Magnetometer in-flight offset accuracy for the BepiColombo spacecraft

Daniel Schmid¹, Ferdinand Plaschke¹, Yasuhiito Narita¹, Daniel Heyner², Johannes Z. D. Mieth², Brian J. Anderson³, Martin Volwerk¹, Ayako Matsuoka⁴,⁵, and Wolfgang Baumjohann⁷

¹Space Research Institute, Austrian Academy of Sciences, Graz, Austria
²Institut für Geophysik und extraterrestrische Physik, Technische Universität Braunschweig, 38106 Braunschweig, Germany
³The Johns Hopkins University Applied Physics Laboratory, Laurel, Maryland, USA
⁴Institute of Space and Astronautical Science, JAXA, Yoshinodai, Kanagawa, Japan
⁵Data Analysis Center for Geomagnetism and Space Magnetism Graduate School of Science, Kyoto University, Kyoto, Japan

Correspondence: Daniel Schmid (daniel.schmid@oeaw.ac.at)

Received: 17 January 2020 – Discussion started: 31 January 2020
Revised: 11 May 2020 – Accepted: 3 June 2020 – Published: 10 July 2020

Abstract. Recently the two-spacecraft mission BepiColombo launched to explore the plasma and magnetic field environment of Mercury. Both spacecraft, the Mercury Planetary Orbiter (MPO) and the Mercury Magnetospheric Orbiter (MMO, also referred to as Mio), are equipped with fluxgate magnetometers, which have proven to be well-suited to measure the magnetic field in space with high precision. Nevertheless, accurate magnetic field measurements require proper in-flight calibration. In particular the magnetometer offset, which relates relative fluxgate readings into an absolute value, needs to be determined with high accuracy. Usually, the offsets are evaluated from observations of Alfvénic fluctuations in the pristine solar wind, if those are available. An alternative offset determination method, which is based on the observation of highly compressional fluctuations instead of incompressible Alfvénic fluctuations, is the so-called mirror mode technique. To evaluate the method performance in the Hermean environment, we analyze four years of MESSENGER (MErcury Surface, Space ENvironment, G EOphysics and Ranging) magnetometer data, which are calibrated by the Alfvénic fluctuation method, and compare it with the accuracy and error of the offsets determined by the mirror mode method in different plasma environments around Mercury. We show that the mirror mode method yields the same offset estimates and thereby confirms its applicability. Furthermore, we evaluate the spacecraft observation time within different regions necessary to obtain reliable offset estimates. Although the lowest percentage of strong compressional fluctuations are observed in the solar wind, this region is most suitable for an accurate offset determination with the mirror mode method. 132 h of solar wind data are sufficient to determine the offset to within 0.5 nT, while thousands of hours are necessary to reach this accuracy in the magnetosheath or within the magnetosphere. We conclude that in the solar wind the mirror mode method might be a good complementary approach to the Alfvénic fluctuation method to determine the (spin-axis) offset of the Mio magnetometer.

1 Introduction

In October 2018, BepiColombo, a two-spacecraft mission of the European Space Agency (ESA) and the Japan Aerospace Exploration Agency (JAXA), was launched to explore Mercury (Benkhoff et al., 2010). One of the spacecraft is the Mercury Planetary Orbiter (MPO), which is a three-axis-stabilized satellite (quasi nadir pointing) to study the surface and interior of the planet (e.g., Glassmeier et al., 2010). The other is Mio (or Mercury Magnetospheric Orbiter, MMO), a spin-stabilized spacecraft (spin period of about 4 s) to investigate the magnetic field environment of Mercury (e.g., Hayakawa et al., 2004; Baumjohann et al., 2006). During the 7.2-year cruise phase, both orbiters are transported by the Mercury Transfer Module (MTM) as a single composite spacecraft. In late 2025, the composite spacecraft will ap-
proach Mercury, where the MTM separates from the other two spacecraft, which are captured into a polar orbit around the planet. As soon as Mio reaches its initial operational orbit of 590 km by 11 640 km above the surface, MPO also separates and lowers its altitude to its 480 km by 1500 km orbit.

The BepiColombo Mercury Magnetometers (MERMAG) constitute a key experiment of the mission; MERMAG consist of the fluxgate magnetometers onboard both MPO and Mio. The magnetometers will provide in situ data for the characterization of the internal field origin as well as its dynamic interaction with the solar wind (see for example Wicht and Heyner, 2014, for a discussion). To achieve this goal, accurate magnetic field measurements are thus of crucial importance. Therefore, the components of a linear calibration matrix M and an offset vector O need to be obtained, in order to convert raw instrument outputs \( B_{\text{raw}} \) to fully calibrated magnetic field measurements (see for example Kepko et al., 1996; Plaschke and Narita, 2016):

\[
B = M \cdot B_{\text{raw}} - O. \tag{1}
\]

Here, the matrix M transforms \( B_{\text{raw}} \) into a spacecraft-fixed orthogonal coordinate system. It comprises nine parameters: three scaling (gain) values of the sensor and an orthogonalization matrix, which is defined by the three angles that yield the magnetometer sensor directions with respect to the spacecraft reference frame (see for example Plaschke and Narita, 2016). The 3D offset vector \( O \), on the other hand, reflects the magnetometer outputs in vanishing ambient fields. These can be attributed to the instrument and also to the field generated by the spacecraft at the position of the magnetometer sensor. Frequent in-flight calibration of these offsets is necessary, as they are known to change over time. To calibrate the magnetometer, all 12 parameters need to be accurately determined. For spinning spacecraft (i.e., Mio) eight of the calibration parameters can be determined directly by minimizing periodic signatures in the de-spun magnetic field signal at the spin frequency and/or at the second harmonic (Kepko et al., 1996). The remaining four parameters (i.e., the absolute gain in the spin plane and along the spin axis, the rotation angles of the sensor around the spin axis, and the spin-axis offset) need to be determined differently. It should be noted that the gains and rotation angle become important at strong fields. In the solar wind and in the Hermean environment they may play a minor role in the magnetic field measurements accuracy in comparison to the offsets.

The following methods, which can also be applied to non-spinning spacecraft (i.e., MPO), are well established for the offset determination:

1. The first is cross-calibration of the magnetometer offset with independent magnetic field measurements from other instruments. The Magnetospheric Multiscale (MMS, Burch et al., 2016) mission uses independent measurements from the Electron Drift Instrument (EDI, Torbert et al., 2016) to cross-calibrate the spin-axis offset of the magnetometers (Nakamura et al., 2014; Plaschke et al., 2014).

2. The offset may also be determined in a characteristic region where the magnetic field is known. Goetz et al. (2016) used diamagnetic cavities to determine the magnetometer offsets of the Rosetta spacecraft mission (Glassmeier et al., 2007).

3. Spacecraft rotations about the spacecraft principle axes from time to time are also a well-established method to obtain magnetometer offsets. This has been done routinely with many missions, most recently including MAVEN (Connerney et al., 2015) for Mars and the Parker Solar Probe (Bale et al., 2016).

4. A common way to calibrate the magnetometer offset in-flight is to use (a) (nearly) incompressible or (b) compressible fluctuations of the magnetic field:

   a. A well-established method is to minimize the variance of the total magnetic field during the passage of (nearly) incompressible (Alfvénic) variations in the (pristine) solar wind (Belcher, 1973; Heggock, 1975). Pure Alfvénic fluctuations are strictly incompressible and circularly polarized. They are characterized by changes in the magnetic field components while the magnitude of the field stays constant. Particularly in inhomogeneous media such simple classifications are found to be impossible (see Tsurutani et al., 2018, for a review). However, in the solar wind the fluctuations of the magnetic field strength (compressible part) are weak compared to the strong fluctuations of the magnetic field vector direction (Khabibulkhamanov and Summers, 1997). By minimizing the changes of the observed total magnetic field of such fluctuations it is therefore possible to adjust the magnetometer offsets with high precision. In Fig. 1a a schematic illustration of (nearly) incompressible Alfvénic fluctuations is shown. The changes of the total magnetic field, \( \delta |B| \), observed by a virtual spacecraft fluctuations are minimal (shown in the bottom).

   b. Recently Plaschke and Narita (2016) introduced an alternative magnetometer offset calibration technique on the basis of the observation of compressible fluctuation, the so-called mirror mode method, which does not require pristine solar wind measurements. The idea is that for strongly compressible fluctuating fields (e.g., mirror modes) the maximum variance direction of the magnetic field should be nearly parallel to the mean (background) magnetic field (Tsurutani et al., 2011). Mirror mode structures are compressional, non-propagating structures which have been observed in various space plasma environments like in the...
solar wind (e.g., Winterhalter et al., 1994), planetary magnetosheaths (e.g., Tsurutani et al., 1982) and even near the magnetic pileup boundaries of comets (e.g., Glassmeier et al., 1993). They are typically characterized by the anticorrelation between the magnetic field strength and density fluctuations, with little or no change in the magnetic field direction across the structure (see for example Winterhalter et al., 1994). Figure 1b shows a sketch of such a mirror mode structure. A virtual spacecraft crossing the structure observes perturbations along the mean magnetic field direction, \( \delta B \). Any differences between the maximum variance direction (evaluated from minimum variance analysis, MVA; for example; Sonnerup and Scheible, 1998) of the magnetic field and the mean (background) magnetic field direction can thus be used for the magnetometer offset determination. Note that although this method is called the mirror mode method, observations of highly compressional waves other than mirror mode structures with strong \( \delta B / |B| \) ratios without \( B \) direction changes are sufficient for their application. An advantage of this method is that compressible waves are ubiquitous in the magnetosphere and magnetosheath. Therefore this method can also be applied to calibrate the magnetometers of spacecraft which remain within the magnetosphere (e.g., MPO).

In this paper we test the applicability of the mirror mode method in different regions (plasma environments) around Mercury, based on 4 years of MESSENGER (MErcury Surface, Space ENvironment, GEophysics and Ranging, Solomon et al., 2007) fluxgate magnetometer (FGM, Anderson et al., 2007) data, which were calibrated using time intervals of (nearly) incompressible Alfvénic fluctuations in the solar wind. For reasons of simplicity (and because for Mio only the spin-axis offset needs to be evaluated) we apply the 1D mirror mode method (see Plaschke and Narita, 2016) and assume that two offset components are already accurately determined. We test to which degree the 1D mirror mode method yields vanishing offsets as expected when using calibrated data as input. We also address the question of how much time a spacecraft needs to spend in each individual plasma region until the magnetometer offset can be determined to a specific accuracy from intervals containing compressional magnetic field fluctuations. The results obtained in this paper enable us to assess whether the mirror mode method would be a useful tool to accurately determine the offset of BepiColombo’s magnetometer.

2 Data and methodology

We use orbital magnetic field data from the MESSENGER FGM (Anderson et al., 2007) between March 2011 and April 2015. The polar orbit of MESSENGER was highly elliptical; the initial altitude ranged between 200 and 15,000 km from Mercury’s surface. With an initial orbital period of 12 h, MESSENGER crossed the magnetopause and bow shock four times within 24 h. We use the 1 Hz calibrated magnetic field data in spacecraft coordinates, where the \( y \) axis is nominally in the anti-sunward direction (radially away from the Sun), the \( z \) axis points towards the payload adapter ring at the bottom of the spacecraft, and the \( x \) axis completes the right-handed coordinated system. If not noted otherwise, we use the magnetic field components \( \{B_x, B_y, B_z\} \) throughout this paper in these coordinates. MESSENGER was a three-axis-stabilized spacecraft, and its magnetometer offsets were routinely corrected using time intervals of Alfvénic fluctuations in the solar wind. We therefore use the MESSENGER magnetometer data as a calibration standard in our magnetometer offset study. To perform a test of the mirror mode method against the MESSENGER magnetometer data, we determine the 1D offset along the \( z \) axis, \( O_z \), in the same way as introduced by Plaschke and Narita (2016): within strongly compressible mirror mode structures, the magnetic field variation maximum variance direction, \( \ell = [\ell_x, \ell_y, \ell_z] \), obtained from a principle component analysis (MVA; Sonnerup and Scheible, 1998) should be reasonably aligned with the mean field direction, \( B = [B_x, B_y, B_z] \). Under the assumption of alignment between the maximum variance direction and the mean field direction Plaschke and Narita (2016) showed that the offset \( O_z \) can be derived by the following:

\[
O_z = B_{xy}(\tan \theta_B - \tan \theta_x), \tag{2}
\]

where \( \theta_B = \arctan(B_z / B_{xy}) \) is the elevation angle of the magnetic field to the \( x-y \) plane and \( \theta_x = \arctan(\ell_z / \ell_{xy}) \) is the elevation angle of the maximum variance direction to that plane. \( B_{xy} = \sqrt{B_x^2 + B_y^2} \) and \( \ell_{xy} = \sqrt{\ell_x^2 + \ell_y^2} \) are the magnetic field and maximum variance within the \( x-y \) plane.
Since a single offset estimate alone might not be very accurate, a statistically significant offset $O_{zf}$ should be determined by finding the maximum of the probability density function $P$ computed by the kernel density estimator (KDE) method with the Gaussian kernel from a sample of individual offset estimations $O_z$:

$$P(O_z^*) = \frac{1}{\sqrt{2\pi Nh}} \sum_{n=1}^N \exp \left[ -\frac{1}{2} \left( \frac{O_z^* - O_{z,n}}{h} \right)^2 \right].$$ (3)

Here $N$ denotes the number of individual offset estimates, $O_{z,n}$, $h$ is a smoothing parameter which denotes the bandwidth of the KDE, and $O_{zf}$ is the offset value at which the probability $P$ is at maximum, $\max\{P(O_z^*)\} = P(O_{zf})$. We use the method introduced by Silverman (1986), to determine an optimal bandwidth $h = c_1 \cdot \sigma(O_{z,n}) N^{-\frac{1}{5}}$, with $c_1 = 1.06$ and $c_2 = -1/5$. The symbol $\sigma$ denotes the standard deviation of the offset-value distribution.

3 Data analysis and results

A basic condition for the mirror mode method is the availability of compressional magnetic field fluctuations. As a first step we use the magnetometer data to estimate the occurrence rate of compressional fluctuations in the different plasma environments around Mercury. Then we compare the offset determination by the mirror mode method with the Alfvén wave method in terms of accuracy. Finally, we evaluate the number of offset samples that are sufficient for in-flight calibration with the mirror mode method, given a minimum required accuracy. Thus, we subsequently determine how many hours the spacecraft (i.e., Mio and/or MPO) need to spend in different plasma environments to obtain reliable offset values.

3.1 Occurrence rate of compressional fluctuations

Mercury’s plasma environment is highly dynamic and home to a plethora of wave modes and fluctuations (Russell, 1989; Boardson et al., 2009, 2012; Sundberg et al., 2015). We separate the MESSENGER orbit segments into solar wind, magnetosheath, and magnetospheric parts, based on an extended boundary data set from Winslow et al. (2013) (see Supplement in Philpott et al., 2020), in order to distinguish occurrence rates of (compressional) fluctuations and offset estimate accuracies by region. We characterize the observed fluctuations in the following way: the magnetic field data are divided into overlapping 30 s intervals shifted by 15 s. Within each sub-interval the magnetic field measurements are transformed into a mean-field-aligned (MFA) coordinate system, where the parallel component, $\hat{b}_|| = B_0/|B_0|$, is given by the average magnetic field within the 30 s interval, $B_0 = [B_{x,0}, B_{y,0}, B_{z,0}]$, and the perpendicular components in this coordinate system are chosen to be $\hat{b}_\perp,1 = [0, -B_{z,0}, B_{x,0}]/||(0, -B_{z,0}, B_{x,0})||$ and $\hat{b}_\perp,2 = \hat{b}_\perp,1 \times \hat{b}_||$. Then the maximum variance direction, $\ell$, of the transverse (perpendicular) magnetic field components is evaluated by a 2D MVA (Sonnerup and Scheible, 1998). The difference between the maximum and minimum values of the magnetic field in this maximum variance direction yields the measure for the transverse fluctuations $\delta B_{\perp,\ell} = B_{\perp,\ell}^\text{max} - B_{\perp,\ell}^\text{min}$. The compressional fluctuations, on the other hand, are essentially represented by fluctuations in the magnetic field magnitude $\delta |B|$, given by the difference between the maximum and minimum value of the magnetic field magnitude within the 30 s interval. Subsequently, to determine whether the compressional or the transverse part is dominant, we further define the compressibility index $Q^\pm$ as

$$Q^\pm = \log_{10} \left( \frac{\delta |B|}{\delta B_{\perp,\ell}} \right),$$ (4)

which is positive (negative) in case of dominant compressional (transverse) fluctuations.

Figure 2a displays the occurrence rate (normalized to unity when integrated over the whole domain) of fluctuations of the magnetic field magnitude (compressible sense) relative to the mean field in various plasma regions: in the solar wind (in red), in the magnetosheath (in green), and in the Hermean magnetosphere (in blue) for the MESSENGER magnetometer data for 4 years. Most of the large-amplitude fluctuations are observed in the magnetosheath. Such a result is not surprising, since the magnetosheath is characterized generally by a highly turbulent plasma with enhanced magnetic field variations. Panel b shows the normalized occurrence rate of the compressibility index $Q^\pm$ of selected intervals with large-amplitude fluctuations ($\delta |B|/B > 0.3$). The shifts of the maximum occurrence rates to negative $Q^\pm$ values indicate that statistically transverse fluctuations were dominant. However, there is also a significant number of time intervals with compressional fluctuations observed. The integrated occurrence rates are shown in Table 1. The first column of Table 1 shows the percentage of the sub-intervals where the magnetic field magnitude fluctuations $\delta |B|/B$ were larger than 0.3. The percentage in the second column shows how many of the intervals with large-amplitude fluctuations ($\delta |B|/B > 0.3$) are dominated by strongly compressional fluctuations, where $\delta |B|/\delta B_{\perp,\ell} > 2$ ($Q^\pm > 0.3$). The last column reflects the percentage of intervals that include these strongly compressional fluctuations out of all intervals of MESSENGER observations in the respective regions.

About 25 % of the time enhanced magnetic field fluctuations ($\delta |B|/B > 0.3$) are observed in the magnetosheath and 10.6 % of these fluctuations had a compressional nature. That means that from the total observation time in the magnetosheath, 2.6 % of the time MESSENGER observed strongly compressional fluctuations. A smaller fraction of the MESSENGER data set represents compressible dominated time intervals of the solar wind (0.4 %) and the Hermean magnetosphere (1.7 %). This suggests that the MESSENGER mag-
Table 1. The first column gives the percentage of the sub-intervals where the magnetic field magnitude fluctuations, $\delta |B|/\bar{B}$, are larger than 0.3 in each respective region. The percentage in the second column shows how many intervals of the sub-intervals with $\delta |B|/\bar{B} > 0.3$ are dominated by compressional fluctuations ($Q^\pm > 0.3$). The last column reflects the percentage of compressional fluctuations MESSENGER observes in total in each region.

<table>
<thead>
<tr>
<th>Region</th>
<th>Total</th>
<th>$Q^\pm &gt; 0.3$</th>
<th>Compressible fluctuations in total</th>
</tr>
</thead>
<tbody>
<tr>
<td>Solar wind</td>
<td>7.9%</td>
<td>5.4%</td>
<td>0.4%</td>
</tr>
<tr>
<td>Magnetosheath</td>
<td>24.6%</td>
<td>10.6%</td>
<td>2.6%</td>
</tr>
<tr>
<td>Magnetosphere</td>
<td>8.0%</td>
<td>21.6%</td>
<td>1.7%</td>
</tr>
</tbody>
</table>

netosheath data may be best suited for the offset calibration with the mirror mode method.

3.2 Test of the mirror mode method

Using Eq. (2) we determine the offset $O_z$ within each 30 s time interval. To ensure reliable offset estimates, the same requirements introduced by Plaschke and Narita (2016) have to be fulfilled within each window:

$$\frac{\delta B_z}{B_{zy}} = \frac{B_{zxy}^{\text{max}} - B_{zxy}^{\text{min}}}{B_{zxy}^{\text{mean}}} > 0.3,$$

where mirror modes are characterized by large magnetic field fluctuations (see Price et al., 1986; Schmid et al., 2014). $B_{zy}^{\text{max}}$, $B_{zy}^{\text{min}}$, and $B_{zxy}^{\text{mean}}$ are the maximum, minimum, and average magnetic field values in the spacecraft $x-y$ plane, respectively. Note that we only consider the $x-y$ plane components, because the $z$ components are subject to an a priori unknown offset $O_z$.

$- \phi < 20^\circ$, where $\phi$ is the angle between the maximum variance, $\ell$, and magnetic field, $B$, directions in the $x-y$ plane. Note that only the $x-y$ plane components is used. Plaschke and Narita (2016) derived this requirement and threshold from Lucek et al. (1999), who identified mirror modes using the angle between maximum variance and magnetic field direction.

$- |\theta_B| < 30^\circ$ and $|\theta_\ell| < 30^\circ$, so that both maximum variance and magnetic field directions point closer to the $x-y$ plane.

Out of 4 years of MESSENGER data these requirements are met for 3.0%, 2.1%, and 0.4% of the intervals within the magnetosphere, magnetosheath, and the solar wind, respectively (see also first column of Table 2). It should be noted that these numbers do not have to match the numbers shown in Table 1, since the criteria are different. The numbers in Table 1 reflect the occurrences of strongly compressional fluctuations and here they reflect the occurrences of intervals with mirror-mode-like characteristics. Interestingly, however, the numbers in the solar wind and the magnetosheath are similar but differ considerably in the magnetosphere. Using Eq. (3) the probability density function $P$ is computed for each region from all determined best-estimate offset $O_z$ in that region. Under the assumption that the data are perfectly calibrated and the mirror mode method works accurately, the offset $O_{zf}$ should vanish and $P$ should be highly symmetric with the peak around 0 nT.

Figure 3 shows the probability density functions $P$ based on the magnetosphere (red), magnetosheath (green), and solar wind (blue) offset estimates. Indeed, the best estimate of the $z$-component offset $O_{zf}$ is around zero for all three regions. It is noteworthy that the standard deviation of individual offset estimates, $O_{z,n}$, is smallest in the solar wind. This is actually not surprising, since the intervals in the solar wind
where the above requirements are fulfilled are well-suited for this calibration method due to their clear compressional signature in the rather low background magnetic field. In Table 2 the total number of intervals \( \langle O_z \rangle \) and their best-estimate offset \( O_{zf} \) are given. Further, the arithmetic mean of the individual offsets in each interval \( O_{z,n} \) with their standard deviations and error are summarized.

The average \( \langle O_{z,n} \rangle \) is found to be quite close to the final offset estimate \( O_{zf} \). Note that, although the standard deviation of the individual offsets \( O_{z,n} \) might be large, a larger number of samples or events helps lower the value of the standard deviation of the mean offset \( \langle O_{z,n} \rangle \) (standard error in Table 2).

### 3.3 Effect of sample size on offset accuracy

In the following we check how the best-estimate offset \( O_{zf} \) is affected by the number of \( O_{z,n} \) estimates. Therefore, from each region, 1 to 20,000 offset estimates \( N = \{ x \times 10^y \mid x \in [1, 9], y \in [0, 4] \} \), max \( N \) = 20,000 are randomly picked and their best-estimate offset \( O_{zf} \) determined. For each \( N \), this sequence is repeated 1000 times (bootstrapping). The standard deviation of the 1000 \( O_{zf} \) \( \sigma (O_{zf}) \) subsequently reflects the uncertainty of the determined \( O_{zf} \).

Here we use the 2\( \sigma \) interval, to evaluate the uncertainty of the best-estimate offset \( O_{zf} \) with 95% confidence. The standard error \( \epsilon \) of the 2\( \sigma \) uncertainty of the best-estimate offset, \( 2 \sigma (O_{zf}) \) from the bootstrapping is given by \( \epsilon \approx 2 \sigma (O_{zf}) / \sqrt{2(M-1)} \) (see for example Squires, 2001). Here \( M = 1000 \), yielding a relative standard error \( \epsilon \) of 4.3% of the evaluated 2\( \sigma (O_{zf}) \). Figure 4 displays 2\( \sigma (O_{zf}) \) as a function of the number \( N \) of the offset estimates used, \( O_{z,n} \).

The dashed lines in Fig. 4 mark offset accuracies at 0.5 and 1.0 nT. It is visible that the offset accuracy increases with the number of samples following a power law. However, below \( \sim 0.5 \) nT considerably more samples are needed to improve the offset accuracy, which could indicate the lower limit of the offset accuracy of the MESSENGER magnetometer determined by the Alfvénic method. As can been seen in Fig. 4, this divergence is only observed in the solar wind, because it is the only region where a significant number of \( O_{zf} < 0.5 \) nT are obtained. The solid lines in Fig. 4 depict the regression lines of a linear least squares fit of the offset accuracies above 0.5 nT and clearly show the power law behavior between these offset accuracies and the number of samples. Table 3 shows the fitting parameters of the regression lines of the power laws,

\[
\log_{10} [2 \sigma (O_{zf})] = \log_{10} [a] + k \cdot \log_{10} (N).
\]  

(5)

Here \( a \) denotes the 2\( \sigma \) confidence of the best-estimate \( O_{zf} \) determined from only one offset \( O_z \), and \( k \) represents the spectral index of the power law.

From the least squares approximation we calculate the minimum number \( N \) of \( O_{z,n} \) estimates to reach an offset accuracy of 0.5 and 1.0 nT. These numbers are multiplied by 30 s, the sliding window time interval, yielding the time ranges from which reliable \( O_{z,n} \) estimates are obtained (i.e., where the criteria from above are satisfied). These time ranges multiplied with the probability to observe these time intervals (see first column in Table 2), finally give the time lengths that a spacecraft needs to spend in the solar wind, magnetosheath, or magnetosphere in order to determine the offset with an accuracy better than 0.5 nT or 1.0 nT. Table 4 shows the minimum number \( N \) of \( O_{z,n} \) estimates with their corresponding time ranges and the necessary spacecraft observation time required to reach these accuracies.
of the power law.Offsets need to stay constant to within 0.05 nT. However, it is important to note that during this time the instrument offsets need to stay constant to within 0.5 nT, otherwise the intrinsic offset drifts over time would limit the achievable precision, independently from the amount of solar wind data. Specifically in the magnetosheath and within the magnetosphere this might be a more important limiting factor, since several thousands of hours of data in those regions are needed to ensure that the offset uncertainty is less than 0.5 nT (see Table 4). Figure 4 shows that the offset accuracies down to 0.5 nT diminish with the number of necessary offset estimates and follow a power law. However, below 0.5 nT the power law correlation flattens (spectral index becomes smaller), which might indicate that the lower limit of the offset accuracy of the calibrated magnetometer data itself has been reached. To ensure, for example, the offset uncertainty to be of the order of 1.0 nT, 60 h of magnetic field data in the solar wind would be sufficient, but hundreds of hours of magnetosheath or magnetosphere data are needed.

In the following we consider possible implications of these results on the BepiColombo mission:

- **Mio (Mercury Magnetospheric Orbiter, MMO):** Since Mio’s orbit allows the spacecraft to encounter the solar wind (particularly near Mercury’s perihelion), the mirror mode method would serve as a reasonable complementary approach to the Alfvénic fluctuation method. Furthermore, Mio is a spin-stabilized spacecraft (with a period of about 4 s) and thus two of the three offsets in the de-spun plane can be directly determined by minimization of spin tone in the data. Since the two spin-plane offsets can be determined very accurately (Kepko et al., 1996), the mirror mode method presented in this study can thus be directly applied to determine the remaining offset (spin-axis offset) very precisely (see Plaschke and Narita, 2016). However, the mirror mode method assumes the time independence of the offset properties during the fluctuation measurements. Offset drifts over time, e.g., due to temperature changes over orbital periods, will be the limiting factors for the achievable accuracies.

- **MPO (Mercury Planetary Orbiter):** The MPO spacecraft, on the other hand, is three-axis stabilized, which should diminish the offset determination accuracy due to the two additional degrees of freedom, and the 1D mirror mode method presented in this study cannot be applied directly. Based on the principles of the 1D mirror method, Plaschke et al. (2017) extended the method to three dimensions and introduced the so-called 3D mirror mode method. This method has successfully

---

**Table 2.** Number of samples (time intervals) used in the statistical study for the offset determination and various offset estimates in units of nanotesla (best-estimate, mean, standard deviation, and standard error).

<table>
<thead>
<tr>
<th>Region</th>
<th>Number of samples (% of total)</th>
<th>Best-estimate $O_{zf}$ (nT)</th>
<th>Mean $\langle O_{z,n} \rangle$ (nT)</th>
<th>Standard deviation $\sigma(O_{z,n})$ (nT)</th>
<th>Standard error $\sigma(O_{z,n})/\sqrt{N}$ (nT)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Solar wind</td>
<td>21200 (0.4 %)</td>
<td>$-0.04$</td>
<td>$-0.10$</td>
<td>6.5</td>
<td>0.04</td>
</tr>
<tr>
<td>Magnetosheath</td>
<td>29289 (2.1 %)</td>
<td>$-0.17$</td>
<td>$-0.34$</td>
<td>14.0</td>
<td>0.08</td>
</tr>
<tr>
<td>Magnetosphere</td>
<td>36652 (3.0 %)</td>
<td>$-0.01$</td>
<td>$-1.01$</td>
<td>12.4</td>
<td>0.09</td>
</tr>
</tbody>
</table>

**Table 3.** Fitting parameters of Eq. (5) with 95% confidence intervals, exhibited from linear least squares fits of the offset accuracies above 0.5 nT in Fig. 4. $a$ is the 2σ confidence of the best-estimate $O_{zf}$ determined from only one offset $O_{z}$ and $k$ is the spectral index of the power law.

<table>
<thead>
<tr>
<th>Region</th>
<th>$a$ (nT)</th>
<th>$k$</th>
</tr>
</thead>
<tbody>
<tr>
<td>Solar wind</td>
<td>$18.6 \pm 1.1$</td>
<td>$-0.87 \pm 0.03$</td>
</tr>
<tr>
<td>Magnetosheath</td>
<td>$34.8 \pm 1.0$</td>
<td>$-0.44 \pm 0.01$</td>
</tr>
<tr>
<td>Magnetosphere</td>
<td>$25.9 \pm 1.0$</td>
<td>$-0.41 \pm 0.01$</td>
</tr>
</tbody>
</table>

While 132 h of solar wind data are sufficient to determine the offset at an accuracy of 0.5 nT, 4241 h are needed in the magnetosheath and more than 6130 h in the magnetosphere. This is an interesting result, since the magnetosheath is the region with the highest probability to observe large-amplitude compressional fluctuations and the solar wind has the lowest (see Table 1). A possible explanation might be that the solar wind data are observations in the shock-upstream region in which the shock-reflected backstreaming ions excite large-amplitude Alfvén waves that either pitch the plasma (through ponderomotive force) or develop into magnetosonic-type waves. However, the results suggest that the fluctuations in the solar wind, although observed less often, are better suited to determine the offset with the mirror mode method.

**4 Discussion and conclusions**

We find that the offset determination method proposed by Plaschke and Narita (2016) is very applicable to the data from the Hermean environment. The results reveal that the solar wind is the most suitable region to accurately determine the offset by the mirror mode method, although the lowest percentage of highly compressible fluctuations is observed there (see Table 1). As can been seen in Table 4, offset determination with uncertainties better than 0.5 nT can be achieved with less than 132 h of solar wind data. However, it is important to note that during this time the instrument offsets need to stay constant to within 0.5 nT, otherwise the intrinsic offset drifts over time would limit the achievable precision, independently from the amount of solar wind data. Specifically in the magnetosheath and within the magnetosphere.
From the results obtained in this work we conclude that in the solar wind the mirror mode method should be a good complementary approach to the Alfvénic fluctuation method to determine the spin-axis offset of the Mio magnetometer. In the case of MPO we cannot evaluate from this study whether the mirror mode method is sufficient to obtain reliable offset estimates. However, from the results of the 1D mirror mode method, we find that considerably more data are needed to reach the same offset accuracy, since the orbit nominally remains within Mercury’s magnetosphere. In a future work the 3D mirror mode method should be applied on the MESSENGER data to assess whether the mirror mode method might also be a valuable tool to obtain reliable offset estimates for the magnetometers of MPO-MAG.

**References**


