

# Morphology of GPS and DPS-TEC over an equatorial station: validation of IRI and NeQuick 2 models.

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## 0.0 Abstract

We investigated total electron content (TEC) at Ilorin (8.50°N 4.65E, dip lat. 2.95) for the year 2010, a year of low solar activity in 2010 with  $R_z=15.8$ . The investigation involved the use of TEC derived from GPS, estimated TEC from digisonde portable sounder data (DPS), the International Reference Ionosphere (IRI) and NeQuick 2 (NeQ) models. During the sunrise period, we found that the rate of increase in DPS-TEC, IRI-TEC, and NeQ-TEC was higher with compared with GPS-TEC. One reason for this can be alluded to an overestimation of plasmaspheric electron content (PEC) contribution in modeled TEC and DPS-TEC. A correction factor around the sunrise, where our finding showed a significant percentage deviation between the modeled TEC and GPS-TEC, will correct the differences. Our finding revealed that during the daytime when PEC contribution is known to be absent or insignificant, GPS-TEC and DPS-TEC in April, September, and December predict TEC very well. The lowest discrepancies were observed in May, June, and July (June solstice) between the observed and all the model values at all hours. There is an overestimation in DPS-TEC that could be due to extrapolation error while integrating from the peak electron density of F2 (NmF2) to around ~ 1000 km in the Ne profile. The underestimation observed in NeQ-TEC must have come from the inadequate representation of contribution from PEC on the topside of the NeQ model profile, whereas the exaggeration of PEC contribution in IRI-TEC amount to overestimation in GPS-TEC. The excess bite-out observed in DPS-TEC, and modeled-TEC shows the indication of over-prediction of fountain effect in these models. Therefore, the daytime bite-out observed in these models requires a modifier that could moderate the perceived fountain effect morphology in the models accordingly. The daytime DPS-TEC performs better than the daytime IRI-TEC and NeQ-TEC in all the months. However, the dusk period requires attention due to highest percentage deviation recorded especially for the models in March, November, and December. Seasonally, we found

33 that all the TECs maximize and minimize during the March equinox and June solstice,  
34 respectively. Therefore, GPS-TEC and modeled TEC reveal the semi-annual variations in TEC.

35 *Keywords: (Total Electron Content (TEC); International Reference Ionosphere (IRI ) and NeQuick 2 Models)*

36

## 37 **1.0 Introduction**

38 Total electron content (TEC) is the total number of free electrons in a columnar of one square  
39 meter along the radio path from the satellite to the receiver on the Earth. TEC exhibits diurnal,  
40 seasonal, solar cycle and geographical variations. Therefore, the physical and dynamical  
41 morphology of the TEC over a given location is of great importance in trans-ionospheric  
42 communications during both quiet and disturbed geomagnetic conditions (Jesus et al., 2016;  
43 Tariku, 2015; Akala et al., 2012; Aravindan and Iyer, 1990 and Olawepo et al., 2015). GPS-TEC  
44 is quantified from the GPS orbiting satellites to the GPS receiver station on the Earth, with an  
45 approximate distance of 20200 km (Liu et al., 2006). Thus, a typical GPS-TEC measurement  
46 incorporates the complete plasmaspheric electron content (PEC). The digisonde portable sounder  
47 (DPS) estimates the bottomside and topside TEC to obtain the total TEC from the electron  
48 density (Ne) profile. The topside DPS-TEC is extrapolated from the peak electron density of the  
49 F2 region (NmF2) to around ~ 1000 km thus, the significant PEC contribution from the higher  
50 altitudes is omitted from DPS-TEC measurement (Belehaki et al., 2004; Zhang et al., 2006 and  
51 Reinisch and Huang, 2001).

52

53 The International Reference Ionosphere (IRI) model depends on worldwide data from various  
54 measurements (Bilitza, 2001; Bilitza, 1986; Bilitza and Rawer, 1998). The IRI model provides  
55 reliable ionospheric densities, composition, temperatures, and composition in the ionospheric  
56 altitude range (Bilitza, 2001; Radicella et al., 1998 and Coisson et al., 2009). The latest version  
57 of the IRI model can be found at all time on the web  
58 (<http://nssdc.gsfc.nasa.gov/space/model/ions/iri.html>) with improvements on earlier versions of  
59 the model. The NeQuick 2 (NeQ) models makes use of the position, time and solar flux or  
60 sunspot number over a given location are variables in the NeQ model code (Coisson et al., 2006;  
61 Andreeva and Lokota, 2013 and Bidaine and Warnant, 2011). The output of the NeQ program  
62 and corresponding TEC are by the electron density along any ray-path and numerical integration  
63 in space and time respectively.

64           The availability of ionospheric parameters for global ionospheric models is deficient  
65 over the African sector compared to the consistent input of the data from the Asian and  
66 American sectors. Therefore, the continuing investigations of the parameters over Africa are  
67 required to improve the global ionospheric model. For example, Bagiya et al. (2009) studied  
68 TEC around equatorial-low latitude region at Rajkot ( $22.29^{\circ}$  N,  $70.74^{\circ}$  E, dip  $14.03^{\circ}$  N ) during  
69 low solar activity, Olwendo et al. (2012) and Karia and Pathak (2011) investigated the TEC data  
70 at Kenyan and Surat (India), respectively. They all noticed a semi-annual variation with  
71 minimum and maximum TEC in June solstice and March equinox, respectively. Using Faraday  
72 rotational technique, Olatunji (1967) investigated TEC variation over the equatorial latitude at  
73 Ibadan. He observed no daytime bite-out and seasonal anomaly over the region. Rastogi et al.  
74 (1975) observed the diurnal variation of TEC using Faraday rotation over the magnetic equator.  
75 They noticed that TEC at the topside was higher than TEC at the bottomside during the  
76 nighttime, however during the daytime; they observed a uniform distribution of the TEC, on the  
77 topside and the bottomside of Ne profile.

78  
79           Regarding the DPS-TEC measurement, Barbas et al. (2010) examined GPS-TEC and DPS-  
80 TEC at Tucuman ( $26.69^{\circ}$ S,  $65.23^{\circ}$ W) during different seasons. They inferred that the DPS-TEC  
81 represented the GPS-TEC with a minimal discrepancy in all seasons. Reinisch et al. (2004)  
82 investigated GPS-TEC and DPS-TEC at mid-latitude and equatorial region. They observed that  
83 the variations of GPS-TEC and DPS-TEC appeared similar, but the daytime values of GPS-TEC  
84 were higher than daytime DPS-TEC. Zhang et al. (2004) studied the variations of DPS-TEC and  
85 GPS-TEC over Hainan and reported that the daytime DPS-TEC and GPS-TEC were close in  
86 values during the daytime, but during the dusk period, they observed a significant discrepancy  
87 between DPS-TEC and GPS-TEC. Belehaki et al. (2004) extracted the plasmaspheric electron  
88 content (PEC) from the GPS-TEC at Athens ( $38^{\circ}$ N,  $23.5^{\circ}$ E) for over a year. They reported a  
89 maximum and minimum contribution of PEC in the morning and evening, respectively. Mosert  
90 et al. (2007), Jodogne et al. (2004) and Mckinnell et al. (2007) concluded that approximated PEC  
91 from the GPS-TEC and DPS-TEC is possible in colocated GPS and DPS station. Adewale et al.  
92 (2012), Okoh et al. (2015), Jee and Scherliess (2005), Kenpankho et al (2013), Sulungu et al.  
93 (2017), and Migoya Orué et al. (2008) validated the IRI-TEC with GPS-TEC at different regions

94 and found high discrepancies between the IRI-TEC and GPS-TEC when compared different  
95 IRI-model options.

96

97 Concerning NeQ model, Cherniak and Zakharenkova (2016) validated NeQ model. They  
98 established underestimation of the topside ionosphere above ~ 500km in the NeQ model, due to  
99 inaccurate representation of topside Ne profile. Rabiou et al. (2014) validated NeQ model using  
100 GPS-TEC over the equatorial region of Africa. They reported that the upper boundary of NeQ  
101 model, up to 20,000 km needed to be adjusted to accommodate the PEC-TEC in NeQ model.  
102 Leong et al. (2013) investigated TEC and NeQ models. They found that the observed and NeQ  
103 TEC were close in values during dusk periods, but the changed TEC revealed higher  
104 discrepancies during the post-sunset. Yu et al. (2012) investigated the monthly average of NeQ-  
105 TEC model over three stations in China (Changchun, Beijing, and Chongqing) during the  
106 quietest period. They revealed that NeQ correctly predicted GPS-TEC. However, the NeQ-TEC  
107 underestimated the GPS-TEC during the dusk period. Rios et al. (2007) investigated the  
108 variations of DPS-TEC and IRI-TEC and found that DPS-TEC was smaller compared to IRI  
109 TEC. McNamara (1985) observed discrepancies between DPS-TEC and IRI-TEC and found that  
110 the IRI underestimated the DPS-TEC during the daytime. Obrou et al. (2008) compared the DPS-  
111 TEC and IRI-TEC at Korhogo during high and low solar activity. They found that the variations  
112 of DPS-TEC and IRI-TEC were close in values during high solar activity (HSA) and low solar  
113 activity (LSA), but the performance of IRI-TEC was better during HSA compared to LSA.

114

115 The current contributions of Africa on the improvement of ionospheric models (IRI and  
116 NeQuick) are not adequate compared with the continuous support received from Asia and South  
117 America. The insufficient instrumentation at the equatorial region of Africa has a considerable  
118 effect on the shortcoming. Therefore, the constant validation of IRI and models with the  
119 observed parameter is necessary for an improved ionospheric model. Furthermore, the  
120 investigation on DPS-TEC has not been reported extensively for comparison purpose over the  
121 equatorial region of Africa. Therefore, this study investigates the linked morphologies between  
122 the variations of GPS-TEC and DPS-TEC, and validations of IRI-TEC, and NeQ-TEC models  
123 with the observed parameters. Our finding will inform the suitability of modeled TEC in place of  
124 GPS-TEC. The result will also determine the appropriate model for the equatorial latitude in

125 Africa. Thus, the deviations in TEC obtained from the combined relationship between GPS-TEC,  
 126 DPS-TEC, IRI-TEC, and NeQ-TEC could be used to correct the discrepancy in the models.

127

128 **2.0 Methods of Analysis of GPS and DPS Data**

129 Data used for this study are those of the five quietest days of each month of the year  
 130 2010. The five quietest days are days ( with  $A_p \leq 4$ ) for which geomagnetic activities are quiet,  
 131 the are obtained from the international quiet days (IQD) table available on the website of  
 132 Australia Geosciences. The data are for Ilorin (8.50°N 4.65E, dip lat. 2.95) during the year 2010,  
 133 a year of low solar activity. TEC data were obtained with GPS receiver and Digisonde Portable  
 134 Sounder (DPS) both of which are located at the Ionospheric Laboratory of the University of  
 135 Ilorin. The methods of data processing are described in the sections below.

136

137 **2.1 GPS-TEC**

138 The slant TEC records from GPS have errors due to satellite differential delay (satellite  
 139 bias (bs)) and receiver differential delay (receiver bias (br)) and receiver inter-channel bias (b<sub>SR</sub>).  
 140 This uncorrected slant GPS-TEC measured at every one-minute interval from the GPS receiver  
 141 derived from all the visible satellites at the Ilorin station are converted to vertical GPS-TEC  
 142 using the relation below in equation (1).

143  $(\text{GPS} - \text{TEC})_V = (\text{GPS} - \text{TEC})_S - [b_S + b_R + b_{SR}]/S(E)$  1

144 Where  $(\text{GPS} - \text{TEC})_S$  is the uncorrected slant GPS-TEC measured by the receiver,  $S(E)$  is the  
 145 obliquity factor with zenith angle (z) at the Ionospheric Pierce Point (IPP), E is the elevation  
 146 angle of the satellites in degrees and  $(\text{GPS} - \text{TEC})_V$  is the vertical GPS-TEC at the IPP. The  
 147 equation two below provides  $S(E)$  as

148

149  $S(E) = \frac{1}{\cos(z)} = \left[ 1 - \left( \frac{R_E \times \cos(E)}{R_E + h_s} \right)^2 \right]^{-1/2}$  2

150 Where  $R_E$  is the mean radius of the Earth measured in kilometer (km), and  $h_s$  is the height of the  
 151 ionosphere from the surface of the Earth, which is approximately equal to 400 km according to  
 152 Langley et al. (2002) Rama Rao et al., (2006a) and Mannucci et al. (1993) . The five quietest  
 153 slant GPS-TEC data for each month in the year 2010 were interpreted using Krishna software  
 154 (Global positioning system total electron content analysis application user's manual, 2009,  
 155 Institute for Scientific Research, Boston College, Chestnut Hill, Massachusetts). This software  
 156 reads raw data and corrects all source of errors mentioned above from Global Navigation  
 157 Satellite System service (IGS) code file. A minimum elevation angle of 20 degrees is used to

158 avoid multipath errors. The estimated vertical GPS-TEC data is a function of a two sigma ( $2\sigma$ )  
159 iteration. This sigma is a measure of GPS point positioning accuracy. We converted the average  
160 one-minute VTEC data to hourly averages.

161

## 162 **2.2 DPS-TEC**

163 Regarding the total electron content (TEC) from the digisonde portable sounder (DPS),  
164 the Standard Archive Output (SAO) files obtained from the DPS at the University of Ilorin were  
165 edited to remove magnetically disturbed days. Huang and Reinisch (2001) technique was used  
166 to compute the DPS-TEC. The vertical DPS-TEC computation by the technique is based on the  
167 application of the integration over the vertical electron density  $[N_e(h)]$  profile as shown in the  
168 equation (3) below.

$$169 \text{TEC} = \int_0^{h_{mF2}} N_{eB}(dh) + \int_{h_{mF2}}^{1000} N_{eT}(dh) \quad 3$$

170 Where  $N_{eB}$  and  $N_{eT}$  are the bottomside and topside  $N_e$  profiles, respectively. We computed the  
171  $N_{eB}$  from the recorded ionograms by using the inversion technique developed by Huang and  
172 Reinisch (1996). The information above the peak of the F2 layer is absent from the record of the  
173 ionogram. Thus, the  $N_{eT}$  is measured by approximating the exponential functions with suitable  
174 scale height (Bent et al., 1972) with a less estimated error of 5%. The ionograms were manually  
175 scaled and inverted into electron density profile using the NHPC software and later processed  
176 with the SAO explorer software based on the technique described above to obtain the TEC  
177 (Reinisch et al., 2005). We estimated an average of TEC for each hour over the selected days.  
178 The universal time (UT) is the time standard for the record of GPS and DPS data, but we  
179 converted UT to local time (LT) by adding one hour to corresponding UT. Nigeria is 1 hour in  
180 advance of Greenwich Mean Time (GMT) thus, 0100 UT is the same as 0200 LT in Ilorin,  
181 Nigeria. The available months of the year were grouped into seasons in order to study the  
182 seasonal variation of TEC and the performances of some of the options in the IRI model..The  
183 four seasons are grouped as March equinox or MEQU (March, and April), June solstice or  
184 JSOL (June, and July), September equinox or SEQU (September, and October) and December  
185 solstice or DSOL (November, December). The monthly median of the five quietest days were  
186 deduced and the average of the monthly median under a particular season as defined above to  
187 infer seasonal variations under GPS-TEC, DPS-TEC, IRI-TEC, and  $N_eQ$ -TEC. The DPS in  
188 Ilorin was installed in March, 2010, as a result data were not available for the months of January

189 to late March, 2010. Therefore, this study does not include the days for which DPS data were not  
 190 available.

191

### 192 **2.3 Validation of IRI - 2016 and NeQuick 2 Models**

193 We correlated the observed TEC with modeled TEC in the IRI-2016 model. The website  
 194 [http://www.ccmc.gsfc.nasa.gov/modelweb/models/iri\\_vitmo.php](http://www.ccmc.gsfc.nasa.gov/modelweb/models/iri_vitmo.php) provides the modeled TEC  
 195 values. We selected the upper boundary height 2000 km and the B0 table option for the  
 196 bottomside shape parameter. The equations 3a, 3b, and 3c represent the difference between GPS-  
 197 TEC and DPS-TEC, GPS-TEC and IRI-TEC and GPS-TEC and NeQ-TEC while equations 4a,  
 198 4b, and 4c below show the percentage change between GPS-TEC and DPS-TEC, GPS-TEC and  
 199 IRI-TEC, and GPS-TEC and NeQ-TEC.

200

$$201 \quad \Delta_{\text{DPS-GPS}} = \text{DPS}_{\text{TEC}} - \text{GPS}_{\text{TEC}} \quad 3a$$

$$202 \quad \Delta_{\text{IRI-GPS}} = \text{IRI}_{\text{TEC}} - \text{GPS}_{\text{TEC}} \quad 3b$$

$$203 \quad \Delta_{\text{NeQ-GPS}} = \text{NeQ}_{\text{TEC}} - \text{GPS}_{\text{TEC}} \quad 3c$$

$$204 \quad \%(\Delta_{\text{DPS-GPS}}) = \frac{\text{DPS}_{\text{TEC}} - \text{GPS}_{\text{TEC}}}{\text{DPS}_{\text{TEC}}} \times 100 \quad 4a$$

$$205 \quad \%(\Delta_{\text{IRI-GPS}}) = \frac{\text{IRI}_{\text{TEC}} - \text{GPS}_{\text{TEC}}}{\text{IRI}_{\text{TEC}}} \times 100 \quad 4b$$

$$206 \quad \%(\Delta_{\text{NeQ-GPS}}) = \frac{\text{NeQ}_{\text{TEC}} - \text{GPS}_{\text{TEC}}}{\text{NeQ}_{\text{TEC}}} \times 100 \quad 4c$$

207  $\Delta_{\text{DPS-GPS}}$ ,  $\Delta_{\text{IRI-GPS}}$ , and  $\Delta_{\text{NeQ-GPS}}$  represent the difference between GPS-TEC and DPS-TEC,  
 208 GPS-TEC and IRI-TEC, and GPS-TEC and NeQ-TEC, respectively while  $\%(\Delta_{\text{DPS-GPS}})$ ,  
 209  $\%(\Delta_{\text{IRI-GPS}})$ , and  $\%(\Delta_{\text{NeQ-GPS}})$ , represent the percentage deviation between GPS-TEC and  
 210 DPS-TEC, GPS-TEC and IRI-TEC, and GPS-TEC and NeQ-TEC, respectively.

211

212 The Abdus Salam International Centre for Theoretical Physics (ICTP) - Trieste, Italy in  
 213 collaboration with the Institute for Geophysics, Astrophysics and Meteorology (IGAM) of the  
 214 University of Graz, Austria developed the web front-end of NeQuick. This quick-run ionospheric  
 215 electron density model developed at the Aeronomy and Radiopropagation Laboratory modeled  
 216 TEC along any ground-to-satellite straight line ray-path. Therefore, we validated the NeQ  
 217 obtained from <https://t-ict4d.ictp.it/nequick2/nequick-2-web-model>.

218

## 219 **3.0 Result**

### 220 **3.1 Monthly Median Variations of GPS and modeled TEC**

221 Figure 1a shows the simultaneous plots of hourly variations of the monthly median of TEC  
222 obtained from GPS-, DPS-, IRI-, and NeQ- TEC during the quiet period. The GPS-TEC is in  
223 black line with the star symbol; the DPS-TEC is in green line with the diamond symbol, IRI-  
224 TEC is in red line with zero symbols, and finally, the NeQ-TEC is in blue line with  
225 multiplication symbol. All the TEC plots are regulated by the same local time (LT) on the  
226 horizontal axis. The result reveals that the morphologies of GPS-, DPS-, modeled-TEC increase  
227 gradually from the sunrise period (0700 - 0900 LT) and reach the daytime maximum, mostly  
228 around (1200 - 1700 LT), and then later decay steadily until a minimum value around 0600 LT.  
229 Therefore, our result suggests that the diurnal variations of the observed and modeled TEC  
230 capture the well known solar zenith angle dependence of TEC since both observed and modeled  
231 TEC characterize pre-sunrise minimum, daytime maximum, daytime depression (modeled TEC)  
232 and post-sunset decay. The lowest and highest pre-sunrise minimum ranged from  $\sim 0.66$  TECU  
233 (DPS) -  $\sim 4.49$  TECU (DPS) while the lowest and highest daytime maximum found between  $\sim$   
234  $17.75$  TECU (NeQ) -  $\sim 38.0$  TECU (DPS). The result shows noontime bite-out in modeled TEC  
235 around 1200 LT and 1500 LT except in GPS-TEC where the bite-out was obscure except that a  
236 slight shift in daytime maximum within 1500 and 1700 LT in all months. We observed two  
237 moderate peaks (pre-noon and post-noon peaks) in DPS-TEC and modeled TEC indicating the  
238 bite-out effect on the modeled and DPS- TEC signatures. We also found around the sunrise  
239 period, the model TEC rises faster than the GPS-TEC, but IRI-TEC rises faster compared to  
240 DPS-TEC and IRI-TEC. Between 0600 and 0900 LT, the lowest and highest difference in the  
241 rises of IRI-TEC compared to GPS-TEC were  $\sim 5.0$  TECU (March) and  $\sim 15.3$  TECU  
242 (November), respectively. The post noontime decay was faster in DPS-TEC compared to GPS-  
243 TEC and modeled TEC in all months. Figure 1b reveals the coincident seasonal variations of  
244 GPS-; DPS-; and modeled-TEC during a quiet period of (i) March Equinox, (ii) June solstice,  
245 (iii) September equinox and (iv) December solstice. The daytime maximum ranges are between  
246  $\sim 24.8$  TECU (NeQ) -  $\sim 34$  TECU (DPS),  $\sim 19.2$  TECU (NeQ) -  $\sim 22.6$  TECU (DPS),  $\sim 24.9$   
247 TECU (NeQ) -  $\sim 33.5$  TECU (DPS) and  $\sim 24.55$  TECU (NeQ) -  $\sim 31$  TECU (DPS), in March  
248 equinox, June solstice, September equinox, and December solstice, respectively. We observed



249 that the morphologies of GPS-TEC and modeled TEC maximize and minimize at March equinox  
250 and June solstice, thus indicating semi-annual variation in observed and modeled TEC.

251

### 252 **3.2 Percentage deviation of DPS-TEC; IRI-TEC; and NeQ-TEC**

253 Figures 2(a), 3(a), and 4(a), are hourly variations of deviation in TEC ( $\Delta$ TEC) between GPS,  
254 DPS, IRI and NeQ derived TEC whereas Figures 2(b), 3(b), and 4(b) depict the mass plots of  
255 hourly variations in the percentage deviation ( $\% \Delta$ TEC) during a quiet period from March -  
256 December. In Figure 2a and 2b, the overestimation by DPS-TEC as given by  $\Delta$ TEC<sub>DPS-GPS</sub> is  
257 within the range of  $\sim 5.13$  TECU (March) -  $\sim 19.12$  TECU (July) around 0700 - 1600 LT while  
258 the underestimation  $\Delta$ TEC<sub>DPS-GPS</sub> fluctuated between  $\sim 3.2$  TECU (June) -  $\sim 16.4$  TECU  
259 (November) around 1700 - 2400 LT. The overestimation and underestimation of  $\% \Delta$ IRI-GPS  
260 ranged from  $\sim 2\%$  -  $\sim 49\%$  and  $\sim -1.36\%$  -  $\sim -306\%$ , respectively. From Figures 3a and 3b, the  
261 overestimation occurred regularly around 0400 - 1200 LT in all months. The overestimated and  
262 underestimated  $\Delta$ TEC<sub>IRI-GPS</sub> were between  $\sim 9.13$  TECU (July) -  $\sim 15.3$  TECU (November) and  
263  $\sim 0.15$  TECU (October) -  $\sim 0.95$  TECU (July), respectively. However, a few underestimation and  
264 overestimation of  $\Delta$ TEC<sub>IRI-GPS</sub> still occurred irregularly around 1300 - 0300 LT in all months.  
265 The result also shows that IRI-TEC completely overestimated GPS-TEC in May and June within  
266 0100 and 2400 LT. The overestimation of  $\% \Delta$ TEC<sub>IRI-GPS</sub> ranged between  $\sim 0.1\%$  to  $\sim 86\%$  in all  
267 months. In Figures 4a and 4b, NeQ-TEC overestimated GPS-TEC within 0100 - 1100 LT and  
268 2000 - 2400 LT with  $\Delta$ TEC<sub>NeQ-GPS</sub> ranged from  $\sim 9.72$  (September) and  $\sim 0.01$  (April). We also  
269 found that NeQ-TEC underestimated  $\Delta$ TEC<sub>NeQ-GPS</sub> was between  $\sim 9.72$  (Nov) -  $\sim 0.11$  (May). The  
270 overestimation and underestimation of  $\% \Delta$ TEC<sub>NeQ-GPS</sub> are within  $\sim 0.02\%$  -  $\sim 81\%$  and  $\sim -0.3\%$   
271 -  $\sim -75\%$  respectively.

272

### 273 **3.3 Comparisons of the percentage deviations from GPS-TEC**

274 From Figure 2b, 3b, and 4b, the percentage deviation between GPS- and DPS-TEC are  
275 more significant; greater than 100% in March-August, September, November, and December  
276 between 0400 - 0500 LT and around 2200 - 2400 LT in June and July. The percentage deviation  
277 between GPS- and IRI-TEC are also lower than 100% except in March around 0400 LT whereas  
278 the difference between GPS- and IRI-TEC is greater than 100%. The percentage deviations in  
279 DPS and modeled-TEC during dusk periods are always higher than their corresponding

280 deviations during the daytime. During the daytime, the deviations are smaller in DPS and NeQ-  
281 TEC compared to IRI-TEC.

282

#### 283 **4.0 Discussion of Result**

284 An investigation into the variations of GPS-TEC, DPS-TEC, and the validations of modeled-  
285 TECs at an equatorial region (8.50N 4.650 E) in Africa during low solar activity in the year 2010  
286 has been carried out. The TEC increases gradually from the sunrise period, then slowly reaches  
287 the daytime maximum, and later decays to the pre-sunrise minimum. This result indicates that  
288 the observed and modeled-TEC are a solar zenith angle dependence showing peak and least TEC  
289 values during the noontime and dusk time, respectively (Wu et al.2008; Aravindan and Iyer  
290 1990; and Kumar and Singh 2009). Interestingly, our result that reveals the faster rise in the  
291 DPS-TEC compared to GPS-TEC during sunrise is not consistent with the findings of Ezquer et  
292 al. (1992) at Tucumán (26.9° S; 65.4° W), Belehaki et al. (2004) at Athens, McNamara (1985) at  
293 low latitude and Obrou et al. (2008) at Korhogo (9.33°N, 5.43°W, Dip = 0.67°S). They all found  
294 that the GPS-TEC increased faster than the DPS-TEC during the sunrise. The enrichment of  
295 plasmaspheric electron content (PEC) on TEC latterly reported by Belehaki et al. (2004)  
296 indicated a significant PEC increase in the morning and dusk time. Recently, Jodogne et al.  
297 (2004), Mosert et al. (2007), and Mckinnell et al. (2007) also obtained a rough estimation of PEC  
298 from the GPS and DPS-TEC variations. They inferred that the combined GPS-TEC and DPS-  
299 TEC could give the PEC contribution in TEC of a given location. Therefore, the higher rise in  
300 DPS-TEC compared to GPS-TEC during the sunrise in our study could be attributed to  
301 inaccurate representation of PEC in the topside DPS-TEC profile while extrapolation from the  
302 peak of F2 region (NmF2) to around ~ 1000 km of the Ne profile. Therefore, a typical TEC  
303 measurement naturally includes a meaningful PEC contribution (Belehaki et al. 2003; Balan and  
304 Iyer, 1983; Carlson, 1996; and Breed et al., 1997).

305

306 The higher values in DPS-TEC compared with IRI-TEC around sunrise is not consistent  
307 with Rios et al. (2007) who investigated the comparison of DPS-TEC and IRI-TEC. They found  
308 that DPS-TEC is smaller than IRI-TEC at all hours. They assumed that the prediction of IRI-  
309 TEC had included the high topside Ne profile. Thus, our observation may suggest that the IRI-

310 TEC has incorporated low topside Ne profile in the IRI model or the excessive enhancement of  
311 PEC contribution in the topside Ne profile in the DPS-TEC.

312  
313 The closeness observed during daytime between GPS-TEC and DPS-TEC in April,  
314 August and December may also suggest that the topside Ne profile in GPS-TEC is accurate in  
315 the DPS-TEC topside profile due to the absence of or negligible PEC contribution in DPS-TEC  
316 values. The insignificant daytime PEC observed in this study is consistent with Rastogi et al.  
317 (1971) and Belehaki et al. (2004). Higher daytime DPS-TEC compared with daytime IRI-TEC  
318 is consistent with the result of McNamara (1985). However, Obrou et al. (2008) at the equatorial  
319 latitude, found higher IRI-TEC relative to DPS-TEC at the low solar activity. Therefore, the  
320 reduced daytime IRI-TEC compared to GPS-TEC values indicates the excessive PEC removal  
321 from the model values that its PEC contribution had been raised initially during the sunrise.  
322 Also, the reduced NeQ-TEC compared to GPS-TEC values in all months is consistent with the  
323 report of Migoya-Orue et al. (2017), Zakharenkova (2016), Rabiou et al., (2014) and Nava and  
324 Radicella (2009). They recommended an added PEC contribution on topside NeQ profile for an  
325 accurate prediction of NeQ model.

326  
327 The daytime bite-out in TEC is due to the occurrence of the most active fountain effect  
328 during the noontime at the magnetic equator. The bite-out results from the vertical plasma drift  
329 due to the combined consequence of mutually perpendicular electric and magnetic fields on the  
330 plasma. The drift lifts the plasma at the magnetic equator and diffuses along geomagnetic field  
331 lines into the high latitudes, therefore, leaving the reduced TEC at the magnetic equator  
332 (Bandyopadhyay, 1970; Olwendo et al., 2013; Skinner, 1966; Bolaji et al., 2012). However, the  
333 absence of daytime bite-out (Olatunji, 1967) in GPS-TEC in our finding may be due to the more  
334 great productions at the bottomside and topside electron content that are enhanced quickly to  
335 replenish the loss of the ionization that occurs through the fountain effect during the noontime.

336  
337 The percentage difference between observed and modeled-TEC reveal that the pre-  
338 sunrise values in DPS-TEC, IRI-TEC, and NeQ-TEC require modifications especially during the  
339 month of March for DPS-TEC and the models, and November and December for DPS-TEC  
340 only. The daytime DPS-TEC is closer to the GPS-TEC value compared to the daytime IRI-TEC

341 and NeQ-TEC values. The nighttime NeQ-TEC and IRI-TEC perform better with GPS-TEC  
342 compared with DPS-TEC in all months. There is also the need to minimize the discrepancies  
343 observed during the dusk periods.

344 Seasonally, we found that TEC maximizes and minimizes during the equinoxes and the  
345 solstices, respectively. Our report is consistent with Mala et al. (2009), Wu et al. (2008), Kumar  
346 and Singh (2009), and Balan and Rao, (1984) who investigated TEC in various regions. They  
347 attributed the seasonal variation in TEC to the seasonal differences in thermospheric  
348 composition. Moreover, the sub-solar point is around the equator during the equinox.  
349 Consequently, the sun shines directly over the equatorial latitude, and in addition to the high  
350 ratio of O/N<sub>2</sub> around the region, this translates to stronger ionization, and generates a semi-  
351 annual variation in TEC. The finding from our study is consistent with the reports of Ross  
352 Skinner, (1966), Bolaji et al. (2012), and Scherliess and Fejer (1999) who obtained semi-annual  
353 variation in TEC. Scherliess and Fejer (1999) also concluded that daytime  $E \times B$  drift velocity  
354 could result in semi-annual variation because the drift is more and less significant in the  
355 equinoctial months and June solstice, respectively.

356

## 357 **5.0 Conclusion**

358 (i) We have examined the variations of observed and modeled TEC over an equatorial region in  
359 Africa during a year of low solar activity. Our findings showed:

360 (i) that GPS-TEC and modeled TEC are solar zenith angle dependence.

361 (ii) a faster sunrise increase in the modeled TEC relative to GPS-TEC which suggest a  
362 overestimation of the topside Ne profile of the modeled TEC due to plasmaspheric electron  
363 content (PEC) into the models.

364 (iii) a good representation of the daytime measured TEC by the and models, suggesting that the  
365 model TEC could represent GPS-TEC in the absence of plasmaspheric TEC contribution.

366 (iv) the  $\Delta\text{TEC}_{\text{IRI-GPS}}$  and  $\% \Delta\text{TEC}_{\text{IRI-GPS}}$  in May and June consistently show overestimations  
367 within 0100 - 2400 LT indicating the enhanced contribution of PEC at all hours in May and June.

368 (v) the percentage deviations in DPS and modeled-TEC relative to GPS-TEC during dusk periods  
369 is always higher than their corresponding differences during the daytime, and the values of  
370 daytime deviation in DPS and NeQ-TEC are smaller compared to daytime deviation in IRI-TEC.

371 This study was carried out during a low solar activity in the year 2010; it will be of advantage to

372 investigate and compare similar reviews during high solar activity with our results.

373

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380

## 381 **7.0 References**

382

383 Adewale, A.O., Oyeyemi, E.O., and Olwendo, J. (2012): Solar activity dependence of total  
384 electron content derived from GPS observations over Mbarara. *Advances in Space*  
385 *Research*. **50**, 415– 426.

386 Akala, A.O., E O Oyeyemi, E O Somoye, A B Adeloye, and A.O., Adewale. 2010. Variability of  
387 foF2 in the African Equatorial Ionosphere. *Advances in Space Research* 45 (11).  
388 COSPAR: 1311–14. doi:10.1016/j.asr.2010.01.003.

389 Andreeva, E. S., and M. V Lokota. 2013. Analysis of the Parameters of the Upper Atmosphere  
390 and Ionosphere Based on Radio Occultation, Ionosonde Measurements, IRI and NeQuick  
391 Model Data. *52(10):1820–26*. doi.org/10.1016/j.asr.2013.08.012.

392 Aravindan, P., and Iyer, K. N. (1990): Day-to-day variability in ionospheric electron content at  
393 low latitudes, *Pmet. Space Sci.*, **38(6)**,743-750.

394 Bagiya, Mala S, H.P Joshi, K N Iyer, M Aggarwal, S Ravindran, and B M Pathan. 2009. "TEC  
395 Variations during Low Solar Activity Period (2005-2007) near the Equatorial Ionospheric  
396 Anomaly Crest Region in India." *Annales Geophysicae* 27: 1047–57. doi:10.5194/angeo-  
397 27-1047-2009.

398 Balan, N., and Rao, P. B. (1984): Relationship Between Nighttime Total Electron Content  
399 Enhancements, *Journal of Geophysical Research*, **89(10)**, 9009-9013.

400 Balan, N., Iyer, K.N. Equatorial anomaly in ionospheric electron content and its relation to  
401 dynamo currents. *J. Geophys. Res.* 88 (A12), 10259– 10262, 1983.

402 Bandyopadhyay, P. (1970): Measurement of total electron content at Huancayo, Peru, *Planet.*  
403 *Space Sci.*, **18**, 129–135, doi:10.1016/0032-0633 (70)90150-9.

404 Barbas, Haro B. De, C. Medina, and V. H. Rios. 2010. "differences between GPS and digisonde  
405 measurements of total electron content." *45(4):403–16*.

406 Belehaki, A Jankowski, N.Reinisch B. W. 2004. "Plasmaspheric Electron Content Derived from  
407 GPS TEC and Digisonde Ionograms." *33:833–37*.

408 Belehaki, A., and Kersley, L. (2003): Statistical validation of the ITEC parameter, Third  
409 Workshop of the COST271 Action, 23-27 September 2003, Spetses, Greece.

410 Bent, R.B., Llewellyn, S.K. and Schmid, P.E., 1972. A highly successful empirical model for the  
411 worldwide ionospheric electron density profile. DBA Systems, Melbourne, Florida.

412 Bidaine, B. and R. Warnant. 2011. "Ionosphere Modelling for Galileo Single Frequency Users :

413 Illustration of the Combination of the NeQuick Model and GNSS Data Ingestion.”  
414 Advances in Space Research 47(2):312–22. doi.org/10.1016/j.asr.2010.09.001).

415 Bilitza, D. (2001): ‘International Reference Ionosphere 2000’, Radio Science, **36**(2), 261-275.

416 Bilitza, D. (1986): International reference ionosphere: Recent developments, Radio Science, **21**,  
417 343-346.

418 Bilitza, D., and Rawer, K. (1998): “International Reference Ionosphere Model (IRI- 93,”  
419 <http://envnet.gsfc.nasa.gov/Models/EnviroNET-Models.html> Adv. Space Res., **69**, 520–  
420 829.

421 Bolaji, O. S., Adeniyi, J. O., Radicella, S. M., and Doherty, P. H. (2012): Variability of total  
422 electron content over an equatorial West African station during low solar activity. Radio  
423 Science, **47**, RS1001 doi:10.1029/2011RS004812.

424 Breed, A.M, G L Goodwin, A-m Vandenberg, E A Essex, K J W Lynn, and Abstract  
425 Ionospheric. 1997. Ionospheric Total Electron Content and Slab Thickness J . H . Silby 32  
426 (4): 1635–43. [10.1029/97RS00454](https://doi.org/10.1029/97RS00454) 2006.

427 Carlson, H.C. Incoherent scatter radar mapping of polar electrodynamics. J. Atmos. Solar-Terr.  
428 Phys. 58 (1–4), 37–56, 1996.

429 Cherniak, I., and Zakharenkova, I. (2016). NeQuick and IRI-Plas model performance on topside  
430 electron content representation: Spaceborne GPS measurements. *Radio Science*, **51**(6),  
431 752-766.

432 Coisson, P., Radicella, S.M., Leitinger, R. and Nava, B., 2006. Topside electron density in IRI  
433 and NeQuick: features and limitations. Advances in Space Research, 37(5), pp.937-942.

434 Ezquer, R. G., Adler, N.O., Radicella, S.M., Gonzalez, M .M., and Manza, J . R. (1992): Total  
435 electron content obtained from ionogram data alone, Radio Science, **27**(3), 429-434.

436 Fejer, B. G. and L. Scherliess. 2001. “On the Variability of Equatorial F-Region Vertical Plasma  
437 Drifts.” 63:893–97.

438 Huang, X., and Reinisch, B. W. (2001). Vertical electron content from ionograms in real time.  
439 *Radio Science*, **36**(2), 335-342.

440 Huang and Reinisch 1996 Huang, Xueqin and B. W. Reinisch. 1996. “Vertical Electron Density  
441 Profiles from the Digisonde Network.” 18(6).

442 Jee, G., Schunk, R.W. and Scherliess, L., 2005. On the sensitivity of total electron content (TEC)  
443 to upper atmospheric/ionospheric parameters. Journal of atmospheric and solar-terrestrial  
444 physics, 67(11), pp.1040-1052.

445 Jesus, R De, P R Fagundes, A Coster, O S Bolaji, J H A Sobral, I S Batista, AJ De Abreu, et al.  
446 2016. Effects of the Intense Geomagnetic Storm of September – October 2012 on the  
447 Equatorial, Low- and Mid-Latitude F Region in the American and African Sector during  
448 the Unusual 24th Solar Cycle." Journal of Atmospheric and Solar-Terrestrial Physics  
449 138–139. Elsevier: 93–105. doi:10.1016/j.jastp.2015.12.015.

450 Jodogne J.-C., H. Nebdi, and R.Warnan. 2004. Advances in Radio Science GPS TEC and ITEC  
451 from Digisonde Data Compared with NeQuick Model. 269–73.

452 Karia, S.P., and Pathak, K.N.(2011): GPS based Tec measurement for a period Aug 2008-Dec  
453 2009 near the northern crest of India equatorial ionospheric anomaly region. Journal of  
454 earth system science, **120.5**, 851-858.

455 Kenpankho, P., P. Supnithi, and T. Nagatsuma. 2013. ScienceDirect Comparison of Observed  
456 TEC Values with IRI-2007 TEC and IRI-2007 TEC with Optional Fo F2 Measurements  
457 Predictions at an Equatorial Region, Chumphon, Thailand. Advances in Space Research

458 Kumar, S., and Singh, A.K. (2009): Variation of ionospheric total electron content in Indian low  
459 latitude region of the equatorial anomaly during May 2007–April 2008, *Advances*  
460 *in Space Research* **43**, 1555–1562.

461 Langley, R., M. Fedrizzi, E. Paula, M. Santos, and A. Komjathy (2002), Mapping the low latitude  
462 ionosphere with GPS, *GPS World*, 13(2), 41–46.

463 Leong, S. K. et al. 2014. Assessment of Ionosphere Models at Banting: Performance of IRI-  
464 2007, IRI-2012, and NeQuick 2 Models during the Ascending Phase of Solar Cycle 24.  
465 *Advances in Space Research* (2013) <http://dx.doi.org/10.1016/j.asr.2014.01.026>.

466 Liu, J. Y., H. F. Tsai, and T. K. Jung (1996b), Total electron content obtained by using the global  
467 positioning system, *Terr. Atmos. Oceanic Sci.*, 7, 107.

468 Mala, S., Bagiya, H. P., Joshi, K. N., Iyer, M., Aggarwal, S., Ravindran, and Pathan, B.M.  
469 (2009): TEC variations during low solar activity period (2005–2007) near the equatorial  
470 Ionospheric Anomaly Crest region in India, *Ann. Geophys.*, **27**, 1047–1057.

471 Mannucci, A. J., B. D. Wilson, and C. D. Edwards (1993), A new method for monitoring the  
472 Earth’s ionospheric total electron content using the GPS global network, paper  
473 presented at ION GPS-93, Inst. of Navigation., pp. 1323–1332, Salt Lake City, Utah,  
474 22–24 Sept.

475 McKinnell, L. A., Opperman, B., and Cilliers, P. J. (2007). GPS TEC and ionosonde TEC over  
476 Grahamstown, South Africa: First comparisons. *Advances in Space Research*, 39(5), 816-  
477 820.

478 McNamara, L.F (1985). The use of total electron content measurements to validate empirical  
479 models of the ionosphere. *Adv. Space Res.* **5** (7), 81– 90.

480 Migoya Orué, Y. O., S. M. Radicella, P. Coïsson, R. G. Ezquer, and B. Nava. 2008. Comparing  
481 TOPEX TEC Measurements with IRI Predictions.” *Advances in Space Research*  
482 42(4):757–62.

483 Migoya-Orué, Y., Folarin-Olufunmilayo, O., Radicella, S., Alazo-Cuartas, K., and Rabiú, A. B.  
484 (2017). Evaluation of NeQuick as a model to characterize the Equatorial Ionization  
485 Anomaly over Africa using data ingestion. *Advances in Space Research*, 60(8), 1732-  
486 1738.

487 Mosert, M, L A McKinnell, M Gender, C Brunini, J Araujo, R G Ezquer, and M Cabrera. 2007.  
488 "Variations of F O F 2 and GPS Total Electron Content over the Antarctic Sector," 327–  
489 33. doi:10.5047/eps.2011.01.006.

490 Mosert, M., Gende, M., Brunini, C., Ezquer, R., and Altadill, D. (2007). Comparisons of IRI  
491 TEC predictions with GPS and digisonde measurements at the Ebro. *Advances in Space*  
492 *Research*, 39(5), 841-847.

493 Nava, B. and S. M. Radicella. 2009. “On the Use of NeQuick Topside Option in IRI-2007.”  
494 43:1688–93.

495 Obrou, O.K., Mene, M.N., Koba, A.T., and Zaka, K.Z. (2008): Equatorial total electron content  
496 (TEC) at low and high solar activity, *Advances in Space Research* **43**, 1757–1761.

497 Okoh, D., McKinnell, L.A., Cilliers, P., Okere, B., Okonkwo, C. and Rabiú, B., 2015. IRI-vTEC  
498 versus GPS-vTEC for Nigerian SCINDA GPS stations. *Advances in Space Research*,  
499 55(8), 1941-1947.

500 Olatunji, E. O. (1967). The total columnar electron content of the equatorial ionosphere. *Journal*  
501 *of Atmospheric and Terrestrial Physics*, 29(3), 277-285.

502 Olawepo, A.O.; Oladipo, O.A; Adeniyi, J.O; and Doherty, P.H (2015). TEC response at two  
503 equatorial stations in the African sector to geomagnetic storms. *Advances in Space*

506 Olwendo, O. J., P. Baki, P. J. Cilliers, C. Mito, and P. Doherty. 2013. "Comparison of GPS TEC  
507 Variations with IRI-2007 TEC Prediction at Equatorial Latitudes during a Low Solar  
508 Activity (2009-2011) Phase over the Kenyan Region." *Advances in Space Research*  
509 52(10).

510 Olwendo, O. J., P. Baki, P. J. Cilliers, C. Mito, and P. Doherty. 2012. "Comparison of GPS TEC  
511 Measurements with IRI-2007 TEC Prediction over the Kenyan Region during the  
512 Descending Phase of Solar Cycle 23." *Advances in Space Research* 49(5):914–21.  
513 Retrieved (<http://dx.doi.org/10.1016/j.asr.2011.12.007>).

514 Olwendo, O.J., Baki, P., Cilliers, P.J., Mito, C., and Doherty, P. (2012): Comparison of GPS  
515 TEC measurements with IRI-2007 TEC prediction over the Kenyan region during the  
516 descending phase of solar cycle 23, *Advances in Space Research* **49**, 914–921.

517 Rabi, A. B., A. O. Adewale, R. B. Abdulrahim, and E. O. Oyeyemi. 2014. "ScienceDirect TEC  
518 Derived from Some GPS Stations in Nigeria and Comparison with the IRI and NeQuick  
519 Models." *Advances in Space Research* 53(9):1290–1303. Retrieved  
520 (<http://dx.doi.org/10.1016/j.asr.2014.02.009>).

521 Radicella, S.M., Bilitza, D., Reinisch, B.W., and Adeniyi, J.O., Mosert Gonzalez, M.E., Zolesi,  
522 B., Zhang, M.L., Zhang, S. (1998): IRI Task Force Activity At ICTP: Proposed  
523 Improvements For The IRI Region Below The F Peak, *Adv. Space Res.* **22**(6)  
524 731-739.

525 Rama Rao, P.V.S., Gopi Krishna, S., Niranjana, K., Prasad, D.S.V.V.D. Temporal and spatial  
526 variations in TEC using simultaneous measurements from the Indian GPS network of  
527 receivers during the low solar activity period of 2004–2005. *Ann. Geophys.* 24, 3279–  
528 3292, 2006b.

529 Rama Rao, P.V.S., Niranjana, K., Prasad, D.S.V.V.D., Gopi Krishna, S., Uma, G. On the validity  
530 of the ionospheric pierce point (IPP) altitude of 350 km in the equatorial and low latitude  
531 sector. *Ann. Geophys.* 24, 2159–2168, 2006a.

532 Rastogi, R G., and Sharma, R. P. (1971): Ionospheric electron content at Ahmedabad (near the  
533 crest of the equatorial anomaly) by using beacon satellite transmissions during half a  
534 solar cycle; *Planet. Space Sci.* **19** 1505–1517.

535 Rastogi, R.G., Iyer, K.N. and Bhattacharyya, J.C., 1975. Total electron content of the ionosphere  
536 over the magnetic equator. *Current Science*, pp.531-533.

537 Rawer, K. and Bilitza, D., 1990. International Reference Ionosphere—plasma densities: status  
538 1988. *Advances in space research*, 10(8), pp.5-14

539 Rawer, K., Lincoln, J. V., and Conkright, R. O. (1981): International Reference Ionosphere—  
540 IRI 79, World Data Center A for Solar-Terrestrial Physics, Report UAG-82, Boulder,  
541 Colorado. 52. 223-232.

542 Reinisch, B.W., Huang, X. Deducing topside profile and total electron content from bottomside  
543 ionograms. *Adv. Space Res.* 27 (1), 23–30, 2001.

544 Reinisch, B. W. Å., X. Huang, I. A. Galkin, V. Paznukhov, and A. Kozlov. 2005. "Recent  
545 Advances in Real-Time Analysis of Ionograms and Ionospheric Drift Measurements with  
546 Digisondes." 67:1054–62.

547 Reinisch, B. W., X. Huang, A. Belehaki, and R. Ilma. 2004. "Advances in Radio Science Using  
548 Scale Heights Derived from Bottomside Ionograms for Modelling the IRI Topside  
549 Profile." (June 2002):293–97.

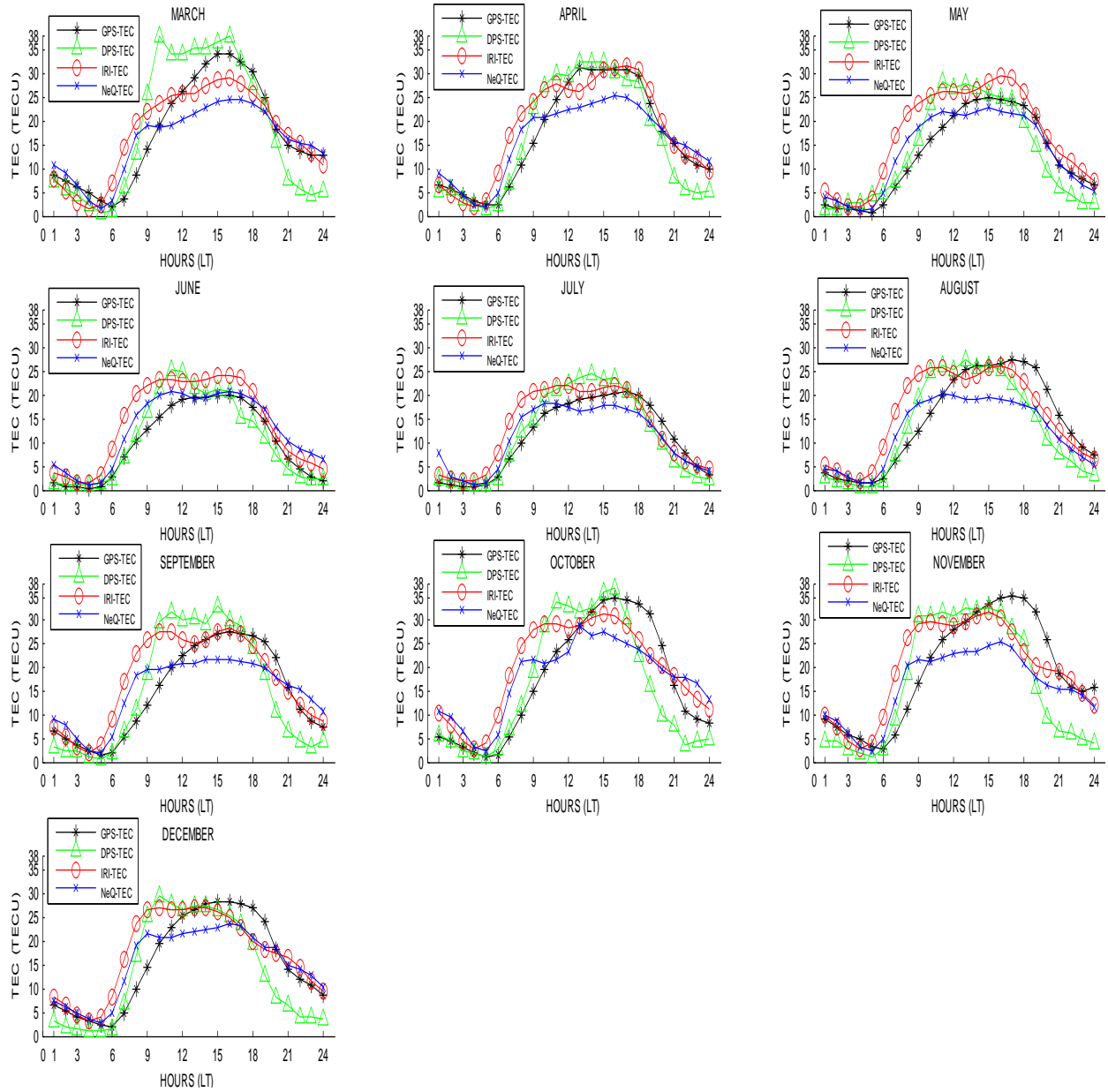


- 550 Rios, V.H., Medina, C.F. and Alvarez, P., 2007. Comparison between IRI predictions and  
551 digisonde measurements at Tucuman. *Journal of Atmospheric and Solar-Terrestrial*  
552 *Physics*, 69(4-5), pp.569-577.
- 553 Skinner, N. J. (1966): Measurements of Total Electron Content Near The Magnetic Equator,  
554 *Planet. Space Sci. Vol. 14*, Pp. 1123 - 1129.
- 555 Scherliess, L., and B. Fejer, Radar and satellite global equatorial F region vertical drift model, *J.*  
556 *Geophys. Res.*, 104,6829-6842,1999.
- 557 Sulungu, Emmanuel Daudi, Christian B. S. Uiso, and Patrick Sibanda. 2017. "Comparison of  
558 GPS Derived TEC with the TEC Predicted by IRI 2012 Model in the Southern Equatorial  
559 Ionization Anomaly Crest within the Eastern Africa Region." *Advances in Space*  
560 *Research*. Retrieved (<http://dx.doi.org/10.1016/j.asr.2017.07.040>).
- 561 Tariku, Yekoye. 2015. "Patterns of GPS-TEC Variation over Low-Latitude Regions (African  
562 Sector) during the Deep Solar Minimum (2008 to 2009) and Solar Maximum (2012 to  
563 2013) Phases." *Earth, Planets, and Space* 67 (1). doi:10.1186/s40623-015-0206-2.
- 564 Wu C-C, Liou K., Shan, S. J., and Tseng, C. L. (2008): Variation of ionospheric total electron  
565 content in Taiwan region of the equatorial anomaly from 1994 to 2003, *Advances in*  
566 *Space Research* 41, 611–616.
- 567 Yu, Xiao, Chengli Shi, Dun Liu, and Weimin Zhen. 2012. "A Preliminary Study of the NeQuick  
568 Model over China Using GPS TEC and Ionosonde Data." 2012 10th International  
569 Symposium on Antennas, Propagation and EM Theory, ISAPE 2012 (36):627–30.
- 570 Cherniak and Zakharenkova 2016 Cherniak, Iurii and Irina Zakharenkova. 2016. "NeQuick and  
571 IRI-Plas Model Performance on Topside Electron Content Representation: Spaceborne  
572 GPS Measurements." 1–15.
- 573 Huang, Xueqin and B. W. Reinisch. 1996. "Vertical Electron Density Profiles from the  
574 Digisonde Network." 18(6).
- 575 Zhang, Man Lian, Sandro M. Radicella, Jian Kui Shi, Xiao Wang, and Shun Zhi Wu. 2006.  
576 "Comparison among IRI, GPS-IGS and Ionogram-Derived Total Electron Contents."  
577 *Advances in Space Research* 37(5):972–77.
- 578 Zhang, M.I., J.K. Shi, X. Wang and S.M. Radicella (2004), Ionospheric variability at low latitude  
579 station: Hainan, China, *Advances in Space Science*, 34, 1860-1868

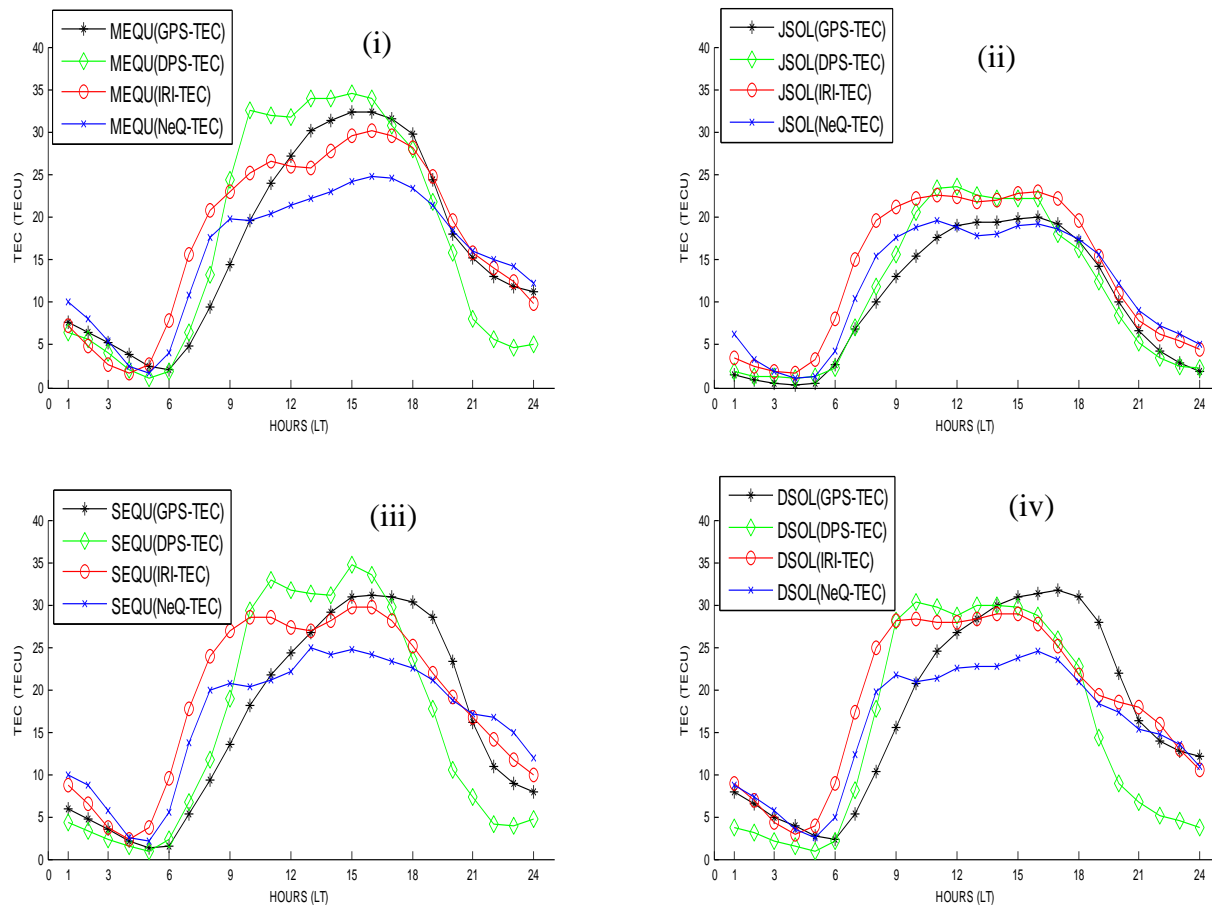
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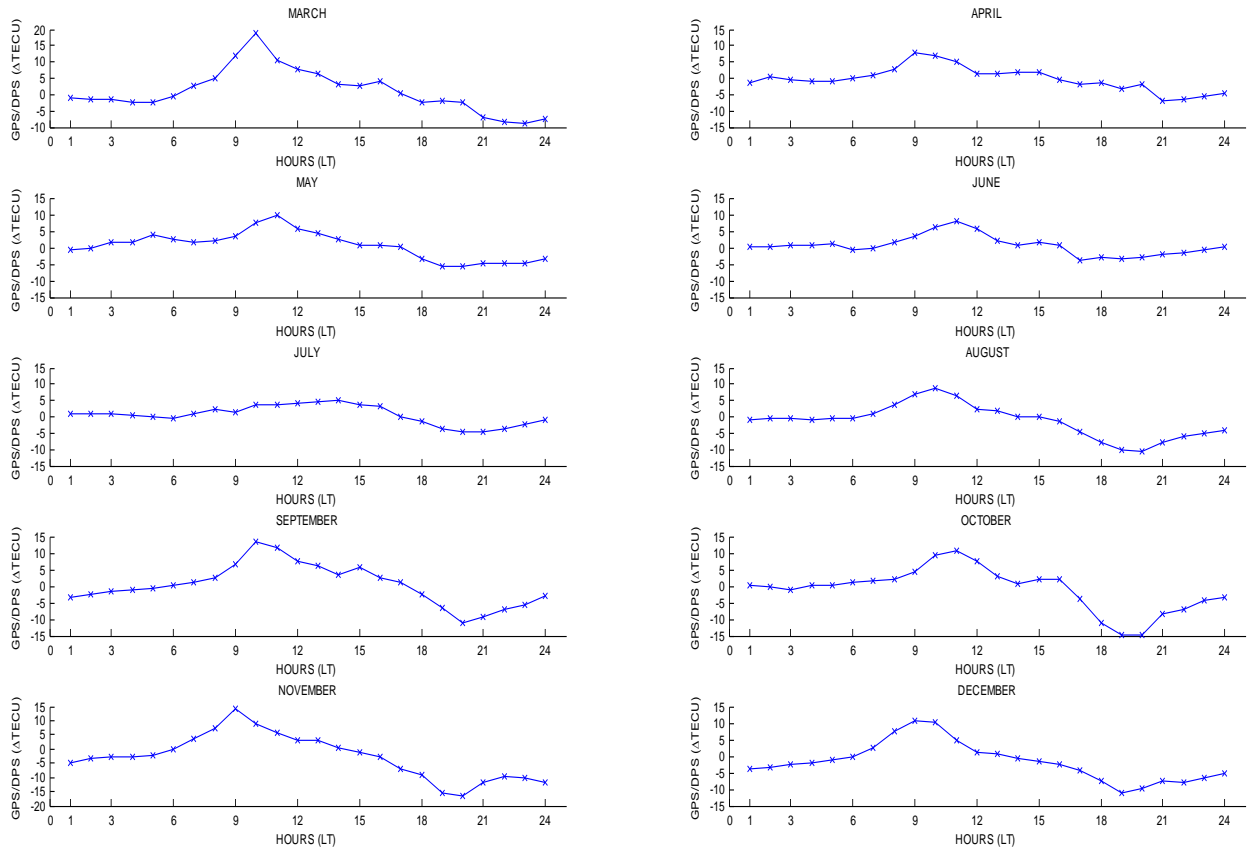
## 582 **8.0 Figures**



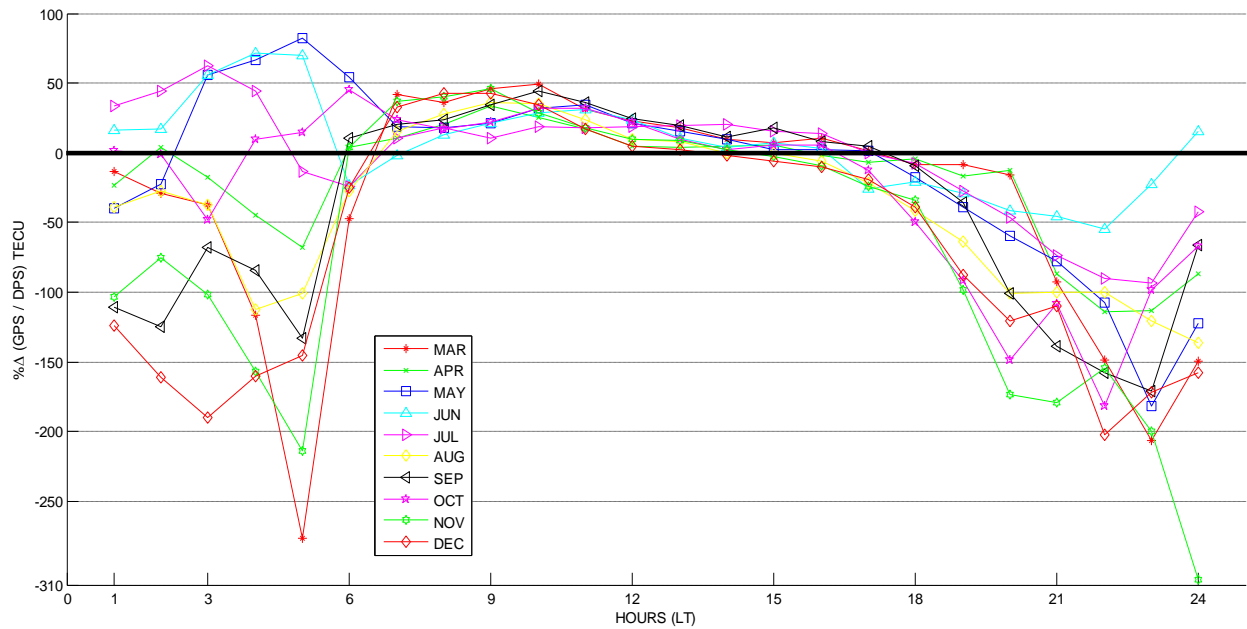
583  
 584 Figure 1a Hourly variations of monthly median of five quiet days of GPS, DPS, IRI, and NeQ-  
 585 TEC in March-December during quiet period. GPS-TEC is in black line with star symbol, DPS-  
 586 TEC is in green line with diamond symbol, IRI-TEC is in red line with zero line with star symbol  
 587 and NeQ-TEC is in blue line with star symbol.



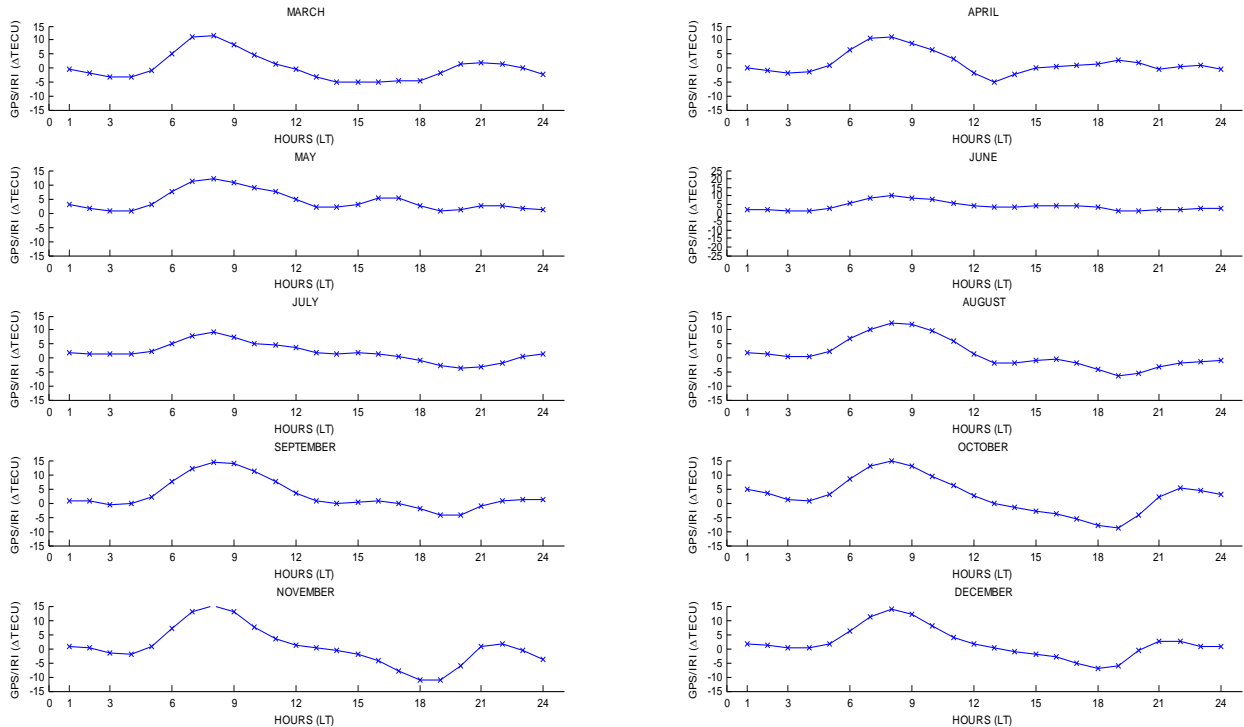
588  
 589 Fig. 1b Seasonal variations of GPS-TEC, DPS-TEC, IRI-TEC and NeQ-TEC for: (i) March  
 590 Equinox, (ii) June Solstice, (iii) September Equinox, and (iv) December Solstice over Ilorin  
 591 during quiet periods in 2010. The line colors and symbols are the same as for diurnal variation in  
 592 Figure 1a for all seasons.



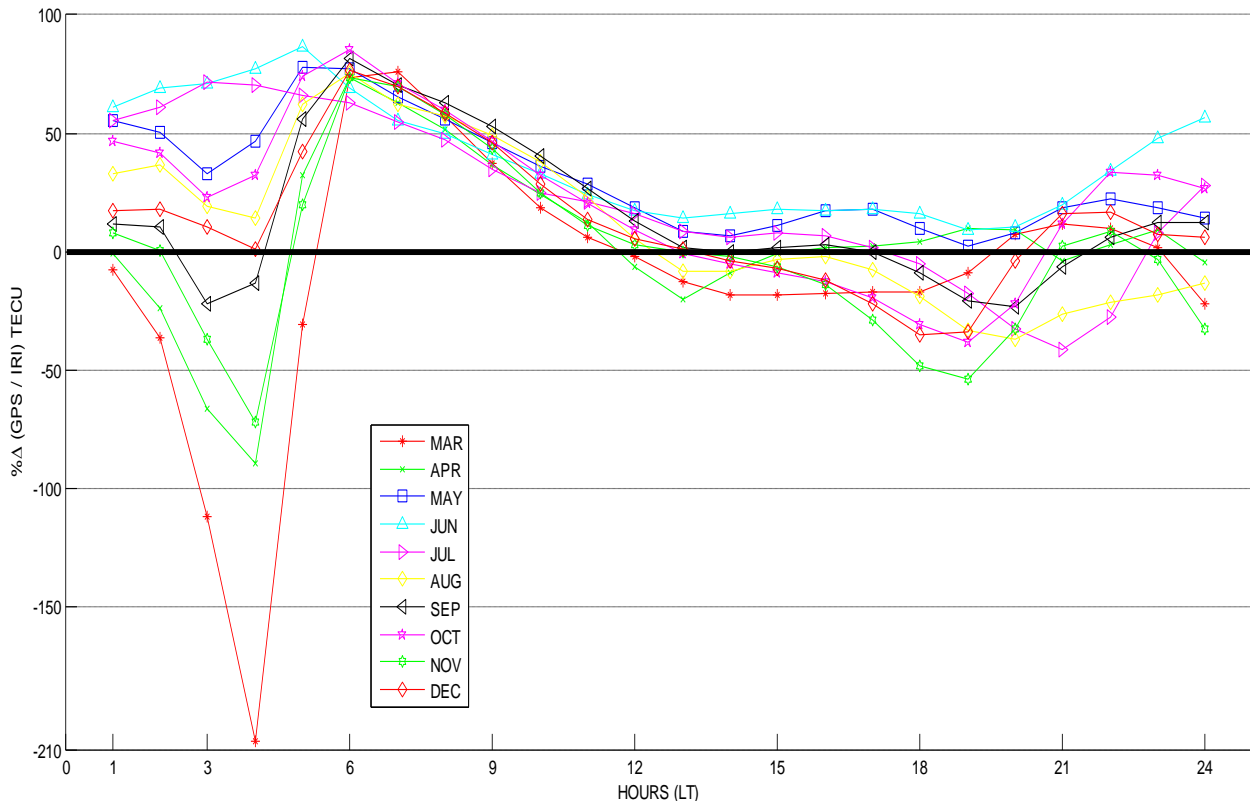
593  
 594 Figure 2a Hourly variations of  $\Delta$ TEC between the GPS-TEC and DPS-TEC from March -  
 595 December during quiet period.  
 596



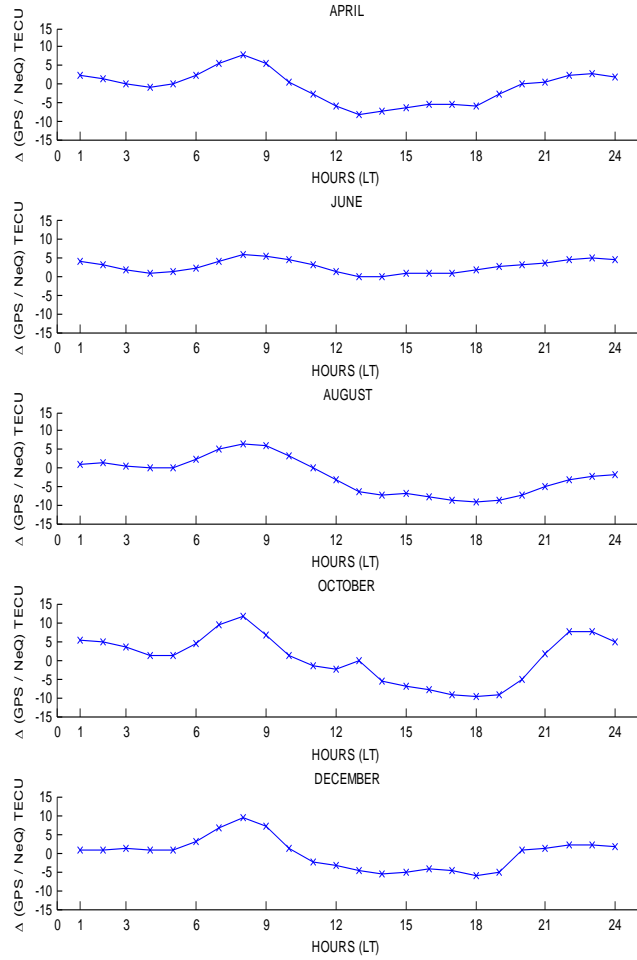
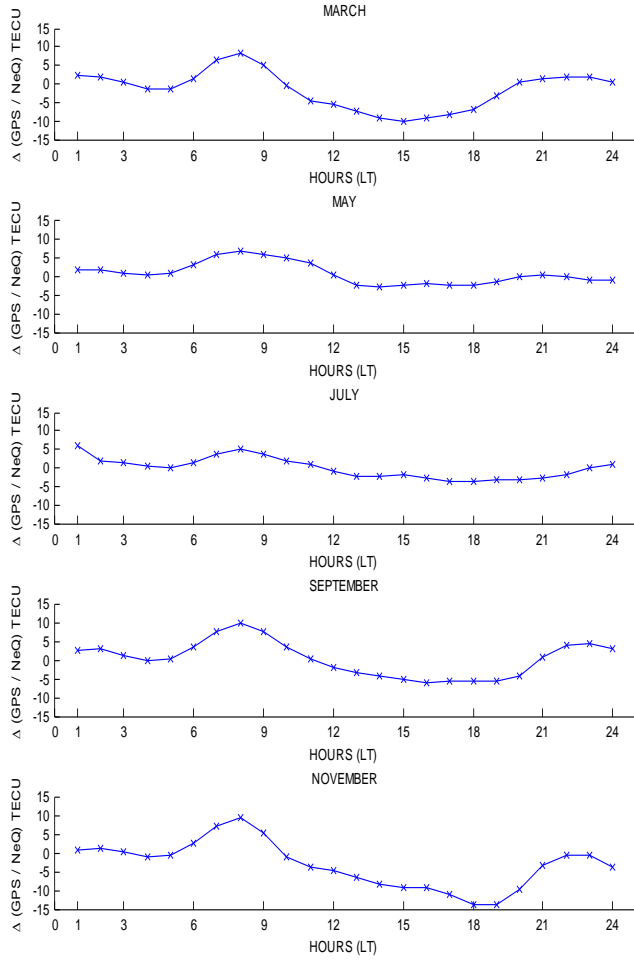
597  
 598 Figure 2b Mass plot of  $\% \Delta$  TEC between the GPS-TEC and DPS-TEC from March - December  
 599 during quiet period. The legend represents line colors and symbols of each deviation in all  
 600 months.



601  
 602 Figure 3a Hourly variations of  $\Delta$ TEC between the GPS-TEC and IRI-TEC from March -  
 603 December during quiet period.  
 604

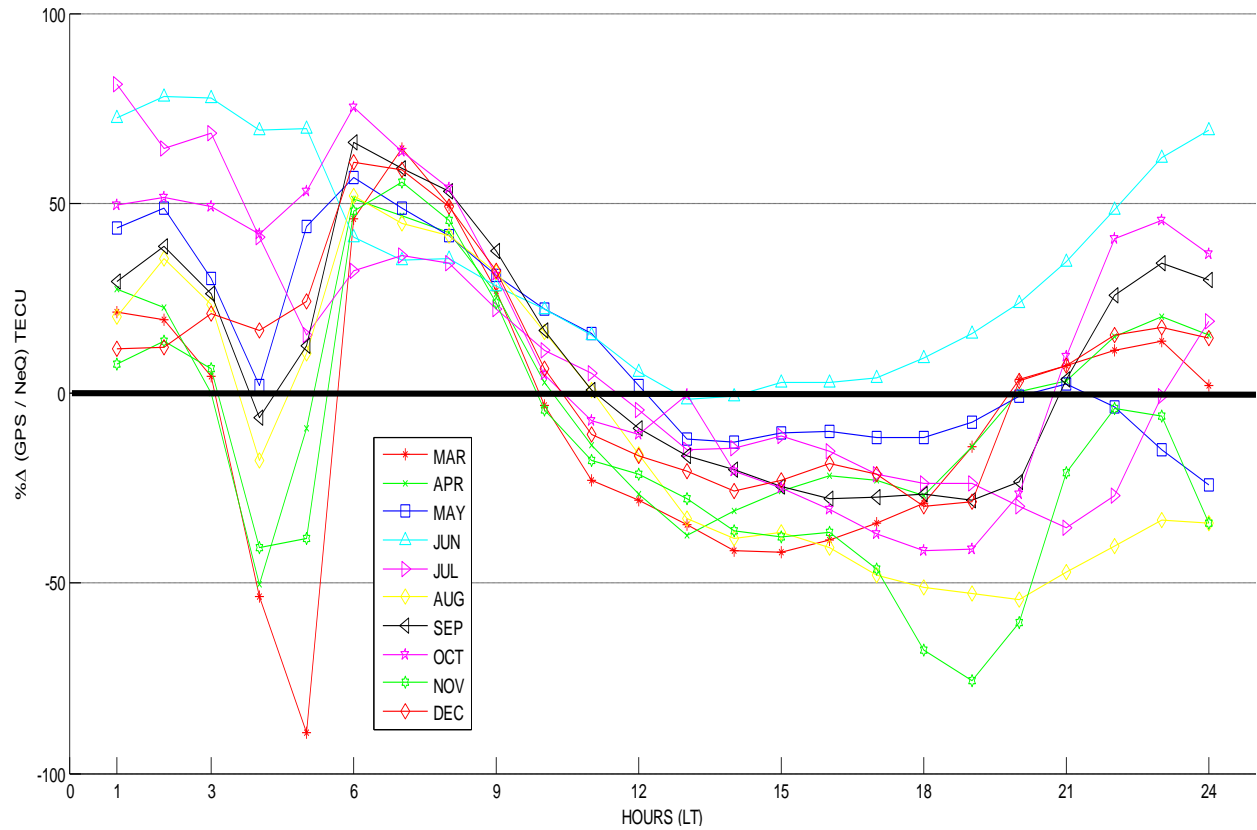


605  
 606 Figure 3b Mass plot of  $\% \Delta$  TEC between the GPS-TEC and IRI-TEC from March to December  
 607 during quiet period. The line colors and symbols are the same as in Figure 2b



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Figure 4a Hourly variations of  $\Delta\text{TEC}$  between the GPS-TEC and NeQ-TEC from March - December during quiet period.



611  
 612 Figure 4b Mass plot of %  $\Delta$  TEC between the GPS-TEC and NeQ-TEC from March - December  
 613 during quiet period. The line colors and symbols are the same as in Figure 2b.  
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