



1 **Ionospheric Total Electron Content responses to HILDCAAs intervals**

2 Regia Pereira Silva ^{*1,2}, Clezio Marcos Denardini¹, Manilo Soares Marques³, Laysa
3 Cristina Araújo Resende^{1,4}, Juliano Moro^{4,5}, Giorgio Arlan da Silva Picanço¹, Gilvan
4 Luiz Borba³, Marcos Aurélio Ferreira dos Santos².

5

6

7 ¹National Institute for Space Research – INPE, São José dos Campos-SP, Brazil.

8 ²Northeast Regional Center – CRN/INPE, Natal-RN, Brazil.

9 ³Geophysics Department (DGEF), Federal University of Rio Grande do Norte
10 (UFRN), Natal-RN, Brazil.

11 ⁴National Space Science Center, China Academy of Science, CAS, Beijing, China;

12 ⁵Southern Regional Space Research Center – CRS/COCRE/INPE, Santa Maria-RS,
13 Brazil.

14

15

16

17 ^{*}Corresponding author: INPE/DIDAE, Av. dos Astronautas, 1758, Jd. da Granja,

18 12227-010, São José dos Campos, SP, Brazil; Tel.: +55 12 3208-7184; e-mail:

19 regia.pereira@inpe.br; regiapereira@gmail.com

20

21

22



23 **Abstract**

24 The High-Intensity Long-Duration and Continuous AE Activities (HILDCAA)
25 intervals are capable of causing a global disturbance in the terrestrial ionosphere.
26 However, the ionospheric storms' behavior due to these geomagnetic activity forms is
27 still not widely understood. In this study, we seek to comprise the HILDCAAs
28 disturbance time effects in the Total Electron Content (TEC) values with respect to
29 the quiet days' pattern analyzing local time and seasonal dependences, and the
30 influences of the solar wind velocity to a sample of ten intervals occurred in 2015 and
31 2016 years. The main results showed that the hourly distribution of the disturbance
32 TEC may vary substantially between one interval and another. Doing a comparative
33 to geomagnetic storms, while the positive ionospheric storms are more pronounced in
34 the winter, this season presents less geoeffectiveness or almost none to HILDCAA
35 intervals. It was find an equinoctial anomaly, since the equinoxes represent more
36 ionospheric TEC responses during HILDCAA intervals than the solstices. Regarding
37 to the solar wind velocities, although HILDCAA intervals are associated to High
38 Speed Streams, this association does not present a direct relation regards to TEC
39 disturbances in low and equatorial latitudes.

40

41

42

43

44

45 *Keywords:* HILDCAA, TEC, Equatorial Ionosphere

46



47 **1. Introduction**

48 As similar to geomagnetic storms, High-Intensity Long-Duration and Continuous AE
49 Activities (HILDCAA) intervals can influence the ionosphere, leading to disturbances
50 in the ionospheric F2-region. It is well known that these intervals can change the F2-
51 region peak height being, generally, less intense than those observed during typical
52 geomagnetic storm events (Sobral et al., 2006; Koga et al., 2011, Silva et al., 2017).

53 In fact, HILDCAAs are characterized by present some criteria: i) the AE index must
54 reach an intensity peak greater than or equal to 1000 nT; ii) The AE index needs to be
55 almost continuous and never drops below 200 nT for more than two hours at a time;
56 iii) The event must have a duration of at least two days, and iv) The event occurred
57 after the main phase of magnetic storms. However, the same physical process may
58 occur whether one of the four criteria are not strictly followed (Tsurutani and
59 Gonzalez, 1987; Tsurutani et al., 2004; Sobral et al., 2006, Tsurutani et al., 2006;
60 Hajra et al., 2013, Silva et al., 2017). As the main feature is the high AE index levels,
61 in this study we have considered drops below 200 nT for more than two hours as long
62 as the AE index value returns in high activity for prolonged hours.

63 The electron density perturbation in the ionosphere during HILDCAA events is
64 different from that one occurred during geomagnetic storms in the equatorial and low
65 latitudes stations. Since the HILDCAA presents a weak/moderate geoeffectiveness
66 when it compares to the other forms of space disturbances, it is expected that the
67 ionosphere response presents a differential behavior.

68 The Total Electron Content (TEC) is an important ionospheric parameter to several
69 studies and technologic applications. As HILDCAAs can cause F2-region peak
70 alterations, it can be observed the enhancements/depletions in TEC profile. In fact,
71 the TEC response to the geomagnetic storms is a well-known issue in the space



72 physics field (Lu et al., 2001; Kutiev et al., 2005; Mendillo, 2006; Maruyama and
73 Nakamura, 2007; Biqiang et al., 2007; de Siqueira et al., 2011). However, only few
74 studies about TEC pattern during HILDCAAs intervals have been found in the
75 literature.

76 Ionospheric storms are manifestations of space weather events, which are caused by
77 energy inputs in the upper atmosphere in the form of enhanced electric fields,
78 currents, and energetic particle precipitation (Buonsanto, 1999; Mendillo, 2006).
79 Usually, ionospheric storms are associated with ionosphere responses to geomagnetic
80 storm events. However, in a broader way, these responses happen due to
81 magnetospheric energy inputs to the Earth's upper atmosphere, and this can occur to
82 all kind of geomagnetic activity form. Park (1974) pointed that ionospheric storms
83 can be understood in terms of the superposed effects of many substorm. In view of
84 the foregoing and considering that the development of ionospheric storms during
85 HILDCAAs intervals has not been dealt with in depth, in the current study we have
86 focused the TEC pattern during this kind of event.

87 Recently, Verkhoglyadova et al. (2013) suggested that HILDCAAs associated with
88 High Speed Streams (HSS) can be one of the external driving TEC variabilities.
89 Indeed, the continuous energy injection and energetic particles precipitation into the
90 polar upper atmosphere during HILDCAA intervals could modify the dynamic and
91 chemical coupling process of the thermosphere-ionosphere system resulting in
92 changes in the electron density. These modifications, beyond to change the auroral
93 electron density, can be mapped to low latitudes involving electric fields
94 disturbances, as prompt penetration electric fields (PPEF) and disturbance dynamo
95 (DD) (Koga et al., 2011; Silva et al., 2017).



96 Therefore, in the current study we have focused the TEC pattern during HILDCAAs
97 intervals, taking account local time dependence, seasonal dependence and high/slow
98 speed streams influences in the equatorial and low latitude ionosphere. This paper is
99 structured as followed: in the next section we present the HILDCAA intervals chosen
100 to support this study as well as the GNSS receivers locations over the Brazilian
101 region. In section 3 we show the results and discussion of the analysis and the
102 conclusions are presented in the last section.

103

104 **2. Data and Methodology**

105 In this study was possible to construct an overall perception of the ionospheric storms
106 occurred during HILDCAA disturbance time intervals that affect the TEC values with
107 respect to the expected behavior for quiet days. The features studied are local time
108 and seasonal dependences, and solar wind velocity influences.

109 We have selected ten HILDCAA intervals occurred during the 2015 – 2016 period.
110 These intervals are listed in Table 1, where the two columns present the identification
111 and the data range of each interval. The geomagnetic indices and interplanetary data
112 used to classify the HILDCAA events were obtained from OMNIWeb
113 (<https://omniweb.gsfc.nasa.gov/ow.html>). The Kp index data were obtained from the
114 World Data Center for Geomagnetism, Kyoto, Japan ([http://wdc.kugi.kyoto-](http://wdc.kugi.kyoto-u.ac.jp/kp/index.html)
115 [u.ac.jp/kp/index.html](http://wdc.kugi.kyoto-u.ac.jp/kp/index.html)). In this work it was used the daily Kp sum value.

116 The TEC mean was initially processed by a program developed at the Institute for
117 Space Research, Boston College, USA (Krishna, 2017). The mean values of vertical
118 TEC (VTEC) were obtained from two Brazilian GNSS stations, São Luís (SL) (2,59
119 S; 44,21 W) and Cachoeira Paulista (CP) (22,68 S; 44,98 W), representing the station
120 closest to the equator and the low latitude station, respectively. The Rinex files used



121 in this study were obtained from Brazilian Network for Continuous Monitoring of the
122 GNSS-RBMC Systems (RBMC) ([https://www.ibge.gov.br/en/geosciences/geodetic-](https://www.ibge.gov.br/en/geosciences/geodetic-positioning/geodetic-networks/20079-brazilian-network-for-continuous-monitoring-of-the-gnss-systems-2?=&t=o-que-e)
123 [positioning/geodetic-networks/20079-brazilian-network-for-continuous-monitoring-](https://www.ibge.gov.br/en/geosciences/geodetic-positioning/geodetic-networks/20079-brazilian-network-for-continuous-monitoring-of-the-gnss-systems-2?=&t=o-que-e)
124 [of-the-gnss-systems-2?=&t=o-que-e](https://www.ibge.gov.br/en/geosciences/geodetic-positioning/geodetic-networks/20079-brazilian-network-for-continuous-monitoring-of-the-gnss-systems-2?=&t=o-que-e)). Besides that, the TEC data during HILDCAA
125 events were analyzed and then compared with a set of three days average belonging
126 to a quiet period, in which it refers to the three days less disturbed ($\Sigma Kp < 24$) of the
127 month of the occurrence of each HILDCAA interval.

128 Figure 1 shows a map with the location of each GNSS station, which is represented
129 by a red triangle. The dashed line represents the magnetic equator. The TEC data
130 obtained during the HILDCAA intervals were analyzed and then compared to the
131 TEC data during the selected quiet days, resulting in dTEC (dTEC = TEC mean –
132 TEC quiet days). All the analyses done in this work took into account the dTEC
133 values.

134

135 **3. Results and Discussions**

136 In this section, we will present the ionospheric TEC responses observed during ten
137 HILDCAA intervals focusing on local time dependence and seasonal features and the
138 solar wind velocity influences.

139

140 **3.1 Local time dependence**

141 A common feature of ionospheric storms is to be associated with dependence on local
142 time, mainly when they are caused by geomagnetic storms (Titheridge and
143 Buonsanto, 1988; Pedatella et al., 2010). However, to the best of the authors'
144 knowledge, no study has been found analyzing this aspect when regarding HILDCAA
145 intervals.



146 Figures 2 and 3 show the mean dTEC hourly values related to all HILDCAA intervals
147 for São Luís and Cachoeira Paulista, respectively. Each panel represents a single
148 interval from the bottom (H01) to the top (H10). The x axis is given in the Universal
149 Time ($LT = UT - 3$) and the color scale represents the dTEC values in TEC units
150 (TECu).

151 Notice that the dTEC values have a greater magnitude for the low latitude GNSS
152 station to the detriment of the closer equatorial GNSS station. The minimum and
153 maximum values are, respectively, -16.00 TECu and 27.40 TECu to São Luís, and -
154 37.60 TECu and 48.80 TECu to Cachoeira Paulista. It was considered the same
155 minimum and maximum values occurred to all intervals, for each station. This fact
156 explains why some intervals appear too close to the quiet time pattern. We believed
157 that since the HILDCAA events has low/moderate geoeffectiveness it was not
158 expected high values of the dTEC.

159 The distribution of the dTEC effects hour-to-hour during HILDCAA intervals shows
160 a substantial variability from one event to another. Habarulema et al. (2013) found
161 that the negative storms effects are observed during geomagnetic storms recovery
162 phases that over equatorial latitudes. However, since HILDCAAs intervals are
163 characterized by a long continuous phase of Dst index recovery, this does not apply.
164 The HILDCAA intervals present the positive dTEC predominance. In a more
165 simplified definition, HILDCAA means an interval where there is always energy
166 injection (Søraas et al., 2004; Sandanger et al., 2005). Silva et al. (2017) observed that
167 during HILDCAA intervals it was seen the uplift of the equatorial F2 region peak
168 height, probably due to prompt penetration electric fields. One of the main
169 mechanisms of TEC enhancements is the rise of the ionosphere to higher altitudes
170 where the recombination rates are small. Besides that, our results are in agreement



171 with the results found by de Siqueira et al. (2017). They did a study comparing the
172 TEC responses between two magnetic storms and two HILDCAAs intervals
173 following by them, and found a great TEC variability pattern from one to another
174 event. Hereupon, it was not possible to find a response pattern to the HILDCAA
175 effects in the equatorial and low latitude TEC considering only the local time. There
176 is great variability, and it is important to consider the day-to-day ionospheric
177 variabilities as well as the separate effect of each electric fields disturbance
178 (PPEF/DD).

179 Comparing both stations, Cachoeira Paulista GNSS station presented higher values
180 both to positive as negative ionospheric storms. During the daytime hours, the latitude
181 is responsible for the different ionospheric responses due to the presence of
182 photoionization. This probably explains the dTEC higher sensibility to low latitude
183 station in detriment of the closer equatorial latitude station.

184 Analyzing the hourly behavior of each interval from Figures 2 and 3, we observed
185 more intensity in TEC disturbances, both for positive and negative storms, during
186 some specific intervals. This aspect led us to make a seasonal analysis, which will be
187 presented in the next section.

188

189 3.2 Seasonal Dependence

190 It is well known for geomagnetic storms that the influence of the season entails on
191 positive/negative ionospheric storms is more pronounced in winter/summer than in
192 equinox months (Matsushita, 1959; Prölss and Najita, 1975; Mendillo, 2006, among
193 others). However, has not yet been established whether the occurrence of HILDCAA
194 interval in different seasons can do different TEC disturbances.



195 In a recent study involving more than one hundred HILDCAA events, Hajra et al.
196 (2013) reported no seasonal dependence, in what regards to predominant occurrence
197 rate in any specific epoch of the year due to the solar cycle influences. They
198 announced the HILDCAAs may occur during any month and any year, with increases
199 in the numbers of events occurring during the solar cycle descending phase. In the
200 current study, it was considered as seasonal dependence feature the TEC disturbances
201 responses at HILDCAA intervals already classified in a seasonal way. The years
202 2015 and 2016 years comprise the descending phase of the 24th solar cycle, which
203 made it possible to catalog an expressive number of HILDCAAs events in a short
204 time. Among the ten intervals chosen for this study, we have separated eight ones to
205 represent the seasonal variability, being two events for each station, taking into
206 account the month of occurrence of each interval, and considering the seasons as they
207 occur in South Hemisphere. The intervals are distributed according to the Table 2.
208 Figure 4 shows the disturbed TEC according to the seasonal classification which the
209 blue and coral colors refer to São Luís and Cachoeira Paulista, respectively. The solid
210 lines show an estimate of the central tendency for all values, minute-to-minute, for all
211 days of the events belongs to the season, while the shaded area represents the
212 confidence interval for that estimate. While the positive storms are more pronounced
213 in the winter for geomagnetic storms, to HILDCAA intervals this season presents less
214 geoeffectiveness, or almost none. Our results show that the equinoxes represent more
215 ionospheric TEC responses during HILDCAA intervals than the solstices. Both
216 equatorial and low latitude stations present positive storms during the autumn, while
217 the spring presents a negative behavior, mainly. This equinoctial anomaly may be
218 originated from the equinoctial differences in neutral winds, thermospheric



219 composition, and electric fields. Additional studies are necessary to quantify how
220 each factor can play an important role in HILDCAA seasonal TEC disturbances.

221

222 3.3 Solar wind velocities analysis

223 During the solar cycle descending phase, polar coronal holes migrate to lower
224 latitudes emanating intense magnetic fields. When HSS from these low latitudinal
225 coronal holes interact with slow speed streams (SSS) a region called Corotating
226 Interaction Regions (CIR) is formed and it is well characterized by compressions of
227 the magnetic field and plasma.

228 There are considerable works whose show how HILDCAA is well associate with
229 HSS and CIRs (Tsurutani et al., 2006; Verkhoglyadova et al., 2013). However, to be
230 associated not necessarily means that the degree of geoeffectiveness is directly related
231 to high speeds.

232 Figure 5 shows the solar wind velocities (V_{SW}) during each HILDCAA interval. As
233 the Figure 4, the blue and coral colors refer to São Luís and Cachoeira Paulista,
234 respectively. The diameter of the bubble is related to the velocity. The results showed
235 great variability from one interval to another, even considering the intervals that
236 occurred in the same year. In our first analysis (not shown here) we did not find a
237 direct association or cross-correlation between the VSW magnitude and the dTEC in
238 the equatorial and low latitude GNSS stations. Kim (2007) indicated that HILDCAA
239 intervals can be accompanied by HSS as well as SSS. It is possible to see in our
240 results that the dTEC responses to some intervals present similar behavior to both
241 HSS and SSS (e.g. H03, H07 and H08). This means that HILDCAA intervals can
242 affect the ionospheric TEC, but not in a direct correlation.

243



244 **4. Conclusions**

245 For this work, the ionospheric TEC response to a sample of ten HILDCAA intervals
246 has been studied. We have used two GNSS stations from RBMC network
247 representing equatorial and low latitude locations. As HILDCAA can affect the
248 equatorial ionospheric F2 region, some disturbed TEC from its quiet time pattern is
249 found. Addressing how the ionospheric storms behave during the HILDCAA
250 intervals is our main goal.

251 Summarizing, HILDCAAs geoeffectiveness in Earth is mainly associated with CIRs,
252 for this reason, the HILDCAA occurrence is more recurrent in the solar cycle
253 descending phase since CIRs play a major role during this phase. Their effects occur
254 during magnetic reconnection due to association with southward z component of the
255 interplanetary magnetic field and Alfvén waves present in it (Tsurutani et al., 2004).
256 These long-lasting intervals are due to continuous injection of energy and
257 precipitation of particles, which disturb the high latitude ionosphere. The mainly
258 disturbs are changes in thermospheric neutral composition, temperature, winds and
259 electric fields. Similar to geomagnetic storms, these disturbs can be mapped to low
260 and equatorial latitude and alter the quiet time ionosphere. However, generally, they
261 are less intense because in one astronomical unit the CIRs are not fully developed. In
262 this study we seek to understand the behavior of the ionospheric storm during
263 HILDCAA intervals. The main results are highlighted below:

- 264 • The hourly distribution of the dTEC during HILDCAAs intervals may vary
265 substantially between low and equatorial latitude. Probably, the photoionization
266 associated with latitude is responsible for these variations;



- 267 • Despite the geomagnetic storms recovery phase presents negative ionospheric
268 storms, this pattern do not occur during HILDCAA intervals. There is great
269 variability from one interval to another, but, predominantly, occurs positive phase;
- 270 • Regarding seasonal features, while the positive storms are more pronounced in the
271 winter for geomagnetic storms, this season present less geoeffectiveness, or almost
272 none to HILDCAA intervals. The equinoxes represent more ionospheric responses
273 to HILDCAA intervals presenting positive/negative phase predominance during
274 the autumn/spring;
- 275 • A well-known HILDCAA feature is its association with HSS present in the solar
276 wind. However, this association does not present a direct relation regards to TEC
277 disturbances in low and equatorial latitudes.
- 278 To conclude, the upshot of this study is the possibility to understand how ionospheric
279 storms behave during some HILDCAA intervals and to contribute to improving the
280 discussions about this issue.
- 281
- 282



283 **Data availability**

284 The data used in this work are made publicly available on the following sites:

285 <https://omniweb.gsfc.nasa.gov/ow.html> , <http://wdc.kugi.kyoto-u.ac.jp/kp/index.html>,

286 and <https://www.ibge.gov.br/en/geosciences/geodetic-positioning/geodetic->

287 [networks/20079-brazilian-network-for-continuous-monitoring-of-the-gnss-systems-](https://www.ibge.gov.br/en/geosciences/geodetic-positioning/geodetic-networks/20079-brazilian-network-for-continuous-monitoring-of-the-gnss-systems-)

288 [2?=&t=o-que-e](#) . The GPS-TEC program used in this work is available in

289 <http://seemala.blogspot.com/>

290

291 **Author contributions**

292 R. P. Silva conceived the study, designed the data analysis, discussed the results and
293 leaded writing this manuscript.

294 C. M. Denardini assisted to conceive the study, to design the GNSS data analysis and
295 discuss the final results.

296 M. S. Marques assisted with the GNSS data analysis and with designing the figures.

297 L. C. A. Resende assisted to design the study and discuss the results of the study.

298 J. Moro assisted to design the study and discuss the results of the study.

299 G. A. S. Picanço assisted to discuss the results of the study and review the
300 manuscript.

301 G. L. Borba assisted to discuss the results of the study and review the manuscript.

302 M. A. F. Santos assisted to discuss the results of the study and review the manuscript.

303 All the authors helped to write and to revise the manuscript.

304

305 **Competing interests**

306 The authors declare that they have no conflict of interest.

307



308 **Special issue statement**

309 This article is part of the special issue “7th Brazilian meeting on space geophysics
310 and aeronomy”. It is a result of the Brazilian meeting on Space Geophysics and
311 Aeronomy, Santa Maria/RS, Brazil, 5–9 November 2018.

312

313 **Acknowledgements**

314 R. P. Silva acknowledges the supports from Conselho Nacional de Desenvolvimento
315 Científico e Tecnológico (CNPq) through the Grant No. 300329/2019-9. C. M.
316 Denardini thanks to CNPq/MCTIC (Grant 303643/2017-0). L. C. A. Resende would
317 like to thank the National Space Science Center (NSSC), Chinese Academy of
318 Sciences (CAS) for supporting her postdoctoral. J. Moro would like to acknowledge
319 the China-Brazil Joint Laboratory for Space Weather (CBJLSW), National Space
320 Science Center (NSSC), Chinese Academy of Sciences (CAS) for supporting his
321 Postdoctoral fellowship, and the National Council for Scientific and Technological
322 Development (CNPq) for the grant 429517/2018-01. G. A. S. Picanço thanks CAPES
323 for supporting his Ph.D. (Grant 88887.351778/2019-00). We also would like to thank
324 the OMNI data (<https://omniweb.gsfc.nasa.gov/ow.html>), and the World Data Center
325 for Geomagnetism, Kyoto (<http://wdc.kugi.kyoto-u.ac.jp/kp/index.html>). The Rinex
326 files were obtained from Brazilian Network for Continuous Monitoring of the GNSS-
327 RBMC Systems (RBMC) at interface
328 [https://www.ibge.gov.br/en/geosciences/geodetic-positioning/geodetic-
329 networks/20079-brazilian-network-for-continuous-monitoring-of-the-gnss-systems-
330 2?=&t=o-que-e](https://www.ibge.gov.br/en/geosciences/geodetic-positioning/geodetic-networks/20079-brazilian-network-for-continuous-monitoring-of-the-gnss-systems-2?=&t=o-que-e). The authors acknowledge Gopi Seemala for making available the
331 GPS-TEC program (<http://seemala.blogspot.com/>).

332

333



334 **References**

- 335 Biqiang, Z., Weixing, W., Libo, L., and Tian, M.: Morphology in the total electron
336 content under geomagnetic disturbed conditions: results from global ionosphere
337 maps, *Ann. Geophys.*, 25, 1555-1568, <https://doi.org/10.5194/angeo-25-1555-2007>,
338 2007.
- 339 Buonsanto, M. J.: Ionospheric storms - A review. *Space Science Reviews*, 88(3-4),
340 563-601, <https://doi.org/10.1023/A:100510753>, 1999.
- 341 de Siqueira, P. M., de Paula, E. R., Muella, M. T. A. H., Rezende, L. F. C., Abdu, M.
342 A., and Gonzalez, W. D.: Storm-time total electron content and its response to
343 penetration electric fields over South America, *Ann. Geophys.*, 29, 1765-1778,
344 <https://doi.org/10.5194/angeo-29-1765-2011>, 2011.
- 345 de Siqueira Negreti, P. M., de Paula, E. R., and Candido, C. M. N.: Total electron
346 content responses to HILDCAAs and geomagnetic storms over South America, *Ann.*
347 *Geophys.*, 35, 1309-1326, <https://doi.org/10.5194/angeo-35-1309-2017>, 2017.
- 348 Habarulema, J. B., McKinnell, L. A., Burešová, D., Zhang, Y., Seemala, G., Ngwira,
349 C., Chum, J., Opperman, B.: A comparative study of TEC response for the African
350 equatorial and mid-latitudes during storm conditions *J. Atmos. Sol. Terr. Phys.* 102,
351 105-114, <https://doi.org/10.1016/j.jastp.2013.05.008>, 2013.
- 352 Hajra, R., Echer, E., Tsurutani, B. T., Gonzalez, W. D.: Solar cycle dependence of
353 High-Intensity Long-Duration Continuous AE Activity (HILDCAA) events,
354 relativistic electron predictors?: *J. Geophys. Res.* 118, 5626-5638,
355 <https://doi.org/10.1002/jgra.50530>, 2013.
- 356 Kim, H.: Study on the particle injections during HILDCAA intervals. *Journal of*
357 *Astronomy and Space Sciences*, v. 24, <https://doi.org/10.5140/JASS.2007.24.2.119>,
358 2007.
- 359 Koga, D., Sobral, J. H. A., Gonzalez, W. D., Arruda, D. C. S., Abdu, M. A., Castilho,
360 V. M., Mascarenhas, M., Gonzalez, A. C., Tsurutani, B. T., Denardini, C. M.,
361 Zamlutti, C. J.: Electrodynamic coupling process between the magnetosphere and the



- 362 equatorial ionosphere during a 5-day HILDCAA event. *J. Atmos. Sol. Terr. Phys.* 73,
363 148-155, <https://doi.org/10.1016/j.jastp.2010.09.002>, 2011.
- 364 Krishna, S. G., GPS-TEC Analysis Software Version 2.9.5, Sep. 2017, [online]
365 Available: <http://seemala.blogspot.in/>.
- 366 Kutiev, I., Watanabe, S., Otsuka, Y., Saito, A.: Total electron content behavior over
367 Japan during geomagnetic storms. *J. Geophys. Res.*, 110(A1),
368 <https://doi.org/10.1029/2004JA010586>, 2005.
- 369 Lu, G., Richmond, A. D., Roble, R. G., Emery, B. A.: Coexistence of ionospheric
370 positive and negative storm phases under northern winter conditions: A case study. *J.*
371 *Geophys. Res.*, 106(A11), 24493-24504, <https://doi.org/10.1029/2001JA000003>, 2001
- 372 Maruyama, T., & Nakamura, M.: Conditions for intense ionospheric storms
373 expanding to lower midlatitudes. *J. Geophys. Res.*, 112(A5),
374 <https://doi.org/10.1029/2006JA012226>, 2007.
- 375 Matsushita, S.: A study of the morphology of ionospheric storms. *J. Geophys. Res.* v.
376 64, n. 3, p. 305-321, <https://doi.org/10.1029/JZ064i003p00305>, 1959.
- 377 Mendillo, M.: Storms in the ionosphere: Patterns and processes for total electron
378 content. *Reviews of Geophysics*, v. 44, n. 4, <https://doi.org/10.1029/2005RG000193>,
379 2006.
- 380 Park, C. G.: A morphological study of substorm-associated disturbances in the
381 ionosphere. *Journal of Geophysical Research*, 79(19), 2821-2827,
382 <https://doi.org/10.1029/JA079i019p02821>, 1974.
- 383 Pedatella, N. M., Lei, J., Thayer, J. P., & Forbes, J. M.: Ionosphere response to
384 recurrent geomagnetic activity: Local time dependency. *Journal of Geophysical*
385 *Research: Space Physics*, 115(A2), <https://doi.org/10.1029/2009JA014712>, 2010.
- 386 Pröls, G. W., Najita, K.: Magnetic storm associated changes in the electron content
387 at low latitudes. *Journal of Atmospheric and Terrestrial Physics*, v. 37, n. 4, p. 635-
388 643, [https://doi.org/10.1016/0021-9169\(75\)90058-6](https://doi.org/10.1016/0021-9169(75)90058-6), 1975.



- 389 Silva, R. P., Sobral, J. H. A., Koga, D., and Souza, J. R.: Evidence of prompt
390 penetration electric fields during HILDCAA events, *Ann. Geophys.*, 35, 1165-1176,
391 <https://doi.org/10.5194/angeo-35-1165-2017>, 2017.
- 392 Sobral, J. H. A., Abdu, M. A., Gonzalez, W. D., Clua De Gonzalez, A. L., Tsurutani,
393 B. T., Da Silva, R. R. L., Barbosa, I. G., Arruda, D. C. S., Denardini, C. M., Zamlutti,
394 C. J., Guarnieri, F.: Equatorial ionospheric responses to high-intensity long-duration
395 auroral electrojet activity (HILDCAA), *J. Geophys. Res.* 111, A07S02,
396 <https://doi.org/10.1029/2005JA011393>, 2006.
- 397 Sandanger, M. I., Søråas, F., Aarsnes, K., Oksavik, K., Evans, D. S., & Greer, M. S.:
398 Proton Injections Into the Ring Current Associated With B_z Variations During
399 HILDCAA Events. *GEOPHYSICAL MONOGRAPH-AMERICAN GEOPHYSICAL*
400 *UNION*, 155, 249, <https://doi.org/10.1029/155GM26>, 2005.
- 401 Søråas, F., Aarsnes, K., Oksavik, K., Sandanger, M. I., Evans, D. S., Greer, M. S.
402 (2004). Evidence for particle injection as the cause of Dst reduction during
403 HILDCAA events. *J. Atmos. Sol. Terr. Phys.* 66(2), 177-186,
404 <https://doi.org/10.1016/j.jastp.2003.05.001>, 2004.
- 405 Titheridge, J. E., Buonsanto, M. J.: A comparison of northern and southern
406 hemisphere TEC storm behaviour. *J. Atmos. Sol. Terr. Phys.* 50(9), 763-780,
407 [https://doi.org/10.1016/0021-9169\(88\)90100-6](https://doi.org/10.1016/0021-9169(88)90100-6), 1988.
- 408 Tsurutani, B. T., Gonzalez, W.D.: The cause of high intensity long-duration
409 continuous AE activity (HILDCAA): interplanetary Alfvén wave trains. *Planet Space*
410 *Sci.* 35, 405-412, [https://doi.org/10.1016/0032-0633\(87\)90097-3](https://doi.org/10.1016/0032-0633(87)90097-3), 1987.
- 411 Tsurutani, B. T., Gonzalez, W.D., Guarnieri, F., Kamide, Y., Zhou, X., Arballo, J. K.:
412 Are high-intensity long-duration continuous AE activity (HILDCAA) events
413 substorm expansion events?, *J. Atmos. Sol. Terr. Phys.* 66,167-176,
414 <https://doi.org/10.1016/j.jastp.2003.08.015>, 2004.
- 415 Tsurutani, B. T., Gonzalez, W.D., Gonzalez, A. L. C., Guarnieri, F. L., Gopalswamy,
416 N., Grande, M., Kamide, Y., Kasahara, Y., Lu, G., Mann, I., McPherron, R., Soraas,
417 F., Vasyliunas, V.: Corotating solar wind streams and recurrent geomagnetic activity:



418 A review. J. Geophys. Res. 111, A07S01, <https://doi.org/10.1029/2005JA011273>,
419 2006.

420 Verkhoglyadova, O. P., Tsurutani, B. T., Mannucci, A. J., Mlynczak, M. G., Hunt, L.
421 A., and Runge, T.: Variability of ionospheric TEC during solar and geomagnetic
422 minima (2008 and 2009): external high speed stream drivers, Ann. Geophys., 31,
423 263-276, <https://doi.org/10.5194/angeo-31-263-2013>, 2013.

424

425



426 **Figure captions**

427 **FIGURE 1** – Map showing the locations of the GNSS stations used in the present
428 study. Both stations are localized in the Brazilian region and are marked by a red triangle,
429 where SL and CP are, respectively, São Luís and Cachoeira Paulista.

430 **FIGURE 2** – dTEC hourly values to all HILDCAA intervals to São Luís (equatorial
431 station).

432 **FIGURE 3** – dTEC hourly values to all HILDCAA intervals to Cachoeira Paulista
433 (low latitude station).

434 **FIGURE 4** – Seasonal dTEC response to HILDCAA intervals. The blue and coral
435 lines refer to São Luís and Cachoeira Paulista, respectively.

436 **FIGURE 5** – Solar wind velocities analysis during HILDCAA intervals. The blue
437 and coral colors refer to São Luís and Cachoeira Paulista stations, respectively, while
438 the bubble diameter is related to velocity (km/s).

439

440



441 **Table captions**

442 **TABLE 1** – The date range for HILDCAA intervals identified during 2015 – 2016

443 years

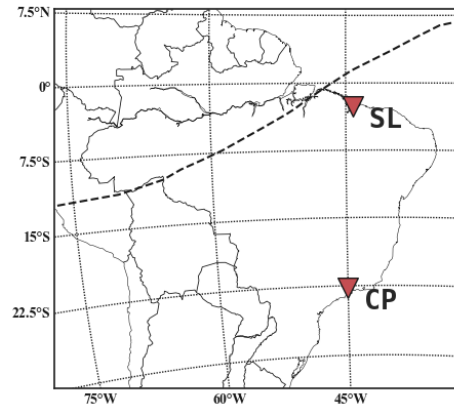
444 **TABLE 2** – Seasonal classification of HILDCAA intervals (according to the seasons

445 in the Southern hemisphere).

446



447 **FIGURE 1 –**



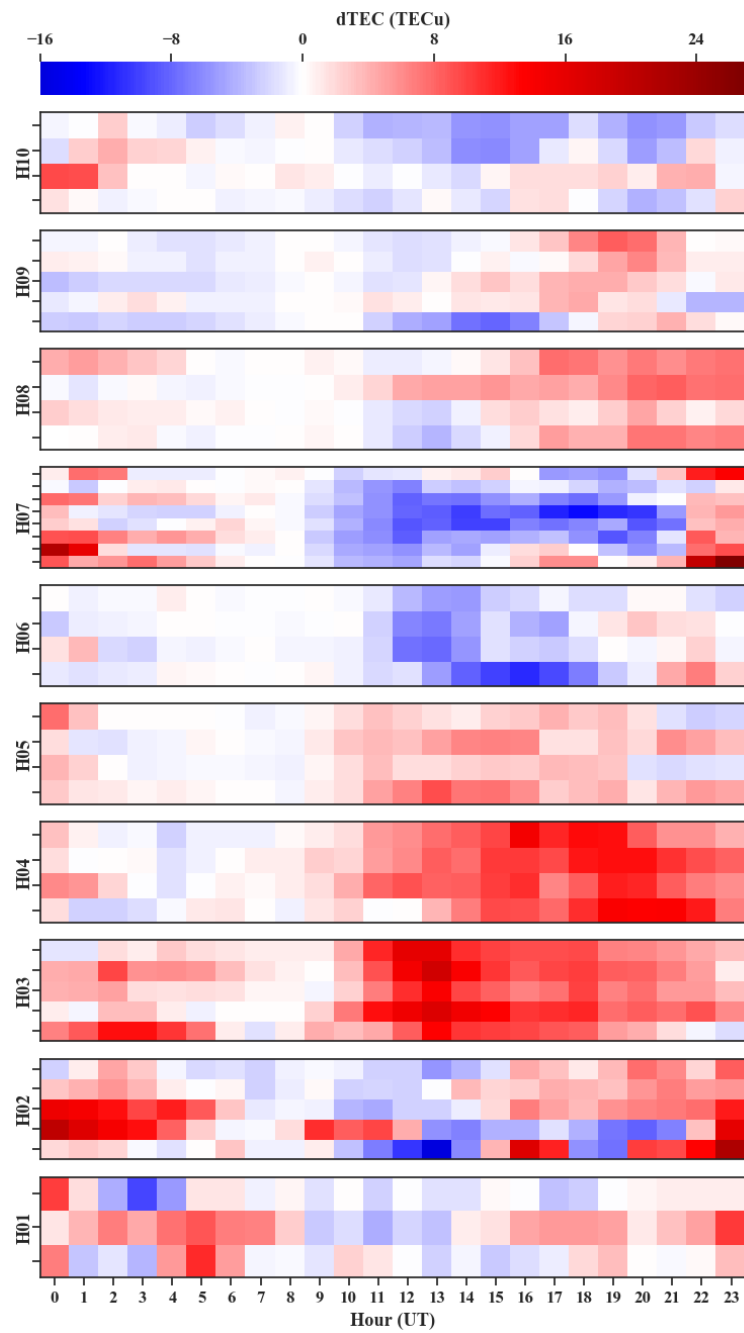
448

449

450



451 **FIGURE 2 -**



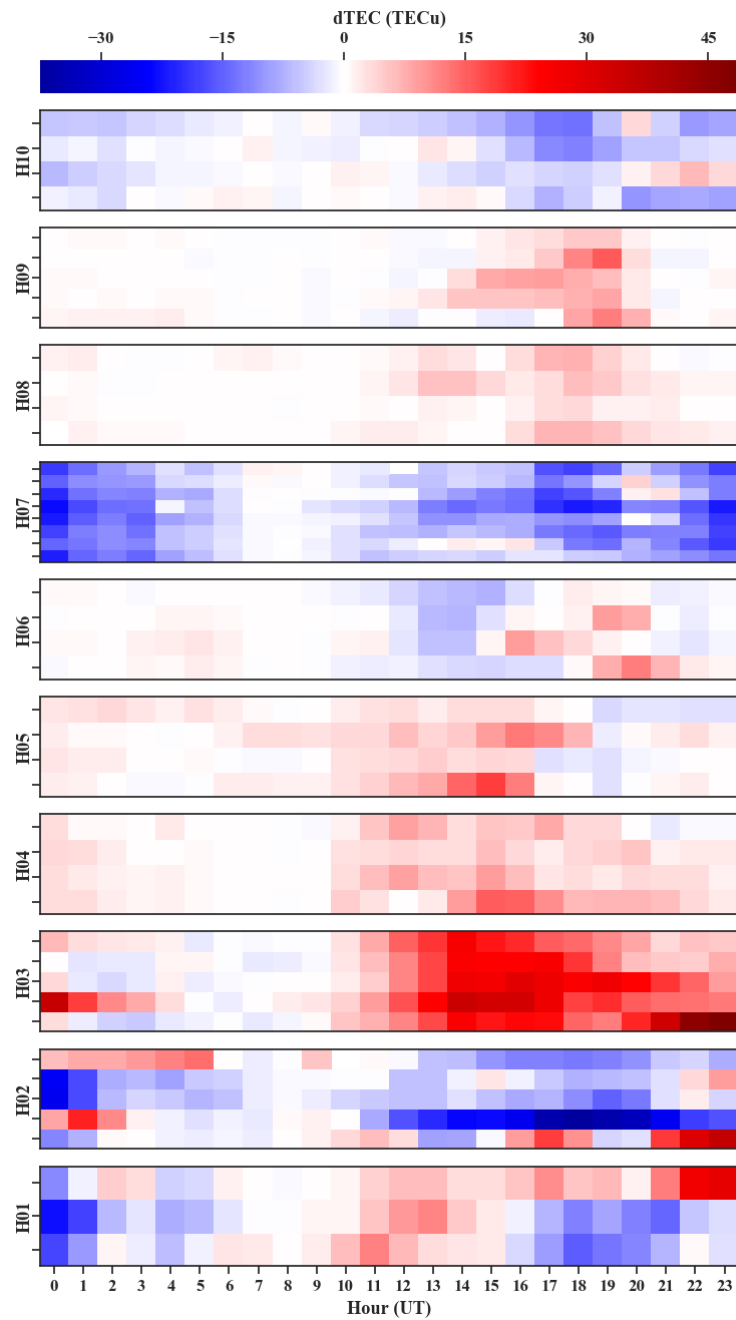
452

453

454



455 **FIGURE 3** –



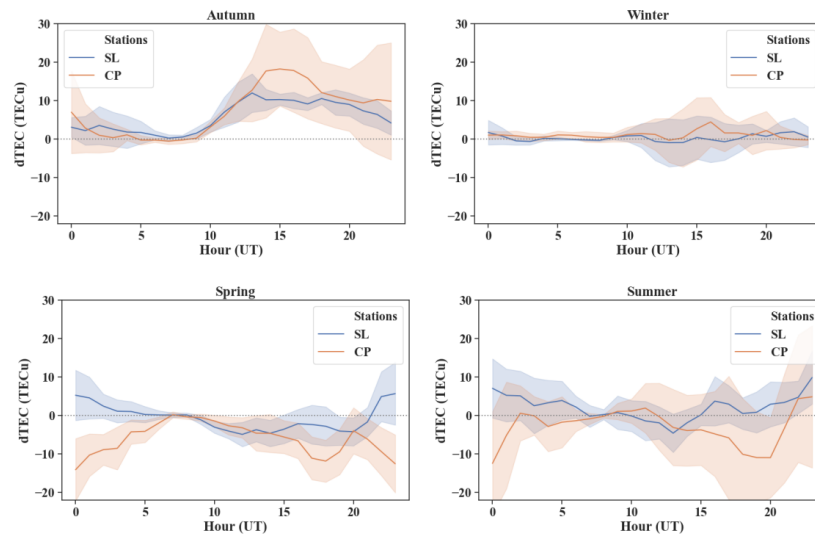
456

457

458



459 **FIGURE 4 –**

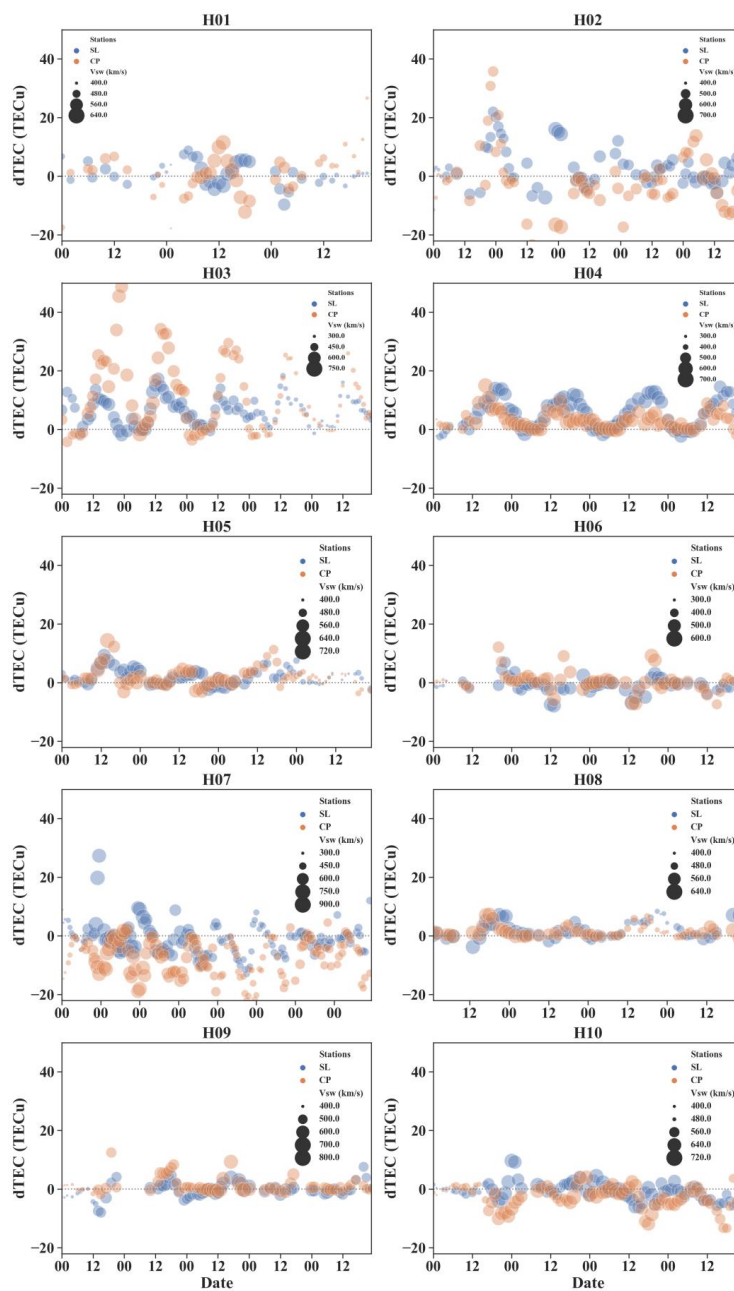


460

461



462 **FIGURE 5** –



463

464



465 **TABLE 1 –**

ID	Date range
H01	2015/03/01 – 03
H02	2015/03/17 – 21
H03	2015/04/16 – 20
H04	2015/06/08 – 11
H05	2015/07/11 – 14
H06	2015/08/15 – 18
H07	2015/10/07 – 14
H08	2016/07/09 – 12
H09	2016/08/03 – 07
H10	2016/12/08 – 11

466

467

468



469 **TABLE 2 –**

Season	HILDCAA Intervals
Autumn	H03 and H04
Winter	H05 and H06
Spring	H07 and H10
Summer	H01 and H02

470