Interactive comment on "Assessing the role of planetary and gravity waves on the vertical
 structure of ozone over central Europe" by Peter Križan –referee 1

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4 The manuscript describes statistics of the lamina appearance in the ozone vertical distribution 5 in dependence of the lamina origin (due to planetary or gravity waves). Thus the subject is 6 well suited to the journal scientific profile. The author uses the methodology elaborated by 7 Teitelbaum et al (1995) to classify the lamina based on the correlation coefficient between 8 vertical profiles of ozone and potential temperature. The reviewer has found interesting and 9 worth publishing results. However, there is a serious problem with selection of the profile 10 data. Thus, the manuscript is not ready for publishing. It may have potential after additional work and resubmission. Table 3 clearly shows that the vertical resolution of the profile should 11 12 be lower than 100 m for proper identification of the lamina with size less than 1 mPa and less than 500 m for the lamina size in the range 1-4 mPa. Figure 12 illustrates strong 13 14 inhomogeneity of the vertical resolution for all the stations. The same is also seen from Table 15 2. Lindenberg profiles should be excluded from the analysis because of the large and variable vertical resolution. Thus, the analysed data are not homogeneous that may influence the 16 17 results. A scale of this effect needs to be evaluated in the revised paper or only the latest 18 results with the high resolution of the ozone profiles should be a subject of analysis. It means 19 that the results shown in Fig.6 should be valid for only two stations since 1990 for the lamina 20 size < 1mPa. For laminae in the range 1-4mPa the analyses will be possible for 3 stations 21 since 1970. Thus in present form Fig. 6 is wrong especially for Lindenberg. 22 23 We excluded the station Lindenberg from the paper and we use only the stations Payerne, 24 Uccle and Legionowo in the period 1995-2016 where the vertical resolution of the ozone 25 profile is about 100 m. 26 27 Minor problems: 1.1-2The title is not proper: Hohenpeissenberg, Payern, and Uccle are 28 located in the western part of Europe. It is better to change the title to "the midlatitudinal 29 Europe". 30 The title of the paper was changed 31 32 1.112-116. Have you excluded from the analyses evidently wrong profiles with the correction 33 factor far from 1 (a case for early Legionowo and Lindenberg ozone profiles)? 34 35 These profiles were excluded from the analyses. 36 37 1.158-185. This section should be rewritten. In fact, Hohenpeissenberg profiles are not proper 38 for analyses of laminae with size <2 mPa as for almost the whole period the vertical 39 resolution is~500 m (see Fig.12). The Hohenpeissenberg data are proper for analysis of the laminae with the size > 2 mPa. The author could not state that similar results were derived for 40 other stations, as for Lindenberg (all observations) and Legionowo (early observations before 41 42 1990) were not possible to identify correctly lamina with the size <2 mPa. 43 44 We use here the station Uccle in the period 1995-2016, so this problem is solved. 45 46 1. 190 -197. Trend values should appear (% for 10 yr.) with their error estimates to discuss the trend significance. The two-joint lines trend model with the turning point in the mid1990s 47 48 needs to apply also for the gravity waves laminae for better comparison with PL laminae. If

- 49 you calculate the trend based on single line approach for the PL laminae you will probably
- 50 result with small negative trend as you discussed for the case of the GL lamina trend.

51	From figure 11 we see principally different trends for PL and GL. So the piecewise regression
52	is suitable only for Pl laminae. This regression is not suitable for GL. In this case it gives
53	insignificant trend before 1995 and insignificant change in 1995. On the other hand, the
54	classical regression is erroneous for PL and the most suitable for GL where it gives significant
55	negative trend.
56	
57	l. 215-220. The discussion is not correct for Payern as this station is located in the valley
58	between the Jura Mountains and Alps.
59	
60	This sentence was changed.
61	Thank you for all your comments. They make my paper better.
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68	Interactive comment on "Assessing the role of planetary and gravity waves on the vertical
69	structure of ozone over central Europe" by Peter Križan –referee 2
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71	Anonymous Referee #2
72	Based on the Teitelbaum method, this manuscript studies the characteristics of ozone lamina
73	under the influence of planetary and gravity waves. This article seems to have done a lot of
74	work. Even though I'm not an expert in this area (ozone lamina), there are a few things that
75	make me confused.
76	Major comments: 1 The formation mechanism of ozone lamina. Tomikawa et al. (2002)
//	reported that the formation of the ozone laminae is closely related to the vertical shear of the
/8	subtropical jet. I strongly suggest the authors to discuss in the introduction about the
/9	Tormation mechanisms of the ozone famina and in which the role of wave activities play. Tormitory, $N = Soto K$, Kito K, Eujiyyara M, Yamamari M, & Sono T, (2002)
80 81	Formation of an azona lamina due to differential advaction revealed by intensive
81 82	characteristic L Goophys Res. 107(D10)
82 83	observations. J. Geophys. Res., 107(D10).
83 84	This paper is referred to in Discussion
85	This paper is referred to in Discussion.
86	2 Lines 101-108 need some references (at least one) or make some explanation: e.g. Why
87	choose 5-15km and 17-22km height area to distinguish. If the identification process is
88	proposed by the author, it appears from the description that the author only uses ozone thin
89	layers at different height region to define whether the ozone laminae is caused by gravity or
90	planetary wayes. This makes it very puzzling because gravity wayes almost exist anywhere in
91	the earth's atmosphere.
92	
93	I did the research for all heights from the ground to the 5 km below the highest profile point.
94	The intervals 5-15 km and 17-22 km were chosen because the correlation here is sufficiently
95	high (above 0.7) or low (below 0.3) for detection of gravity and planetary waves. In these
96	intervals the ozone profile is strongly influenced by atmospheric waves. Outside these
97	intervals the profile is not so strongly influenced as in these intervals. The atmospheric
98	waves can occur outside intervals, but they do not influence the ozone vertical profile.
99	3. The authors have only mentioned the thin layer of ozone caused by gravitational and
100	planetary waves, but I think that some other meso-and small-scale atmospheric processes

101 102	(such as strong convection, tropopause folding, strong wind shear, stratospheric streamers, etc.) may also responsible for the formation of ozone laminae.
103 104 105	Various mechanisms of lamina formation are described in the Discussion
105 106 107 108 109	4. Gravity and planetary waves run through the title and the paper, but there is no evidence of their existence in the manuscript (even though the authors indicate that the ozone profile can be used to detect fluctuations)
110 111	This title of paper was recommended by editor.
112 113 114	5. Reading the manuscript, I still didn't understand how gravity and planetary waves affect and lead to ozone laminae. Personally, a detailed case is necessary.
115 116	This problem is theoretically solved in Teitelbaum paper.
117 118 119 120	Minor comments: 1. Lines 44-46, as you mentioned, it is the large lamina that has a close correlation with the total ozone content, not the narrow lamina. The actual significance of narrow lamina still not clear throughout the manuscript.
120 121 122 123	We were interested in laminae of various sizes because according to theory gravitational waves produce predominantly small size laminae. On the other hand, planetary waves are able to form also the large laminae.
124 125 126 127	2. Lines 47-48, needs relevant references (at least one), especially about the influence of waves on the laminae.
128 129	The references were given in the Introduction.
130	3. Line 75 from-> on ?
131 132	4. Line 76 for the approximating-> for approximating
133 134	This grammar mistakes were corrected
135 136	5. Line 125 partitioning of laminae-> partitioning laminae?
137 138	We can say partitioning of laminae or lamina partitioning but not partitioning laminae
139 140 141	6. Conclusion: as mentioned in the introduction, if the Teitelbaum method is suitable for central Europe? And how well?
142 143 144 145 146	Teitelbaum method was demonstrated for data at the Sodankyla (northern Finland) and this method was able to detect atmospheric waves in the ozone profile. Grant et al. (1998) used the same method for the tropical stations and this method brought reasonable results. So we suppose this method is suitable also for the stations in Europe, because we obtained results which were expected in the case of well working method,
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153	Assessing the role of planetary and gravity waves on the vertical structure of ozone over
154	midlatitudinal Europe
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162	Abstract
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164	Planetary and gravity waves play an important role in the dynamics of the atmosphere. They
165	are present in the atmospheric distribution of temperature, wind and ozone content. These
166	waves are detectable also in the vertical profile of ozone and they cause its undulation. One of
167	the structures occurring in the vertical ozone profile is laminae, which are narrow layers of
168	enhanced or depleted ozone concentration in the vertical ozone profile. They are connected
169	with the total amount of ozone in the atmosphere and with the activity of the planetary and the
170	gravity waves. The aim of this paper is quantifying these processes in the central Europe. We
171	compare the occurrence of laminae induced by planetary waves (PL) with the occurrence of
172	these induced by gravity waves (GL). We show that the PL are $3-5$ times more frequent than
173	the gravity wave ones. There is a strong annual variation of PL, while GL exhibit only a very
174	weak variation. With the increasing lamina size the share of GL decreases and the share of
175	PL increases. The vertical profile of lamina occurrence is different for small planetary wave
176	and gravity wave laminae. The trend of large lamina occurrence frequency is given by the
177	trend in PL, not by GL.
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179	Key words: ozone lamina; vertical ozone profile, planetary wave activity, gravity waves
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182	1. Introduction
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184	I nere are various structures in the vertical profile of ozone affected by the activity of the

planetary and gravity waves. Ones of them are narrow layers of the enhanced or depleted ozone 185 concentration in the ozone vertical profile, which are called ozone laminae. The first 186 investigation of these structures was made by Dobson (1973), who found that they occur 187 predominantly in a cold half of the year. The existence of laminae was confirmed by lidar and 188 189 satellite measurements (Bird et al., 1997, Orsolini et al., 1997, Kar et al., 2002). They were found also in water vapour in the stratosphere (Teitelbaum et al., 2000). The dynamics of the 190 191 stratosphere plays a crucial role in a lamina formation. This finding was confirmed by the ability 192 of dynamical models to capture these narrow layers (Manney et al., 2000, Orsolini et al., 2001). 193 The number of large laminae is strongly correlated with the total ozone content and it is the 194 reason why we have been interested in laminae (Krizan and Lastovicka, 2005).

The laminae are not only the indicator of the atmospheric ozone content but also they are connected with the gravity and planetary wave activity. Teitelbaum et al. (1995) developed a identification procedure which enable us to detect the planetary and gravity wave activity in the ozone vertical profile. In this paper we apply this method to ozone laminae and each lamina we sort to the one of the following groups: laminae induced by gravity wave activity (GL), by planetary wave activity (PL) and laminae which are neither induced by the gravity waves nor 201 by the planetary waves. Similar method was used by Grant et al., (1998) and Pierce and Grant 202 (1998) but only for the Wallops Island station. The aim of this paper is finding the 203 characteristics of GL and PL in central Europe in the period 1970-2016. At first we test if the 204 Teitelbaum method is suitable for central Europe. Next the annual variation of GL and PL is 205 examined. Then we explore the dependence of lamina composition on their size. We also 206 compare the vertical distribution of GL and PL. We deal with their trends. The content of this 207 paper is as follows: section 2 describes methods and data, section 3 gives results, in section 4 208 the results are discussed and the last section is conclusions.

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2. Methods and data

213 Now we shortly describe the lamina searching procedure. Each positive lamina consists of the 214 three main points: the lower minimum, the main maximum and the upper minimum. The depth 215 of lamina must be between 500 and 3500 m due to the vertical resolution of the ozonosondes 216 (lower limit) and due to the fact that the ozone lamina is a narrow layer of the enhanced ozone 217 concentration (upper limit). The size of laminae is given as a difference between the ozone 218 concentration in the main maximum and the average concentration from both minima. More 219 about the lamina searching procedure can be found in (Krizan and Lastovicka, 2004) and 220 (Lastovicka and Krizan, 2005).

221 The method used in this paper for the searching the activity of gravity and planetary waves in 222 the ozone profile is a modification of the methods given by Teitelbaum et al. (1995). Figure 1 223 (upper panel) shows the real ozone profile at Hohenpeissenberg on February 2, 1970. We use 224 the linear interpolation with the step 50 m for approximating the ozone profile with the high 225 vertical resolution. Then the 50 point moving average (2500 m in vertical) is applied to this real 226 profile to obtain the smooth profile. This smooth profile is also displayed in fig.1 (upper panel). 227 The same procedure is applied to the potential temperature and the results are given in fig. 1 228 (lower panel). In the next step we compute the differences between the high resolution profile 229 and the smooth profile for the ozone partial pressure (fig 2 upper panel) and the potential 230 temperature (fig 2 lower panel). The differences are much higher for the ozone profile than for 231 the potential temperature profile. The differences in the vertical gradients of the ozone partial 232 pressure and the potential temperature must be taken into account. So we must apply the 233 following correction factor to the potential temperature perturbations:

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 $R(z) = [(1/O_{3avg})^*(dO_3/dz)]^*[(1/O_{avg})^*(dO_3/dz)]$ (1, 1)

238 where $O_{3avg}(\Theta_{avg})$ is the average ozone partial pressure (potential temperature) profile in the layer with the width dz. The vertical distribution of this correction is given in fig.3 (upper 239 240 panel). The correction is the highest in the lower stratosphere where the vertical gradient of 241 ozone is strong. Above 20 km we observe the negative values of this factor, which is 242 predominantly given by the negative gradient of the ozone partial pressure and the strong 243 positive gradient of the potential temperature. When we multiply the potential temperature 244 perturbations with this correction, we obtain the perturbations, which are shown in fig. 3 (lower 245 panel). These new perturbations are not similar to that given in fig.2 –lower panel. In each point 246 of the high resolution ozone profile we compute the correlation coefficient between the ozone 247 perturbations and the scaled potential temperature perturbation up to 5 km above this point. 248 The vertical dependence of this correlation coefficient from the ground to the point which is 249 situated 5km below the highest ozone profile point is seen in fig.4. If the correlation coefficient 250 is greater than 0.7, the vertical ozone profile in this point is influenced by the gravity waves. In 251 fig 4 the correlations are higher than 0.7 at some altitudes above 5 km and below 15 km. If the lamina maximum is situated in this high correlation area, we conclude this lamina is induced 252 by the gravity waves. On the other hand, if these correlations are low (between -0.3 and 0.3), 253 254 we consider the ozone profile to be influenced by the planetary waves in this point (from 17 to 255 22 km on fig. 4) and again if there is a lamina maximum there we consider this lamina as the 256 one induced by the planetary waves. When the correlation coefficient is above 0.3 and below 257 0.7 or below -0.3 we are not able to evaluate what type of laminae is present and call them 258 indistinguishable laminae. The boundary values of correlation coefficients were taken from Teitelbaum et al. (1995) 259

We are going to apply this procedure to the following European midlatitudes stations: Hohenpeissenberg (Germany, 1970-2016, 5166 files), Payerne (Switzerland, 1970-2016, 5998 files), Uccle (Belgium, 1970-2015, 6221 files), Lindenberg (Germany, 1975-2013, 2380 files) and Legionowo (Poland, 1979-2016, 1728 files). These data were taken from WOUDC Toronto (http://woudc.org/archive/Archive-NewFormat/). During the research some problems with a vertical resolution of ozone profile were occurred and so at the end we exclude the data from the station Lindenberg. The Hohenpeissenberg data was used only for large laminae.

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270 **3. Results**

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3.1. Performance of method273

274 At first we must answer the question if the procedure used in the paper is successful in 275 partitioning of laminae to the groups. If the procedure is suitable, the number of the 276 indistinguishable laminae cannot be very high. The performance of this procedure is given in tab.1 for Hohenpeissenberg for each month and for all laminae regardless the size. The results 277 278 at the other stations are very similar. From this table we see that approximately 47 % of all 279 laminae are PL, while GL laminae formed about 10 % and the share of indistinguishable 280 laminae is about 43 %. It means more than 50 % of all laminae can be divided into the laminae 281 induced by the gravity or the planetary wave activity. So we can conclude this procedure is 282 successful in lamina partitioning, because nobody can expect only GL and PL will be present 283 and no indistinguishable laminae. Practically there is no yearly course in the lamina 284 composition.

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3.2. Vertical resolution and number of laminae287

288 At first we must look at the homogeneity of the sonde vertical resolution used in this 289 paper. The results are given in fig. 5. We see the resolution is not homogenous and the resolution 290 increases (vertical distance decreases) in the period 1970-2016. And thus we must ask the 291 question if this resolution change has effect on a number of laminae detected in the profile. We 292 have computed correlation coefficient between the yearly values of lamina number and vertical 293 resolution. If these correlations are significant the resolution influences the lamina number and 294 vice versa. We did the correlations for the following groups of laminae: small (<1 mPa), 295 medium (1-4 mPa) and large (>4 mPa). The results are shown in tab.1. The number of small 296 laminae is strongly correlated with vertical resolution. It means the numbers of small laminae 297 are affected by the resolution. With increasing size of laminae these correlations decrease. For 298 large laminae the results are station dependant. These results are a bit surprising because one 299 expects negative correlations of lamina number with resolution and these negative correlations 300 were observed only for small laminae. For the explanation of these results we must look at the

301 average lamina depth in small, medium, and large laminae (table 2), which was obtained for 302 the best vertical resolution (below 100 m). We can see the increase of lamina depth with 303 increasing size. When the depth of laminae is small (small laminae), the vertical resolution 304 strongly influences the lamina number, because with decreasing resolution the number of 305 detected laminae decreases. On the other hand, the average depth of large laminae is above the 306 worst vertical resolution (800 m- fig.5) and so the increasing resolution does not influence 307 significantly the number of detected laminae.

308 According to the resolution results we can select periods used in this paper because we 309 are interested also in small and medium laminae number of which is resolution dependant. To 310 exclude this dependence, the vertical resolution of sonde must be comparable or smaller than 311 the average depth of laminae and thus one can see (table 2) the maximal vertical resolution in 312 the case of small laminae must be 100 m and for medium laminae 500 m. The depth of large laminae is above the worst vertical resolution so the large lamina results are not resolution 313 314 dependant. Originally we considered also the station Lindenberg but it had to be excluded due 315 to large and variable vertical resolution. The station Hohenpeissenberg is suitable only for 316 several years after 2010. Only the stations Payerne and Uccle have suitable vertical resolution 317 in the period 1990-2016 and the station Legionowo in the period 1995-2016. Because we must 318 do compromise between the quality and amount of data we take into account only these three stations in the period 1995-2016 for the small and medium laminae and the Hoheinpeissenberg 319 320 data for the large ones.

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323 **3.3.** Annual variation of laminae induced by the gravity and the planetary wave activity 324

Figure 6 shows the annual variation of the number of laminae larger than 2 mPa for GL and PL at all stations used in this paper. The group of lines with the strong annual variation with maximum in winter and minimum in summer/autumn are PL while the lines with the only very weak variation belong to GL. This different behaviour of the annual variation is the evidence that the both type of laminae are formed by different processes.

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332 3.3. Dependence of lamina type on the size of laminae

333 334 In this section we deal with the lamina type occurrence frequency in the selected classes of 335 lamina size. The laminae were sorted to the following groups: small (<1 mPa), medium size (1-4 mPa) and large (>4 mPa) and in each group we found the occurrence frequency of different 336 337 types of laminae. The results are presented in fig.7. The results are almost identical for all 338 stations. The share of GL is decreasing with the increasing size and the opposite is true for PL. 339 The performance of used procedure increases with the increasing lamina size (the share of 340 indistinguishable laminae decreases). The gravity waves are able to produce predominantly 341 small laminae, while the planetary waves produce also the large ones. Similar results were also 342 obtained by Teitelbaum et al. (1995).

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3.4. Vertical dependence of the occurrence of advection and gravity wave laminae

Now we examine the altitudinal dependence of occurrence of GL and PL at the stations
used in this paper for all seasons. March, April and May form spring, June, July, August are
summer months, September, October and November are the autumn ones and December,
January and February is winter. We divided the ozone vertical profile into 2 km wide intervals
and in each interval we search for the lamina occurrence. The results are displayed as the

351 percentage of all laminae which occur in the individual altitude interval. We grouped laminae 352 into two groups: small (<2 mPa) and large (>2 mPa) and in each group we are searching for the 353 lamina occurrence. The results are displayed only for the station Uccle, because at the other 354 stations the results are similar. The winter results are given in fig. 8 for the large (upper panel) 355 and the small (lower panel) laminae. The large laminae have similar behaviour both for GL and 356 PL. Their maximal occurrence is observed in the lower stratosphere and there are no large 357 laminae in the troposphere. On the other hand, the occurrence of the small laminae is different. 358 GL have maximal occurrence in the troposphere where the occurrence of PL is small. Small 359 PL have the maximal occurrence in the lower stratosphere, where the small gravity wave 360 laminae are rare. In the troposphere there is local minimum in small PL and the main maximum 361 in the small gravity wave occurrence. Spring (fig.9) behaviour of the lamina occurrence is 362 similar to the winter one. In summer (fig.10) the large GL have broad stratospheric maximum and the smaller maximum is observed in the troposphere. Large GL have sharper stratospheric 363 364 maximum and they are very little present in the troposphere. Small PL maximum is observed 365 in the stratosphere, while the small GL have tropospheric maximum. At the nearly same height we observe local minimum in small GL and maximum in the gravity wave ones. In autumn 366 (fig.11) the behaviour of large laminae is a bit similar to the summer one and the main maximum 367 368 in occurrence of small PL is higher than that of the laminae induced by the planetary wave.

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3.5 Trend of the large laminae

373 Now the long-term of the large laminae occurrence (larger than 4 mPa) is investigated. 374 The results are shown in fig. 12. A change in the trend of the PL in the mid-1990s is seen. 375 Before the mid-1990s the negative trend is observed, while after this point the positive one is 376 present. This fact confirms the findings of Krizan and Lastovicka (2006). But this is not the 377 main massage of this paper. The main massage of this paper concerning the trend is the 378 following: we observe a huge difference in the long- term trend between GL and PL: trend of 379 PL has the sharp change in the mid-1990s, which is confirmed by significant negative trend of 380 the lamina number before 1995 and the significant change of this trend in 1995 (tab.4), while 381 GL has a bit smaller insignificant negative trend before 1995 and insignificant trend change in 382 1995. So in the case of PL we observe change in trend in 1995, but in the case of GL no change 383 occurred. The piecewise regression model is suitable only for PL. On the other hand, the most 384 proper model for GL is the standard regression for the period 1970-2016. In this case we obtain 385 significant negative trend of lamina number about 17 % / decade.

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3894. Discussion390

391 We found the occurrence frequency of PL to be about 4-6 times larger than that of GL. 392 The most frequent way of formation of the laminae induced by planetary waves is vertically 393 different advection of air with the various ozone content (Manney et al., 2000). Tomikawa et 394 al. (2002) proposed as one of lamina formation mechanism vertical shear of the subtropical jet. 395 In these processes we observe transformation of the horizontal gradient of the ozone 396 concentration into the vertical one. The air with the high ozone concentration comes to the 397 central Europe in winter from the edge of the polar vortex (Orsolini et al., 2001). On the other 398 hand, the low ozone air has its origin inside the polar vortex and it is transported to the mid 399 latitudes (Reid and Vaughan, 1991) or it is the air from the low latitudes where ozone 400 concentration is low (Orsolini et al., 1995).

401 The strong source of gravity waves is orography (Smith et al., 2008), especially passing 402 the air through a mountain range when the gravity waves occur in the downwind side of the 403 ridge. For stations used in this paper the most important mountains are the Alps. These stations 404 are situated in a such way during prevailing west winds they are not on the leeward side of the 405 Alps and the share of gravity wave laminae are practically the same for all stations. The same 406 is true for the laminae induced by planetary waves. In this case all stations are practically under 407 the same conditions. So we cannot expect large interstation differences in lamina partitioning. 408 It will be reasonable to do this investigation at the stations which lie on the leeward side of 409 mountains or at stations which are in hot spots of the gravity wave activity (Sacha et al., 2016). 410 The other sources of the gravity waves are jet stream and convection (Guest et al. 2000; Yoshiki 411 et al. 2004). Their conditions are the same for all stations used in this study. In the troposphere 412 the stratosphere-troposphere exchange may cause the positive laminae and in the stratosphere 413 this exchange may lead to formation of negative laminae (Kritz, 1991).

414 Laminae greater than 2 mPa occur very predominantly in the stratosphere where the 415 ozone concentration is high. When the ozone concentration is high, the probability of large 416 lamina formation increases. The confirmation of this rule is also the yearly course of PL where 417 the maximal occurrence is observed when the ozone concentration is the highest (winter and 418 spring). On the other hand, in the troposphere we observe neither the PL large laminae nor the large GL due to small ozone concentration. Similarly, we observe less large PL in the 419 420 stratosphere in summer and fall. This dependence of the lamina occurrence on the background 421 ozone concentration is valid only for PL, not for the gravity wave ones.

For the laminae smaller than 2 mPa the situation is different. We observe the differences in the vertical distribution of PL and GL. In winter the maximal occurrence is observed in the lower stratosphere in the case of PL, while gravity wave laminae have its occurrence maximum in the tropopause. In spring the small GL maximum lies lower than in winter. In summer the occurrence distribution has bimodal structure with one maximum in the troposphere and the other one in the stratosphere. In fall the stratospheric mode is dominant.

In summer and fall there is no polar vortex. Vortex remnants (Durry et al., 2005) may
form the positive laminae in the stratosphere while the advection of air from low latitudes (Koch
et al., 2002) creates layers with the low ozone concentration.

In the troposphere the situation is different. Positive laminae are created by various processes: the stratosphere-troposphere exchange (Manney et al., 2000), the advection of polluted air from the boundary layer (Oltmans et al, 2004; Collete et al., 2005) or in situ ozone production (Li et al., 2002). Tropospheric gravity waves occur predominantly in the transition region from the troposphere to the stratosphere where there is a strong change in the atmospheric stability

437 Our paper is based on the lamina searching procedure introduced by Teitelbaum et al. 438 (1995). In their paper no climatological results are presented. They illustrated the method for 439 partitioning of laminae for several case studies. The goal of our paper is to use this method for 440 obtaining the climatological results from the mid-Europe ozonosonde stations. Similar 441 searching method was used by Grant et al. (1998) and Pierce and Grant (1998) but for tropical 442 and low latitudes stations. The authors found rare occurrence of PL and majority of laminae 443 was induced by gravity waves. We found more PL compared to the gravity induced ones, 444 because our investigation was done in middle latitudes, not in the low and tropical ones. The 445 activity of planetary waves is stronger in mid latitudes compared to the low and equatorial ones. 446

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449 **5.** Conclusions

451	The main results of this paper are:
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453	• The most often the laminae are induced by the planetary wave activity (45-50 %),
454	following by the indistinguishable ones (about 40 %). The share of the gravity wave
455	laminae is about 10 %.
456	• There is a pronounced annual variation in the occurrence frequency of PL, while there
457	is no such variation for GL
458	• With increasing lamina size the share of gravity wave and indistinguishable laminae
459	decreases while the share of the planetary wave laminae increases.
460	• The vertical distribution of lamina number for large laminae has maximum in the
461	stratosphere while the distribution of small laminae is type and season dependant.
462	• There are huge differences in trend patterns of PL and GL in the period 1970-2016
463	• There are huge unreferees in trend patterns of the and Oh in the period 1970 2010.
464	Competing interests
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466	The author declare that he has no conflict of interest
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474	Acknowledgement
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476	Support by the Grant Agency of the Czech Republic via Grant 18-01625S is acknowledged.
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479	References
480	
481	Bird, J.C., Pal, S.R., Carswell, A.I., Donovan, D. P., Manney, G.L., Harris, J.M., and Uchino,
482	O.: Observations of ozone structures in the Artic polar vortex. J. Geophys. Res., 102, D9,
483	10,785-10, 800, 1997.
484	Collete, A. and Ancellet, G.: Impact of vertical transport processes on the tropospheric ozone
485	lavering above Europe. Part II: Climatological analysis of the past 30 years. Atmos. Environ.
486	39, 5423-5435, 2005.
487	Dobson, G., M., B.: The laminated structure of the ozone in the atmosphere. Quart. J. R. Met.
488	Soc. 99, 599-607, 1973.
489	Durry, G. and Haucherone, A.: Evidence for long-lived polar vortex air in the mid-latitude
490	summer stratosphere from in situ laser diode CH4 and H2O measurements. Atmos. Chem.
491	Phys. 5, 1697-1472, 2005.
492	Guest, F. M., M. J. Reeder, C. J. Marks, and D. J. Karoly: Inertiagravity waves observed in
493	the lower stratosphere over Macquarie Island, J. Atmos. Sci., 57, 737–752, 2000.
494	Grant, W.B., Pierce, R.B., Oltmans, S.J. and Edward, W.: Seasonal evolution of total and
495	gravity waves induced laminae in ozonosonde data in the tropics and subtropics, GRL., 25,11,
496	1863-1866, 1998.
497	Kar, J., Trepte, C.R., Thomason, L.W. and Zawodny, J. M.: Observations of layers in ozone
498	vertical profiles from SAGE II(v 6.0) measurements, Geoph. Res. Lett, 29, NO 10,
499	10.1029/2001GL014230, 2002.

- 500 Koch, G., Wernli, H., Staehelin, J. and Peter, T.: A Langrangian analysis of stratospheric
- 501 ozone variability and long-term trends above Payerne (Switzerland) during 1970-2001. JGR,
- 502 107, D19, 437, doi: 10.1029/2001JD001550, 2002.
- 503 Kritz, M.A., Rosner, S.W., Danielsen, E.F. and Selkirk, H.B.: Air mass origins and
- 504 troposphere to stratosphere exchange associated with mid-latitude cyclogenesis and
- tropopause folding inferred from ⁷Be measurements. J. Geophys. Res., 96, D9, 17,405-17,414,
 1991.
- 507 Križan, P. and Laštovička, J.: Definition and determination of laminae in ozone profiles.
- 508 Studia geoph. et geod., 48, 777-789, 2004.
- 509 Križan, P and Laštovička, J.: Trends in positive and negative ozone laminae in the Northern
- 510 Hemisphere. J. Geophys. Res., D 10107, doi: 10.1029/2004JD005477, 2005.
- 511 Laštovička, J. and Križan, P. : Trends in laminae in ozone profiles in relation to trends in
- 512 someother middle atmospheric parameters., Physics and Chemistry of the Earth, 31, 46-53, 513 2006
- 513 2006.
- 514 Li, Q. et al.: Stratospheric versus pollution influences on ozone at Bermuda: Reconciling past
- 515 analyses. JGR, 107, D 22, 4611, doi: 10.1029/2002JD002138, 2002.
- 516 Manney, G. L., Michelsen, H. A., Irion, F. W., Toon, G. C., Gunson, M.R. and Roche, A. E.:
- 517 Lamination and polar vortex development in fall from ATMOS long-lived trace gases
- 518 observed during November 1994. J. Geophys. Res., 105, D23, 29,023-29,038, 2000.
- 519 Oltmans, S. J., Johnson, B. J., Harris, J. M., Thompson, A. M., Liu, H. Y., Chan, C. Y.,
- 520 Vömel, H., Fujimoto, T., Brackett, V. G., Chang, W. L., Chen, J. P., Kim, J. H.,
- 521 Chan, L. Y. and Chang, H.W.,: Tropospheric ozone over the North Pacific from ozonosonde 522 observations: JGR, 109, D15801, doi: 10.1029/2003JD003466, 2004.
- 523 Orsolini, Y., Simon, P. and Cariolle, D.: Filamentation and layering of an idealized tracer by
- 524 observed winds in the lower stratosphere. Geoph. Res. Lett, 22, No. 7, 839-842, 1995.
- 525 Orsolini, Y.J., Hansen, G., Hoppe, U. P., Manney, G.L. and Fricke, K.H., : Dynamical
- 526 modeling of wintertime lidar observations in the Artic: Ozone laminae and ozone depletion.
- 527 Q.J.R. Meteorol. Soc., 123, 785-800, 1997.
- 528 Orsolini, Y.J., Hansen, G., Manney, G.L., Livesey, N. and Hoppe U.P.: Lagrangian
- 529 reconstruction of ozone column and profile at the Artic Lidar Observatory for Middle
- 530 Atmosphere Research (ALOMAR) throughout the winter and spring of 1997-1998. J.
- 531 Geophys. Res., 106, D 9, 10011-10021, 2001.
- 532 Pierce, R.B. and Grant, W.B.: Seasonal evolution of Rossby and gravity wave induced
- 533 laminae in oznosonde data obtained from Wallops Island, Virginia, Geoph. Res. Lett. 25,11,
- 534 1859-1862,1998.
- 535 Reid, S.J. and Vaughan, G.: Lamination in ozone profiles in the lower stratosphere, Q.J. R.
- 536 Met. Soc., 117, 825-844, 1991.
- 537 Sacha, P., Lilienthal, F., Jacobi, C., and Pisoft, P.: Influence of the spatial distribution of
- 538 gravity wave activity on the middle atmospheric dynamics, Atmos. Chem. Phys., 16, 15755-
- 539 15775, doi:10.5194/acp-16-15755-2016, 2016.
- 540 Smith, R.B., B.K. Woods, J. Jensen, W.A. Cooper, J.D. Doyle, Q. Jiang, and V. Grubišić:
- 541 Mountain Waves Entering the Stratosphere. *J. Atmos. Sci.*, **65**, 2543–2562, 542 https://doi.org/10.1175/2007JAS2598.1, 2008.
- 543 Teitelbaum, H., Moustaoui, M., Ovarlez, J. and Kelder, H.: The role of atmospheric waves in
- the laminated structures of ozone profiles at high latitude. Tellus, 48A, 442-455, 1995.
- 545 Teitelbaum, H., C. Basdevant, and M. Moustaoui: Explanations for simultaneous laminae in
- 546 water vapor and aerosol profiles found during the SESAME experiment, *Tellus*, 52A, 190-
- 547 202, 2000.

- 548 Tomikawa, Y., Sato, K., Kita, K., Fujiwara, M., Yamamori, M. and Sano, T.: Formation of an
- ozone lamina due to differential advection revealed by intensive observations. J. Geophys.
 Res., 107, D 10, 10.1029/2001/JD000386,2002.
- 551 Yoshiki, M., N. Kizu, and K. Sato: Energy enhancements of gravity waves in the Antarctic
- 552 lower stratosphere associated with variations in the polar vortex and tropospheric
- 553 disturbances, J. Geophys. Res., 109, D23104, doi:10.1029/2004JD004870,2004.



Figure 1: Real and smooth ozone (upper panel) and potential temperature (lower panel)
vertical profile at the Hohenpeissenberg from February 2, 1970.



Figure 2: Differences between real and smooth vertical profile from February 2 , 1970 for
 ozone (upper panel) and potential temperature (lower panel)





Figure 3: Vertical profile of potential temperature correction factor (upper panel) and vertical
 profile of differences between real and smooth potential temperature profile (lower panel) after
 correction.



Figure 4: The vertical profile of correlations between the corrected potential temperature 583 differences and the ozone differences from February 2, 1970 at Hohenpeissenberg. The red 584 vertical lines are the borders for the laminae induced by the planetary waves and the blue 585 vertical line is the border for gravity wave ones.







592 Figure 5: Long term evolution of average vertical resolution of profiles at the European593 ozonesonde stations.





598 Figure 6: The annual variation of the lamina number per ozone profile for PL (group of lines
599 with the strong variation) and for GL (group of lines with the weak variation) at the European
600 ozonosonde stations.



Figure 7: The dependence of the lamina composition on a lamina size for PL (upper panel), GL (middle panel) and indistinguishable laminae (lower panel) at the European stations (1-

607 GL (middle panel) and indistinguisha608 Payerne, 2 – Uccle, 3 – Legionowo)



612 Figure 8: The vertical dependence of the occurrence of the laminae induced by the gravity
613 waves and the ones induced by planetary waves at Uccle the period 1995-2016 in winter in
614 terms of percentage of all GL and all PL.

















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Figure 11: Trend of the number of laminae per ozone profile induced by the planetary waves (PL) and by the gravity waves (GL) for laminae greater than 4 mPa at Hoheinpeissenberg in

641 the period 1970-2016.

642

	January	February	March	April	May	June	July	August	Sept	Oct	Nov	Dec
advect	48	49	48	48	45	41	44	46	47	46	47	48
gravity	10	10	11	10	11	11	10	11	10	11	9	10
undist	42	41	41	42	44	48	46	43	43	43	44	42

644

Table 1: Monthly composition of laminae (%) at Hohenpeissenberg in the period 1970-2016
 (advect- advection laminae, gravity – gravity waves laminae, undist- undistinguishable
 laminae)

648

	<1mPa	1-4 mPa	>4 mPa
Hohenpeissenberg	-0.95 /-0.68	-0.57/0.55	-0.09/ 0.25
Payerne	-0.49/-0.37	-0.50/0.29	0.32/0.58
Uccle	-0.66/-0.61	0.57 /-0.07	0.00/0.16
Lindenberg	-0.79/-0.51	-0.88/-0.54	-0.76 /0.14
Legionowo	-0.81/-0.80	-0.77 /-0.07	0.31 /0.19

649

Table 2: Correlation coefficient of lamina number and average vertical resolution at the European mid latitudes stations from the period 1970-2016 (before slash - advective laminae, after slash gravity wave laminae). Significant correlation coefficient values are in hold

after slash – gravity wave laminae). Significant correlation coefficient values are in bold.

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654

655

	<1 mPa	1-4 mPa	>4 mPa
Hohenpeissenberg	198/203	733/1021	1895/2057
Payerne	112/144	486/597	1874/1803
Uccle	121/206	486/761	1832/1775
Legionowo	104/142	535/702	1909/1983

656

Table 3: Average lamina depth (m) in the selected lamina size intervals at the European midle latitude stations for the vertical resolution below 100m (before slash - advective laminae, after

659 slash – gravity wave laminae).

660

661

662

	till 1995	change in 1995	after 1995
PL	-18	36	18
GL	-16	-4	-20

663

Table 4: Results of piecewise linear regression applied to the yearly means of the large lamina
 number in the period 1970-2016 at the station Hoheinpeissenberg. The trend is expressed as a
 percentage per 10 years. Significant trends are expressed in bold.