POINT-BY-POINT CORRECTIONS:

<u>General Remarks</u>: We have reviewed the entire document based on the most recent reviewer's comment(s), with the below details. In general, we have reviewed presentation of grammar etc. as shown in the marked up version of the manuscript also attached for your kind perusal.

 Query: The comments refer to the author version on Suplement of AC5 of 18 April. Lines 99-105: The TEC global investigations have been done by many other researchers not only by the three cited in the paper.

Response to query 1: Line 101 – 103: We agree that there are a plethora of authors who had previously examined the subject and not just the 3 previously mentioned. We have shown more previous authors in lines 101-103 ((e.g. Bhuyan and Borah, 2007; Maruyama et al., 100 2004; Jee et al., 2004; Balan et al., 1994; Rama Rao et al., 2006a, b, c; Rama Rao et al., 1985; 101 Bolaji et al. 2012; Wanninger, 1993; Akala et al., 2013; Komjathy et al., 1998; Langley et al., 2002 102 Sunda et al., 2013; Torr and Torr, 1973; Tsai et al., 2001 and references therein). Whereas there may be more authors than the aforementioned, we feel that the enlarged references should be adequate for a paper like ours. If there are specific references that the reviewer will want to see and are not yet included, we will be happy to accommodate such if they are specifically highlighted.

 Query: Line 106: please say in what sector each TEC information was obtained. Rama Rao et al. results refer to Indian sector while Wanninger is in the Asian sector.

Response to query 2: In Lines 106 – 107, we have included the sectors of Rama Rao and Wanninger. In Lines 114-115, we endeavoured to include the sectors wherein those investigations were carried out as suggested by the reviewer. (Studies of Rama Rao et al. (2006a, b) in the Indian sector and Wanninger (1993) in the 106 Asian sector reported.....)

Query: Lines 113-116: Again, the results were obtained at different sectors, you must specify eachone.
 Response to query 3: Lines 114 – 115, 120 of the revised manuscript now has the sector where the investigations were conducted (Bhuyan and Borah (2007) working in the Indian sector and Komjathy et al. (1998) 114 and Lee and Reinisch (2006) while studying in the American sector compared TEC).

4. Query: Lines 206-208: Must clarify. The TEC increase is due to the upward vertical elevation of the equatorial ionosphere during the day, resulting in an additional ionization.

Response to query 4: Lines 210 - 214 now better explains TEC increase because of the **E** × **B** drift at the equator. (Dabas et al. (2003); Somoye et al. (2011); Hajra et al. (2016); D'ujanga et al. (2017) attributed the 208 steep increase in TEC to the upward **E** × **B** vertical plasma drift and the rapid filling up of the 209 magnetic field tube at sunrise resulting to solar EUV ionization. During daytime, an eastward 210 electric field at the equator causes plasma to be lifted to greater heights. This dynamo-generated 211 eastward electric field combined with the northward geomagnetic field lifts the equatorial 212 ionosphere from 700 km to 1000 km, resulting in additional ionization (D'ujanga et al., 2017; 213 Somoye et al., 2011). Suranya et al., 2015 further mentioned that upward vertical **E** × **B** drift could 214 lead to equatorial ionization anomaly (EIA) and meridional winds.)

- 5. Query: Line 209: Use a proper basic reference about the Rayleigh_Taylor Instability. Response to query 5: Basic reference on Rayleigh Taylor instability has been cited.
- 6. Query: Line 304-310: Specify the researchers you referred in line 305. It means that your results are not in agreement with previous results? Rewrite all lines to clarify what you mean.

Response to query 6: The statement in Lines 310 - 312 is now better clarified. In order to be clear, we deleted the lines that may be ambiguous in the explanation.

 Query: Lines 326-328: I understood that your results are in agreement with D'ujanga et al. (2017) results obtained in the African sector and with Bagiya et al. (2009) in the Indian sector. OK? Explain better the correlations in the different sectors.

Response to query 7: Line 334 – 337 now explains the correlation of our result with those of Bagiya et al and D'ujanga et al.

Query: Lines 329-333: Explain better the TEC increases related with the decrease of O/N2 ratio.
 The decrease of O/N2 ratio increases the electron density, and this explains the TEC increase.

Response to query 8: Lines 340 - 341 now explains better TEC increase in relation to O/ N₂ ratio.

9. Query: Lines 338-345: Figure 10 must be better explained. The figure intends to show the comparison between the monthly values of TEC and Rz from 2011 to 2014. What is the OBS-TEC are you using in this figure for each month? I suggest you consider the maximum diurnal value of TEC observed at each month presented in figures 1-4, since you are comparing with solar cycle variation.

Response to query 9: Figure 10: In this figure, we agreed to use the monthly mean TEC values and monthly mean sunspot number to illustrate solar cycle dependency instead of using the highest diurnal value for each month. We trust our approach is less tedious and more appropriate. In addition, the Figure has been better explained, please See lines 351 - 358 and lines 363 - 371.

10. Query: Your Conclusions are just a repetition of what you presented in the results and discussions. Here you must shortly resume the results of TEC variations you obtained at Birnin-Kebbi, and compare with was obtained from previous works near the same region at different periods, and with the results obtained at other sectors. Taking into account the results that also evaluated the OBS-TEC versus the IRI Models.

Response to query 10: We agree with the reviewer's observation and as such, we rewrote the conclusion to accommodate all issues raised by the reviewer.

OTHERS

- 11. The introductory aspect of the abstract have been rewritten to suite the title of the paper.
- 12. The last line of the last paragraph of Figures 9 was rewritten.
- 13. In the cause of review, new references were cited while some where deleted.

DIURNAL, SEASONAL AND SOLAR CYCLE VARIATION OF TOTAL ELECTRON CONTENT AND COMPARISON WITH IRI-2016 MODEL AT BIRNIN-KEBBI

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8 ABSTRACT

9 Ionospheric error is the major error source for the signals of Global Positioning System (GPS) 10 satellites. In the analysis of GPS measurements, ionospheric error is somewhat assumed to be a 11 nuisance. The error induced by the ionosphere which is proportional to the number of electrons 12 along the line of sight (LOS) from the satellite to receiver and can be determined in order to study 13 the diurnal, seasonal, solar cycle and spatial variations of the ionosphere during quiet and disturbed 14 conditions. In this study, we characterize the diurnal, seasonal and solar cycle variation of observed 15 (OBS) total electron content (TEC) and compare the result with the International Reference Ionosphere (IRI-2016) model. We obtained TEC from a dual frequency GPS receiver located at 16 17 Birnin-Kebbi Federal Polytechnic (BKFP) in Northern Nigeria (geographic location: 12.64° N; 18 4.22° E; 2.68° N dip) for the period 2011 – 2014 TEC is modified in the ionosphere by changing 19 solar Extreme Ultra-Violet (EUV) radiation, geomagnetic storms, and the atmospheric waves that 20 propagate up from the lower atmosphere. Therefore, TEC depends on local time, latitude, longitude, season, geomagnetic conditions, solar cycle activity, and condition of the troposphere. 21 A dual frequency GPS receiver located at an equatorial station, Birnin Kebbi in Northern Nigeria 22

(geographic location: 12.64° N; 4.22° E; 2.68° N dip), has been used to investigate variation of 23 TEC during the period of 2011 to 2014. We investigate the diurnal, seasonal and solar cycle 24 dependence of observed (OBS) TEC and comparison with latest version of International Reference 25 Ionosphere (IRI-2016) model. We observed differences between the diurnal variation OBS-TEC 26 27 and IRI-2016 model for all hours of the day except during the post-midnight hours. Slight post-28 noon peaks in the daytime maximum and post-sunset decrease and enhancement are observed in the diurnal variation of OBS-TEC of some months. On a seasonal scale, we observed that OBS-29 30 TEC values were higher in the equinoxes than the solstices only in 2012. Whereas in 2011, 31 September equinox and December solstice recorded higher magnitude followed by March equinox and lowest in June solstice. In 2013, December solstice magnitude was highest, followed by the 32 equinoxes and lowest in June solstice. In 2014, March equinox and December solstice magnitude 33 were higher than September equinox and June solstice magnitude. June solstice consistently 34 35 recorded the lowest values for all the years. OBS-TEC is found to increase from 2011 to 2014, 36 thus revealing solar cycle dependence. TEC is an important ionospheric parameter used to monitor possible space weather impacts on satellite to ground communication and satellite navigation 37 system. 38

KEYWORDS: TEC; diurnal; seasonal; variation; solar cycle 24; IRI-2016.

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46 1 INTRODUCTION

47 Ionospheric irregularities as a result of inhomogeneity in electron density leads to variations in the intensity of radio signals. <u>fading</u> as a result of irregularities (inhomogeneity in electron density) 48 (Somoye, 2010; Ogwala et al. 2018, Ogunmodimu et al. 2018). Akala et al., (2011) reported 49 50 posited that the variable nature of the equatorial/ low latitude ionosphere do adversely affect is a 51 threat to threatens communication and navigation/ satellite systems in the region. The equatorial/ low latitude ionosphere exhibits many unique features such as the seasonal anomaly, semi-annual 52 anomaly, equinoctial anomaly, noon bite-out, spread-F, equatorial electrojet (EEJ) and equatorial 53 plasma bubbles (EPB) (Stankov, 2009; Maruyama et al., 2004; Jee et al., 2004; Codrescu et al., 54 1999). 55

For many decades, scientists have been studying these peculiar ionospheric features and their roles 56 57 they play in trans-ionospheric electromagnetic radio wave propagation These studies are carried 58 out using different techniques and instruments. One of the instruments often used is the GPS 59 receiver. The GPS receiver provide direct measurements from satellites. Their sounding capacity 60 extends to the topside of the ionosphere, but is and affected by time and space constraints (Ciraolo 61 and Spalla, 2002). Recently, GPS receiver is the most efficient method used to eliminate the effect 62 of the ionosphere on radio signals. This method combine signals in different L band frequencies, L1 (1575 MHz) and L2 (1228 MHz) (Bolaji et al., (2012; Alizadeh et al., 2013). 63

Almost all space geodetic techniques transmit signals in at least two different frequencies for better accuracy (Alizadeh et al., 2013). These are The signals are then combined linearly in order to and can greatly eliminate the effect of the ionosphere on radio signals. The ionospheric effect on radio signal is proportional to the total electron content (TEC), which is defined as the number of electrons per square meter from satellite in space to receiver on ground is as shown in Eq. (1).

$$69 \quad TEC = \int n_e(s) ds \tag{1}$$

It is measured in multiples of TEC units (1 TECU = 10^{16} el/m²). Due to the dispersive nature of the ionosphere, there is a time delay between the two frequencies of a GNSS signal as it propagate s through the ionosphere as shown in Eq. (2) as $\Delta t = t_2 - t_1$. Thus,

73
$$\Delta t = \left(\frac{40.3}{C}\right) \times \frac{TEC}{\left[\left(\frac{1}{f_2^2}\right) - \left(\frac{1}{f_1^2}\right)\right]}$$
(2)

74 Where C is speed of light and f is frequency. Hence, Δt measured between the L1 and L2 75 frequencies is used to evaluate TEC along the ray path.

When Global Navigation Satellite System (GNSS) signals propagate through the ionosphere, the 76 carrier experiences phase advance and the code experiences a group delay due to the electron 77 78 density along the line of sight (LOS) from the satellite to the receiver (Bagiya et al., 2009; Tariku, 2015). Thus, the carrier phase pseudo ranges are measured too short, and the code pseudo ranges 79 are measured too long compared to the geometric range between the satellite and the receiver. This 80 results in a range error of the positioning accuracy provided by a GPS receiver. The range error 81 due to TEC in the ionosphere varies from hundreds of meters at mid-day, during high solar activity 82 when the satellite is near the horizon of the observer, to a few meters at night during low solar 83 activity, with the satellite positioned at zenith angle (Bagiya et al., 2009). By measuring this delay 84 85 using dual frequency GPS receivers, properties of the ionosphere can be inferred and used to 86 monitor space weather events such as GNSS, HF communications, Space Based Observation Radar and Situational Awareness Radar, etc. It is documented that Ionospheric delay which is 87 (proportional to TEC) is the highest contributor to GPS positioning error (Alizadeh et al., 2013; 88 89 Akala et al., 2013).

TEC in the ionosphere can also be studied using empirical ionospheric model such as the
International Reference Ionosphere (IRI). IRI is a joint undertaking by the Committee on Space
Research (COSPAR) and International Union of Radio Science (URSI) with the goal of developing
and improving an international standard for the parameters in earth's ionosphere (Bilitza et al.,
2014). An updated version has been recently developed recently to cater for lapses of previous
models. IRI provides the vertical TEC (VTEC) from the lower boundary (60 – 80 km) to a userspecific upper boundary (Bilitza et al., 2016).

97 In the past few decades, studies on the temporal and spatial variations of TEC have gained popularity in the scientific community (Wu et al., 2008). However, understanding the variability 98 of TEC will also go a long way in obtaining the positioning accuracy of GNSS under disturbed 99 100 and quiet conditions. As such, previous studies (e.g. Bhuyan and Borah, 2007; Maruyama et al., 101 2004; Jee et al., 2004; Balan et al., 1994; Rama Rao et al., 2006a, b, c; Rama Rao et al., 1985; Bolaji et al. 2012; Wanninger, 1993; Akala et al., 2013; Komjathy et al., 1998; Langley et al., 2002 102 103 Sunda et al., 2013; Torr and Torr, 1973; Tsai et al., 2001 and references therein) investigated the global distribution of TEC variations and its characteristics at all latitudes, during different solar 104 cycle phases under disturbed and quiet conditions. have been investigated by some researchers 105

Studies of Rama Rao et al. (2006a, b) in the Indian sector and Wanninger (1993) in the Asian sector reported maximum day-to-day variability in TEC at the Equatorial Ionization Anomaly (EIA) crest regions, increasing peak value of TEC with increase in integrated equatorial electrojet (IEEJ) strength, maximum monthly average diurnal variations during equinox months followed by winter months and lowest during summer months. They also reported positive correlation of TEC and EEJ and the spatial variation of TEC in the equatorial region. Titheridge (1974) and Langley et al. (2002) attributed the lowest TEC values during the summer seasons to 113 low ionization density resulting from reduced O/ N_2 ratio (production rates) which is as a result of increased scale height. Bhuyan and Borah (2007) working in the Indian sector and Komjathy et al. 114 (1998) and Lee and Reinisch (2006) while studying in the American sector compared TEC derived 115 from GPS receivers with IRI model in the equatorial/ low latitude sector and inferred that the 116 117 diurnal amplitude of TEC is higher during the equinoxes followed by December solstices and 118 lowest in June solstice, i.e., observing winter anomaly in seasonal variation. They further reported discrepancies between IRI model and their measured values during most hours of the day at their 119 120 various locations. Malik et al. (2016) on their studies over the Malaysian peninsular reported 121 higher IRI values than observed maximum useable frequency (MUF) values but behaves similarly diurnally and seasonally with no clear trend. Akala et al. (2013) on the comparison of equatorial 122 GPS-TEC observations over an African station and an American station during the minimum and 123 ascending phases of solar cycle 24 reported that seasonal VTEC values were maximum and 124 125 minimum during March equinox and June solstice respectively, during minimum solar cycle phase 126 at both stations. They also reported that during the ascending phase of solar cycle 24, minimum 127 and maximum seasonal VTEC values were recorded during December solstice and June solstice respectively. They further showed that IRI-2007 model predicted better in the American sector 128 129 than the African sector.

The aim of this paper is (i) to characterize TEC on diurnal, seasonal and solar cycle scales in the Nigerian Equatorial ionosphere (ii) to compare OBS-TEC with IRI-2016 model in order to find if the model underestimates or overestimates TEC values at the African longitudinal sectors. In section 2, we describe the data and methodology. Section 3 shows the result and discussion while concluding remarks are in section 4.

135 2 DATA AND METHODOLOGY

136 **2.1 DATA**

The Receiver Independence Exchange (RINEX) Observation GPS data files were 137 downloaded daily from NIGNET website (www.nignet.net) and processed using Bernese software 138 139 and GPS TEC analysis software. The RINEX file contains 60 iteration data (i.e. in 1 minute time resolution). The GPS-TEC analysis software was designed by Gopi Seemala of the Indian Institute 140 141 of Geomagnetism. The summary of this application is to read raw data, processes cycle slips in phase data, reads satellite biases from the International GNSS services (IGS) code files (and 142 143 calculates them if unavailable), and calculates receiver bias, inter-channel biases for different 144 satellites in the constellation, and finally plots the VTEC values on the screen and writes the ASCII output files (*CMN) for STEC and (*STD) for VTEC in the same directory of the data files. Effect 145 due to multipath is eliminated by using a minimum elevation angle of 50° . 146

Observation GPS-TEC obtained from the TEC analysis software is the slant TEC (STEC)
and vertical TEC (VTEC). STEC is polluted with several biases that must be eliminated to get
VTEC. VTEC is calculated from the daily values of STEC using Eq. (3).

150
$$VTEC = (STEC - [b_R + b_S + b_{RX}])/S(E)$$
 (3)

151 Where b_R , b_S , and b_{RX} are receiver bias, satellite bias and receiver interchannel bias respectively. 152 S(E), which is the oblique factor with zenith angle, z at IPP (Ionospheric Pierce Point) is expressed 153 in Eq. (4).

154
$$S(E) = \frac{1}{\cos(z)} = \left\{ 1 - \left(\frac{R_E \times \cos(E)}{R_E + h_S} \right)^2 \right\}^{-0.5}$$
 Bolaji et al. (2012) (4)

155 R_E = the mean radius of the earth in km and h_S = ionospheric height from the surface of the earth. 156 According to Rama Rao et al., (2006c), ionospheric shell height of approximately 350km is 157 appropriate for the equatorial/ low latitude region of the ionosphere for elevation cut off angle of 158 > 50°. This is valid in this study.

159	Hourly VTEC data obtained from these processing software are averaged to da	aily TEC
160	values in TEC units (1 TECU = 10^{16} el/m ²). OBS-TEC from Birnin-kebbi, on geographic	Latitude
161	12.47° N and geographic Longitude 4.23° E located in Northern Nigeria, obtained de	uring the
162	period $2011 - 2014$, which corresponds to the ascending ($2011 - 2013$) and maximum	m (2014)
163	phases of solar cycle 24 were compared with derived TEC obtained from International R	Reference
164	Ionosphere (IRI-2016) model	website
165	(https://ccmc.gsfc.nasa.gov/modelweb/models/iri2016_vitmo.php). The 2016 version	of IRI
166	provides important changes and improvements on previous IRI versions (Bilitza et al., 201	16). Solar
167	cycle 24 is regarded as a quiet solar cycle which peaked in 2014 with maximum sunspor	t number
168	(103) occurring in February. Values of sunspot number, Rz, in Text format were obtain	ned from
169	Space Physics Interactive Data Resource (SPIDR) website (www.ionosonde.spidr.com	<u>ı)</u> shortly
170	before it became unavailable. Table 1 shows the years used in this study and their corre	sponding
171	sunspot number, Rz.	

173 Table I: Table of years, solar cycle phase and sunspot number, Rz [Source: Aut	thor]
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Years	Solar Cycle Phase	Sunspot Number, Rz
2011	Ascending	55.7
2012	Ascending	57.6
2013	Ascending	64.7
2014	Maximum	79.6

175 2.2 METHODOLOGY

Diurnal variations of hourly OBS-TEC and hourly IRI-2016 model (NeQuick topside option) were plotted using the monthly mean values of OBS-TEC and monthly mean of IRI-2016 model against local time (LT) on the same Figure. The corresponding percentage deviation (percentage Dev or % DEV) of IRI-2016 from OBS-TEC were also analysed using the monthly mean values of OBS-TEC and monthly mean values of IRI-2016 against local time (LT). Percentage Dev is obtained using Eq. (5) below:

182
$$\% DEV = \left(\frac{OBS - IRI}{OBS}\right) \times 100$$
 (5)

where OBS represents Observation-TEC values and IRI represents TEC derived by IRI-2016model.

The OBS-TEC data was grouped following Onwumechilli and Ogbuehi (1964) into four seasons namely: March equinox (February, March and April), June solstice (May, June and July), September equinox (August, September and October) and December solstice (November, December and January), in order to investigate seasonal variation. Finally, Annual variation of OBS-TEC and sunspot number, Rz were also analysed by plotting mean OBS-TEC and mean Rz against each month of the year.

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3

RESULT AND DISCUSSIONS

Figures 1 to 4 show s the diurnal variation of OBS-TEC and IRI-2016 model in the Nigerian Equatorial Ionosphere (NEI) for the years 2011 to 2014 respectively. OBS-TEC were obtained from the GPS receiver installed at Birnin-Kebbi station. The diurnal variation of OBS-TEC and IRI-2016 model TEC reveals the typical characteristics of an equatorial/ low latitude ionosphere. We show Generally, the day-to-day variation of OBS-TEC have been indicated using with error bar showing the standard deviation from mean values. The study reveals that day-to-day variation of OBS-TEC is higher during the daytime than night time for all the years. It well established 199 known fact that during the day, the sun causes variations in temperature, neutral wind, electron density and electric field thereby modulating the structure and evolution of the ionosphere and 200 thermosphere (Gorney, 1990; Forbes et al. (2006). These Figs. show s a steep rise in OBS-TEC 201 202 from a minimum of ~2 TECU between 03:00 - 05:00 LT in 2011, ~3 TECU (04:00 - 05 LT) in 2012, ~3 TECU (03:00 – 05:00 LT) in 2013 and 2014. OBS-TEC increased to a broad daytime 203 204 maximum between 12:00 LT - 14:00 LT for all years before falling to a minimum after sunset. 205 The diurnal variation of IRI-2016 model shows TEC increasing from a minimum of ~ 2 TECU in 206 2011, ~ 4 TECU in 2012 and 2013, and ~ 5 TECU in 2014 between 03:00 – 04:00 LT for all years, 207 to a broad daytime peak between 08:00 - 14:00 LT, before falling steeply to minimum before sunset. Hence the IRI-2016 model attained its peak before OBS-TEC. Dabas et al. (2003); Somoye 208 209 et al. (2011); Hajra et al. (2016); D'ujanga et al. (2017) attributed the steep increase in TEC to the 210 and upward $\mathbf{E} \times \mathbf{B}$ vertical plasma drift resulting from and the rapid filling up of the magnetic field 211 tube at sunrise resulting at to solar EUV ionization. During daytime, an eastward electric field at 212 the equator causes plasma to be lifted to greater heights. This dynamo-generated eastward electric field combined with the northward geomagnetic field lifts the equatorial ionosphere from 700 km 213 to 1000 km, resulting in additional ionization (D'ujanga et al., 2017; Somoye et al., 2011). Suranya 214 215 et al., 2015 further mentioned that upward vertical $\mathbf{E} \times \mathbf{B}$ drift could lead to equatorial ionization anomaly (EIA) and meridional winds. The se magnetic field tubes then collapse after sunset due 216 217 to low thermospheric temperature and Rayleigh Taylor Instability (RTI) (Berkner and Wells, 1934) 218 giving rise to the minimum TEC values after sunset. These results are similar to findings of Bolaji et al., (2012), Fayose et al., (2012), Okoh et al., (2014), Eyelade et al., (2017) who have explored 219 220 the NEI.



Fig. 1: Diurnal variation of OBS-TEC showing error bar and IRI-2016 model of each month during January – December 2011 at Birnin-Kebbi



Fig. 2: Same as Figure 1 for 2012.



Fig. 3: Same as Figure 1 for 2013.



Fig. 4: Same as Figure 1 for 2014.

It can be seen that OBS-TEC is much higher in 2014 with maximum value up to 70 TECU in March compared with IRI-2016 maximum of 54 TECU in the month of October, 2014. The diurnal variation reveals that the peak of OBS-TEC of majority of the months for all years shifted to slightly post-noon hours (13:00 – 14:00 LT). This type of peak shifting is peculiar to equatorial/ low latitude regions and the Polar Regions of the ionosphere and it is found to depend on the equatorial ionization anomaly and solar zenith angle respectively (Rama Rao et al., 2009; D'ujanga

238 et al., 2017). Another key observation major phenomenon seen in the diurnal variation of OBS-TEC is the post-sunset decrease and slight enhancement in some months. The night time 239 240 enhancement of TEC, for example, March, April and October of the year 2011, March and April 241 of the year 2012, March, April, September and October of the year 2013, January, April and September of the year 2014 was documented by previous researchers like Rama Rao et al., 2009; 242 243 D'ujanga et al., 2017. They attributed it to the product of eastward and westward directed electric field which produces an upward and downward motion of ionospheric plasma during the day and 244 245 night respectively.

246 Figures 5 to 8 show s the diurnal variation of percentage deviation of IRI-2016 model from OBS-TEC in the Nigerian Equatorial Ionosphere (NEI) for all years respectively. On a general 247 248 note, IRI-2016 model only presented suitable predictions for the post-midnight hour between 00 - 100249 03 LT of the day for all years. All other hours from 04 - 23 LT show s some discrepancies. In fact, for some of the months: October, November and December of 2012, October and December, 2013, 250 251 and September and October of 2014 these discrepancies lasted throughout the day. Whereas in 252 some other months: June, July and August of 2011, June, July and August of 2012, June and 253 August of 2013 and February, June and July of 2014 these discrepancies collapsed during the pre-254 midnight hours (18 - 23 h). It is also important to mention that IRI-2016 model either over estimated or under estimated TEC in the NEI especially during daytime hours as shown in the 255 plots. 256

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Fig. 5: Percentage deviation of IRI-2016 from OBS-TEC for year 2011



Fig. 6: Percentage deviation of IRI-2016 from OBS-TEC for year 2012



Fig. 7: Percentage deviation of IRI-2016 from OBS-TEC for year 2013



Fig. 8: Percentage deviation of IRI-2016 from OBS-TEC for year 2014

The mass plots in Figs. 5 - 8 further reveal that negative percentage deviation shows higher values of IRI-2016 than OBS-TEC values. The reverse is the case for positive percentage deviation. Highest negative percentage deviations are seen between 04 – 05 LT for all months throughout the years in this study. Highest Negative percentage deviation of ~ 300% was recorded in the month of October, 2014 at 05 LT. Table II shows the summary of months with daytime over- or under-estimate of IRI-2016 in the NEI.

YEAR	OVER ESTIMATE	UNDER ESTIMATE	SAME RANGE
2011	January, July, August	February – April, September -	May - June
		December	
2012		January - December	
2013	September	January – August, October -	
		December	
2014		January - December	

Therefore, it is clear from the Fig. 1-4, Fig. 5 – 8 and Table II that IRI-2016 model did not predict well in the NEI. This may be attributed to insufficient data which is as a result of the sparse distribution of GPS infrastructure in this region. Our result agree with those Komjathy et al. (1998); Lee and Reinisch (2006); Malik et al. (2016). Bhuyan and Borah (2007) reported higher IRI TEC than their measured values at about all local time in their location. Mosert et al. (2007) and Sethi et al. (2010) also reported discrepancies between of IRI TEC predictions and GPS TEC during high solar activity (HSA) and low solar activity (LSA) respectively at equatorial/ low latitudes.

Figure 9 show plots of the seasonal variations of OBS-TEC for the four years under investigated. The change in concentration of Oxygen and molecular Nitrogen has been reported to be the main cause of seasonal variation of ionospheric parameters. Seasonal variation of OBS-TEC in this study depicts semi-annual variation with equinoctial maximum (~ 52 TECU) and solsticial minimum (~ 44 TECU) in 2012. D'ujanga et al., (2017) reported that since the sun passes through the equator during the equinox, both March and September equinox experience the same solar radiation. It is also a well-established fact that March 20 and September 23 are the only times

296	in the year when the solar terminator is perpendicular to the equator, giving rise to the equinoctial
297	maximum. The semi-annual variation resulting from the effect of equatorial ionization anomaly
298	(EIA) in the ionosphere has been attributed to the effect of solar zenith angle and magnetic field
299	geometry (Wu et al., 2008; Rama Rao et al., 2006a). Another important feature of ionospheric
300	parameters (known as equinoctial asymmetry) which is reported in the work of Bolaji et al., (2012);
301	Akala et al., (2013); Eyelade et al., (2017); D'ujanga et al., (2017); Aggarwal et al., (2017), is
302	clearly seen in all years used in this work. Akala et al., (2013) also reported minimum and
303	maximum seasonal VTEC values during June solstice and December solstice respectively, during
304	ascending phase of solar cycle 24. Equinoctial asymmetry is a strong phenomenon in low latitudes
305	(Aggarwal et al., 2017). The equinoctial asymmetry has been explained in terms of the differences
306	in the meridional winds leading to changes in the neutral gas composition during the equinoxes.
307	



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Fig. 9: Seasonal variation of observed OBS-TEC during (a) 2011 (b) 2012 (c) 2013 and (d) 2014

In 2011 and 2014, the seasonal variation of OBS-TEC in the ionosphere did not follow the pattern reported by these researchers. In 2011, September equinox and December solstice recorded higher magnitude, followed by March equinox; the lowest was in June solstice. In 2013, December solstice magnitude was highest, followed by the equinoxes, March and September respectively and lowest in June solstice. This corresponds to result obtained by Akala et al. (2013), which they attributed to increase in ion production rate in winter season and anti-correlation between December and June Solstice pre-reversal velocity enhancement. In 2014, March equinox magnitude was highest, December solstice and September equinox magnitudes were about same range while June solstice magnitudes were least. December solstice magnitude is found to occur between the magnitudes of the equinoxes in 2011 and 2014. The September equinox magnitude and March equinox magnitude are observed to interchange in 2011 and 2014. Overall, June solstice magnitudes were lowest during all the years. Titheridge (1974) attributed the June solstice least magnitudes to low ionization resulting from reduced production rates, i.e. O/ N₂ ratio.

325 Also, for all seasons, pre-midnight (18 - 23 LT) values of TEC are higher than post-326 midnight (00 - 05 LT) TEC values for all years. In 2011, pre-midnight TEC values are in the 327 range of 8 – 30 TECU while post-midnight TEC values ranges from 3 to 13 TECU. In 2012, premidnight TEC values are in the range of 9-35 TECU while post-midnight TEC values are between 328 329 3 to 17 TECU. In 2013, the pre-midnight TEC values are between 9 - 35 TECU while postmidnight TEC values ranges from 3 – 15 TECU. Finally in 2014, pre-midnight TEC values are 330 between 9 to 47 TECU while the post-midnight TEC ranges from 3 to 18 TECU. Furthermore, the 331 332 maximum OBS-TEC values in 2011 (49 TECU) and 2012 (52 TECU) were recorded in September Equinox season. In 2013, OBS-TEC reached a maximum of 53 TECU in December solstice while 333 in 2014, the maximum OBS-TEC (70 TECU) was recorded in March Equinox season in the NEI 334 335 (West African sector). This result agrees in general with those of D'ujanga et al. (2017) who obtained higher TEC values during the equinoxes than during the solstices in Ethiopia (East 336 African sector). This same result was observed by Bagiya et al. (2009) that TEC values are high 337 338 in equinoctial months than solsticial months in the Indian sector. While the former authors reported maximum TEC of ~ 58 TECU during the equinox months, the latter authors reported maximum 339 340 TEC ~ 50 TECU during the equinoxes. Seasonal variation of TEC is dependent on thermospheric 341 neutral compositions since during the day the equator is hotter than the pole. Meridional winds

therefore flows from the equator towards the pole. This flow cause a change in the neutral composition resulting in the decrease of the ratio of O/N_2 at the equator. The decrease of O/N_2 ratio increases the electron density, and thus resulting in result in higher TEC increase values during the equinoxes (Bagiya et al., 2009). The corresponding annual range error (meters) of the season with maximum OBS-TEC using 1 TECU variation to represent an error of 0.16 m in position is summarized in Tables III.

YEAR	SEASON OF	VALUE	CORRESPONDING
	MAXIMUM OBS-TEC	(TECU)	ERROR (m)
2011	September Equinox	49	8
2012	September Equinox	52	8
2013	December Solstice	53	8
2014	March Equinox	70	11

Table III: Season of maximum OBS-TEC and their corresponding range error.

349

350 Figure 10 shows the comparison of the monthly mean OBS-TEC and monthly mean sunspot number, Rz from 2011 – 2014, showing an increase and decrease of TEC with solar cycle. 351 Our result is in good agreement with those of Chakrabarty et al. (2012) and D'ujanga et al. (2017). 352 353 The former authors reported a direct solar effect on TEC while the latter authors observed that the 354 trend in TEC follow that of solar parameters. Solar cycle dependency of ionospheric parameters 355 such as TEC provides useful information to study the behavior and variations of the physical and 356 photochemical processes in the ionosphere (Liu et al. 2006). Solar activity indicates the intensity of the solar electromagnetic radiations like wavelengths of X-rays and extreme ultraviolet (EUV) 357 radiations. These radiations vary regularly and irregularly over time, larger variability tends to 358

359 occur at very short wave length. It is well documented that the variability of solar activity results

360 in huge variations in the temperature, neutral wind, neutral density, ion and electron densities and

361 electric fields in the ionosphere (Forbes et al. 2006).

362







Figure 10: Annual variation of OBS-TEC with sunspot number, Rz at Birnin-Kebbi

365 The present result also agrees with that of Rama Rao et al. (1985) who reported a direct solar control on TEC. Balan et al. (1994); Liu et al. (2011) and Liu et al. (2006) and many more 366 reported that during low and moderate solar activity, TEC and N_mF_2 increases linearly with solar 367 368 proxies, but the linearity collapse during high solar activity. This agrees with our result except for July month of 2012 and 2014 which shows saturation effect on TEC i.e decrease in TEC with 369 370 increase in solar activity. The saturation effect on TEC was reported in the work of Balan et al. (1994) and Balan et al. (1996). Both authors in their research concluded that the saturation effect 371 has not been clarified and hence might be due to other factors near the earth's environment and 372 not as a result of the influence of solar activity. We could not establish the cause of the saturation 373 374 effect on TEC in this study, however, saturation effect will be further investigated in future studies.

375 Ionospheric climatology especially solar activity and the equatorial ionization anomaly (EIA)
376 effects of the `ionosphere provides an insight into space weather events (Liu et al., 2011).

377

378 4 CONCLUSIONS

Studies on OBS-TEC and IRI-2016 model at Birnin-Kebbi in Northern Nigeria during the
 ascending and maximum phases of solar cycle 24 have been carried out. The Our result shows
 obtained reveals the following:

382 1. Day-to-day variations of TEC were observed to be higher during the daytime than nighttime for all the years. During the day, the sun causes variations in temperature, neutral 383 384 wind, electron density and electric field in the ionosphere. The diurnal variation shows a steep rise in that OBS-TEC and IRI-2016 model rising from a minimum in the early hours 385 386 of the day of ~2 TECU between 03:00 05:00 LT in 2011, ~3 TECU (04:00 05 LT) in 387 2012, ~3 TECU (03:00 - 05:00 LT) in 2013 and 2014. OBS-TEC increased to a broad 388 daytime maximum between 12:00 LT - 14:00 LT before falling steeply to a minimum after sunset for all years. Magnetic field tubes are rapidly filled up at dawn resulting in the 389 increase of extreme ultraviolet (EUV) ionization (Anderson et al., 2004; D'ujanga et al., 390 391 2017). The diurnal variation of IRI-2016 model shows TEC rising from a minimum of ~ 2 TECU in 2011, ~ 4 TECU in 2012 and 2013, and ~ 5 TECU in 2014 between 03:00 - 04:00 392 LT, to a broad daytime peak between 08:00 - 14:00 LT, before falling to a minimum before 393 394 sunset.

The diurnal variation reveals that the peak of OBS-TEC of many of the some months were
 delayed till after-noon, showing dome-like shape while noon bite-out, a special feature

397 observed in equatorial/ low latitude is can be seen in the peak of majority of the plots of
 398 IRI-2016 model. Post-sunset decrease and enhancement due to pre-reversal zonal electric
 399 field after sunset, were also observed in the diurnal variation of OBS-TEC in some months.

- 3. On a general note, we concluded that IRI-2016 model cannot be used as proxy for TEC
 measurements for most hours of the day for the years investigated. Our result agree with
 those of Komjathy et al. (1998); Lee and Reinisch (2006); Malik et al. (2016); Bhuyan and
 Borah (2007); Mosert et al. (2007) and Sethi et al. (2010) at their respective locations. and
 OBS-TEC show some discrepancies throughout the day except during the post-midnight
 hours (00 05 LT) in the NEI.
- 4. For all seasons, pre-midnight (18 23 h) values of TEC are higher than post-midnight (00 406 -05 h) TEC values during all years. In 2011, pre-midnight TEC values are in the range of 407 8-30 TECU while post-midnight TEC values ranges from 3 to 13 TECU. In 2012, pre-408 midnight TEC values are in the range of 9 - 35 TECU while post-midnight TEC values are 409 410 between 3 to 17 TECU. In 2013, the pre-midnight TEC values are between 9 - 35 TECU while post-midnight TEC values ranges from 3-15 TECU. Finally in 2014, pre-midnight 411 TEC values are between 9 to 47 TECU while the post-midnight TEC ranges from 3 to 18 412 413 TECU. Seasonal variation shows an asymmetry in the equinoxes and solstices as reported in the research of Fayose et al. (2012) working at Ilorin in the NEI and Eyelade et al. (2017) 414 also working in the NEI. 415
- Maximum OBS-TEC values in 2011 (49 TECU) and 2012 (52 TECU) were recorded in
 September Equinox season. In 2013, OBS-TEC reached it's a maximum of 53 TECU in
 during the December solstice season while in 2014, the maximum OBS-TEC (70 TECU)
 was recorded in March Equinox season. This result agrees in general with those of

D'ujanga et al. (2017) in the East African sector and Bagiya et al. (2009) in the Indian
sector, who obtained higher TEC values during the equinoxes than during the solstices.
Thermospheric neutral compositions is a major cause of seasonal variation of ionospheric
parameters such as TEC, is dependent on since during the day the equator is hotter than the
poles.

- 6. Finally, annual variation of OBS-TEC varies linearly with and annual sunspot number, Rz,
 thus revealing strong dependence of TEC on solar activity (sunspot number). This linearity
 collapsed in the month of July of 2012 and 2014. Solar activity and the equatorial ionization
 anomaly (EIA) effects of the ionosphere provides an insight into space weather events.
 OBS-TEC and sunspot number were found to increase gradually from 2011 to 2014 in
 agreement with Rama Rao et al. (1985) that there is reported a direct solar control on TEC.
- 431 AUTHORS' CONTRIBUTIONS
- 432 Author AA designed the study, obtained the data and wrote the first draft of the manuscript. Author
 433 BB and CC analysed the data and wrote the protocol. All other authors managed the literature
 434 searches, read and approved the final manuscript.

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442 **REFERENCES**

- Aggarwal, M., Bardhan, A., Sharma, D.K. Equinoctial asymmetry in ionosphere over Indian
 region during 2006 2013 using COSMIC measurements. Advances in Space Res., 60, 999 –
 1014, 2017.
- 446 Akala, A.O., Somoye, E.O, Adeloye, A.B., Rabiu, A.B. Ionospheric f_0F_2 variability at equatorial 447 and low latitudes during high, moderate and low solar activity. Indian Journal of Radio and Space 448 Physics.Vol. 40, 124 – 129, 2011.
- Akala, A.O., Seemala, G.K., Doherty, P.H., Valladares, C.E., Carrano, C.S., Espinoza, J., and
 Oluyo, K.S. Comparison of equatorial GPS-TEC observations over an African station and an
 American station during the minimum and ascending phases of solar cycle 24. Ann. Geophys., 31,
- 452 2085, 2013.
- Alizadeh, M.M., Wijaya, D.D., Hobiger, T., Weber, R., Schuh, H. Ionospheric effects on
 microwave signals in J. Bohm and H. Schuh (eds). Atmospheric Effect in Space Geodesy. Springer
 atmospheric sciences. Doi: 10.1007/978-3-642-36932-2_2, © Springer-Verlag Berlin Heidelberg,
 2013.
- 457 Anderson, D., Anghel, A., Chau, J., Veliz, O. Daytime vertical $\mathbf{E} \times \mathbf{B}$ drift velocities inferred from 458 ground-based magnetometer observations at low latitudes, Space Weather, 2, S11001; doi: 459 10.1029/2004SW000095, 2004.
- 460 Ayorinde, T.T., Rabiu, A.B., and Amory-Mazaudier, C. Inter-hourly variability of Total Electron
- 461 Content during the quiet condition over Nigeria within the Equatorial Ionization Anomaly region.
- 462 J. Atmos. Solar Terr. Phys., 145, 21 33, 2016.

- Berkner, L.V. and Wells, H.W. F-region ionosphere investigation at low latitude. Terres. Magn.,
 39, 215, 1934.
- 465 Bagiya, M.S., Joshi, H.P., Iyer, K.N., Aggarwal, M., Ravin-dran, S., and Pathan, B.M. TEC
- 466 variations during low solar activity period (2005 2007) near the Equatorial Ionization Anomaly
- 467 Crest region in India. Ann Geophys., 27, 1047 1057, 2009.
- Balan, N., Bailey, G.J., Moffett, R.J. Modelling studies of ionospheric variations during an intense
 solar cycle. J. Geophys. Res., 99, 17467 17475, 1994.
- 470 Balan, N., Bailey, G.J., Su, Y.Z. variations of the ionosphere and related solar fluxes during solar
- 471 cycle 21 and 22. Advances in Space Res., 18, 11 14, 1996
- Bhuyan, P.K. and Borah, R.R. TEC derived from GPS network in India and comparison with the
 IRI. Advances in Space Res., 39, 830 840, 2007.
- 474 Bilitza, D., Altadill, D., Zhang, Y., Mertens, C., Truhlik, V., Richards, P., McKinnell, L.A.,
- 475 Reinisch, B. International reference ionosphere 2012 A model of international collaboration. J.
- 476 Space Weather Space Clim., 4, 1 12, doi: 10.1002/201JA018009, 2014.
- Bilitza, D., Altadill, D., Truhlik, V., Shubin, V., Galkin, I., Reinisch, B., and Huang, X.
 International reference ionosphere 2016: from ionospheric climate to real-time weather
 predictions. Space weather, 15, 418 429, doi: 10.1002/20165SW001593, 2016.
- 480 Bolaji, O. S., Adeniyi, J.O., Radicella, S.M., and Doherty, P.H. Variability of total electron content
- 481 over an equatorial West African station during low solar activity. Radio Sci., 47, RS1001, doi:
- 482 10.1029/2011RS004812, 2012.

483 Charkrabarty, D., Bagiya, M.S., Thampi, S.V., Iyer, K.N. Solar EUV flux (0.1 - 50 nm), $F_{10.7}$ cm 484 flux, sunspot number and total electron content in the crest region of the ionization anomaly during 485 the deep minimum between solar cycle 23 and 24. Indian Radio and Space Phys., 41, 110 – 120, 486 2012.

- 487 Ciraolo, L., and Spalla, P. TEC analysis of IRI simulated data. Adv. Space Res., 29, 6, 959 966,
 488 2002.
- Codrescu, M. V., Palo, S. E., Zhang, X., Fuller-Rowell, T. J., Poppe, C. TEC climatology derived
 from TOPEX/POSEIDON measurements, Journal of Atmospheric Solution, 61, 281-298, 1999.
- 491 Dabas, R.S., Singh, L., Lakshmi, D.R., Subramanyam, P., Chopra, P., Garg, S.C. Evolution and 492 dynamics of equatorial plasma bubbles: relationships to $\mathbf{E} \times \mathbf{B}$ drifts, post-sunset total electron 493 content enhancements, and equatorial electrojet strength. Radio Sci., 38. doi: 10.1029/2001RS002586, 2003. 494

D'ujanga, F.M., Opio, P. Twinomugisha, F. Variation of total electron content with solar activity
during the ascending phase of solar cycle 24 observed at Makerere University, Kampala. Space
Weather: Longitude and Hemispheric Dependences and Lower Atmosphere Forcing, Geophysical
Monograph 220, First Edition. Edited by Timothy Fuller-Rowell, Endawoke Yizengaw, Patricia
H. Doherty, and Sunanda Basu. © 2017 American Geophysical Union. Published 2017 by John
Wiley & Sons, Inc., 2017.

Eyelade, V.A., Adewale, A.O., Akala, A.O., Bolaji, O.S. and Rabiu, A.B. Studying the variability
in the diurnal and seasonal variations in GPS TEC over Nigeria. Ann. Geophys., 35, 701 – 710,
2017.

- 504 Fayose, R.S., Rabiu, B., Oladosu, O., Groves, K. Variation of total electron content (TEC) and
- their effect on GNSS over Akure. Nigeria, Applied Physics Research, 4, 2, 2012.
- 506 Forbes, J.M., Bruinsma, S., Lemoine, F.G. Solar rotation effects in the thermospheres of Mars
- 507 and Earth. Science, 312, 1366–1368, 2006.
- Gorney, D. J. Solar cycle effects on the near-earth space environment. Rev Geophys, 28, 315–
 336, 1990.
- 510 Hajra, R., Chakraborty, S.K., Tsurutani, B.T., DasGupta, A., Echer, E., Brum, C.G.M., Gonzalez,
- 511 W.D., Sobral, H.A. An empirical model of ionospheric total electron content (TEC) near the crest
- 512 of the equatorial ionization anomaly (EIA). J. Space Weather Space Clim., 6, A29, doi:
- 513 10.1051/swsc/2016023, 2016.
- Jee, G., Schunk, R. W., Scherliess, L. Analysis of TEC data from the TOPEX/Poseidon mission,
- 515 Journal of Geophysical Research, 109, A01301, doi:10.1029/2003JA010058, 2004.
- 516 Komjathy, A., Langley, R., Bilitza, D. ingesting GPS-derived data into the IRI for single frequency
- radar altimeter ionospheric delay corrections. Adv. Space Res., 22, 6, 793 802, 1998.
- Langley, R., Fedrizzi, M., Paula, E., Santos, M., Komjathy, A. Mapping the low latitude
 ionosphere with GPS. GPS World, 13, 2, 41 46, 2002.
- 520 Lee, C.C. and Reinisch, B.W. Quiet condition hmF2, NmF2 and Bo variations at Jicamarca and
- comparison with IRI-2001 during solar maximum. J. Atmos. Solar Terr. Phys., 68, 2138 2146,
 2006.
- 523 Liu, L., Wang, W., Chen, Y., Le, H. Solar activity effects on the ionosphere: A brief review. Space
- 524 Physics and Space Weather Geophysics. Chinese science Bulletin. 56, 12, 1202 1211, 2011.

- Liu, L., Wang, W., Ning, B., Pirog, O.M., and Kurkin, V.I. Solar activity variations of the
 ionospheric peak electron density. J. Geophys. Res., vol 111, A08304, doi:
 10.1029/2006JA011598, 2006
- 528 Maruyama, T., Ma, G., and Nakamura, M. Signature of TEC storm on 6 November 2001 derived
- 529 from dense GPS receiver network and ionosonde chain over Japan. Journal of Geophysical
- 530 Research, 109, A10302, doi: 10.1029/2004JA010451, 2004.
- 531 Mosert, M., Gende, M., Brunini, C., Ezquer, R. and Altadill, D. Comparisons of IRI TEC with
- 532 GPS and Digisonde measurements at Ebro. Advances in Space Res., 39, 841 847, 2007.
- 533 Okoh, D., Lee-Anne McKinnell, L., Cilliers, P., Okere, B., Okonkwo, C., Rabiu, A.B. IRI-VTEC
- versus GPS-vTEC for Nigerian SCINDA GPS stations. Advances in Space Research,
 http://dx.doi.org/10.1016/j.asr.2014.06.037, 2014.
- 536 Ogunmodimu, O., Rogers, N.C., Falayi, E., Bolaji, S. Solar Flare induced cosmic noise absorption,
- 537 NRIAG Journal of Astronomy and Geophysics. 7, 1, 31-39, 2018.
- 538 Ogwala, A., Somoye, E.O., Oyedokun, O., Adeniji-Adele, R.A., Onori, E.O., Ogungbe, A.S.,
- 539 Ogabi, C.O., Adejo, O., Oluyo, K.S., Sode, A.T. Analyses of Total Electron Content over Northern
- and Southern Nigeria. J. Res. and Review in Sci., 21 27, 2018.
- 541 Onwumechilli, C.A., and Ogbuehi, P.O. Journal Atmos. Terr. Phys., 26, 894, 1964.
- 542 Rama Rao, P.V.S., Niranjan, K., Rama Rao, B.V., Rama Rao, B.V.P.S., Prasad, D.S.V.V.D. Proc.
- 543 URSI/ IPS Conference on the ionosphere and Radio wave Propagation. Sydney, Australia, 1985.
- 544 Rama Rao, P.V.S., Krishna, S.G., Prasad, J.V., Prasad, S.N.V.S., Prasad, D.S.V.V.D., Niranjan,
- 545 K. Geomagnetic storm effects on GPS based navigation. Ann. Geophys., 27, 2101 2110, 2009.

- Rama Rao, P.V.S., Krishna, S.G., Niranjan, K., Prasad, D.S.V.V.D. Study of temporal and spatial
- 547 characteristics of L-band scintillation over the Indian low-latitude region and their possible effects
- 548 on GPS navigation. Ann. Geophys., 24, 1567 1580, 2006a.
- 549 Rama Rao, P.V.S., Krishna, S.G., Niranjan, K., Prasad, D.S.V.V.D. Temporal and spatial
- variations in TEC using simultaneous measurements from indian GPS network of receivers during
- 551 low solar activity period of 2004 2005. Ann. Geophys., 24, 3279 3292, 2006b.
- 552 Rama Rao, P.V.S., Niranjan, K., Prasad, D.S.V.V.D., Krishna, S.G., Uma, G. On the validity of
- the ionospheric pierce point (IPP) altitude of 350km in the Indian equatorial and low-latitude
- sector. Ann. Geophys., 24, 2159 2168, 2006c.
- Sethi, N. K., Pandey, V. K., Mahajan, K. K. Comparative study of TEC with IRI model for solar
 minimum period at low latitude. Advances in Space Research, 27, 45 48, 2010.
- 557 Suranya, P.L., Prasad, D.S.V.V.D., Niranjan, K., Rama Rao, P.S.V. Short term variability in foF2
- and TEC over low latitude stations in the Indian sector. Indian J. of Radio and Space Phys., 44, 14
 27, 2015.
- 560 Somoye, E.O. Diurnal and seasonal variation of fading rates of E- and F-region echoes during IGY
- and IQSY at the equatorial station of Ibadan. Indian Journal of Radio and space Physics, 38, 194
 202, 2010.
- Somoye, E.O., Akala, A.O., Ogwala, A. Day-to-day variability of h'F and foF2 during some solar
 cycle epochs. Journal Atmos. Solar Terr. Physics, 73, 1915 1922, 2011.
- 565 Stankov, S. M. Trans-ionospheric GPS signal delay gradients observed over mid-latitude Europe.
- 566 Advances in Space Research, 43, 1314–1324, 2009.

- Sunda, S. and Vyas, B.M. Local time, seasonal and solar cycle dependency of longitudinal
 variations of TEC along the crest of EIA over India. J. Geophys. Res., vol 118, 6777 6785, 2013.
- 569 Tariku, Y.A. Pattern of GPS-TEC variability over low-latitude regions (African sector) during the
- 570 deep solar minimum (2008 to 2009) and solar maximum (2012 to 2013) phases. Earth, Planets,
- 571 and space. 67, 35, 2015.
- 572 Titheridge, J.E. Changes in atmospheric composition inferred from ionospheric production rates.
- 573 J. Atmos. Terr. Phys., 36, 1249 1257, 1974.
- 574 Torr, M.R. and Torr, D.G. The seasonal behavior of the F2 layer of the ionosphere. J. Atmos. Terr.
 575 Phys., 35, 2237, 1973.
- 576 Tsai Ho-Fang, Liu Jann-Yenq, Tsai Wei-Hsiung and Liu Chao-Han, Tseng Ching-Liang and Wu
- 577 Chin-Chun. Seasonal variations of the ionospheric TEC in Asian equatorial anomaly regions. J.
- 578 Geophysical Res., vol 106, A12, 30,363 369, 2001.
- 579 Wanninger, L. Effects of the equatorial ionosphere on GPS. GPS World, 2, 48, 1993.
- 580 Wu, C.C., Liou, K., Shan, S.J., Tseng, C.L. Variation of ionospheric total electron content in
- Taiwan region of the equatorial anomaly from 1994 2003. Adv. Space Res., 41, 611 616, 2008.