

1 **SOLAR ECLIPSE-INDUCED PERTURBATIONS AT MID-LATITUDE DURING**
2 **THE 21 AUGUST 2017 EVENT**

3
4 Bolarinwa J. Adekoya¹, Babatunde O. Adebeyi², Timothy W. David¹, Stephen O. Ikubanni², Shola J.
5 Adebiyi², Olawale. S. Bolaji^{3,4} and Victor. U. Chukwuma¹

6 ¹Department of Physics, Olabisi Onabanjo University, P.M.B. 2002, Ago Iwoye, Nigeria

7 ²Space Weather Group, Department of Physical Sciences, Landmark University, P.M.B 1001, Omu-Aran,
8 Kwara State, Nigeria

9 ³Department of Physics, University of Lagos, Akoko – Yaba, Lagos, Nigeria

10 ⁴Department of Physics, University of Tasmania, Hobart, Australia

11
12 Correspondence to: Bolarinwa J. Adekoya (adekoyabolrinwa@yahoo.com; adekoya.bolarinwa@ouagoiwoye.edu.ng)

13
14 **Abstract**

15 A study of the response of some ionospheric parameters and their relationship in describing the behaviour
16 of ionospheric mechanisms during the solar eclipse of 21 August 2017 is presented. Mid-latitude stations
17 located along the eclipse path and with data availability on the Global Ionospheric radio Observatory (GIRO)
18 database were selected. The percentage of obscuration at these stations range between 63-100%.
19 Decrease in electron density during the eclipse is attributed to reduction in solar radiation and natural gas
20 heating. The maximum magnitude of the eclipse coincided with $hmF2$ increase and with a lagged maximum
21 decrease in $NmF2$ consistently at the stations investigated. The results revealed that the horizontal neutral
22 wind flow is as a consequence of the changes in the thermospheric and diffusion processes. The unusual
23 increase/decrease in the shape/thickness parameters during the eclipse period relative to the control days
24 points to the perturbation caused by the solar eclipse. The relationship of the bottomside ionosphere and
25 the F2 layer parameters with respect to the scale height are shown in the present work as viable
26 parameters for probing the topside ionosphere during eclipse. Furthermore, this study shows that in
27 addition to traditional ways of analysing the thermospheric composition and neutral wind flow, proper
28 relation of standardized $NmF2$ and $hmF2$ can be conveniently used to describe the mechanisms.

29

30 **Keywords:** solar eclipse; solar radiation; bottomside profile parameters; $NmF2$ and $hmF2$; Topside
31 ionosphere; GIRO database.

32

33 **1 Introduction**

34 Solar eclipse provides opportunity to study the causes of drastic changes in the atmosphere arising from
35 reduction in solar radiation and plasma flux. The atmosphere responded to these changes by modifying the
36 electrodynamic processes and ionization supply of its species to the nighttime-like characteristics during
37 the daytime. Different physical mechanisms (e.g. neutral wind, thermospheric composition, diffusion
38 process etc.) that explain the distribution of plasma at the different ionospheric layers are well established.
39 However, these mechanisms do compete with themselves in explaining the ionosphere, especially the
40 topside ionosphere (see Gulyaeva, 2011).

41

42 At mid-latitudes, the effect of diffusion processes and its relationship with the thermospheric compositions
43 has been extensively studied during episodes of solar eclipse (Muller-Wodarg et al., 1998; Jakowski et al.,
44 2008; Le et al., 2009; Wang et al., 2010; Chuo, 2013). At equatorial and low-latitude regions, the $E \times B$
45 plasma drift had been used to explain the circumstances of solar eclipse on transport processes (Adeniyi et
46 al., 2007; Adekoya et al., 2015). Recently, attention has been drawn to the study of the topside ionosphere
47 during an eclipse for improved prediction and modelling (Huba and Drob, 2017; Chrniak and Zakharenkova,
48 2018). Reinisch et al., (2018) compared the modelled and measured studies of electron densities at the
49 altitude range of about 150 - 400 km during the eclipse. They found that at lower altitude (at about 150 km)
50 the modelled and the measured agreed well to the changes in the altitude profile of electron density
51 compared to at higher altitudes. The authors however posited that it would be improved if the model
52 $NmF2$ peak falls more slowly to better match the data. Consequently, the present study investigates the
53 effects of solar eclipse of August 21, 2017 on the constituents of the ionosphere at mid-latitudes using
54 some ionosonde data (bottomside parameters, scale height (H) estimated from the fitted α -Chapman layer)
55 which have not been given much attention in previous works especially in analysing solar eclipse effect.
56 Using these parameters to analyse the circumstances of solar eclipse at the topside ionosphere and its
57 plasma distribution mechanisms make this paper significantly different from previous studies. **Thus**, we
58 intend to achieve by analysing the ionospheric parameters that controls the distribution of plasma at the
59 topside and bottomside layers of the F2 region. To shed light on these analysis, section 2 highlights the data
60 source, methodology, and path of the eclipse. The **results and discussion** were presented in section 3, while
61 section 4 presents the summary and concluding remark of the result.

62

63 **2 The solar eclipse path and Data source**

64 With regards to the eclipse of 21 August 2017, the totality of the eclipse is visible from within a narrow
65 corridor that traverses the United States of America. However, in the surrounding areas, which include all
66 of mainland United States and Canada, the eclipse was partial. From the footprint of the Moon's shadow as
67 seen from some locations, the eclipse started **from around 17:00 UT and ended around 20:00 UT. Figure 1**
68 **shows the detail coverage area and circumstances of the solar eclipse.** More details of its path can be seen
69 via NASA – Total solar eclipse of 2017 August 21 (<https://eclipse.gsfc.nasa.gov/>). The details on the local
70 circumstances of the eclipse, the time of the first, mid and last contact of the eclipse over the ionosphere of
71 the investigated stations was highlighted in table 1. More details on the total solar eclipse event and its
72 partiality, the circumstances surrounding its progression and its magnitude of obscuration can be obtained
73 through the link http://xjubier.free.fr/en/index_en.html. The path of the eclipse informed the choice of
74 stations. The ionospheric data used for this study for the selected mid-latitude stations were obtained from
75 the Global Ionospheric Radio Observatory (GIRO) networks, <http://giro.uml.edu/> (Reinisch and Galkin 2011)
76 and manually validate. The calculated daily average of summation Kp, Ap and solar flux indices was

77 obtained from the National Space Science Data Centres (NSSDC's) OMNI database
78 <https://omniweb.gsfc.nasa.gov/>.

79

80 **3 Methods of data analysis**

81 $NmF2$ values for both the eclipse and control days were obtained from their corresponding critical
82 frequencies ($foF2$) using the expression: $NmF2 = ((foF2)^2 / 80.5) e/m^3$. The control day value is the average
83 value of the two days before/after the eclipse day (i.e. 6, 12, 24 and 27). These reference days were chosen
84 such that they have similar geomagnetic, interplanetary and solar properties with the eclipse day. The daily
85 average value of control days and eclipse day interplanetary index (Ap and Kp), and solar flux unit index
86 (F10.7) ranges from 8 – 12 nT for Ap, 2 – 3 for Kp index and 75.6 – 89.1 sfu (1 solar flux unit (sfu) = 10^{-22}
87 Wm $^{-2}$ Hz $^{-1}$) for F10.7, indicating that geomagnetic and solar activities of these days is unsettled (see
88 Adekoya et al., 2015 for classification of geomagnetic activity). The typical behaviour of the $NmF2$ and
89 $hmF2$ on the eclipse day (i.e. $NmF2e$ and $hmF2e$) was compared with that of the control day ($NmF2c$ and
90 $hmF2c$) to observe the changes brought by the short period of loss of photoionization in the ionosphere.
91 This will measure the direct consequence of the solar radiation disruption (due to the eclipse) on the
92 ionospheric chemical, transport and thermal processes in the F2 layer. The ionized layer depends majorly
93 on three parameters, viz: $NmF2$, $hmF2$, and the plasma scale height (H_m).
94

95 The GIRO provides access autoscaled values of ionospheric parameters generated by Automatic Real-Time
96 Ionogram Scaler with True height (ARTIST) algorithm, which is inherent in the UMLCAR-SAO Explorer
97 (Reinisch and Huang 1983; Galkin et al., 2008; Reinisch and Galkin 2011), facilitates the derivation of
98 bottomside profiles. From the ULMCAR-SAO Explorer, the manually scaled ionogram with high accuracy are
99 calculated from the standard true-height inversion program (Reinisch and Huang, 1983; Huang and
100 Reinisch, 1996). The parameters obtained include the critical frequency ($foF2$, Hz), and its height ($hmF2$,
101 km) of the F layer and the shape ($B1$), and thickness ($B0$) parameters. Likewise, the scale height (H_m) of the
102 F2 layer is obtained from the bottomside. It is estimated from the fitted α -Chapman function with a
103 variable scale height, $H(h)$, to the measured bottomside profile $N(h)$, which then determined as the
104 Chapman scale height at $hmF2$ (i.e. $H(h > hmF2) \approx H_m (hmF2)$) (Huang and Reinisch 2001; Reinisch and
105 Huang 2001; Reinisch et al., 2004). The topside profile is then related to the scale height at the layer, from
106 the bottomside profile, represented with α -Chapman function (Reinisch and Huang, 2001). This is because
107 the Chapman function described the electron density profile, $N(h)$ aptly. Also, H_m provides a linkage
108 between the bottomside ionosphere and the topside profiles of the F region (Liu et al., 2007). Therefore,
109 H_m describes the constituents of the ionospheric plasma, which decreases with increasing altitude. The
110 fitting formulas of α -Chapman function are provided in equation 1 below.

$$N_e = N_m F2 \exp \left\{ \frac{1}{a} [1 - z - \exp(-z)] \right\}; \quad z = \frac{h - hmF2}{H_m}$$

1

112 where all the parameters have their usual meaning.
113
114 However, Xu et al. (2013) and Gulyaeva (2011) related ionospheric F2 - layer scale height, H to the topside
115 base scale height, H_{sc} , given by $H_{sc} = hsc-hmF2 \approx 3 \times H_m$). Where hsc is the height at which the electron
116 density of the F2-layer falls by a factor of an exponent, at an upper limit of 400 km altitude (i.e. $NmF2/e$)
117 (see Xu et al., 2013). That is, the region where electron density profile gradient is relatively low. Gulyaeva
118 (2011) showed theoretically that H_{sc} increase over H_m by a factor of approximately three (3) and is a
119 consequence of the $Ne/NmF2$ ratio (Ne – plasma density), which corresponds to H_m in the Chapman layer.
120 At altitudes very close to $hmF2$, the ratio equals 0.832, while it is 0.368 at altitudes beyond the $hmF2$.
121 Therefore, we adopted the definition of Gulyaeva (2011) for the topside base scale height as the region of
122 the ionosphere between the F2-peak and 400 km altitude. Summarily, the topside based scale height
123 ionosphere here is defined as the region between the F2 peak and hsc or $3H_m$. It is thus evident that H is a
124 key and essential parameter in the continuity equation for deriving the production rate at different
125 altitudes, a pointer to the F2 topside electron profiler, as well as a good parameter for evaluating the
126 transport term (Yonezawa, 1966; Huang and Reinisch, 2001; Reinisch and Huang, 2001; Belehaki et al.,
127 2006; Reinisch et al., 2004). Consequently, the parameter H_m can be used as a proxy for observation
128 relating to the topmost side electron density profile. Furthermore, the division of the topsides and the
129 bottomside ionosphere may be related to the difference in the effective physical mechanisms in the
130 regions. Hence, the bottomside parameters $B1$ and $B0$ of the ionosphere, as presented in this work, helped
131 in examining the perturbation of solar eclipse in the bottomside ionospheric F2 layer.

132
133 **4 Results and Discussion**
134 This section presents the temporal evolution of the maximum electron density ($NmF2$), and its
135 corresponding height ($hmF2$) over the ionosphere at the selected mid-latitude stations along the path of
136 solar eclipse of 21 August 2017. The control day variation relative to the eclipse day is also presented.
137 Figure 2 presents the variation of maximum electron density and the corresponding peak height, during
138 both the eclipse and control days. Figure 3 depicts the variation of scale height and the bottomside
139 parameters ($B0$ and $B1$) due to the eclipse by superposing plots for both the eclipse and control days.
140 Analysis of these parameters during an eclipse event may help in the modelling of the ionospheric profiles
141 (the topsides and bottomside electron density distribution profile) during the short nighttime-like period of
142 the day.

143
144 Figure 2a presents the $NmF2$ and $hmF2$ variations during the eclipse event and the control day over the
145 Idaho National Lab; having an obscuration magnitude of 100% around the daytime period. The effect of the
146 disruption of solar radiation was evident as the $NmF2$ started decreasing at the first contact of the eclipse
147 compared to an incessant increase on the control day in Fig. 2ai. The start time or first contact (08:43:31

148 LT), the maximum magnitude period (10:01:53 LT) and the end time or the last contact (11:25:46 LT) of the
149 eclipse are marked with the vertical lines S, M and E respectively. The decrement in $NmF2$ during the
150 eclipse phase was due to reduction in the ionization. This reduction caused changes in the photochemical
151 and transport process of the atmosphere during the daytime, thus exhibiting nighttime characteristics. It
152 should be noted that the maximum decrease in $NmF2$ did not coincide with the maximum magnitude of the
153 eclipse obscuration, rather with a time lag of few minutes, i.e., 1030 LT. This lag period fell within the
154 relaxation period over Idaho ionosphere, with $NmF2$ and $hmF2$ simultaneously attaining their peak
155 magnitudes of 1.67 e/m^3 and $\sim 239 \text{ km}$. Hence, the ionosphere returned to its pre-eclipse state. Contrary to
156 the decrease in the $NmF2$ amplitude at the recovery phase of the eclipse, the $hmF2$ increases, attained 239
157 km peak around 1030 LT and then decreases depicting the eclipse caused morphology.

158

159 The ionosphere over Boulder, Eglin AFB, Austin, Millstone Hill and Point Arguello did not show any contrary
160 variation to that observed over Idaho during the eclipse event. The decrease and increase in $NmF2$ and
161 $hmF2$ after the maximum magnitude are simultaneous. The only exception was that the local time at which
162 each station observed the effects were different. Their obscuration percentage ranged from 62.5 – 93.37%.
163 This did not cause any significant change in the way they responded to the reduction in solar heating. The
164 ionosphere over Boulder experienced the totality of the eclipse with 93.37 % magnitude, which is next to
165 Idaho (100%) in obscuration, the $hmF2$ was observed to increase few minutes after the maximum
166 magnitude of the obscuration. This behaviour is typical for other stations at the eclipse window, but the
167 time of $NmF2$ minimum decrease did not always coincides with the $hmF2$ enhancement after the maximum
168 obscuration. These observations posit that the minimum rate of electron production does not necessarily
169 translate to the peak electron density of the molecular gases formed. This is because the electron
170 concentration depends on the loss rate by dissociative recombination, too.

171

172 At mid-latitudes, the ionospheric F2 plasma distribution is controlled by diffusion processes (Rishbeth
173 1968). There are two basic mechanisms that define the diffusion process during an eclipse: First is the
174 coolness brought by the partial removal of photoionization (Müller-Wodarg et al., 1998), which is believed
175 to instigates the downward diffusion process, and the atmospheric expansion due to the gradual increase
176 in the temperature after the totality. The downward diffusion process was related to the increase in the
177 molecular gas (N_2) concentration during the cooling process. However, the aftermath of the coolness was
178 related to the upward diffusion process. These mechanisms were proxy to the electron density distribution
179 during the eclipse window. Our analysis suggests that the observed decrease in $NmF2$ is due to the
180 downward diffusion flux of the plasma while the increase that followed is by upward diffusion (e.g. Le et al.,
181 2009; Adekoya and Chukwuma 2016). Several works on eclipse (Müller-Wodarg et al., 1998; Grigorenko et
182 al., 2008; Adekoya and Chukwuma 2016; Hoque et al., 2016) have shown that it was not just the electron

183 density that is being affected during an eclipse window, but the thermospheric wind as well, since the
184 thermospheric wind emanating from the ratio of gas species is related to the variation in electron density.
185 It has been observed that the increase in the mean molecular gas of thermospheric composition decreases
186 the electron density and vice versa. Le et al. (2010) related the valley of electron density distribution during
187 the eclipse phases to the contraction/compression and expansion of the atmosphere brought by the
188 decrease and increase in temperature. Chukwuma and Adekoya (2016) attributed the decrease in the
189 electron temperature to the downward vertical transport process and the decrease in the cooling process
190 to the upward vertical transport process.

191

192 Figure 3 describes the variation of H_m , $B1$ and $B0$ in three columns respectively for all the stations. Looking
193 at the H_m plots, one can see that there was a define morphological description of H_m at the eclipse window.
194 From the first contact of the eclipse, there was an incessant increase in peak variation that maximized some
195 minutes after the maximum contact of the eclipse, i.e., about 15 – 45 mins later. Following the peak
196 magnitude after the maximum contact of the eclipse, the H_m sharply decreases, reaching the minimum
197 peak before its rather increase throughout the remaining period of the eclipse second phase. It was further
198 observed that the minimum decrease in $NmF2$ amplitude corresponds to increase in H_m at all stations;
199 implying the upward lifting of the topside electron to the region of higher altitude at the eclipse window.
200 Hence, the scale height variation highlights the decrease in electron production and the vertical distance
201 through which the pressure gradient falls at the topside during the eclipse activity. The observation
202 illustrates the mutual relationship between the $NmF2$ and H_m , which may aid in extrapolating the topside
203 ionospheric profile (Gulyaeva, 2011). In essence, scale height changes observed during the eclipse window
204 can be used to explain the pressure gradient, electron density distribution and transport processes. In this
205 sense, the diffusion coefficients are expressed as ratio of determinants (determinant here refers to the
206 concentration of species ($[O]$ and $[N_2]$), with the size of the determinants depending upon both the number
207 of species in the gas mixture and the level of approximation. Therefore, the increase (decrease) in the scale
208 height can be used as a proxy for downward (upward) diffusion process at the topside ionosphere.
209 Consequently, the thermospheric wind, which causes plasma distribution in the topside ionosphere, is
210 induced by solar radiation. Moreover, the significant changes observed in the scale height variation during
211 the eclipse window also indicated that transport processes are affected as they are temperature
212 dependent. Therefore, changes in the thermospheric compositions due to the solar eclipse at the topside
213 layer will affect the density profiles of the ionosphere.

214

215 It is noteworthy that the increase (decrease) in the scale height decreases (increases) the electron density
216 during the eclipse window. The sensitivity of electron density to temperature at the topside directly affects
217 the electron density profile (e.g. Wang et al., 2010); as cooling due to decrease in temperature results in

218 decrease in the electron density via reduced ionization. This indicates that the decrease (increase) in
219 electron temperature at the topside ionosphere causes the increase (decrease) in the scale height, which is
220 related to the diffusion and transport processes and subsequently affect the pressure gradient of the
221 plasma. From plots of H_m (fig. 3) and $NmF2$ (fig. 2), it was observed that the minimum decrease in $NmF2$
222 corresponded with peak increase in scale height. This implies that the topside ionosphere is more sensitive
223 (than the bottomside) to any changes in the solar radiation. Thus, the pressure gradients can be analysed in
224 terms of either the scale height or electron density during solar eclipse.

225

226 From column 2 and 3 of Figure 3, we observed that the measured shape ($B1$) and thickness ($B0$) parameters
227 of the ionosphere over these stations exhibit significant variations during the eclipse event. $B1$ responded
228 with a decrease at the first contact of the eclipse compared to the control day. This decrease was gradual
229 throughout the eclipse window and followed the variation of solar ionizing radiation. However, $B0$ variation
230 differs to that of the $B1$ observation. The $B0$ increases from the first contact and reached the maximum
231 peak few minutes after the maximum obscuration magnitude, which coincided with the minimum decrease
232 in $B0$. Generally, the pattern of the day to day variation of the bottomside parameters was the average
233 morphology, but the increase in the $B0$ and the decrease in the $B1$ parameters during the eclipse period
234 compared to the control day was a notable one and can be related to the perturbation caused by the solar
235 eclipse. During the eclipse, the solar radiation was lost; trapped atomic ions O^+ was converted into
236 molecular ion (NO^+ and O_2^+) by charge transfer, owing to the sufficient concentration of molecular gasses
237 (N_2 and O_2) (Rishbeth, 1988). The height of the ionospheric slab indeed increased with reduced width,
238 which is attributable to compression due to loss of solar heating.

239

240 The behaviour of the ionosphere can be explained during solar eclipse with any of the components that
241 constitute the topside and the bottomside ionosphere and can be looked at, from the angle of the
242 percentage of concentration of the components. In this regard, the deviation percentage of $NmF2$ ($\delta NmF2$)
243 and $hmF2$ ($\delta hmF2$) during the eclipse day away from the control day were plotted in Figure 4. This is done
244 to describe the contribution of the thermospheric wind and compositions. Although observing the variation
245 of $NmF2$ and $hmF2$ alone can be used for observing the changes in the behaviour of the thermospheric
246 compositions and wind flow, if properly analysed, but it is more convenient to describe these mechanisms
247 by standardizing the original variables used during the event. The normalization effort (with the use of
248 $\delta NmF2$ and $\delta hmF2$) presents the original variation of $NmF2$ and $hmF2$ onto directions which maximize the
249 variance. Consequently, the result can be used for analyses of any mechanisms that drive the ionospheric
250 plasma, if properly related.

251

252 The deviation percentage in Figure 4 was defined as the ratio of $((NmF2e - NmF2c)/NmF2c) \times 100$. The
253 same relation is defined for the $hmF2$ parameter. As earlier pointed out, during eclipse period, neutral
254 composition becomes the dominant chemical process arising from diffusion activities. The increase in the
255 neutral composition leads to the increase in the molecular gas concentration and compete with diffusion
256 process. Hence the deviation percentage discusses the neutral composition changes and delineate how
257 these changes may affect the electron densities as well as its profiles in the atmosphere during the eclipse.
258 The respective maximum and minimum peak response of the deviation percentage is attributed to the
259 enhancement and depletion of $\delta NmF2$. One can sees from the plots, the deviation percentage started
260 increasing at the first contact of the eclipse (the first dashed vertical line) and reached the maximum,
261 appearing few minutes after the maximum magnitude of the eclipse (the second dashed vertical line). This
262 behaviour is similar to the conditions of the neutral compositions during the eclipse event reported by
263 Muller-Wodarg et al. (1998).

264

265 Another important process observed in this study is the neutral wind flow effect. To identify the direction
266 of the wind, the $\delta NmF2$ colour legend in the contour plots was used in Figure 4. The negative values
267 represent a westward wind contribution and the positive values is for the eastward wind. Looking at the
268 marked eclipse region in the figure, it was revealed that the $\delta NmF2$ started decreasing from the first
269 contact of the eclipse, maximized few minutes after the maximum contact mark and, thereafter decreases.
270 It has been established that at daytime, the peak height of the plasma will be reduced due to lost in
271 recombination. At nighttime, equatorward neutral wind drives the F2-layer plasma to higher altitudes
272 where recombination rate is slower. The ionospheric processes during solar eclipse is said to represent a
273 partial nighttime/sunset ionospheric process (Adekoya et al., 2015; Adekoya and Chukwuma, 2016). Thus,
274 the F2 plasma behaviour at the eclipse window is induced by the equatorward neutral wind flow. The
275 neutral wind acts jointly with the plasma flows from the topside ionosphere, resulting in F2 region plasma
276 density variation. Therefore, the westward/eastward neutral wind flow is related to the
277 depletion/enhancement in the deviation, which was clearly shown in the marked eclipse region of the
278 figure. The plots in Figure 3 had established the ionospheric dynamics of diffusion processes, neutral
279 compositions and the flow of neutral wind caused by the eclipse perturbation, which can invariably reduce
280 the effectiveness and reliability of radio wave propagation.

281

282 Relative to the mutual relationship between the topside and bottomside ionosphere, we considered the
283 linear correlation coefficient (R) of H_m versus $hmF2$ and H versus $B0$ during the eclipse window, In Fig. 5., R
284 ranges from (0.80 - 0.90) for $H_m/hmF2$ relationship, and 0.57-0.89 for the $H_m/B0$ connection. This good
285 linear agreement revealed the dependence of $hmF2$ and $B0$ on the scale height. Apart from revealing the
286 dependence between the parameters, the relationship may also provide a convenient way for modelling

287 the topside profile from the knowledge of the bottomside parameter, $B0$, during the eclipse period.
288 Further, fig. 6 illustrates the relationship between the bottomside (continuous line) and the topside
289 (dashed line) ionosphere over Idaho National Lab during solar eclipse compared to the non-eclipse period.
290 On the left side was the ionospheric profile during the first contact of the eclipse, the middle and right-side
291 profiles are during the maximum contact and last contact of the eclipse respectively. The black curve
292 represents the profile for the eclipse day (August 21) and the red curve is for the one of the selected
293 reference days, August 27. It is clear from the plots that the ionospheric profiles vary with the solar ionizing
294 radiation at the eclipse window and shows the suitability of using the bottomside F-region for probing the
295 topside ionosphere. This behaviour was typical for the ionospheric profiles from other stations along the
296 path of the eclipse. Also, the strong correlation between $hmF2$ and H_m indicates that there may be some
297 interrelated physical mechanisms controlling the behaviour of the plasma at the topside ionosphere during
298 solar eclipse. That is, $hmF2$ is strongly depends on neutral wind flow and explain the state of thermospheric
299 compositions (e. g. Liu et al., 2006; Fisher et al., 2015). Since all these parameters competes during the
300 eclipse, one can argue that with the accessibility of one, in place of the other (as a consequence of their
301 relationship), the prediction and modelling of the ionosphere can be conveniently achieved.

302

303 **5 Conclusions**

304 This paper presents the induced perturbation of solar eclipse of 21 August 2017 on the ionospheric F
305 parameters and how they describe the mechanisms of the ionosphere at mid-latitude. The perturbation
306 effects and dynamics during a solar eclipse episode using ionospheric F2 parameters ($NmF2$ and $hmF2$), the
307 bottomside profile thickness ($B0$) and shape ($B1$) parameters of electron density and the plasma scale
308 height (H_m), which are not often used for eclipse study, were investigated. These parameters represent the
309 state of the F-region ionosphere. The changes observed during the eclipse phase is related to the reduction
310 in solar radiation and natural gas heating. The $NmF2$ minimum was attained around 30 - 45 minutes after
311 the totality of the eclipse when it decreases to about 65% of its control day. This decrease in $NmF2$ was
312 uplifted to the higher altitude where recombinational rate is reduce compared to the non-eclipse day. The
313 thickness and shape parameters which are often limited to the bottomside F-region were seen as viable
314 parameters for probing the topside ionosphere, relative to the scale height during the eclipse. Therefore,
315 their relationship in describing one another is established. The implication is that eclipse-caused
316 perturbation could have been better explained using some ionosonde parameters. The changes in the
317 neutral wind flow, thermospheric compositions and diffusion processes found their explanation in the
318 behaviour of the F region plasma during eclipse. In addition, it can be concluded that the behaviour of
319 $\delta NmF2$ and $\delta hmF2$ during eclipse can be conveniently used to describe the mechanisms of thermospheric
320 composition and wind flow.

321

322 **Acknowledgements**
323 We acknowledge use of global ionospheric Radio Observatory data provided by ULMCAR
324 (<http://ulcar.uml.edu/DIDBase/>) and the World Data Center for Geomagnetism, Kyoto
325 (<http://wdc.kugi.kyoto-u.ac.jp/index.html>) for geomagnetic activity data. We thank the management team
326 of the national Aeronautics and Space Administration (NASA) service (<http://eclipse.gsfc.nasa.gov>) and
327 http://xjubier.free.fr/en/site_pages/SolarEclipseCalc_Diagram.html for progression and eclipse local
328 circumstances information. The authors thank Professor Ljiljana R, Cander and the **anonymous reviewers**
329 for their constructive corrections and suggestions that tremendously improved the structure and quality of
330 the paper.

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477 **Table Caption**

478 **Table 1:** List of ionosonde station, geographic coordinate, eclipse progression time and percentage of
479 maximum obscuration.

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482 **Figure Captions**

483 **Figure 1:** The orthographic map showing the coverage area and circumstances of the solar eclipse, and the
484 observatory stations of the total solar eclipse event of August 21, 2017 . The thick blue line region of
485 represents the path of the maximum magnitude of the eclipse and the pale blue lines mark the region of
486 where the partial eclipse is experienced, with the magnitude of partiality.

487

488 **Figure 2:** Ionospheric $NmF2$ and $hmF2$ variations during the eclipse day (black continuous line) and the
489 control day (dash blue line). The three vertical lines represents the different phases of the eclipse (S - start
490 time of the initial phase, M - the period of the maximum magnitude of the eclipse, and E - the end time of
491 the recovery phase or the last contact of the eclipse progression). The local time of the respective eclipse
492 contact points for each station are given in table 1.

493

494 **Figure 3:** The local time variation of the ionospheric scale height and the bottomside ($B0$ and $B1$). The other
495 features are the same as in Fig. 1.

496

497 **Figure 4:** Variation of the deviation percentage of $NmF2$ ($\delta NmF2$) and $hmF2$ ($\delta hmF2$) magnitudes for
498 observing the changes in the behaviour of the thermospheric composition and wind flow related to the loss
499 rate during the eclipse phase. The three vertical dashed lines marked the eclipse start time, the time of
500 maximum obscuration and the last contact time of the eclipse (i.e. eclipse phase). Table 1 highlights the
501 local time contact point of the eclipse corresponding the international standard time (IST) eclipse
502 progression. The direction of wind was identify using the $\delta NmF2$ colour legend, the negative values
503 represents the westward wind direction and the positive values is for the eastward wind.

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505 **Figure 5:** Linear relationship of H versus $hmF2$ and H versus $B0$ during the eclipse of 21 August 2017
506 progression phase.

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508 **Figure 6:** Example of the ionospheric profile at the eclipse window of Idaho National Lab showing the
509 bottomside profile (continuous line) and the modelled topside profile shown as a dashed line. The
510 maximum point of the continuous line is the point in which the peak value of the measured $foF2$ and $hmF2$
511 are obtained. The respective measured values $foF2$, $hmF2$ and the corresponding $B1$, $B0$, and H_m
512 parameters values are provided in the plot. The black curve represents the profile for the eclipse day
513 (August 21) and the red curve is for the one of the selected reference days, August 27. On the left side, was
514 the profile during the first contact of the eclipse, the middle and the right profiles are for the maximum
contact and the last contact of the eclipse respectively.

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534 **Table 1:** List of ionosonde station, geographic coordinate, eclipse progression time (Universal time/ Local
535 time) and percentage of maximum obscuration.

Station	GLat	GLong	Eclipse Start time (UT)/(LT)	Eclipse Max Time (UT)/(LT)	Eclipse End Time (UT)/(LT)	% of max obscuration	UT to LT difference
IDAHO NATIONAL LAB	43.81	247.32	16:14:15/ 08:43:31	17:32:37/ 10:01:53	18:56:30/ 11:25:46	100	16:29:17
BOULDER	40	254.7	16:22:33/ 09:21:21	17:46:10/ 10:44:58	19:13:46/ 12:12:34	93.37	16:58:48
EGLIN AFB	30.5	273.5	17:04:41/ 11:18:29	18:37:08/ 12:50:56	20:03:48/ 14:17:36	83.322	18:13:48
AUSTIN	30.4	262.3	16:40:45/ 10:09:55	18:10:10/ 11:39:20	19:39:35/ 13:08:45	65.93	17:29:10
POINT ARGUELLO	34.8	239.5	16:02:39/ 08:00:15	17:16:55/ 09:14:31	18:39:36/ 10:37:12	64.608	15:57:36
MILLSTONE HILL	42.6	288.5	17:27:28/ 12:41:16	18:45:53/ 13:59:41	19:58:38/ 15:12:26	62.533	19:13:48

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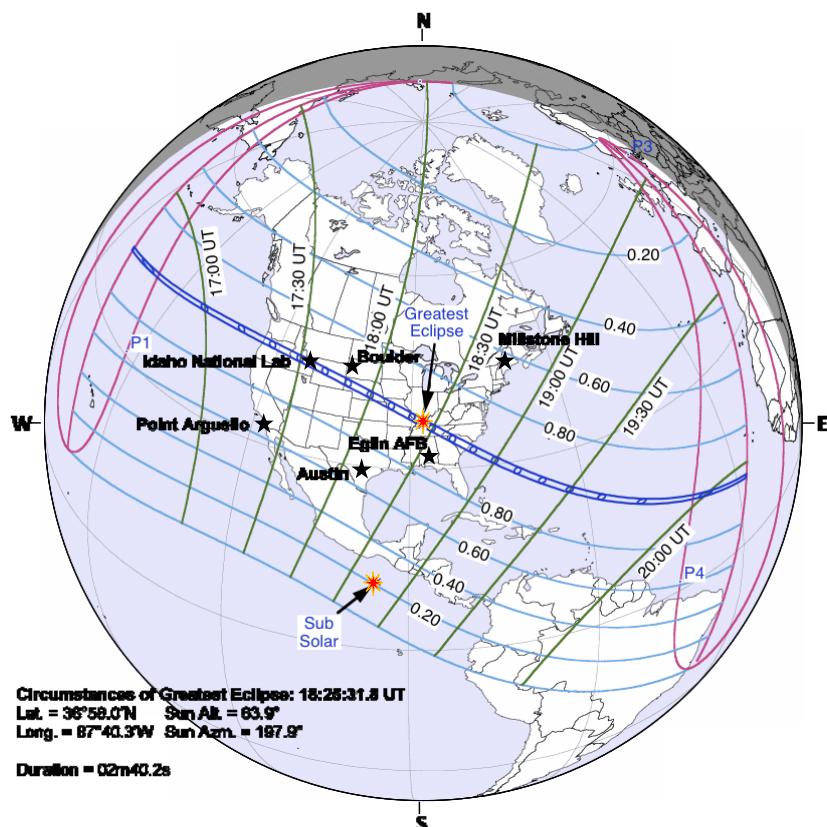
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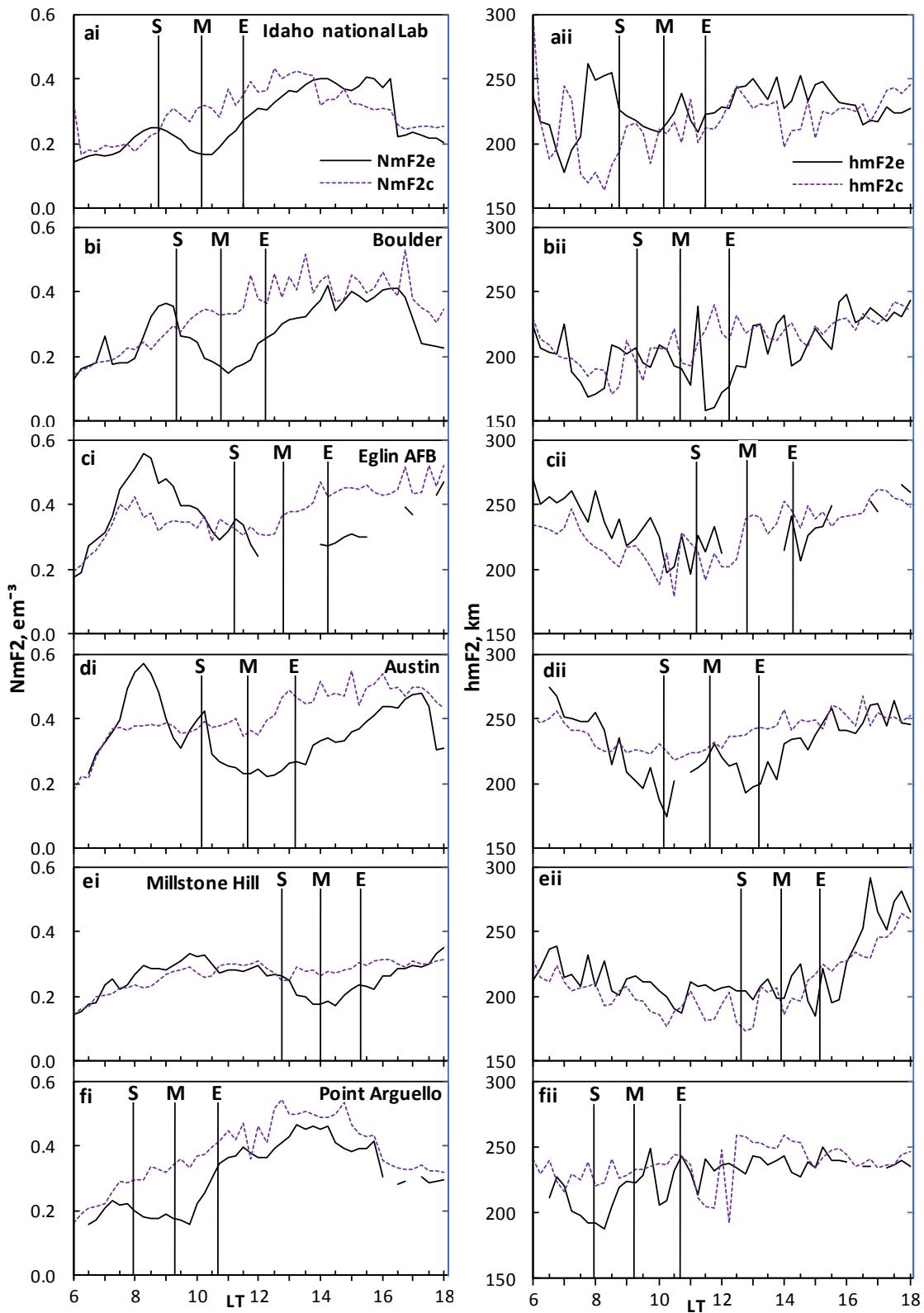


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553 **Figure 1:** The orthographic map showing the coverage area and circumstances of the solar eclipse, and the
554 observatory stations of the total solar eclipse event of August 21, 2017 . The thick blue line region of
555 represents the path of the maximum magnitude of the eclipse and the pale blue lines mark the region of
556 where the partial eclipse is experienced, with the magnitude of partiality.

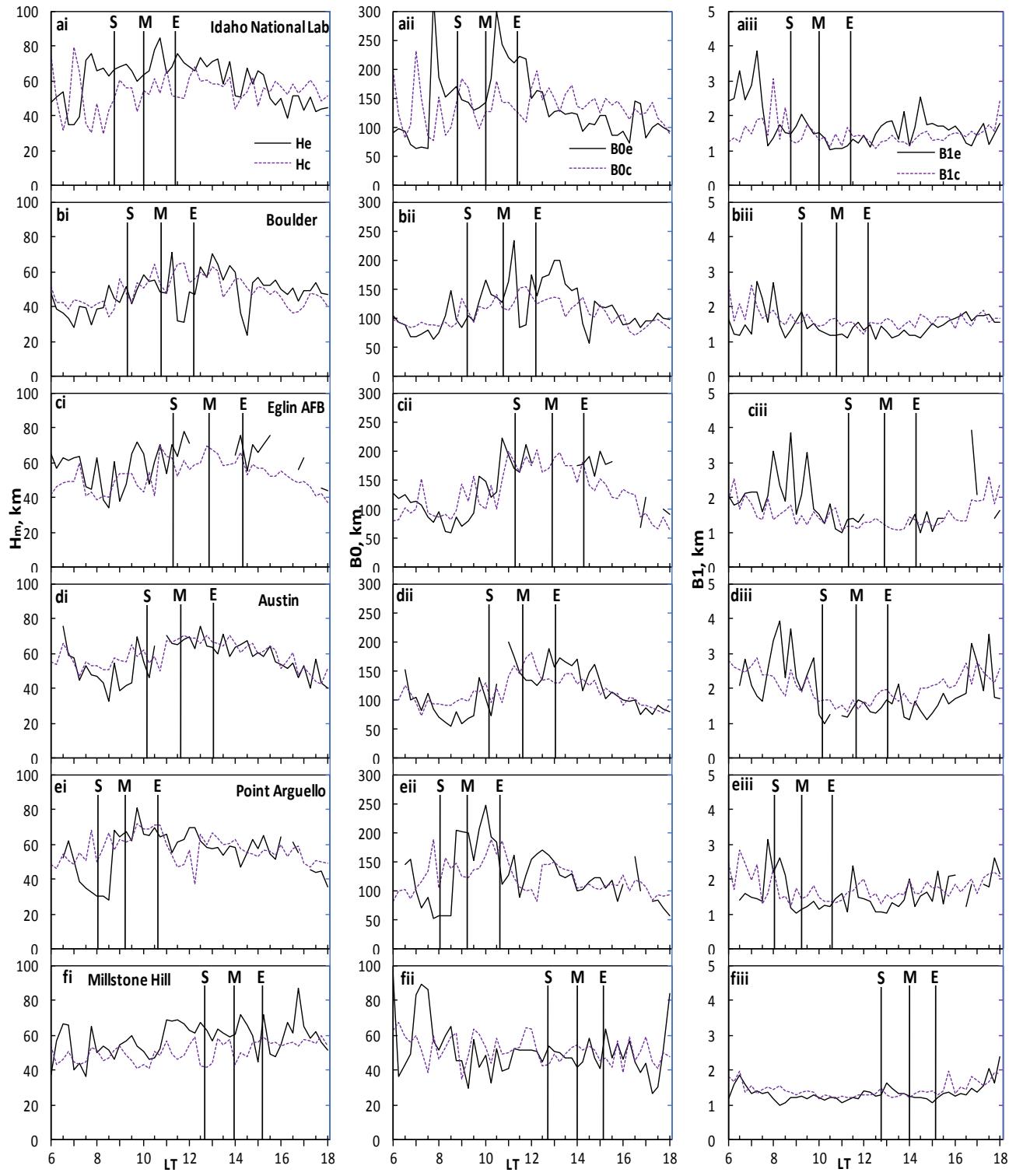
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560 **Figure 2:** Ionospheric $NmF2$ and $hmF2$ variations during the eclipse day (black continuous line) and the
 561 control day (dash blue line) was presented to delineate effect of solar eclipse of August 21, 2017 on the
 562 ionosphere. The three vertical lines represents the different phases of the eclipse (S - start time of the
 563 initial phase, M - the period of the maximum magnitude of the eclipse, and E - the end time of the recovery
 564 phase or the last contact of the eclipse progression). The local time of the respective eclipse contact points
 565 for each station are given in table 1.

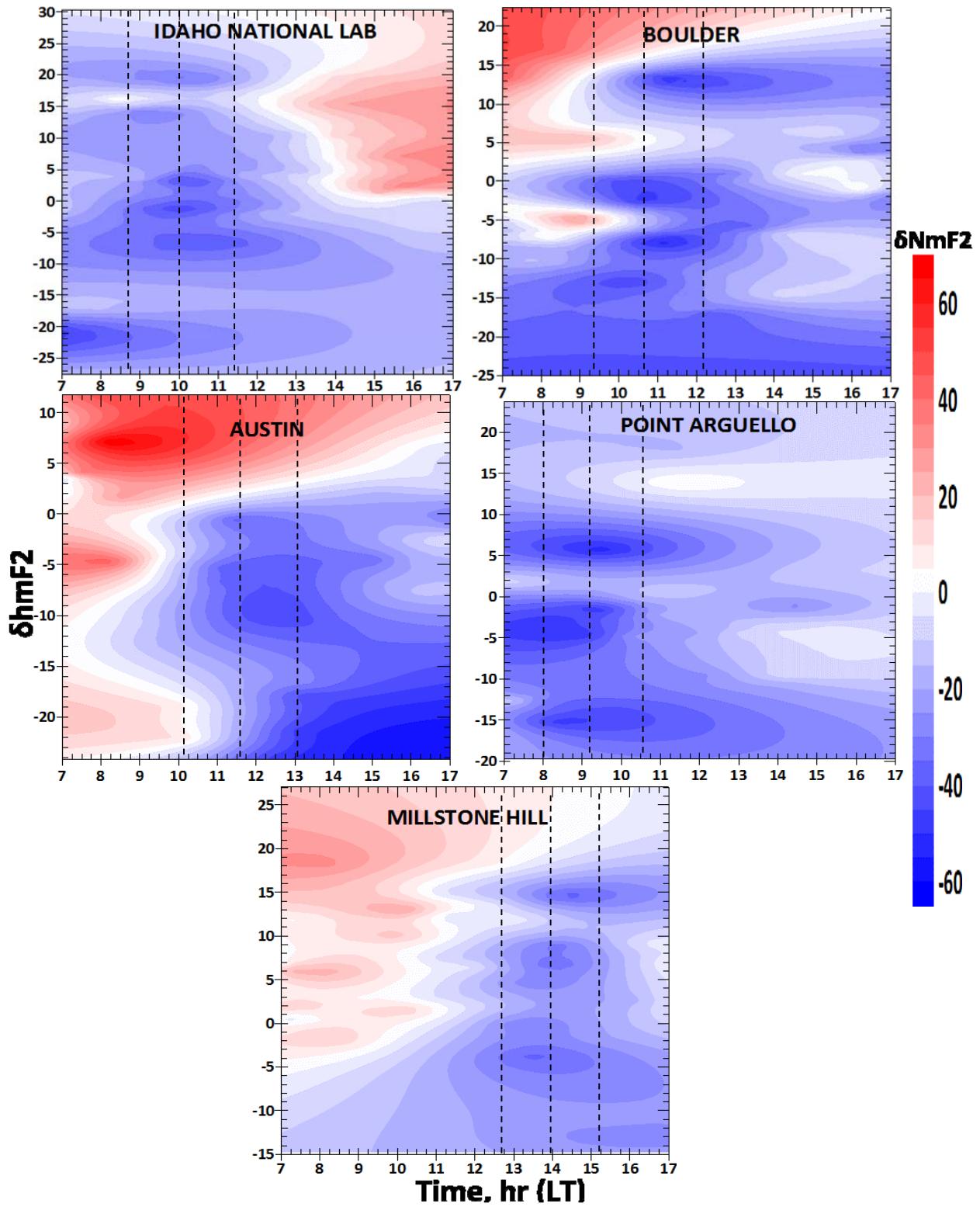


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567 **Figure 3:** The local time variation of the ionospheric scale height and the bottomside (BO and $B1$). The other
568 features are the same as in Fig. 1.

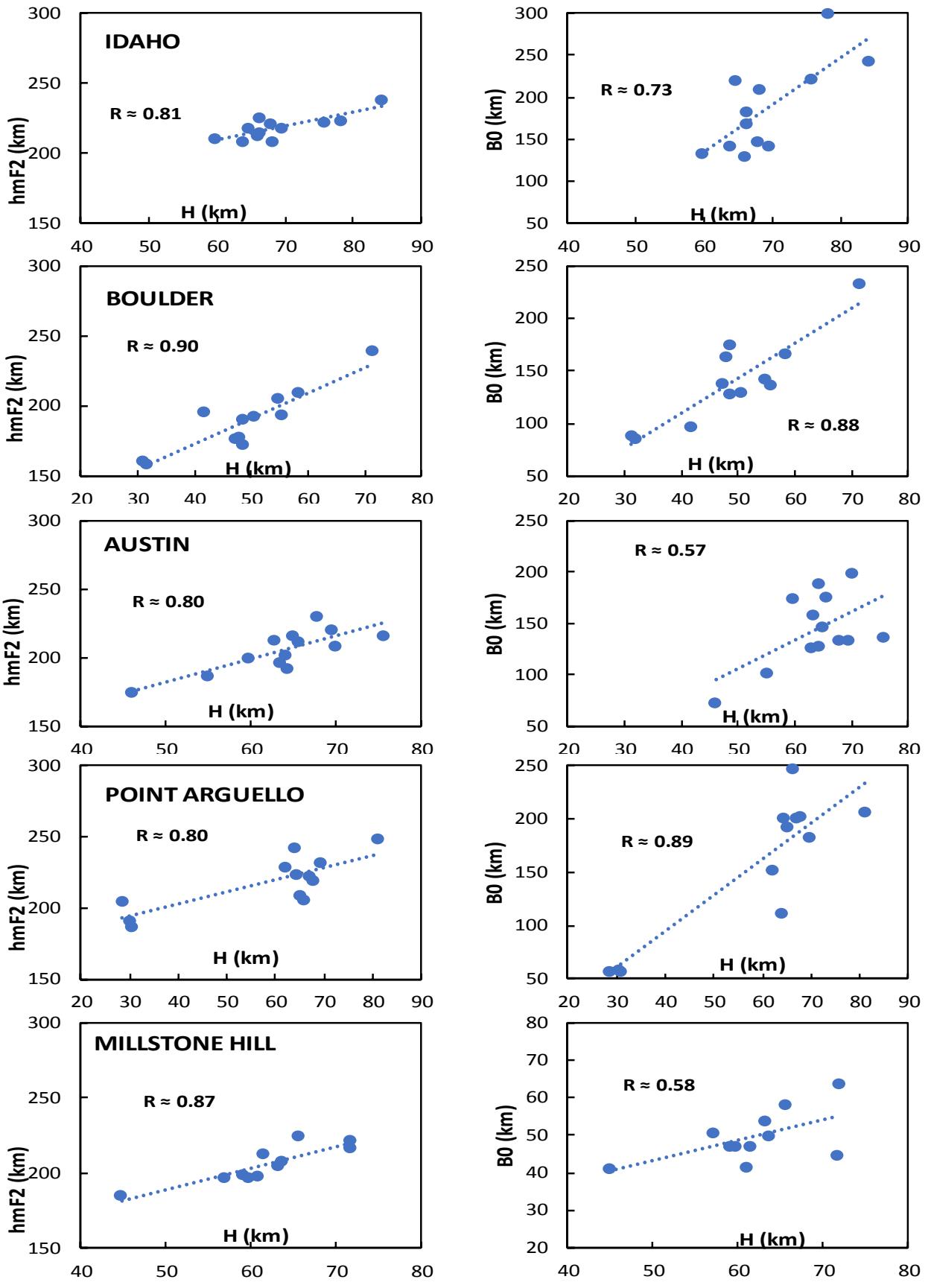
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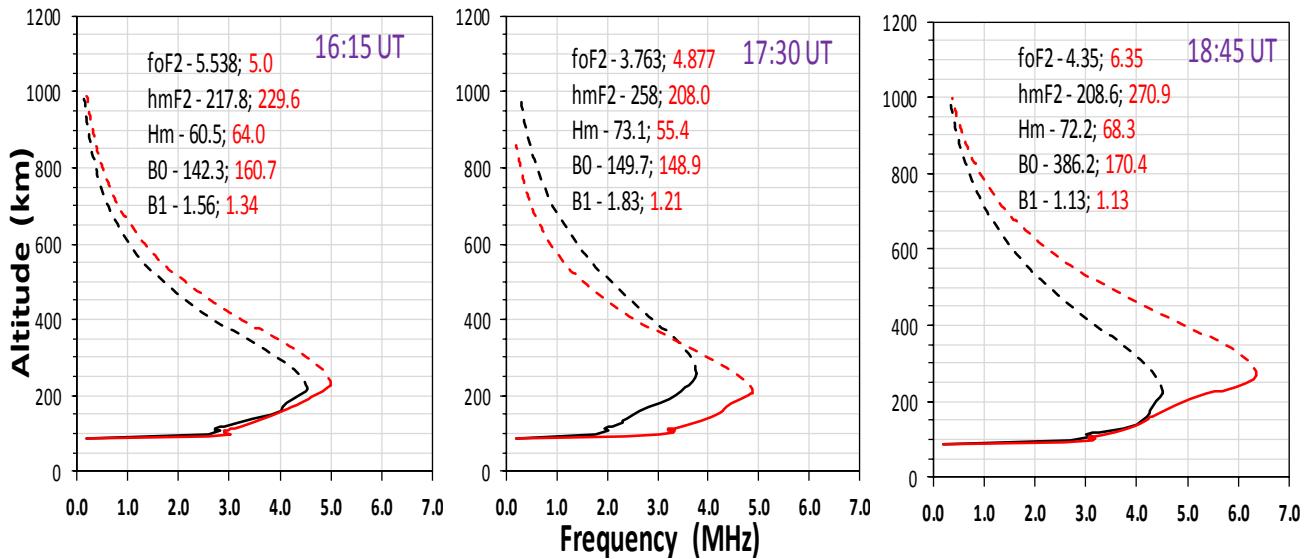


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 573 observing the changes in the behaviour of the thermospheric composition and wind flow related to the loss
 574 rate during the eclipse phase. The three vertical dashed lines marked the eclipse start time, the time of
 575 maximum obscuration and the last contact time of the eclipse (i.e. eclipse phase). Table 1 highlights the
 576 local time contact point of the eclipse corresponding the international standard time (IST) eclipse
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581 progression phase.



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