SOLAR ECLIPSE-INDUCED PERTURBATIONS AT MID-LATITUDE DURING
THE 21 AUGUST 2017 EVENT

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Abstract

A study of the response of some ionospheric parameters and their relationship in describing the behaviour
of ionospheric mechanisms during the solar eclipse of 21 August 2017 is presented. Mid-latitude stations
located along the eclipse path and with data availability on the Global Ionospheric radio Observatory (GIRO)
database were selected. The percentage of obscuration at these stations range between 63-100%.
Decrease in electron density during the eclipse is attributed to reduction in solar radiation and natural gas
heating. The maximum magnitude of the eclipse coincided with $h_{\text{mF}2}$ increase and with a lagged maximum
decrease in $N_{\text{mF}2}$ consistently at the stations investigated. The results revealed that the horizontal neutral
wind flow is as a consequence of the changes in the thermospheric and diffusion processes. The unusual
increase/decrease in the shape/thickness parameters during the eclipse period relative to the control days
points to the perturbation caused by the solar eclipse. The relationship of the bottomside ionosphere and
the F2 layer parameters with respect to the scale height are shown in the present work as viable
parameters for probing the topside ionosphere during eclipse. Furthermore, this study shows that in
addition to traditional ways of analysing the thermospheric composition and neutral wind flow, proper
relation of standardized $N_{\text{mF}2}$ and $h_{\text{mF}2}$ can be conveniently used to describe the mechanisms.

Keywords: solar eclipse; solar radiation; bottomside profile parameters; $N_{\text{mF}2}$ and $h_{\text{mF}2}$; Topside
ionosphere; GIRO database.

1 Introduction

Solar eclipse provides opportunity to study the causes of drastic changes in the atmosphere arising from
reduction in solar radiation and plasma flux. The atmosphere responded to these changes by modifying the
electrodynamic processes and ionization supply of its species to the nighttime-like characteristics during
the daytime. Different physical mechanisms (e.g. neutral wind, thermospheric composition, diffusion
process etc.) that explain the distribution of plasma at the different ionospheric layers are well established.
However, these mechanisms do compete with themselves in explaining the ionosphere, especially the topside ionosphere (see Gulyaeva, 2011).

At mid-latitudes, the effect of diffusion processes and its relationship with the thermospheric compositions has been extensively studied during episodes of solar eclipse (Muller-Wodarg et al., 1998; Jakowski et al., 2008; Le et al., 2009; Wang et al., 2010; Chuo, 2013). At equatorial and low-latitude regions, the $E \times B$ plasma drift had been used to explain the circumstances of solar eclipse on transport processes (Adeniyi et al., 2007; Adekoya et al., 2015). Recently, attention has been drawn to the study of the topside ionosphere during an eclipse for improved prediction and modelling (Huba and Drob, 2017; Chrniak and Zakharenkova, 2018). Reinisch et al., (2018) compared the modelled and measured studies of electron densities at the altitude range of about 150 - 400 km during the eclipse. They found that at lower altitude (at about 150 km) the modelled and the measured agreed well to the changes in the altitude profile of electron density compared to at higher altitudes. The authors however posited that it would be improved if the model $NmF2$ peak falls more slowly to better match the data. Consequently, the present study investigates the effects of solar eclipse of August 21, 2017 on the constituents of the ionosphere at mid-latitudes using some ionosonde data (bottomside parameters, scale height ($H$) estimated from the fitted α-Chapman layer) which have not been given much attention in previous works especially in analysing solar eclipse effect. Using these parameters to analyse the circumstances of solar eclipse at the topside ionosphere and its plasma distribution mechanisms make this paper significantly different from previous studies. This, we intend to achieve by analysing the ionospheric parameters that controls the distribution of plasma at the topside and bottomside layers of the F2 region. To shed light on these analysis, section 2 highlights the data source, methodology, and path of the eclipse. The result and discussion were presented in section 3, while section 4 presents the summary and concluding remark of the result.

2 Data source, methodology, and the path of the eclipse

With regards to the eclipse of 21 August 2017, the totality of the eclipse is visible from within a narrow corridor that traverses the United States of America. However, in the surrounding areas, which include all of mainland United States and Canada, the eclipse was partial. More details of its path can be seen via NASA – Total solar eclipse of 2017 August 21 (https://eclipse.gsfc.nasa.gov/). From the footprint of the Moon’s shadow as seen from some locations, the eclipse started from around 08:00 LT and ended around 14:30 LT (not shown). The details on the local circumstances of the eclipse, the time of the first, mid and last contact of the eclipse over the ionosphere of the investigated stations were highlighted in table 1. More details on the total solar eclipse event and its partiality, the circumstances surrounding its progression and its magnitude of obscuration can be obtained through the link...
http://xjubier.free.fr/en/index_en.html. The ionospheric parameters data used for this study for the selected mid-latitude stations were obtained from the Global Ionospheric Radio Observatory (GIRO) networks (Reinisch and Galkin 2011) and manually validated. The parameters include the maximum electron density of the F2-layer (\( NmF2, m^{-3} \)), and its height (\( hmF2, km \)), the shape parameter (\( B1 \)), the thickness parameter (\( B0 \)), and the Chapman scale height (\( H \)) of the F2 layer. The path of the eclipse informed the choice of stations.

\( NmF2 \) values for both the eclipse and control days were obtained from their corresponding critical frequencies (\( foF2 \)) using the expression: \( NmF2 = ((foF2)^2 / 80.5) e/m^3 \). The control day value is the average value of the two days before/after the eclipse day (i.e. 6, 12, 24 and 27). These reference days were chosen such that they have similar geomagnetic, interplanetary and solar properties with the eclipse day. The daily average value of control days and eclipse day interplanetary index (Ap and \( \Sigma Kp \)), and solar flux unit index (F10.7) ranges from 8 – 12 nT for Ap, 20 – 27 nT for \( \Sigma Kp \) and 75.6 – 89.1 sfu (1 solar flux unit (sfu) = 10⁻²² Wm⁻² Hz⁻¹) for F10.7, indicating that geomagnetic and solar activities of these days is unsettled (see Adekoya et al., 2015 for classification of geomagnetic activity). The calculated daily average of summation Kp, Ap and solar flux indices was obtained from the National Space Science data Centres (NSSDC’s) OMNI database https://omniweb.gsfc.nasa.gov/. The typical behaviour of the \( NmF2 \) and \( hmF2 \) on the eclipse day (i.e. \( NmF2e \) and \( hmF2e \)) was compared with that of the control day (\( NmF2c \) and \( hmF2c \)) to observe the changes brought by the short period of loss of photoionization in the ionosphere. This will measure the direct consequence of the solar radiation disruption (due to the eclipse) on the ionospheric chemical, transport and thermal processes in the F2 layer. The ionized layer depends majorly on three parameters, viz: \( NmF2 \), \( hmF2 \), and the ionospheric scale height (\( H \)). The H describes the constituents of the ionospheric plasma, which decreases with increasing altitude. It is estimated from the fitted \( \alpha \)-Chapman layer with a variable scale height, \( H(h) \), to the measured bottomside profile \( N(h) \), which then determined as the Chapman scale height at \( hmF2 \) (i.e. \( H(hmF2) = H \)) (Huang and Renisch 2001; Reinisch and Huang 2001). Together with the information of \( NmF2 \) and \( hmF2 \), the topside profile can be best represented, which is assumed to follow the \( \alpha \)-Chapman function (Huang and Reinisch 2001). Also, H provides a linkage between the bottomside ionosphere and the topside profiles of the F region (Liu et al., 2007).

However, Xu et al. (2013) and Gulyaeva (2011) related ionospheric F2 - layer scale height, \( H \) to the topside base scale height, \( Hsc \), given by \( Hsc = hsc-hmF2 = 3 \times H \). Where \( hsc \) is the height at which the electron density of the F2-layer falls by a factor of an exponent, at an upper limit of 400 km altitude (i.e. \( NmF2/e \)) (see Xu et al., 2013). That is, the region where electron density profile gradient is relatively low. Gulyaeva (2011) showed theoretically that \( Hsc \) increase over Hm by a factor of approximately three (3) and is a
consequence of the $Ne/NmF2$ ratio ($Ne$ – plasma density), which corresponds to $H$ in the Chapman layer. At altitudes very close to $hmF2$, the ratio equals 0.832, while it is 0.368 at altitudes beyond the $hmF2$. Therefore, we adopted the definition of Gulyaeva (2011) for the topside base scale height as the region of the ionosphere between the F2-peak and 400 km altitude. Summarily, the topside based scale height ionosphere here is defined as the region between the F2 peak and $hsc$ or $3H$. It is thus evident that $H$ is a key and essential parameter in the continuity equation for deriving the production rate at different altitudes, a pointer to the F2 topside electron profiler, as well as a good parameter for evaluating the transport term (Yonezawa, 1966; Huang and Reinisch, 2001; Reinisch and Huang, 2001; Belehaki et al., 2006; Reinisch et al., 2004). Consequently, the parameter $H$ can be used as a proxy for observation relating to the topmost side electron density profile. Furthermore, the division of the topsides and the bottomside ionosphere may be related to the difference in the effective physical mechanisms in the regions. Hence, the bottomside parameters $B1$ and $B0$ of the ionosphere, as presented in this work, helped in examining the perturbation of solar eclipse in the bottomside ionospheric F2 layer.

3 Result and Discussion
This section presents the temporal evolution of the maximum electron density ($NmF2$), and its corresponding height ($hmF2$) over the ionosphere at the selected mid-latitude stations along the path of solar eclipse of 21 August 2017. The control day variation relative to the eclipse day is also presented. Figure 1 presents the variation of maximum electron density and the corresponding peak height, during both the eclipse and control days. Figure 2 depicts the variation of scale height and the bottomside parameters ($B0$ and $B1$) due to the eclipse by superposing plots for both the eclipse and control days. Analysis of these parameters during an eclipse event may help in the modelling of the ionospheric profiles (the topsides and bottomside electron density distribution profile) during the short nighttime-like period of the day. Figure 1a presents the $NmF2$ and $hmF2$ variations during the eclipse event and the control day over Austin; having an obscuration magnitude of 65.93% around the daytime period. The effect of the disruption of solar radiation was evident as the $NmF2$ started decreasing at the first contact of the eclipse in Fig. 1ai. The start time or first contact, the maximum magnitude period and the end time or the last contact of the eclipse are marked with the vertical lines S, M and E respectively. The decrement in $NmF2$ during the eclipse phase was due to reduction in the ionization. This reduction caused changes in the photochemical and transport process of the atmosphere during the daytime, thus exhibiting nighttime characteristics. It should be noted that the maximum decrease in $NmF2$ did not coincide with the maximum magnitude of the eclipse obscuration, rather with a time lag of few minutes. This lag period fell within the relaxation period over Austin ionosphere, with $NmF2$ and $hmF2$ simultaneously attaining their peak
magnitudes. Hence, the ionosphere returned to its pre-eclipse state. Contrary to the decrease in the \( NmF2 \) amplitude, the \( hmF2 \) increased at the total obscuration of the eclipse window.

The ionosphere over Eglin AFB, Boulder, Point Arguello, Millstone Hill and Idaho National Lab, did not show any contrary variation to that observed at Austin during the eclipse event. The decrease and increase in \( NmF2 \) and \( hmF2 \) after the maximum magnitude was simultaneous. The only exception was that the local time at which each station observed the effects were different. Their obscuration percentage ranged from 62.5 – 100%. This did not cause any significant change in the way they responded to the reduction in solar heating. The ionosphere over Idaho National Lab experienced the totality of the eclipse with 100 % magnitude, the \( hmF2 \) was observed to increase few minutes before the maximum magnitude of the obscuration. However, other stations responded differently, their \( hmF2 \) peak enhancement was observed after the maximum obscuration. All these observations may be linked with the fact that the level of minimum rate of electron production does not necessarily coincide with peak electron density of the molecular gases formed. This is because the electron concentration depends on the loss rate by dissociative recombination too.

At mid-latitudes, the ionospheric F2 plasma distribution is controlled by diffusion processes (Rishbeth 1968). There are two basic mechanisms that define the diffusion process during an eclipse: First is the coolness brought by the partial removal of photoionization (Müller-Wodarg et al., 1998), which is believed to be the originator of the downward diffusion process, and the atmospheric expansion due to the gradual increase in the temperature after the totality. The downward diffusion process was related to the increase in the molecular gas (\( N_2 \)) concentration during the cooling process. However, the aftermath of the coolness was related to the upward diffusion process. These mechanisms were proxy to the electron density distribution during the eclipse window. Our analysis suggests that the observed decrease in \( NmF2 \) is due to the downward diffusion flux of the plasma while the increase that followed is by upward diffusion (e.g. Le et al., 2009; Adekoya and Chukwuma 2016). Several works on eclipse (Müller-Wodarg et al., 1998; Grigorenko et al., 2008; Adekoya and Chukwuma 2016; Hoque et al., 2016) have shown that it was not just the electron density that is being affected during an eclipse window, but the thermospheric wind as well, since the thermospheric wind emanating from the ratio of gas species is related to the variation in electron density. It has been observed that the increase in the mean molecular gas of thermospheric composition decreases the electron density and vice versa. Le et al. (2010) related the trough of electron density distribution during the eclipse phases to the contraction/compression and expansion of the atmosphere brought by the decrease and increase in temperature; leading to the downward drift of the plasma during the eclipse window. Chukwuma and Adekoya (2016) attributed the decrease in the electron temperature to
the downward vertical transport process and the decrease in the cooling process to the upward vertical transport process.

Figure 2 describes H, $B_1$ and $B_0$ in three columns respectively for all six stations. It was observed from the plots that the minimum decrease in $NmF_2$ amplitude corresponds to increase in H at all stations; implying the upward lifting of the topside electron to the region of higher altitude at the eclipse window. Hence, the scale height variation highlights the decrease in electron production and the vertical distance through which the pressure gradient falls at the topside during the eclipse activity. The observation illustrates the mutual relationship between the $NmF_2$ and H, which may aid in extrapolating the topside ionospheric profile (Gulyaeva, 2011). In essence, scale height changes observed during the eclipse window can be used to explain the pressure gradient, electron density distribution and transport processes. In this sense, the diffusion coefficients are expressed as ratio of determinants (determinant here refers to the concentration of species ([O] and [N$_2$]), with the size of the determinants depending upon both the number of species in the gas mixture and the level of approximation. Therefore, the increase (decrease) in the scale height can be used as a proxy for downward (upward) diffusion process at the topside ionosphere. Consequently, the thermospheric wind, which causes plasma distribution in the topside ionosphere, is induced by solar radiation. Moreover, the significant changes observed in the scale height variation during the eclipse window also indicated that transport processes are affected as they are temperature dependent. Therefore, changes in the thermospheric compositions due to the solar eclipse at the topside layer will affect the density profiles of the ionosphere (Müller-Wodarg et al., 1998).

It is noteworthy that the increase (decrease) in the scale height decreases (increases) the electron density during the eclipse window. The sensitivity of electron density to temperature at the topside directly affects the electron density profile (e.g. Wang et al., 2010); as cooling due to decrease in temperature results in decrease in the electron density via reduced ionization. This indicates that the decrease (increase) in electron temperature at the topside ionosphere causes the increase (decrease) in the scale height, which is related to the diffusion and transport processes and subsequently affect the pressure gradient of the plasma. From plots of H (fig. 2) and $NmF_2$ (fig. 1), it was observed that the minimum decrease in $NmF_2$ corresponded with peak increase in scale height. This imply that the topside ionosphere is more sensitive (than the bottomside) to any change in the solar radiation. Thus, the pressure gradients can be analysed in terms of either the scale height or electron density.

From column 2 and 3 of Figure 2, we observed that the measured shape ($B_1$) and thickness ($B_0$) parameters of the ionosphere over these stations exhibit significant variations during the eclipse event. $B_1$ responded
with a decrease at the first contact of the eclipse compared to the control day. This behaviour differs from
that of the B0 observation. B0 parameter from the first contact increases and reached the maximum peak
few minutes after the maximum obscuration magnitude, which coincided with the minimum decrease in
B0. Generally, the pattern of the day to day variation of the bottomside parameters is the average
morphology, but the increase in the B0 and the decrease in the B1 parameters during the eclipse period
compared to the control day was a notable one and can be related to the perturbation caused by the solar
eclipse. During the eclipse, the solar radiation was lost; trapped atomic ions O⁺ was converted into
molecular ion (NO⁺ and O₂⁺) by charge transfer, owing to the sufficient concentration of molecular gasses
(N₂ and O₂) (Rishbeth, 1988). The height of the ionospheric slab indeed increased with reduced width,
which is attributable to compression due to loss of solar heating.

The behaviour of the ionosphere can be explained during solar eclipse with any of the components that
constitute the topside and the bottomside ionosphere and can be looked at, from the angle of the
percentage of concentration of the components. In this regard, the percentage of deviations of NmF2
(DNmF2) and hmF2 (DhmF2) during the eclipse day away from the control day were plotted in Figure 3. This
is done to describe the contribution of the thermospheric wind and compositions. Although observing the
variation of NmF2 and hmF2 alone can be used for observing the changes in the behaviour of the
thermospheric compositions and wind flow, if properly analysed, but it is more convenient to describe
these mechanisms by standardizing the original variables used during the event. The normalization effort
(with the use of DNmF2 and DhmF2) presents the original variation of NmF2 and hmF2 onto directions
which maximize the variance. Consequently, the result can be used for analyses of any mechanisms that
drive the ionospheric plasma, if properly related.

The percentage deviation in Figure 3 was defined as the ratio of ((NmF2e – NmF2c)/NmF2c) x 100. The
same relation is defined for the hmF2 parameter. As earlier pointed out, during eclipse period, neutral
composition becomes the dominant chemical process arising from diffusion activities. The increase in the
neutral composition leads to the increase in the molecular gas concentration and compete with diffusion
process. Hence the percentage deviation in Fig. 3 discusses the neutral composition changes and delineate
how these changes may affect the electron densities as well as its profiles in the atmosphere during the
eclipse. The respective maximum and minimum peak response of the percentage deviation is attributed to
the enhancement and depletion of DNmF2. One sees from the plots that the percentage deviation started
increasing at the first contact of the eclipse (the dash vertical line) and reached the maximum, appearing
few minutes after the maximum magnitude of the eclipse was obtained. This behaviour is similar to the
conditions of the neutral compositions during the eclipse event reported by Muller-Wodarg et al. (1998).
Another important process observed in this study is the neutral wind flow effect. To identify the direction of the wind, the DNmF2 colour legend in the contour plots was used in Figure 3. The negative values represent a westward wind contribution and the positive values is for the eastward wind. Looking at the marked eclipse region in the figure, it would be seen that the DNmF2 started decreasing from the first contact of the eclipse and maximized few minutes after the totality mark and started increasing again. It has been established that at daytime, the peak height of the plasma will be reduced due to lost in recombination; but at nighttime, equatorward neutral wind drives the F2-layer plasma to higher altitudes where ion loss rate is slower. The behaviour of the F2 plasma during solar eclipse cannot be completely related to the nighttime period due to the fact that all the processes controlling the nighttime variation are not completely actualised but can be related to partial nighttime/sunset period (see Adekoya et al., 2015). Thus, the slight increase in the peak height and equatorward neutral wind flow is the driver during the solar eclipse phase. The neutral wind acts jointly with the plasma flows from the topside ionosphere, resulting in F2 region plasma density variation. Therefore, the westward/eastward neutral wind flow was related to the depletion/enhancement in the deviation, which was clearly shown in the marked eclipse region of the figure. The plot in Figure 3 had established the ionospheric dynamics of diffusion processes, neutral compositions and the flow of neutral wind caused by the eclipse perturbation, which can invariably reduce the effectiveness and reliability of radio wave propagation.

Relative to the mutual relationship between the topside and bottomside ionosphere, we considered the linear correlation coefficient \( R \) of H versus \( hmf2 \) and H versus \( B0 \) during the eclipse window, Figure 4. \( R \) ranges from (0.52-0.92) for H/\( hmf2 \) relationship, and 0.37-0.92 for the H/\( B0 \) connection. This good linear agreement revealed the dependence of \( hmf2 \) and \( B0 \) on the scale height. The only exception where low correlation was observed was at Idaho (0.47) and Millstone (0.37) with respect to the H versus \( B0 \) relationship. Apart from revealing the dependence between the parameters, the relationship may also provide a convenient way for modelling the topside profile from the knowledge of the bottomside parameter, \( B0 \), during the eclipse period. Also, the strong correlation between \( hmf2 \) and H indicates that there may be some inter-related physical mechanisms controlling the behaviour of the plasma at the topside ionosphere. That is \( hmf2 \) strongly depend on the neutral wind flow and explain the state of thermospheric composition (Liu et al., 2006; Fisher et al., 2015). Since all these parameters competes during the eclipse, one can argue that with the accessibility of one, in place of the other (as a consequence of their relationship), the prediction and modelling of the ionosphere can be conveniently achieved.
4 Conclusions

This paper presents the induced perturbation of solar eclipse of 21 August 2017 on the ionospheric F parameters and how they describe the mechanisms of the ionosphere at mid-latitude. The perturbation effects and dynamics during a solar eclipse episode using ionospheric F2 parameters ($N_{mF2}$ and $h_{mF2}$), the bottomside profile thickness ($B0$) and shape ($B1$) parameters of electron density and the plasma scale height ($H$), which are not often used for eclipse study, were investigated. These parameters represent the state of the F-region ionosphere. The changes observed during the eclipse phase is related to the reduction in solar radiation and natural gas heating. The $N_{mF2}$ minimum was attained at ~30 minutes after the totality of the eclipse when it decreases to about 65% of its control day. This decrease in $N_{mF2}$ was uplifted to the higher altitude compared to the non-eclipse day. The thickness and shape parameters which are often limited to the bottomside F-region were seen as viable parameters for probing the topside ionosphere, relative to the scale height during the eclipse. Hence their relationship in describe one another is established. Implication is that eclipse-caused perturbation could have been better explained using some ionosonde parameters. The changes in the neutral wind flow, thermospheric compositions and diffusion processes found their explanation in the behaviour of the F region plasma during eclipse. In addition, it can be concluded that the behaviour of $DN_{mF2}$ and $Dh_{mF2}$ during eclipse can be conveniently used to describe the mechanisms of thermospheric composition and wind flow.

Acknowledgements

We acknowledge use of global ionospheric Radio Observatory data provided by ULMCAR (http://ulmar.uml.edu/DIDBase/) and the World Data Center for Geomagnetism, Kyoto (http://wdc.kugi.kyoto-u.ac.jp/index.html) for geomagnetic activity data. We thank the management team of the national Aeronatics and Space Administration (NASA) service (http://eclipse.gsfc.nasa.gov) and http://xjubier.free.fr/en/site_pages/SolarEclipseCalc_Diagram.html for progression and eclipse local circumstances information. The authors thank Professor Ljiljana R, Cander and the anonymous reviewer for their constructive corrections and suggestions that tremendously improved the structure and quality of the paper.

References


Table Caption

Table 1: List of ionosonde station, geographic coordinate, eclipse progression time and percentage of maximum obscuration.

Figure Captions

Figure 1: Ionospheric $Nmf2$ and $hmF2$ variations during the eclipse day (black continuous line) and the control day (dash blue line). The three vertical lines represents the different phases of the eclipse (S - start time of the initial phase, M - the period of the maximum magnitude of the eclipse, and E - the end time of the recovery phase or the last contact of the eclipse progression).

Figure 2: The local time variation of the ionospheric scale height and the bottomside ($B0$ and $B1$). The other features are the same as in Fig. 1.

Figure 3: Variation of the percentage deviation of $NmF2$ ($DNmf2$) and $hmF2$ ($DhmF2$) magnitudes for observing the changes in the behaviour of the thermospheric composition and wind flow related to the loss rate during the eclipse phase in comparing to the period before/after the event. The three vertical dashed lines marked the eclipse start time, the time of maximum obscuration and the last contact time of the eclipse (i.e. eclipse phase). The direction of wind was identify using the $DNmf2$ colour legend, the negative values represents the westward wind direction and the positive values is for the eastward wind.
Table 1: List of ionosonde station, geographic coordinate, eclipse progression time and percentage of maximum obscuration.

<table>
<thead>
<tr>
<th>Station</th>
<th>GLat</th>
<th>GLong</th>
<th>Eclipse Start time (UT)</th>
<th>Eclipse Max Time (UT)</th>
<th>Eclipse End Time (UT)</th>
<th>% of max obscuration</th>
<th>UT to LT difference</th>
</tr>
</thead>
<tbody>
<tr>
<td>AUSTIN</td>
<td>30.4</td>
<td>262.3</td>
<td>16:40:45.1</td>
<td>18:10:10.3</td>
<td>19:39:35.0</td>
<td>65.93</td>
<td>17:29.2</td>
</tr>
<tr>
<td>EGLIN AFB</td>
<td>30.5</td>
<td>273.5</td>
<td>17:04:41.1</td>
<td>18:37:07.6</td>
<td>20:03:47.7</td>
<td>83.322</td>
<td>18:13.8</td>
</tr>
<tr>
<td>POINT ARGUELLO</td>
<td>34.8</td>
<td>239.5</td>
<td>16:02:38.5</td>
<td>17:16:54.8</td>
<td>18:39:36.0</td>
<td>64.608</td>
<td>15:57.6</td>
</tr>
<tr>
<td>BOULDER</td>
<td>40</td>
<td>254.7</td>
<td>16:22:33.1</td>
<td>17:46:09.6</td>
<td>19:13:45.9</td>
<td>93.37</td>
<td>16:58.8</td>
</tr>
<tr>
<td>MILLSTONE HILL</td>
<td>42.6</td>
<td>288.5</td>
<td>17:27:28.1</td>
<td>18:45:52.5</td>
<td>19:58:38.3</td>
<td>62.533</td>
<td>19:13.8</td>
</tr>
<tr>
<td>IDAHO NATIONAL LAB</td>
<td>43.81</td>
<td>247.32</td>
<td>16:14:15.2</td>
<td>17:32:36.5</td>
<td>18:56:30.1</td>
<td>100</td>
<td>16:29.3</td>
</tr>
</tbody>
</table>
Figure 1: Ionospheric \( NmF2 \) and \( hmF2 \) variations during the eclipse day (black continuous line) and the control day (dash blue line). The three vertical lines represent the different phases of the eclipse (S - start time of the initial phase, M - the period of the maximum magnitude of the eclipse, and E - the end time of the recovery phase or the last contact of the eclipse progression).
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Figure 4: Linear relationship of H versus \( hmF2 \) and H versus \( B0 \) during the eclipse of 21 August 2017 progression phase.