



1 A new scenario applying traffic flow analogy to poleward expansion of auroras

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9 Abstract

10 An auroral ionosphere is generally incompressive and non-uniform medium with anisotropic
11 conductivities. Compressibility may occur, however, following the onset of field line dipolarization.
12 This behavior can happen when; (1) Westward directing electric fields transmitted from the
13 dipolarization region accumulate both electrons and ions in equatorward latitudes in F region. (2) The
14 mobility difference of electrons and ions in E region produces electrostatic potential in a quasi-neutral
15 condition, positive in higher latitudes and negative in lower latitudes. (3) Density modulation in F
16 region excites ion acoustic wave propagating along the field lines towards the magnetosphere. (4) The
17 ion acoustic wave stops in the ionosphere for about 4 min because of a low phase velocity (~ 1.6 km/s).
18 During this compressive interval, density accumulation in equatorward latitudes expands upstream to
19 form a poleward expansion of auroras analogous to upstream propagation of a shock in traffic flow on
20 crowded roads. Electrostatic potential produced in the E region generates field-aligned currents and
21 closing Pedersen currents to retain electrostatic potential in a quasi-neutral ionosphere. The ion
22 acoustic wave produces upward electric fields along the field lines in accordance with the Boltzmann
23 relation which contributed to the ion upflow at topside ionosphere.

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26 **1. Introduction**

27 “Auroras and solar corona observed at the solar eclipse are optical phenomena unique in space physics.



28 With sufficient knowledge about the underlying physical processes, once auroras have been captured
29 by a highly sensitive imager they provide an unexpected wealth of information about plasma
30 environment of the Earth” (copied from [Oguti, 2010]). As one of several unanswered questions
31 involving poleward expansion, plasma drifts in the ionosphere observed by the balloon-measured
32 electric fields [Kelley et al., 1971], by the Ba releases [Haerendel, 1972] and by radar observations
33 [Nielsen and Greenwald, 1978] were opposite to the expanding direction of auroras. To account for
34 the difference in propagation directions, it was suggested that auroras were directly connected to the
35 reconnecting flux tube moving tailward in plasma sheet. The mismatch in the time history of the aurora
36 has been a continuing object of debate [Lui and Rostoker, 2000].

37 In this report, we present a new scenario to explain differing directions of plasma drifts in the
38 poleward expansion of aurora by assuming compressibility of the ionosphere. It has been suggested
39 that incompressibility and current continuity are inherent electromagnetic properties of the ionosphere.
40 In the incompressive ionosphere, current continuity generates polarization charge and associated field-
41 aligned currents (FACs) along the boundary of the conductivity discontinuity and locally modify the
42 electric fields to enhance the current intensity along a high conductivity strip, referred to as Cowling
43 Channel [Baumjohann, 1983]. The ionospheric currents close not only horizontally in the ionosphere
44 but also in the magnetosphere via FACs through which the magnetosphere and ionosphere is coupled
45 (MI-coupling). The electromagnetic processes associated with auroras have been discussed in this MI-
46 coupling scenario. We will note that incompressibility can be violated for a short interval following
47 the dipolarization onset, and show that the compressibility led to the poleward expansion of auroras
48 and produced parallel electric fields at the topside ionosphere by the excitation of ion acoustic wave.
49 The compressibility also produced electrostatic potential and field-aligned currents.

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52 **2. Transmission of westward electric fields to the ionosphere**

53 It is well known that westward electric fields are produced in the midnight magnetosphere in
54 association with the convection surge [Quinn and Southwood, 1982]. The onset of the convection



55 surge is preceded by the pre-onset interval characterized by the inflow of plasma sheet plasmas
 56 towards the equatorial plane [Saka and Hayashi, 2017]. The inflow is accompanied by the high-m
 57 waves in all-sky image propagating towards the onset latitudes [Nishimura et al., 2014; Saka et al.,
 58 2014]. We assume that strength of the westward electric fields of the surge is of the order of 100 mV/m
 59 in the auroral ionosphere, corresponding to 2 mV/m observed in dipolarization front in plasma sheet
 60 [Runov et al., 2011]. Westward electric fields in the plasma sheet are transmitted along the field lines
 61 to the auroral ionosphere by the guided poloidal mode [Radoski, 1967]. Figure 1 illustrates the
 62 earthward transmission of the westward electric fields along the field lines, where the electric fields
 63 were confined in latitudes between λ_1 and λ_2 , peaked at the center. Drift across the magnetic fields
 64 for the j -th species ($\mathbf{U}_{j\perp}$) can be written in the F region as [Kelley, 1989],

$$65 \quad \mathbf{U}_{j\perp} = \frac{1}{B} \left[\mathbf{E} - \frac{k_B T_j}{q_j} \frac{\nabla n}{n} \right] \times \hat{\mathbf{B}}. \quad (1)$$

66 Here, \mathbf{E} denote electric fields pointing westward and $\hat{\mathbf{B}}$ denotes a unit vector of the magnetic fields
 67 B , downward in auroral ionosphere. Symbols k_B , T_j , q_j , and n are the Boltzmann constant, temperature
 68 of the j -th species, charge of the j -th species, and density of electrons (ions), respectively. The electric
 69 field of the order of 100 mV/m exceeded the diffusion (second term) by three orders of magnitudes in
 70 low temperature ionosphere. The diffusion term can be ignored and $E \times B$ drift predominated in the
 71 F region. We calculated numerically the density perturbations using the mass conservation equation,

$$72 \quad \frac{\partial n}{\partial t} + \frac{\partial}{\partial x} (nU) = 0. \quad (2)$$

73 Here n is plasma density, U denotes $E \times B$ drift in the x direction (southward) defined by $e^{-(x/20)^2}$,
 74 proportional to the electric field profile. We treated this as one-dimensional case as the surge was
 75 assumed to spread wider in longitudes than in latitudes. Density distribution was assumed to be
 76 uniform at $T = T_0$ and developed with time as illustrated in the inset. The results show that the plasmas
 77 were accumulated in equatorward edge of the electric fields and the density hole was produced in the
 78 poleward. Because density accumulation was caused primary by the spatial gradient of the drift



79 velocity, we can approximate the equation (2) as,

$$80 \quad \frac{\Delta n}{\Delta t} + n_0 \frac{\Delta U}{\Delta x} = 0. \quad (3)$$

81 Substituting $\Delta U = 10^3 \text{ m s}^{-1}$ and $\Delta x = 10^4 \text{ m}$, we have $\frac{\Delta n}{\Delta t} = 10^{10} \text{ m}^3 \text{ s}^{-1}$ at the background density

82 $n_0 = 10^{11} \text{ m}^{-3}$. This gives density modulation of the order of $\frac{\Delta n}{n_0} = 100\%$ in ten seconds. The

83 accumulation may continue beyond ten seconds because the lifetime of the dipolarization front seems
 84 to be longer [Runov et al., 2011].

85 In the bottom side of the ionosphere (E region), differences in drift velocity of electrons and
 86 ions produces electrostatic potential. Drift trajectories in the E region may be written [Kelley, 1989]
 87 for electrons by,

$$88 \quad \mathbf{U}_{e\perp} = \frac{1}{B} [\mathbf{E} \times \hat{\mathbf{B}}] \quad (4)$$

89 and for ions by,

$$90 \quad \mathbf{U}_{i\perp} = b_i [\mathbf{E} + \kappa_i \mathbf{E} \times \hat{\mathbf{B}}]. \quad (5)$$

91 Here, b_i is mobility of ions defined as $\Omega_i / (B \nu_{in})$, κ_i is defined as Ω_i / ν_{in} . Symbols Ω_i and
 92 ν_{in} are ion gyrofrequency and ion-neutral collision frequency, respectively. $\hat{\mathbf{B}}$ denotes a unit vector
 93 of the magnetic fields B . To derive equations (4) and (5), pressure gradient term (diffusion) was
 94 again ignored. In the E region ($\kappa_i = 0.1$), although the first term of (5) exceeds the second term by
 95 one order of magnitudes, electron accumulation in equatorward latitudes by the imposed westward
 96 electric fields are produced by the mobility difference of electrons in (4) and the second term of the
 97 ions in (5). However, southward electric fields associated with the electron accumulation in lower
 98 latitudes increased immediately and simultaneously ion drifts in the first term of (5) increase. If the
 99 southward electric fields grew to exceed the westward electric fields by an order of magnitudes, ion
 100 drifts in the first term of (5) and electron drifts in (4) balanced to satisfy the quasi-neutrality. This is
 101 equivalent to the generation of the Pedersen currents in the ionosphere. The Pedersen currents would



102 have closed to the field-aligned current (FAC) and sustain the steady state electrostatic potential
103 produced by the compression. The electrostatic potential thus produced in auroral ionosphere would
104 contribute to the atmospheric electricity in Antarctica [Kondo, 1970; Minamoto and Kadokura, 2011].

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107 3. Poleward expansion of density accumulation

108 If the compressed ionospheric plasmas stop in the ionosphere, then question arises regarding the
109 maximum accumulation in the ionosphere. One possible mechanism to suppress the accumulation may
110 be associated with the ionospheric screening that decreased the amplitudes of penetrated westward
111 electric fields. Total electric fields E , a sum of the incident and reflected westward electric fields, may
112 be written as $E = 2\Sigma_A / (\Sigma_A + \Sigma_p)$, where Σ_A and Σ_p are Alfvén conductance defined by
113 $1/\mu_0 V_A$ and height integrated Pedersen conductance in the ionosphere, respectively [Kan et al., 1982].

114 Symbols μ_0 and V_A denote magnetic permeability in vacuum and Alfvén velocity, respectively.

115 Substituting a typical conductance $\Sigma_A = 10^0 S$, and $\Sigma_p = 10^1 S$, total electric fields decreased at the
116 density peak to 18% of the incident electric fields. Another clue may be found in the polarization
117 electric fields produced by the accumulation itself. The electric fields associated with the accumulation
118 (secondary electric fields), directing southward, grew quickly and increased the amplitudes over the
119 primary westward electric fields. Assuming the convection surge is localized in longitudes, the
120 secondary electric fields generate polarization charge, negative to the east and positive to the west, by
121 the electrons transported towards east. The induced polarization charge decreased westward electric
122 fields and slowed down the flow velocity at the density peak. In addition to the above scenarios, we
123 suggest that excess accumulation of the ionospheric plasmas may be suppressed through the term,
124 $(\mathbf{U} \cdot \nabla)\mathbf{U}$, in the equation of motion. This effect, however, may not be a significant factor in
125 decelerating the equatorward flows, because the electric fields in the ionosphere much exceeded the
126 pressure gradient force. All these possible mechanisms would slow the southward flow at the density
127 peak located to the south of the maximum flow velocity. Latitudinal profile of the resulting southward



128 drifts is illustrated by the dashed line in the upper panel of Figure 2. The slowdown of the drift velocity
129 suppressed the further accumulation while equatorward flow continues. This leads to nonlinear
130 evolution of the density profiles analogous to upstream propagation of a shock in traffic flow [Lighthill
131 and Whitham, 1955].

132 We will study how these plasmas, which pile up at equatorward latitudes, develop over time.
133 Assuming drift velocity U is a function of the density n , the conservation equation (2) may be written
134 as [Farlow, 1982],

$$135 \quad \frac{\partial n}{\partial t} + g(n) \frac{\partial n}{\partial x} = 0. \quad (6)$$

136 Here, $g(n) = \frac{dq}{dn}$ and q is a flux defined by $q = nU(n)$. The flux (q)-density (n) curve as illustrated
137 in Figure 3(A) by $q = An(n_0 - n)$ matches drift in the ionosphere where the drift slowed down at
138 the density peak. Here, A is a constant value related to the average drift velocity. Drift velocity is
139 smooth in the range $n < n_c$ whereas the flow decelerates in the range $n > n_c$ and completely stops
140 at $n = n_0$. The conservation equation (6) and flux-density curve in Figure 3(A) show us nonlinear
141 evolutions of the density profile as is illustrated in Figure 3(B). An initial density profile at $T=0$ shown
142 by solid lines indicates that flow stops completely at x downstream of x_1 . At $T=T_1$, the density profile
143 was deformed to a shock indicated by dashed curve in red. A shock front developed at $n = n_c$ where
144 propagation velocity of the kinematic waves, $g(n)$, [Whitham, 1999] changed its sign from positive
145 to negative (Figure 3(A)). The shock front at $x = x_c$ propagated upstream at velocity $U = -A \cdot n_{in}$.
146 If the number densities and fluxes carried to onset latitudes are sufficiently high, then the poleward
147 expansion may be explosive; with quick formation of the shock and a large expansion velocity. The
148 electrostatic potential produced in the E region would have also expanded poleward. This electrostatic
149 potential may produce vortical flows in the F region and loop currents in the E region, which may
150 expand as a shock.

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153 4. Parallel electric fields associated with density accumulation

154 The transient compression of the ionospheric plasmas associated with the accumulation by the
155 $E \times B$ drift would excite the ion acoustic wave in the ionosphere travelling along the field lines
156 upward and downward directions from the density peak of the F region. Figure 4(A) shows height
157 distribution of the pre-onset density profile of electrons (black), density profile caused by the
158 accumulation (dashed-red) and its decay associated with the travelling ion acoustic wave (dashed-
159 green). It is assumed that accumulation doubled the electron density profile at the equatorward edge
160 of convection surge from 90 km to 1000 km in altitudes. Electron density profile was plotted using
161 sunspot maximum condition given in Prince and Bostic (1964). The travelling ion acoustic waves,
162 upward and downward, are denoted by vertical arrows. Ion acoustic wave propagating downward may
163 be eventually absorbed in the neutrals, while the upward wave may propagate along the field lines. By
164 assuming that electron motions are determined by the Boltzmann relation in our case, we give the
165 electrostatic potential ϕ of the ion acoustic wave by the relation,

$$166 \quad \delta n_e = -n_e \frac{q\phi}{k_B T_e}, \quad (7)$$

167 or by [Chen, 1974],

$$168 \quad E_{\parallel} = -\frac{k_B T_e}{q} \frac{\nabla_{\parallel} n_e}{n_e}. \quad (8)$$

169 Here, k_B is Boltzmann constant, q is electron charge, T_e is electron temperature, n_e is electron
170 density ($n_e = n_i$), and δn_e is perturbed electron density. We will focus only on the upward travelling
171 ion acoustic wave. Equation (8) gives electric field strengths of the order of $0.4 \mu V / m$ and
172 $2.0 \mu V / m$ for $T_e = 1000 K$ and $T_e = 5000 K$, respectively, when the e-folding distance of
173 density dropout along the field lines was 200 km. For ions, however, field-aligned flows in collisional
174 plasmas with gravity force may be given as [Kelley, 1989],



175
$$V_{i//} = b_i E_{i//} - D_i \frac{\nabla_{i//} n}{n} - \frac{g}{v_{in}}. \quad (9)$$

176 Here, b_i and D_i denote mobility and diffusion coefficient of ions defined by $\frac{q_i}{M_i v_{in}}$ and

177 $\frac{k_B T_i}{M_i v_{in}}$, respectively. Symbols, M_i , q_i , v_{in} and g are ion mass, electric charge of ions, ion-

178 neutral collision frequency and gravity, respectively. To evaluate equation (9), we assumed $1000K$
179 for ion temperatures and the same e-folding distance used in equation (8). Ions are oxygen and ion-
180 neutral collision frequencies for the nighttime sunspot maximum condition given in Prince and Bostic
181 (1964) were used. Parallel electric fields were estimated by the equation (8) for different electron
182 temperatures. Snapshot of the velocity profile in altitudes from 400 km to 800 km are shown in Figure
183 4(B) for the two cases of electron temperatures, 5000K for black dots and 1000K for red dots. For
184 the low temperature case (1000k), there occurred no ion upflow because the parallel electric fields
185 could not overcome the gravity. We suggest that electron temperatures over 2700K would be needed
186 to excite ion upflow or for excitation of ion acoustic wave propagating upward. Ion velocity increased
187 rapidly above 600 km altitudes and reached 1369 m/s at 800 km when electron temperature increased
188 to 5000K. The velocity of the ion upflow became comparable to the phase velocity of ion acoustic
189 wave (1.6 km/s) at the altitudes of 800 km. This may suggest that an initial perturbation of the ion
190 acoustic wave would have been localized in altitudes below 800 km in the topside ionosphere and
191 stopped there for 4 min because of its slow propagation velocity (1.6 km/s). We conclude that the ion
192 upflow in topside ionosphere was caused primary by the parallel electric fields excited by the upward
193 travelling ion acoustic wave.

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196 7. Summary

197 Allowing for the compressibility of the ionosphere for a short interval following the
198 dipolarization onset, we proposed a new scenario for the poleward expansion of auroras including



199 generation of field-aligned currents and parallel electric fields of the ionospheric origin. This scenario
200 partly explains the discrepant time history of the aurora: plasmas drift in the ionosphere opposite to
201 the expanding direction of aurora. We emphasize that the ionosphere is not a mere earthward boundary
202 of the magnetosphere but a source region directly producing magnetospheric processes.

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205 **Acknowledgements**

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207 Dynamics Campaign (GADC) [Oguti et al., 1988]. We also acknowledge STEP Polar Network
208 (<http://step-p.dyndns.org/~khay/>).

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260 **Figure Captions**

261 Fig. 1: Transmission of the westward electric fields along the field lines is illustrated in the flux tube
262 localized between latitudes λ_1 and λ_2 . The inset at the lower right corner illustrates the latitudinal
263 profile of the southward drift velocity (U) of the ionospheric plasmas produced by these electric fields
264 and density profiles (n) developed in time from T_0 to T_7 (see text).

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266 Fig.2: Same as inset in Figure 1 but a slowdown of the equatorward drift was considered (see text).
267 The equatorward drift velocity slowed down (dashed curve in upper panel) at the peak of plasma
268 accumulation in lower latitudes (lower panel). The solid curve denotes velocity profile with no
269 slowdown.

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271 Fig. 3: (A) Flux(q)-density(n) curve in ionosphere associated with equatorward drift. Flux was
272 maximum at n_c and diminished at the density peak, n_0 . $\frac{\partial q}{\partial n}$ denotes propagation of kinematic waves,
273 positive towards the concentration section and negative away from the concentration. (B) Black curve
274 denoted by $T=0$ shows the initial density profile. Flow is stopped at x downstream of x_1 . Plasmas with
275 density (n_{in}) move into the slowdown section. Dotted red curve shows shock formation at $T=T_1$.

276 Plasmas in the section $n_{in} - n_c$ are accelerating ($\frac{\partial q}{\partial n} > 0$), while plasmas in $n_c - n_0$ are decelerating

277 ($\frac{\partial q}{\partial n} < 0$) to form the shock front. The shock front propagates upstream at a velocity $-A n_{in}$.



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279 Fig 4: (A) Vertical profile of electron number density in pre-onset (black) and vertical profile
280 associated with accumulation (dashed-red) from 90 km to 1000 km in altitudes. The e-folding distance
281 of density decrease is 200 km in altitudes above 400 km. Dashed-green curve denotes a decay of the
282 density accumulation caused by travelling ion acoustic wave along the field lines upward and
283 downward as marked by vertical arrows. (B) Velocity profile in altitudes for ions (oxygen) produced
284 by parallel electric fields $0.4\mu V / m$ ($T_e=1000K$) in red dots and $2.0\mu V / m$ ($T_e=5000K$) in black
285 dots. Vertical flows in altitudes from 400 km to 800 km are shown. Flow velocity is positive upward
286 and negative downward.

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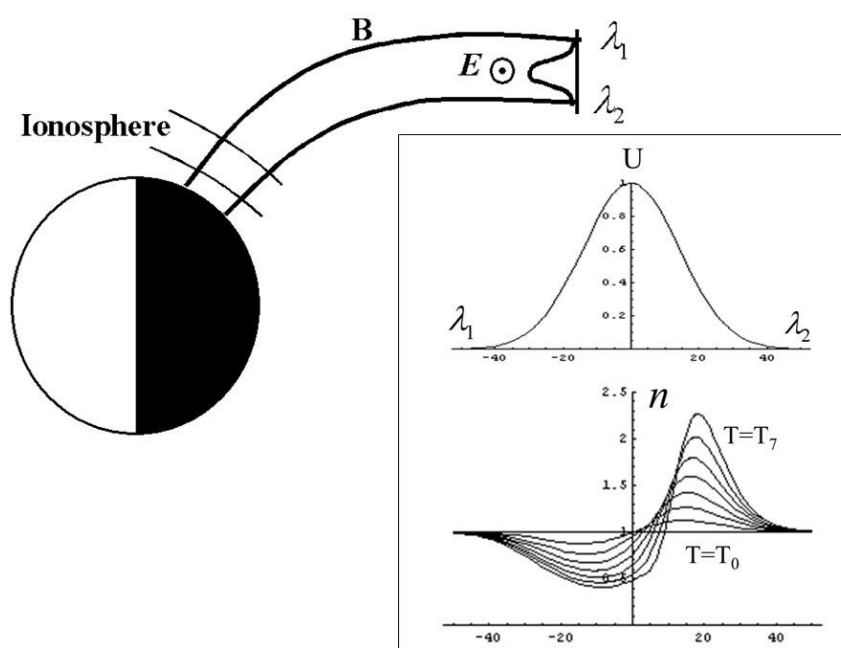
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Figure 1



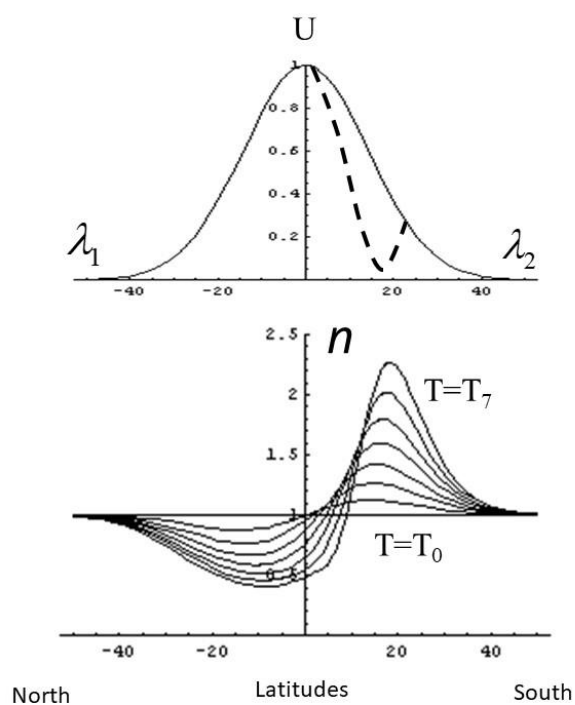
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Figure 2



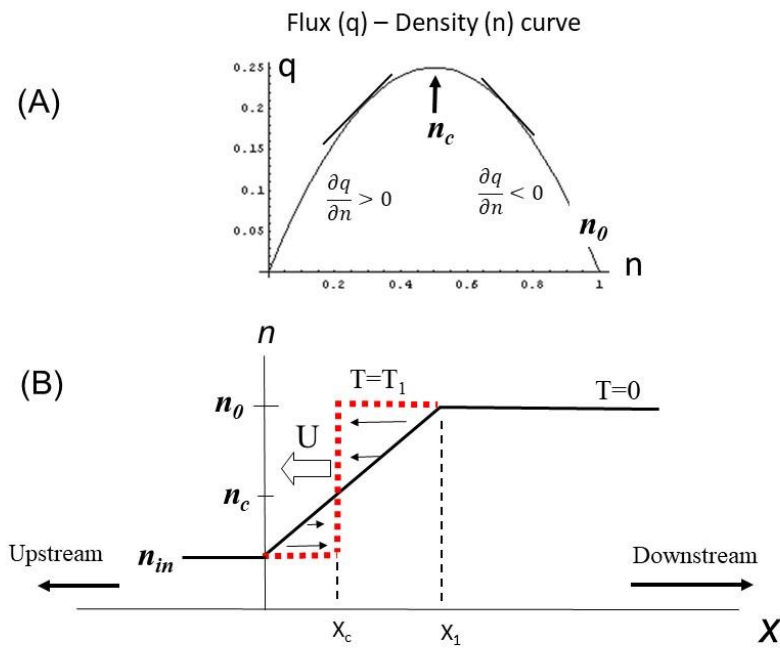
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Figure 3

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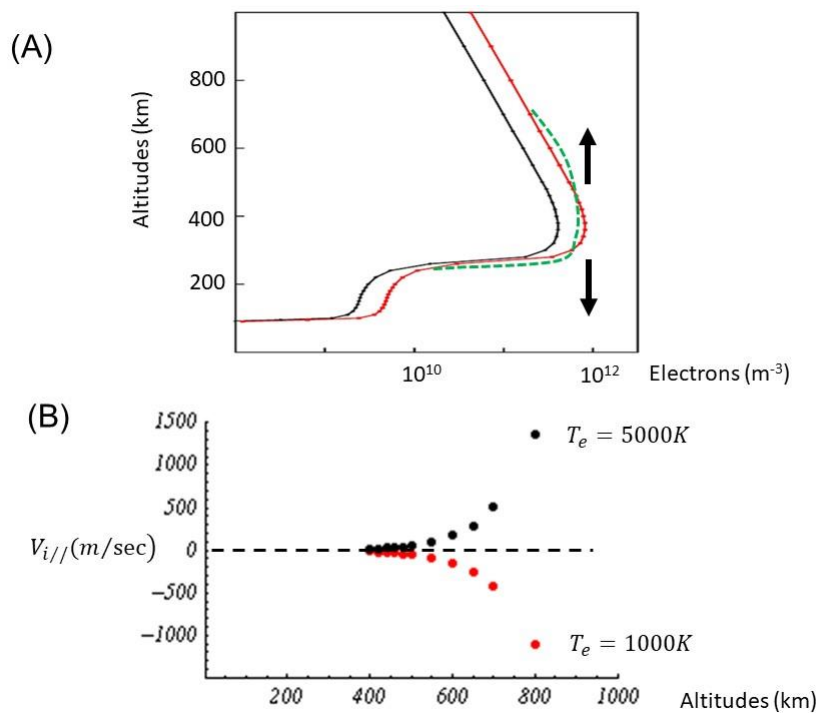
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Figure 4