitude above the Moscow region in Russia (~56°N; 41°E), taking several hundreds of NLC images during the flight that lasted 1.7 hours. The combination of a high-resolution camera and large geographic coverage (~1500 km) have provided a unique technique of NLC observations from the stratosphere, which is impossible to currently achieve either from the ground or space. We have estimated that a horizontal extension of the NLC field as seen from the balloon was about 1450 x 750 km whereas it was about 800 x 550 km as seen from the ground. The NLC field was located in a cold area of the mesopause (136-146 K), which was confirmed by satellite measurements. The southmost edge of the NLC field was modulated by partial ice voids of 150-250 km in diameter. A medium-scale gravity wave had a wavelength of 49.4±2.2 km and amplitude of 1.9±0.1 km. The final state of the NLC evolution was represented by thin parallel gravity wave stripes. Balloon-borne observations provide new horizons in studies of NLCs at various scales from metres to thousands of km. Here we present a review paper on our experiment

describing initial results. Detailed studies on time evolution of the cloud movements will be done in the future.

Keywords: noctilucent clouds, mesospheric dynamics, balloon-borne stratospheric observations, atmospheric gravity waves

1 Introduction

Night-shining clouds or noctilucent clouds (NLCs) are the highest clouds in the Earth's atmosphere observed at the summer mesopause between 80 and 90 km. NLCs can be readily seen from mid- and subpolar latitudes of both hemispheres. NLCs are composed of water-ice particles of 30–100 nm in radius that scatter sunlight and thus NLCs are observed against the dark twilight arc from May until September in the Northern Hemisphere and from November to February in the Southern Hemisphere (Bronshten and Grishin, 1970; Gadsden and Schröder, 1989; Liu et al., 2016). NLCs are also observed from space and in this case they are usually called Polar Mesospheric Clouds (PMCs) (Thomas, 1984).

NLCs are almost always represented by a wave surface having a complex interplay between small-scale turbulence processes of 10-1000 metres, atmospheric gravity waves (GW) with wavelengths of 10-1000 km, planetary waves, solar thermal tides and lunar gravitational tides of about 10000 km (Witt, 1962; Fritts et al., 1993; Rapp et al., 2002; Kirkwood and Stebel, 2003; Chandran et al., 2009; Dalin et al., 2010; Fiedler et al., 2011; Taylor et al., 2011; Pertsev et al., 2015). Sometimes, distinguished non-linear mesospheric phenomena like mesospheric walls or fronts appear at the mesopause which clearly separate two volumes of the mesopause having cold and warm air masses with temperature difference of 20-25 K across a few km (Dubietis et al., 2011; Dalin et al., 2013).

NLCs/PMCs are systematically observed and studied from the ground (optical imagers, lidars), as well as from space (The Aeronomy of the Ice in the Mesosphere (AIM), Odin, Solar Backscatter Ultraviolet Radiometer (SBUV) instruments) (e.g., Karlsson and Gumbel, 2005; Dalin et al., 2008; Bailey et al., 2009; Fiedler et al., 2011; DeLand and Thomas, 2015); there are also irregular (campaign-based) NLC observations conducted by using sounding rockets and aircraft (Zadorozhny et al., 1993; Gumbel and Witt, 2001; Reimuller et al., 2011). These techniques have advantages and disadvantages. In particular, ground-based measurements provide a high horizontal resolution of ~20 m and high temporal resolution of seconds (optical imagers) (Dalin et al., 2010; Baumgarten and Fritts, 2014) and high vertical resolution of 50-150 metres using lidars (Baumgarten et al., 2009) but are limited ~~to~~ by tropospheric weather conditions and restricted to a certain small region on the Earth's surface. Satellite measurements, on the other hand, provide global PMC coverage but have low spatial horizontal resolution (~5 km) as well as large spatial gaps of several hundreds of km between adjacent orbits at middle and subpolar latitudes. Thus, there is no perfect technique to observe and study NLCs/PMCs so far. At the same time, observations made from stratospheric altitudes (20-40 km) are potentially available for comprehensive studies of NLCs/PMCs. So far, there have been conducted three published experiments from stratospheric balloons providing NLC/PMC observations. The first one was performed over Antarctica between 29 December 2012 and 9 January 2013 (Miller et al., 2015). The E and B Experiment (EBEX) was dedicated to another research field concerning polarization in the cosmic microwave background (Reichborn-Kjennerud et al., 2010). At the same time, two star cameras of the EBEX experiment, having a narrow field of view of 4.1° x 2.7°, were able to register fine structures of PMCs and turbulence dynamics, ranging from several km down to 10 m. Another balloon-borne experiment (PMC-Turbo) was conducted between 8 and 14 July 2018 over Sweden-Greenland-Canada territories in order to capture NLCs with seven optical cameras and lidar (Fritts et al., 2019). The PMC-Turbo experiment was launched about 2.5 days after the experiment described in the present paper.

In this paper, we report on scientific results of a new balloon-borne experiment dedicated to studies of NLC middle- and large-scale dynamics at horizontal scales of more than 100 km (Dalin et al., 2019). Such experiment, conducted for the first time, opens new horizons for studies of middle- and large-scale dynamical features in combination with a high spatial resolution at the summer mesopause, currently unachievable for other techniques like ground-based and space measurements.

2 Technique and method

The Stratospheric Observations of Noctilucent Clouds (SONC) experiment is a special balloon-borne experiment dedicated to studies of large-scale dynamical features in NLCs. A high resolution high sensitive camera (Sony Alpha A7S), having a full frame 35 mm 12 megapixel sensor (4240 x 2832 pixels) and equipped with a wide-angle lens (field of view, FoV, is 109.7° x 81.6°), has been installed on a meteorological sounding balloon. This combination of a high resolution sensor and wide FoV yields spatial horizontal resolutions of ~30 m and ~3000 m, when looking at 83 km from 20 km at elevation angles of 90° and 0°, respectively. The horizontal coverage of a mesopause layer is over 2000 km, when viewing along the horizon at low elevation angles. The balloon was launched from the Moscow region, Russia (~56°N; 41°E), on the night of 5-6 July 2018. Since a gondola payload is constantly rotating and shaking during its flight, the NLC camera was installed on a special stabilized platform. The 3-axis motorized gimbal stabilized platform (Fig. 1) was designed and built by the Aerospace laboratory "Stratonautica" (http://stratonautica.ru), which has a wide experience in building such platforms and launching sounding balloons. The platform was designed to rotate in a 60° step in the azimuth angle in order to capture the whole hemisphere (360°) since NLCs can appear in any direction as observed from mid-latitudes, including the southern part of the sky (Hultgren et al., 2011; Suzuki et al., 2016). The NLC camera took images every 6 s during the whole flight, obtained several thousands of images and several hundreds of images capturing NLCs. Besides, automatic exposure bracketing was used to take four images in sequence with different exposures, allowing us to register various NLC brightness from very bright to very faint as well as faint stars, which are important information for the photogrammetric technique and georeference procedure of the images.

The balloon was launched at 21:34 UT on 5 July 2018 and the total flight duration was about 1.7 hours. The ascent speed was around 5 m/s and the balloon reached its maximum altitude of 20.4 km where it burst; then the payload descended with a parachute and the payload was successfully recovered. A GPS receiver was installed onboard in order to obtain information on the balloon trajectory. The flight characteristics of the SONC balloon are shown in Fig. 2.

A ground-based support consisting of three automated NLC cameras was established in the Moscow region in order to launch the balloon at the time of NLC appearance. Also, a number of amateur observers significantly contributed to the NLC observational programme before and during the flight. A launch window was preliminarily chosen at the beginning of July based on long-term statistics of NLC observations conducted in the Moscow region since 1962 to present time. This statistics demonstrate that NLCs appear at the beginning of July with about 65% occurrence probability on a clear night.

3 The observation

During the flight, the balloon-borne camera captured an extended NLC field with a number of interesting features discussed in section 4. One can note the following general characteristics of the NLC display:

a) NLCs were observed between 20:30 and 23:15 UT (23:30 and 02:15 LT) on 5 July 2018.

b) NLCs were located between 82.6 and 85.1 km. The NLC height was estimated by using synchronously taken images obtained from two ground-based cameras located in the Moscow region.

c) NLC field extended along the horizon from NW to NE at low elevation angles from –5° to +11° as seen from the balloon.

d) NLCs were modulated by atmospheric gravity waves of various scales having horizontal wavelength from 9 km to 50 km.

e) NLCs were traveling in a rather unusual direction from the south to north at the observed mean speed of ~43 m/s.

f) NLCs were fading during the balloon ascent and they got very faint and less structured at the maximum balloon altitude of 20.4 km. The brightest and well-developed NLCs were observed when the SONC balloon was between 6 and 13 km, that is why we analyze the most profound features of NLC images obtained at this height range.

Each analyzed NLC image was georeferenced using horizontal coordinates of referenced stars (at least 15 stars are needed). The technique of the NLC georeference, triangulation height estimation and error analysis can be found in Dalin et al. (2004, 2015).

4 Results and discussion

The projection of the NLC field on the surface along with the temperature map obtained with the Aura/MLS spectrometer is shown in Fig. 3. The description on the MLS temperature product and its validation can be found in Froidevaux et al. (2006) and Schwartz et al. (2008). One can see that the NLC field (their actual coverage) extended mostly from the west to east along an area filled with low temperatures of 136-146 K, and the NLCs were located north of 58°N due to rapidly increasing temperature with decreasing latitude. That is why the NLCs were observed at low elevation angles (far to the north as seen from the Moscow region) on this particular night.

Detailed analysis of five consecutive in time balloon-borne images (Figs. 4 and 5) has revealed the following features of the NLC display:

a) The horizontal extent of the NLC field from the western to eastern observable border was about 1450 km, and from the northern to southern border of about 750 km. Such distances are impossible to observe from the ground due to the Earth’s curvature and limited area of the twilight arch. The central part of the NLC field, having extension of about 850 x

550 km, was seen from the ground but the western and eastern wings of the field as well as the northern edge were located below the local ground horizon, making it impossible to observe them. Thus, balloon-borne NLC observations have an obvious great advantage over ground-based observations in terms of larger geographic coverage which is comparable to PMC observations made from space since a PMC observation scene has spatial coverage of about 2000 km along the AIM satellite track and 1000 km across track (Rusch et al., 2009).

b) The southmost edge of the NLC field was modulated by partial circles (something like ice voids but with open southern border), ~~those~~ whose centers are shown by the red arrows in Figures 4 and 5. The diameters of these partial ice voids are estimated to be in the range of 150-250 km. The mechanism of the formation of ice voids in NLCs/PMCs is not clear now, and it is an ongoing topic in atmospheric physics. One can mention three main mechanisms which are currently discussing in the literature. Trubnikov and Skuratova (1967) addressed a theory of cellular convection and demonstrated its principal possibility in the summer mesosphere in relation to NLC occurrences. The authors estimated convective cells to be in the range of 90-250 km in radius, that agrees well with sizes of partial ice voids obtained in the present study. However, the main criterion for the convection to be developed, namely, the height gradient of the potential temperature should have negative values. We have carefully estimated the potential temperature gradient or the static stability (based on Aura/MLS temperature measurements) in the analyzed area and could not find any signatures of its negative values in the mesosphere and mesopause region. It means that in this particular case cellular convection could not be responsible for the observed partial ice voids in the NLCs.

However, satellite measurements can easily miss a negative static stability at local scales due to poor horizontal resolution and local ice voids may be generated by a gravity wave breaking. Rusch et al. (2009) have hypothesized that ice voids could be caused by heating due to the passage of warm crests of a gravity wave. It is possible in the present case. However, we could not find any significant displacement of the partial ice voids (their boundaries) relative to the NLC field, i.e., the partial ice voids traveled with the same speed and direction as the entire NLC field did (~43 m/s from the south to north). One would expect an intrinsic phase speed and intrinsic direction of the movement of the partial ice voids if they were generated by a large-scale gravity wave of a wavelength of several hundreds of km. Thus, it is difficult to prove this hypothesis of the influence of a large-scale gravity wave on the formation of the observed partial ice voids.

Thurairajah et al. (2013b) have proposed another mechanism related to a shock wave generated by a meteorite, which expands and cools the air that in turn leads to the formation of large ice particles which fall out of an NLC field (analogously to hole-punch clouds due to the passage of an aircraft). However, we observe large-scale partial ice voids (150-250 km) in a broad area of the mesopause over 1000 km. It was hardly possible that any big meteorite could produce such large holes in such broad area, and we did not observe any meteor motion in our ground-based and balloon images.

Megner et al. (2018) have recently presented an interesting case study of a quasi-stationary ice void in NLCs which did not follow the general wind, suggesting that it was formed by a localized warming at the summer mesopause. This is not the case in our case study, in which we have observed partial ice voids moving at the general wind speed in the same direction along with the entire NLC field.

In the present case study, the partial ice voids had irregular shape and sizes ranging from 150 to 250 km. Also, these partial voids moved along the wind, having the same speed and direction. Thus, it is difficult to connect these partial voids with regular wave disturbances. At the same time, as shown in Fig. 3, the southmost border of the NLC field was confined to the warm air mass located at sub-polar latitudes of ~58°N and lower. The mesopause temperature at this border was equal to ~147 K at 86 km altitude. The MLS data cannot reproduce the exact shape of this border due to low horizontal resolution (~15°) and temporal resolution of ~1.5 h. However, it is well known that tropospheric frontal systems have a meandering shape, sometimes with intrusions of warm and cold air masses as in case of the formation of a frontal wave cyclone (Ahrens, 1993; Stull, 2000). In our case the warm front at the mesopause and the NLC partial ice voids resemble a tropospheric frontal wave, in which there are intrusions of warm air masses, moving from midlatitudes, into the cold air mass located at sub-polar and polar latitudes (see Fig. 6). Therefore, we consider that the most probable source of these partial ice voids observed in the NLCs in this particular case is the intrusion of warm air masses into the cold air mass with the NLC field, sublimating ice particles. A similar conclusion was proposed by Thurairajah et al. (2013a) who have analyzed a large ice void observed in PMCs (using AIM/CIPS satellite images) and have concluded that "…*warmer temperatures (warmer than the frost point temperature of ~144 K) at the location of the void may be related to increased tidal activity and transport of warm air from low latitudes*." Also, Rusch et al. (2009) and Thurairajah et al. (2013b) have demonstrated that southmost borders of PMCs can be highly modulated by partial ice voids of several hundreds of km in diameter, and the authors have found the structural similarity between PMC images and those seen in tropospheric clouds.

c) Clear vertical modulation of the NLC layer is shown with the red arrow in Fig. 7. This is a unique view on a particular gravity wave seen at the local horizon of the balloon; that is why this wave modulation is viewed almost at the right angle to the line-of-sight. This geometry allows observing a thin layer of NLC modulated in altitude by propagating gravity waves of small and medium scales. Such geometry is almost impossible to obtain from the ground since NLCs seen at the very horizon are usually masked by topography, tropospheric clouds and, most importantly, by tropospheric aerosols, which are constantly present and significantly absorb NLC brightness when looking at the very horizon. We have carefully estimated parameters of this particular wave: its horizontal wavelength was equal to 49.4±2.2 km and its amplitude was 1.9±0.1 km. We define this amplitude as a semi-amplitude A of a monochromatic wave with oscillation frequency ω, which is half of the peak-to-peak wave amplitude between the highest (crest) and lowest (trough) displacement values. In this calculation, the angle of 13.3° between the camera image plane and vertical plane at the NLC altitude was taken into account. A unique altitude determination based on the balloon images is possible under the assumption that the observed wave amplitude occurred in the vertical plane only and due to the fact that the balloon traveled a small horizontal distance (about 2 m) during a short exposure time of 1/8 s. Also note that since NLCs are clearly seen both in the crest and trough of the wave (ice particles did not completely sublimated in the wave trough), we have estimated the wave amplitude both in the wave crest and trough. The amplitude estimations are the same in the wave trough and crest (within the given uncertainty). All this makes us confident in the estimation of the amplitude of this particular wave. We have analysed nine images at various viewing angles in order to deduce the maximum vertical displacement (amplitude) of this particular wave. The nine images showing progressive changes in the wave vertical displacement can be found at the following webpage: ftp://ftp.irf.se/outgoing/pdalin/NLC/SONC_experiment_2018_07_05/WAVE_AMPLITUDE/ This is the most precise estimation of the amplitude of a gravity wave at the mesopause by using NLC observations (Witt, 1962; Haurwitz and Fogle, 1969; Bronshten and Grishin, 1970; Demissie et al., 2014). Since wave amplitude represents wave energy per unit mass ($E \sim 0.5 \cdot A^2 \cdot \omega^2$), this is an important source of information for estimating the wave energy budget at the upper atmosphere, and also can be used for future model studies to estimate a wave source in the lower atmosphere (Demissie et al., 2014).

d) Small-scale billow-type gravity waves were estimated to have horizontal wavelengths of 8-11 km (Fig. 7). Such small-scale gravity waves are well-known to be observed in NLC layers (Witt, 1962; Dalin et al., 2010; Pautet et al., 2011; Baumgarten and Fritts, 2014;

Demissie et al., 2014), but we demonstrate this result in order to emphasize the ability to resolve small-scale NLC structures by using a large FoV camera, having a high resolution sensor, onboard a sounding balloon.

e) Figure 8 illustrates an NLC image taken from altitude of 20.3 km which is very close to the maximum reached altitude of 20.4 km. The NLCs were rather faint by that time that is in line with an idea of the intrusion of warm air masses from mid- to subpolar latitudes. These large-scale warm air masses led to rapid sublimation of ice particles at large scales of about 1500 km. At the same time, on can see a very interesting feature to be considered. There were several thin parallel gravity wave bands (stripes) with lengths of 50-200 km and widths of ~ 3-5 km in cross-section. The reasons of seeing such thin stripes are as follows: (a) The SONC balloon was in the stratosphere, i.e., above the troposphere in which optically strong air turbulence is constantly present (b) The exposure time of this image was very short of 1/125 s. All these made the image free from blurring (as minimum blurring as possible for moving NLCs and balloon motion). This image demonstrates a final stage of the NLC evolution (NLCs disappeared in 20 min since the image was taken), and these thin stripes might represent a final morphological state of the NLC evolution. Further balloon-borne NLC observations of very faint NLCs are required to confirm this consideration.

5 Conclusions

The combination of high resolution images (~30 m) and large geographic coverage (over 1500 km) is a unique property intrinsic to stratospheric balloon-borne NLC observations, which is impossible to achieve either from the ground or space. In general, a balloon-borne NLC observation provides us with the following new opportunities in case of a long duration flight of several days:

a) NLC imaginary can be obtained for 24 hours a day and during several days due to very little Rayleigh atmospheric scattering in the visible subrange of the spectrum above 20 km (Hughes, 1964);
b) Quantitative information on a wide range of waves (gravity and planetary waves, solar tides), propagating through the summer mesopause can be obtained;
c) Neutral wind velocity at the mesopause and large-scale trajectory of NLC fields over 1500 km can be measured;
d) Quantitative information on long mesospheric fronts, solitons and other non-linear processes can be obtained;

e) Quantitative information on small-scale turbulent structures (down to 1 m) can be obtained in case of using a narrow field of view lens.

f) High resolution vertical NLC structure (wave modulation, double layers) can be retrieved by observing NLCs at the very horizon. Absence of any terrain obstacles and tropospheric aerosol loading makes such stratospheric NLC observations unique.

g) Absence of optically strong tropospheric turbulence makes NLC images free from atmospheric blurring that in turn results in well-defined fine structures of gravity waves and turbulence in the mesopause region.

In the present study, we have estimated the following characteristics of the NLC field:

a) The horizontal extent of the NLC field as seen from the SONC balloon was about 1450 x 750 km whereas it was about 800 x 550 km as seen from the ground. This emphasizes the great advantage of making large-scale balloon-borne observations over medium-scale ground-based ones.

b) NLC field was traveling from the south to north at a mean velocity of 43 m/s;

c) The southmost edge of the NLC field was modulated by partial ice voids of 150-250 km in diameter, which were like generated by the intrusion of warm air masses moving from mid- to sub-polar latitudes. The mesopause temperature at this edge was equal to ~147 K, i.e., it was a threshold temperature separating the mesopause region filled with NLCs from the warm area without NLCs.

d) A medium-scale wave had a wavelength of 49.4±2.2 km and vertical amplitude of 1.9±0.1 km. This is the most precise estimation of a gravity wave amplitude ever made.

e) Small-scale billow-type gravity waves had wavelengths of 8-11 km.

f) The final morphology state of the NLC evolution was represented by thin parallel gravity wave stripes with lengths of 50-200 km and widths of ~3-5 km.

Data availability. The reader can access the SONC experiment images and balloon GPS coordinates, used in the paper, via publically available project ftp server at the Swedish Institute of Space Physics:
ftp://ftp.irf.se/outgoing/pdalin/NLC/SONC_experiment_2018_07_05/

Author contributions. PD wrote the paper, made calculations and plotted the figures. NP and VP read and made suggestions appropriated for the paper. DE provided the raw balloon-borne images and balloon GPS coordinates. VR contributed to the image processing. All the authors read and commented regarding the work and agreed with the content and submission of this paper.

Competing interests. The authors declare that they have no conflict of interest.

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Figure captions:

Figure 1. The 3-axis motorized gimbal stabilized platform, holding the NLC camera, designed and built by the Aerospace laboratory "Stratonautica". Photo by Denis Efremov.

Figure. 2. (Left) the altitude of the SONC balloon as a function of time flight. (Right) the vertical-horizontal trajectories of the SONC balloon: the red line is the upleg and the black line is the downleg ~~trajectories~~ trajectory.

Figure 3. The temperature map at the mesopause (geometric height of 86.1 km) as measured by the Aura/MLS spectrometer on 5 July 2018. Nighttime measurements around the globe have been selected to produce the map. Upon the temperature map, the outer borders of the NLC field (the actual NLC coverage) are overplotted: the red line is as seen from the SONC balloon, the black line is as seen from the ground at the launch. The black dots mark the position of the balloon at 7.8 km at the ground and ground-based observers.

Figure 4. The NLC field as observed from the SONC balloon at 4092 m, 4947 m, 7836 m, 9077 m and 13928 m above the ground at 21:46 UT, 21:49 UT, 21:57 UT, 22:01 UT, 22:20 UT on 5 July 2018. The red arrows indicate the centers of large areas free from NLC particles (partial ice voids).

Figure 5. Projection of the NLC fields (shown in Figure 4) as observed from the SONC balloon on the surface. The red arrows indicate the centers of large areas free from NLC particles (partial ice voids).

Figure 6. A schematic representation of the intrusion of warm air masses from mid- to sub-polar latitudes, forming partial ice voids in the observed NLCs. A general concept of this scheme is analogous to the formation of a wave cyclone in the troposphere (see Figs. 8.18 and 8.19 in Ahrens, 1993).

Figure 7. The SONC balloon image taken at 6222 m above the ground at 21:49 UT on 5 July 2018. The red arrow marks the vertical modulation of the NLC layer by a gravity wave of medium scale. The green arrow indicates small-scale billow-type gravity waves.

Figure 8. The SONC balloon image taken at 20.3 km above the ground at 22:48 UT on 5 July 2018 represents the final stage of NLC evolution on that night.