



Validation of Clyde River SuperDARN radar velocity measurements with the RISR-C incoherent scatter radar

Alexander Koustov¹, Robert Gillies², and Peter Bankole¹

¹Department of Physics and Engineering Physics, University of Saskatchewan, Saskatoon, Saskatchewan, Canada

²Department of Physics and Astronomy, University of Calgary, Calgary, Alberta, Canada

Correspondence: Alexander Koustov (sasha.koustov@usask.ca)

Received: 6 September 2018 – Discussion started: 14 September 2018

Accepted: 5 December 2018 – Published: 14 December 2018

Abstract. The study considers simultaneous plasma velocity measurements in the eastward direction carried out by the Clyde River (CLY) Super Dual Auroral Radar Network (SuperDARN) high-frequency (HF) radar and Resolute Bay (RB) incoherent scatter radar – Canada (RISR-C). The HF velocities are found to be in reasonable agreement with RISR velocities up to magnitudes of 700–800 m s^{−1} while, for faster flows, the HF velocity magnitudes are noticeably smaller. The eastward plasma flow component inferred from SuperDARN convection maps (constructed for the area of joint measurements with consideration of velocity data from all the radars of the network) shows the effect of smaller HF velocities more notably. We show that the differences in eastward velocities between the two instruments can be significant and prolonged for observations of strongly sheared plasma flows.

1 Introduction

The Super Dual Auroral Radar Network (SuperDARN) high-frequency (HF) radars have been installed to continuously monitor the $\mathbf{E} \times \mathbf{B}$ plasma drift in the Earth's ionosphere (Greenwald et al., 1995). To achieve this goal, the radars detect coherent ionospheric echoes from the F region and measure their Doppler velocity. It is assumed that the decameter ionospheric irregularities, responsible for SuperDARN echoes, move with the velocity close to the $\mathbf{E} \times \mathbf{B}$ plasma drift. A number of comparisons of SuperDARN velocity measurements with concurrently operating incoherent scatter radars (ISRs) that measure the $\mathbf{E} \times \mathbf{B}$ plasma drift has been performed in the past (Ruohoniemi et al., 1987; Davies

et al., 1999, 2000; Milan et al., 1999; Xu et al., 2001; Gillies et al., 2009, 2010, 2018; Bahcivan et al., 2013; Koustov et al., 2016). These comparisons, overall, supported the above major assumption of the SuperDARN measurements. However, occasional significant differences between HF radar line-of-sight (LOS) velocities and $\mathbf{E} \times \mathbf{B}$ drift component along the beam have been noticed. Initially, these were thought to originate from differences in the echo collecting areas and in the signal integration time (Davies et al., 1999) but the body of the data published so far questions this notion. It is now accepted that the HF velocities of the F region echoes are generally smaller (Gillies et al., 2018). One factor found to lead to this result is an assumption made during SuperDARN velocity measurements, which sets the index of refraction for the ionosphere of unity. However, this explanation cannot account for large differences of more than 20 %–30 %. Koustov et al. (2016) stressed the original finding by Xu et al. (2001) that the HF velocity magnitudes are substantially smaller (up to a factor of 2) than the $\mathbf{E} \times \mathbf{B}$ drift component for high-speed flows exceeding 1000 m s^{−1}. Furthermore, the HF velocity magnitudes are often larger than the $\mathbf{E} \times \mathbf{B}$ flow component along the radar beam (e.g., Ruohoniemi et al., 1987; Koustov et al., 2016; Gillies et al., 2018). Such observations have been interpreted in terms of lateral deviation of HF radar beams (Koustov et al., 2016; Gillies et al., 2018). Other SuperDARN–ISR velocity inconsistencies have been associated with the occurrence of E region echoes at traditionally expected F region ranges for SuperDARN (Bahcivan et al., 2013; Gillies et al., 2018).

Despite obvious progress in measurement interpretation, HF-based $\mathbf{E} \times \mathbf{B}$ measurements require further investigation if one wants to continue improving the quality of the con-

vection mapping with HF radars. In addition, although all SuperDARN radars work on the same principle and often even have identical hardware, validation work for every unit is necessary to be confident in the reliability and consistency of measurements across the network.

In this study, we undertake a validation study for the Clyde River (CLY) SuperDARN radar. In a broader context, this effort complements the previous validation work for the Rankin Inlet (RKN) and Inuvik (INV) SuperDARN radars by Koustov et al. (2009), Mori et al. (2012), Bahcivan et al. (2013), Koustov et al. (2016), and Gillies et al. (2018). Since the CLY radar currently provides a significant contribution to the global-scale convection mapping with SuperDARN such work is of particular importance. We take advantage of the availability of $\mathbf{E} \times \mathbf{B}$ plasma drift measurements made by the recently installed Resolute Bay (RB) incoherent scatter radar – Canada (RISR-C) (e.g., Gillies et al., 2016). In the present work, we compare CLY- and ISR-based velocities in a different way than the previous studies.

Traditionally, gate-by-gate comparison of data from two radar systems that make measurements in roughly the same directions is performed (e.g., Gillies et al., 2018). Such an approach cannot be implemented for the CLY–RISR-C geometry because none of these radar’s beams are close enough in terms of their direction (see map on Fig. 1 in Gillies et al., 2018). For this reason, we consider RISR-C two-dimensional vectors in a certain area (which are inferred by merging data from multiple individual beams using the approach by Heinselman and Nicolls, 2008) and compare them with CLY data averaged over three beams and four gates. Thus, we assess the data in a statistical sense, in terms of the average and median velocities over a large spatial domain.

A validation using highly averaged data is appropriate since the SuperDARN global-scale maps of plasma flow obtained with the Potential Fit technique (Ruohoniemi and Baker, 1998; Shepherd and Ruohoniemi, 2000) are built using median-filtered line-of-sight (LOS) velocities (the so-called gridded velocities). These are inferred from up to 27 LOS velocity values in bins consisting of data in neighboring range gates (one smaller number gate and one larger number gate) and radar beams (one smaller number beam and one larger number beam) and for three consecutive radar scans. This implies that the input to the Potential Fit procedure is a highly smoothed HF velocity covering 3–6 min of raw data and a significant space domain. In this view, there is a sense in considering 2-D RISR-C data and comparing them with HF velocity medians, or vectors from the convection maps, over large spatial areas of overlap.

Although our aim is to validate the CLY velocity measurements, there is additional value from the CLY–RISR comparison. The RISR method of velocity vector estimations also has some limitations (Heinselman and Nicolls, 2008) that need testing. A couple of the limitations we will consider are a lack of velocity measurements along magnetic field lines and the expectation of spatially quasi-uniformity of flows,

which is not always satisfied. Thus, our work can be considered as a mutual validation of both radars’ performance. Compatibility of the vector estimates by RISR and SuperDARN is expected, but the degree of this agreement is not yet known.

2 Geometry of RISR-C and Clyde River radar observations

Figure 1 shows the fields of view (FoVs) of the CLY and RKN SuperDARN radars starting from range gate 5 and the location of the RB incoherent scatter radar RISR-C, which we will simply refer to as the RISR radar hereafter. This radar makes measurements in multiple beams; it uses 11 beams in the so-called “world-day” mode and 51 beams in the so-called “imaging” mode. Measured line-of-sight velocities in all the beams and at all ranges are used to infer 2-D vectors of the $\mathbf{E} \times \mathbf{B}$ plasma flow according to the procedure outlined by Heinselman and Nicolls (2008). The resultant vectors are reported with 0.25° bin size of magnetic/geographic latitude. The points to which the measurements are assigned are shown in Fig. 1 for the height of 300 km. The actual centerline for the points of data merging depends on data availability in specific beams (Gillies et al., 2018).

Figure 1 also shows the orientation of the CLY beams 4, 5, 6 (along their centers), and the area from which data were considered, the shaded rectangle region flanked by beams 4 and 6 between range gates 18 and 22. The monitored ionospheric region is centered at geographic latitude of $\sim 72.5^\circ$. An important feature of this area is that within these range gates the CLY beams 4–6 are almost parallel to the lines of equal geographic latitude at the chosen radar range gates, as shown in Fig. 1. This means that one can directly compare CLY LOS velocities with the eastward component (in geographic coordinates) of a RISR $\mathbf{E} \times \mathbf{B}$ velocity vector. We note that the area of CLY observations was also covered by measurements from the RKN and INV radars (and occasionally by the Saskatoon and Kodiak SuperDARN radars), so that SuperDARN convection maps were usually well constrained.

3 Methodology of the LOS velocity comparison

We consider here an extensive data set comprising of about 1000 h of RISR measurements made over the entire year of 2016. On the days when the radar was operational it typically worked for the whole 24 h, albeit switching, once in a while, its mode of operation, except the world-day mode which usually covered an entire day. The range resolution of measurements in both modes is ~ 50 km. The data are available for winter and both equinoxes, with no measurements made in the summer. We consider 5 min RISR data because they have much smaller errors than the 1 min data that are also available.

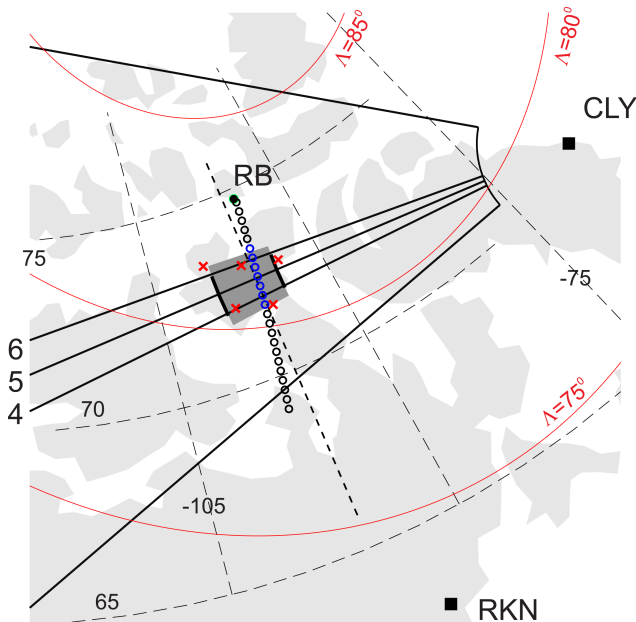


Figure 1. Field of view of the SuperDARN radar at CLY. The black straight lines are the orientation of specific beams (4–6) that were considered in this study. Shaded areas represent areas of HF radar data averaging. RB is the location of the RISR-C ISR. The radar reports the $\mathbf{E} \times \mathbf{B}$ vector with a bin size of 0.25° of geographic latitude for points shown as black circles stretching roughly along the magnetic meridian crossing the RB zenith. The blue-colored circles are those locations whose data were used for comparison with the CLY measurements. The solid red arcs are the magnetic latitudes of 75° , 80° , and 85° .

Our approach to the CLY–RISR velocity comparison is as follows. We first select a 5 min period of RISR velocity measurements at geographic latitudes of $\sim 71.625\text{--}73.125^\circ$ (see blue circles in Fig. 1) and compute the median velocity value for RISR. We then compute the median value of the CLY velocity over matching 5 min interval in three beams and four gates, as mentioned above. The matched data pair is then entered into a common data set.

Figure 2 shows the total number of 5 min intervals of joint RISR–CLY radar measurements, times when RISR and CLY both made measurements in the blue and shaded regions shown in Fig. 1, as a function of UT. This histogram distribution does not include individual events when CLY data were obviously contaminated by ground scatter profoundly affecting the velocity comparison (Gillies et al., 2018). The ground scatter was identified with the conventional selection criteria (e.g., Sect. 4.1 in Ponomarenko et al., 2007). The number of intervals was much larger from noon to dusk (local solar noon is at about 19:00 UT). This is because of the preferential occurrence of CLY echoes at ranges of interest during the daytime (Ghezelbash et al., 2014).

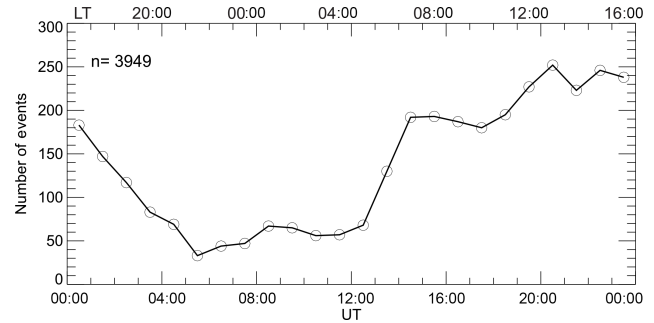


Figure 2. Number of CLY–RISR 5 min intervals of joint observations for all data considered. Total number of available intervals is shown in the top-left corner. For the area of observations, local time (scale at the top) roughly coincides with the magnetic local time.

4 Results for CLY LOS velocity and RISR comparison

Figure 3a shows the CLY LOS velocity versus the RISR eastward $\mathbf{E} \times \mathbf{B}$ component for the entire data set, produced as described above. The total number of points is close to 4000, which is a significant number. Overall, both positive and negative velocities are well represented. Although some spread is present, a significant amount of points are located close to the line of equality. To assess the plot, we binned the data according to the RISR measurements by using 100 m s^{-1} bins of the latter. Binned in this way CLY velocity medians are shown by black–white dots. The vertical black–white bars crossing each dot are the binned CLY velocity value ± 1 standard deviation. We also binned the data of Fig. 3 according to bins of the CLY velocity (pink asterisks, shown by thin symbols in order to not contaminate the plot).

The black–white dots are reasonably close to the line of the perfect agreement. The pink asterisks are actually very close to the line of equality. Good alignment with the line of equality and good correspondence between the location of the black dots and pink asterisks indicate that the velocities are almost linearly related, especially in the range from -500 to $+500 \text{ m s}^{-1}$. One clear departure of the black dots from the line of equality are the RISR velocities with magnitudes greater than $\sim 750 \text{ m s}^{-1}$.

An alternative way of assessing the data trends in Fig. 3a is to make a linear fit to the cloud of points. Parameters of the linear fit are presented in Table 1 for four ranges of the RISR velocity, ± 500 , ± 750 , ± 1000 , and $\pm 1500 \text{ m s}^{-1}$. The slope is 0.73 for the smallest velocities of $\pm 500 \text{ m s}^{-1}$, which includes about 71 % of all the data points. A linear fit to almost all the data has a decreased slope of 0.64.

5 Methodology of vector comparison between SuperDARN and RISR

The approach to the velocity vector comparison between the RISR and SuperDARN data is as follows. We restrict

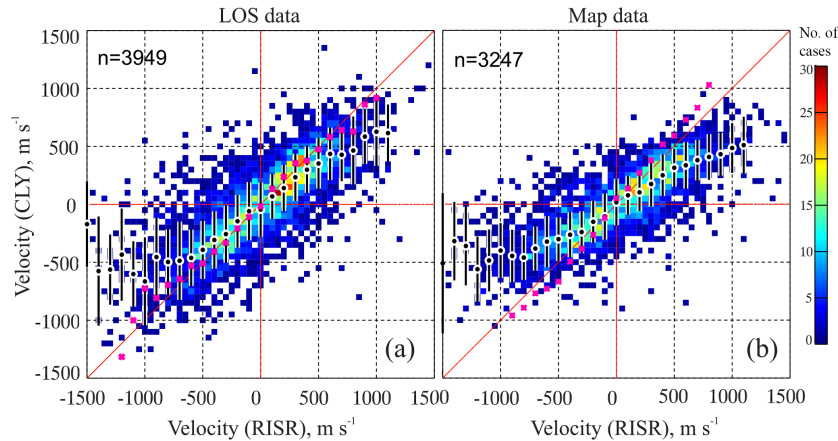


Figure 3. (a) Scatterplot of the CLY LOS velocity versus $E \times B$ eastward velocity component as inferred by RISR. Total number of points n is shown in the top-left corner. The black–white dots are medians of the CLY velocity in 100 m s^{-1} bins of RISR velocity. The black vertical lines are the standard deviations of the CLY velocity in each bin. The pink dots are medians of the RISR velocity in 100 m s^{-1} bins of CLY measurements. (b) The same as (a) but the eastward flow component inferred from SuperDARN flow maps was considered.

Table 1. Parameters of the linear fit line $\text{Velocity}_{\text{CLY}} = a \cdot \text{Velocity}_{\text{RISR}} + b$, the number of points involved in the fitting, and the squared correlation coefficient for various ranges of the RISR velocity. Left three columns are the LOS velocity comparison while right three columns are for the 2-D velocity comparison.

m s^{-1}	LOS comparison				2-D comparison			
	a	$b \text{ (m s}^{-1}\text{)}$	Points	R^2	a	$b \text{ (m s}^{-1}\text{)}$	Points	R^2
± 500	0.73	−9.90	2815	0.46	0.59	−16.32	2202	0.45
± 750	0.71	−10.29	3558	0.58	0.56	−11.31	2851	0.57
± 1000	0.68	−8.27	3823	0.6	0.54	−9.91	3106	0.62
± 1500	0.64	−5.31	3932	0.6	0.502	−7.08	3227	0.61

consideration to the same area of joint CLY–RISR observations as in the LOS comparison, shown in Fig. 1. Here the SuperDARN convection vectors are available at geomagnetic latitudes of $80.5\text{--}81.5^\circ$ and $\sim 7^\circ$ of magnetic longitude. In this area, the convection maps and vectors are mostly based on RKN, INV, and CLY radar measurements with only occasional contributions from other SuperDARN radars. We selected the three grid nodes at 81.5° magnetic latitude (MLAT) that were closest to the area of the CLY LOS velocity assessment and the two closest grid nodes at 80.5° magnetic latitude, marked by red crosses in Fig. 1. For each vector location, the geographic east component of the flow was computed and the median value (out of potentially five values, although for some periods it was as low as one measurement) was calculated to represent the eastward plasma flow component of a 5 min SuperDARN map. This is not a traditional temporal resolution for the SuperDARN mapping (which is usually 2 min); such data processing has been done to avoid the need of additional averaging of 2 min SuperDARN maps. Unfortunately, the start times of RISR measurement intervals were often irregularly spaced while SuperDARN maps were synchronized to exactly correspond

to 5 min boundaries (i.e., 0–5, 10–15 min, etc.). For the comparison, only HF and ISR data that were less than 2 min apart were considered. For this reason, even when both radar systems were operational, the actual number of joint points per hour was below the expected number of 12.

For RISR, the eastward $E \times B$ plasma velocity component was usually available at all points shown by open circles in Fig. 1. For the comparison with SuperDARN vectors, only measurements at geographic latitudes between 71.625 and 73.125° (given with a bin size of 0.25° , blue-colored circles in Fig. 1) were considered, and the median value of the eastward component was computed. The selection criteria produced a slightly shorter (but still statistically significant) data set than was obtained for the LOS velocity comparison. We stress that although the data for the comparison were along one specific direction, geographic east, two-dimensional vectors were used in determination of the velocity component for both systems with the SuperDARN vectors calculated using measurements from all radars including CLY, RKN, and INV, as well as the statistical model by Ruohoniemi and Greenwald (1996).

6 Results for vector comparison between SuperDARN and RISR

Figure 3b plots the eastward component of the plasma flow measured by RISR and SuperDARN. The spread of the data looks similar to that of Fig. 3a (the LOS comparison). We assessed Fig. 3b using the same methods as performed on Fig. 3a (see Sect. 4). Overall agreement of the data clearly holds.

Several results from Fig. 3b are consistent with the data of Fig. 3a. First, the SuperDARN map-based velocities are somewhat smaller than those of RISR. This is recognizable through an obvious deviation of the distribution maxima from the line of equality, especially at RISR positive velocities of $>500 \text{ m s}^{-1}$. Secondly, the tendency for the SuperDARN velocity being smaller is greater for larger RISR magnitudes. This feature is seen for both positive and negative RISR velocities. Finally, consistent with previous reports (Koustov et al., 2016; Gillies et al., 2018), there is a number of points for which the radars show oppositely directed flows. This was more frequent for small RISR velocities. Although Fig. 3 shows good consistency of the data provided by the two radar systems, the differences can be as large as a factor of 2 in individual measurements.

The agreement between the convection vectors given by RISR and SuperDARN is expected. We see that the consistency deteriorates once 2-D data are involved, but mostly at intermediate velocity magnitudes of $300\text{--}600 \text{ m s}^{-1}$. The inconsistencies are characterized by slower SuperDARN velocities. Interestingly, the differences for large velocity magnitudes in Fig. 3b are comparable to those in Fig. 3a.

To assess the data trends in Fig. 3b in an alternative way, linear fits to the scatter of points in Fig. 3b were made for four ranges of the RISR velocity of ± 500 , ± 750 , ± 1000 , and $\pm 1500 \text{ m s}^{-1}$, similar to those for the LOS velocity comparison. The slope of the fitted line, the y intercept, the number of points involved in each fitting, and the squared correlation coefficient are presented in Table 1. The slopes are close to 0.6 for the set of smallest velocities ($\pm 500 \text{ m s}^{-1}$), which includes about 68 % of all the available data. The slope decreases to 0.5 if almost all the data are considered. We think that the deterioration of the agreement at intermediate and large velocity magnitudes is due to the broader area over which the SuperDARN data are averaged for the 2-D comparison. In this case, there is more chance for SuperDARN to include ground-scatter-contaminated measurements, giving effectively slower grid velocities to the fitting procedure.

7 On possible reasons for velocity disagreements

One reason frequently given for the systematic underestimation of the SuperDARN velocity measurements is the assumption that the index of refraction is unity (Gillies et al., 2009; Ponomarenko et al., 2009). We attempted to evaluate

the importance of this effect in our data set. A plot similar to Fig. 3a was produced, but with the CLY velocity being corrected by considering the electron density (at the F region peak) measured by RISR. The plot looked very much similar to Fig. 3a. We assessed the plot by applying the linear fit line to the HF velocity medians in 100 m s^{-1} bins of RISR velocity, considering the range of almost linear dependence, between -1000 and $+1000 \text{ m s}^{-1}$ of RISR velocities. The slope of the best fit line improved to ~ 0.75 (from ~ 0.65). This improvement is consistent with the previous studies though it does not entirely account for the differences between the radar measurements.

We also investigated the diurnal variation of the velocity ratio $R = \text{Vel}_{\text{HF}} / \text{Vel}_{\text{RISR}}$ as done previously by Gillies et al. (2018) to explore possible influences of the refractive index on velocity using typical local time variations in the electron density as a proxy for refractive index. For the winter and equinoctial ionosphere over Resolute Bay, the largest densities are systematically observed near local solar noon and during the afternoon hours (18:00–22:00 UT) (e.g., Ghezelbash et al., 2014; Themens et al., 2017). It is therefore expected that the velocity ratio R would be smallest during these times, as reported by Gillies et al. (2018) for the RKN radar. The nighttime results by Gillies et al. (2018) are more confusing. First, strangely, the ratios here were often above 1 at latitudes southward of RB and systematically below 1 (but not as far below unity as they were near noon) at latitudes poleward of RB. Gillies et al. (2018) indicated that the vertical plasma flow velocities in RISR measurements were, very likely, incorrectly estimated for nighttime observations. Since the observation area in our comparison is close to RB, we expect that this effect will also affect the RISR–CLY comparison.

Figure 4 plots the hourly median ratio R as a function of time for our CLY–RISR data set. One can see that R varies significantly. It is lower during daytime (noon is at about 19:00 UT) than during dawn/prenoon (12:00–18:00 UT), but its values are smallest during nighttime (midnight is at about 07:00 UT). Interestingly, the average ratio over all UTs is 0.83, which is closer to 1 than the slopes of the lines in Fig. 3 (Table 1). This is probably because the infrequent high-velocity data are averaged out by dominating data at low velocities in certain RISR bins of Fig. 4.

We think that the low nighttime R values are caused by overestimation of true plasma drift in a plane perpendicular to the magnetic field by RISR in the midnight sector. We note that this is not quite consistent with Gillies et al. (2018), who interpreted their nighttime data in terms of effectively decreased RISR LOS velocities. Our data suggest effectively increased RISR velocity magnitudes. One of the factors affecting the derivation of the averaged flow pattern at nighttime, for both radar systems, is that the flows in this sector are often very irregular, even in the polar cap (Bristow et al., 2016). Under these conditions, the solution is subject to large uncertainty.

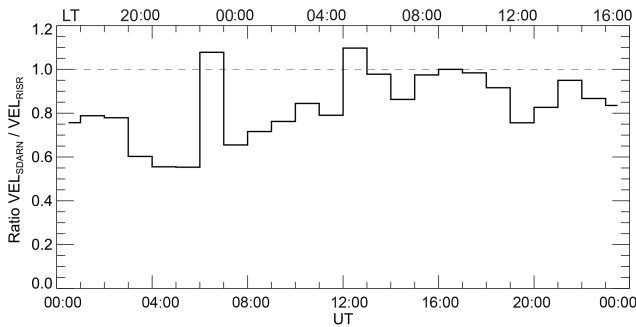


Figure 4. Line plot of the hourly median velocity ratio R versus UT for the CLY radar. The data set is the same as for Fig. 3a. For the area of observations, local time (scale at the top) roughly coincides with the magnetic local time.

Gillies et al. (2018) believe that large nighttime ratios of RKN to RISR velocity could be due to errors in HF measurements because the RKN beams can experience significant lateral deviations so that actual measurements are performed at smaller flow angles with a larger LOS velocity component. This explanation cannot be applied to our observations. This is because the CLY radar observes azimuthally, along the average plasma flow most of the time (except of short periods at near noon and near midnight when the flows are predominantly meridional) so that lateral deviations of the CLY beams would lead to, depending of the orientation of the plasma flow with respect to the CLY beam, either smaller or larger LOS velocities.

We think that the HF–RISR velocity inconsistency can also originate, at least partially, from the nature of HF signal formation. The effect has been discussed in general terms by Uspensky et al. (1989) as applied to E region coherent scatter and by Koustov et al. (2016) for F region coherent backscatter. The flows in the nighttime ionosphere are very likely to be more patchy/grainy with occasional occurrence of regions with enhanced flow magnitude (low electron density) and decreased flow magnitude (high electron density). We argue that in the case of a patchy ionosphere there is a good chance that the ratio R would be smaller than in the case of a uniform ionosphere and homogeneous flow. Flow enhancements and decreases affect both RISR and HF measurements but in a profoundly different manner. The RISR radar would average the velocity in patches with enhanced and depleted electron density together, and it would report what can be classified as the background flow velocity. In the presence of electron density patches with enhanced and decreased $\mathbf{E} \times \mathbf{B}$ plasma flows, HF radars would preferentially detect stronger signals from those areas where the electron density is enhanced, and the electric field (flow magnitude) is decreased, so that they would show somewhat smaller velocity than the background value measured by an incoherent scatter radar.

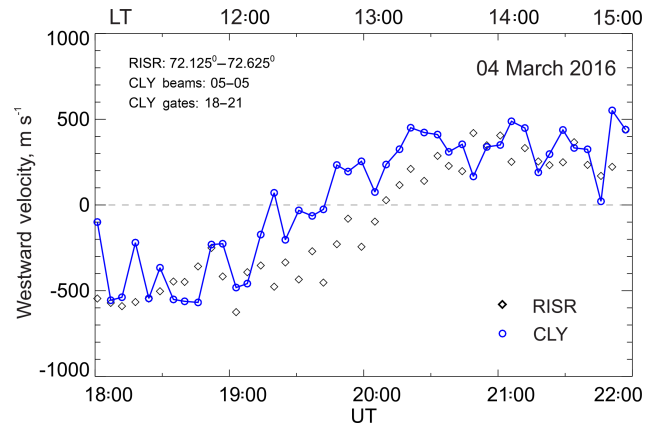


Figure 5. Eastward component of the $\mathbf{E} \times \mathbf{B}$ drift as measured by RISR (diamonds, 5 min resolution data) and matched velocity medians of CLY observations (blue circles, 5 min medians of original 1 min measurements in beams and gates overlapping the region of RISR observations) for the event of 4 March 2016. For the area of observations, local time (scale at the top) roughly coincides with the magnetic local time.

It is conceivable to have the opposite situation with HF velocities above the background flow if regions with enhanced density have a stronger local electric field, as discussed in Uspensky et al. (1989). In this respect, Koustov et al. (2016) and Gillies et al. (2018) noticed that HF velocities could be larger than the $\mathbf{E} \times \mathbf{B}$ plasma drift component measured by ISRs. Such points are occasionally seen in previously published data (Ruohoniemi et al., 1987; Davies et al., 1999). Our data in Fig. 3 also show such points but, in general, the data agree fairly well. Although the work of Koustov et al. (2016) and Gillies et al. (2018) related the larger HF velocity effect to lateral deviations of the HF radar beams from the expected directions, it could partially be due to the aforementioned effect of ionospheric microstructuring.

Potentially, low R values can be related to the occurrence of misidentified ionospheric scatter because some ionospheric echoes with low velocities can actually be ground or mixed ionospheric and ground scatter. Gillies et al. (2018) showed that removal of points that could potentially be affected by ground scatter improves the RKN–RISR velocity agreement significantly. Our analysis showed that ground scatter is rare during winter and equinox nighttime for the CLY radar, which is consistent with low nighttime F region densities (Ghezlbash et al., 2014; Themens et al., 2017). We also have to remind the reader that presumably obvious events with CLY ground-scatter contamination have been removed from our consideration in Fig. 3.

Investigating our database, we identified one special situation when the RISR–SuperDARN velocity disagreements were particularly strong. Figure 5 gives an example of CLY–RISR observations on 4 March 2016 where RISR and CLY

velocities differ consistently by several hundred m s^{-1} over a period of almost 2 h.

Figure 6a illustrates the typical spatial velocity distribution within the radar FoV, for one velocity scan during the above event. A sharp change in the LOS velocity polarity in the poleward and equatorward portions of the FoV is noticeable. The polarity transition occurs in the central beams 5–7. Figure 6b gives a global-scale map of plasma flow inferred from all SuperDARN radar measurements. The flow pattern in Fig. 6b was calculated by applying the new SuperDARN statistical model by Thomas and Shepherd (2018) which became available just recently. The map has a number of vectors originating from the RKN and INV radar measurements as well as those from CLY measurements. The presence of highly curved flows is evident near noon. Under these conditions, both SuperDARN and RISR can have difficulties in the construction of a 2-D vector field.

We comment that the flows seen in Fig. 6 are sunward, roughly along the magnetic meridian near noon, signifying the occurrence of a reverse convection cell. This is expected since the IMF B_z was steady at about +4 nT starting from 18:30 UT all the way until $\sim 22:00$ UT for this event.

Evaluating the extent to which the SuperDARN and RISR vectors are affected by the shear in the flow is difficult. We can see that the centers (foci) of the convection cells, according to RISR and SuperDARN, do not coincide in latitude for many maps in this event. In addition, the agreement between the RISR and SuperDARN map data improves dramatically when only the lower latitude SuperDARN map data are considered.

We investigated this further by determining the location of the convection reversal boundary (CRB) for the reverse convection cell (like that shown in Fig. 6b by the dashed contour). This is done by considering the standard 2 min SuperDARN maps, the CLY LOS velocity maps, and by looking at the reversal in the latitudinal profile of the RISR velocity (these are given for 5 min intervals). The CRB location based on the SuperDARN maps was determined by finding the middle latitude between the two neighboring points on a standard plasma flow map with opposite directions of the flow, toward the Sun and away from the Sun. The CRB location based on the CLY LOS velocity maps was determined by plotting the LOS velocity versus beam number and finding the azimuth and the range of the point at which the LOS velocity is zero. The CRB location from the RISR measurements was found by plotting the azimuthal component of the RISR plasma flow versus latitude and finding the latitude with zero velocity. All the CRB locations were given in terms of the geomagnetic latitude. The accuracy of the CRB determination in all cases is on the order of half of a degree of geomagnetic latitude.

The resulting data are presented in Fig. 7. The CRB inferred from SuperDARN maps is located almost 2° higher in MLAT than that determined from both CLY velocity maps and RISR data at the beginning of the event, and the differ-

ences are minimal toward the end of the event. The fact that the CRB location from CLY velocity maps is closer to that inferred from RISR data hints that perhaps the SuperDARN fitting procedure is the major factor for strong differences between the SuperDARN maps and RISR measurements in this specific event. This is not to say that RISR measurements are exact; they are very likely also subject to errors under these strongly sheared and curved flows. One reason could be that the solution for the 2-D velocity vector field from the original LOS RISR data (Heinselman and Nicolls, 2008) smooths out the true sharp changes of the flow. Having a wider FoV for the RISR radar is expected to improve the quality of the flow pattern derivation under such conditions.

8 Summary and conclusions

In this study, we attempted to validate the CLY SuperDARN radar velocity measurements by comparing them with the data collected by the Resolute Bay incoherent scatter radar. Because no line-of-sight velocity comparison is possible for the geometry of joint observations, we adopted here a different approach. Namely, we considered the eastward component of the $\mathbf{E} \times \mathbf{B}$ flow vector, as inferred from RISR measurements in multiple beams, and compared it to CLY velocities from a number of eastward-oriented beams and with the eastward component of the plasma flow inferred from 2-D SuperDARN maps. The analysis undertaken allows us to draw several conclusions.

The CLY radar velocities measured in beams 4–6 are statistically comparable to the $\mathbf{E} \times \mathbf{B}$ component of the plasma drift along these beams (eastward/azimuthal plasma flows) as measured by the RISR incoherent scatter radar. This implies that the velocity data provided by the CLY radar to the SuperDARN database are reliable and suitable for convection mapping involving all SuperDARN radars. The comparisons performed are an addition to the previous validation work for the RKN and INV SuperDARN radars.

The slope of the best linear fit line to the CLY velocity variation versus $\mathbf{E} \times \mathbf{B}$ component (as measured by RISR) applied to the binned values is on the order of 0.65 if all the available data (removing data with obvious groundscatter contamination) in the range $\pm 1000 \text{ m s}^{-1}$ are considered. Correction of HF velocities on the index-of-refraction effect improves the slope to ~ 0.75 . The slope of the linear fit line for the corrected data is still below 1, implying that additional factors affect the relationship. Additionally, diurnal variations of the ratio of HF velocity to the RISR velocity show their strongest decrease below 1 during nighttime but not daytime. This implies that the deterioration of the RISR–SuperDARN velocity agreement at nighttime is caused not by the index-of-refraction effect but by other factors.

The effect of HF velocity underestimation for the CLY radar becomes progressively stronger for plasma drifts faster than about $\sim 750 \text{ m s}^{-1}$.

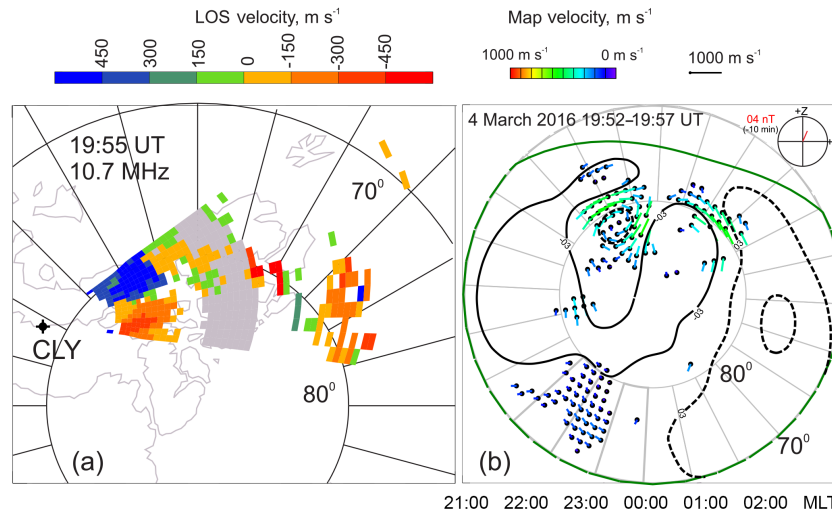


Figure 6. (a) A CLY LOS velocity map at 19:55 UT on 4 March 2016 and (b) a 5 min convection map calculated from all SuperDARN radar measurements for the same period of time. The TS2018 statistical model (Thomas and Shepherd, 2018) of order 8 for the solar wind electric field of 2.1 mV m^{-1} was used. Contours of the electric potential are 6 kV apart.

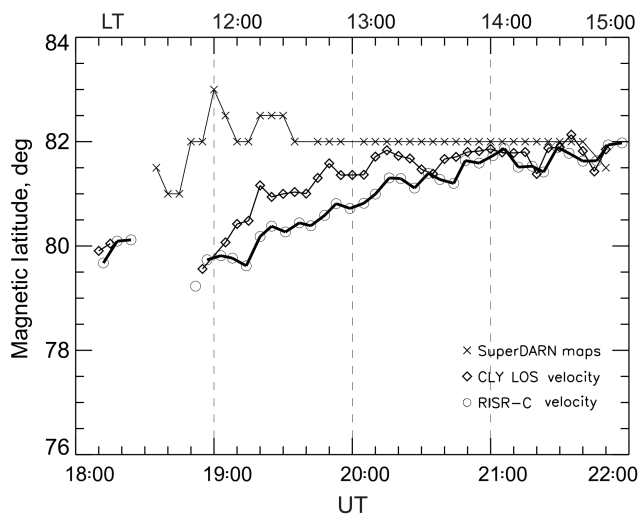


Figure 7. Magnetic latitude of the flow reversal location within the dayside reverse convection cell as inferred from SuperDARN convection maps (crosses), CLY LOS velocity maps (diamonds), and RISR measurements for the event of 4 March 2016. For the area of observations, local time (scale at the top) roughly coincides with the magnetic local time.

One factor that may contribute to slower HF velocities, in addition to the refractive index, is the nature of HF signal collection. HF radars receive stronger signals from ionospheric regions with enhanced electron density where the electric field and/or $\mathbf{E} \times \mathbf{B}$ plasma drift can be decreased compared to the background plasma.

In the case of highly sheared plasma flows, such as near dayside reverse convection cells occurring under strongly

dominant IMF $B_z > 0$, the differences between RISR and SuperDARN velocity vectors can be large.

The reasonable agreement between the velocities of the two systems quantified as the slopes of the linear fit lines at the level of 0.6–0.8 for both the LOS and 2-D comparisons implies that the RISR technique of the $\mathbf{E} \times \mathbf{B}$ derivation from multiple individual radar beams is usually a reliable method most of the time. The comparison suggests that the RISR vectors are less reliable in the midnight sector where the flows are often very irregular, and strong vertical motions occur.

Data availability. SuperDARN data can be obtained from <https://superdarn.ca> (last access: 11 December 2018). RISR-C data are available at <http://data.phys.ualgary.ca> (last access: 11 December 2018).

Author contributions. RG and PB worked on raw data processing and their preliminary analysis. PB prepared some diagrams. AK did most of the comparison work and wrote the initial manuscript. All authors participated in the writing, and all commented on the paper.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. Continuous funding of SuperDARN radars by National Scientific Agencies of Australia, Canada, China, France, Italy, Japan, Norway, South Africa, United Kingdom, and the United States of America is appreciated. The current research would be impossible without ongoing support from the

Canadian Foundation for Innovation, Canadian Space Agency's Geospace Observatory (GO Canada) continuation initiative to the University of Saskatchewan radar group, and NSERC Discovery grant to Alexander Koustov. The University of Calgary RISR-C radar is funded by the Canada Foundation for Innovation and is a partnership with the US National Science Foundation and SRI International. We thank C. Graf for the initial analysis of RISR density data and R. Fiori for the help in software development. Discussions of various aspects of the paper with Pavlo Ponomarenko and his help in software development are appreciated. The authors are indebted to two anonymous reviewers who not only identified weaknesses in the original manuscript but also made numerous constructive suggestions and, in addition, tremendously improved the writing style.

Edited by: Steve Milan

Reviewed by: two anonymous referees

References

- Bahcivan, H., Nicolls, M. J., and Perry, G.: Comparison of SuperDARN irregularity drift measurements and F-region ion velocities from Resolute Bay ISR, *J. Atmos. Sol.-Terr. Phys.*, 105–106, 325–331, <https://doi.org/10.1016/j.jastp.2013.02.002>, 2013.
- Bristow, W. A., Hampton, D. L., and Otto, A.: High-spatial-resolution velocity measurements derived using Local Divergence-Free Fitting of SuperDARN observations, *J. Geophys. Res.-Space*, 121, 1349–1361, <https://doi.org/10.1002/2015JA021862>, 2016.
- Davies, J. A., Lester, M., Milan, S. E., and Yeoman, T. K.: A comparison of velocity measurements from the CUTLASS Finland radar and the EISCAT UHF system, *Ann. Geophys.*, 17, 892–902, <https://doi.org/10.1007/s00585-999-0892-9>, 1999.
- Davies, J. A., Yeoman, T. K., Lester, M., and Milan, S. E.: *Letter to the Editor*: A comparison of F-region ion velocity observations from the EISCAT Svalbard and VHF radars with irregularity drift velocity measurements from the CUTLASS Finland HF radar, *Ann. Geophys.*, 18, 589–594, <https://doi.org/10.1007/s00585-000-0589-6>, 2000.
- Ghezelbash, M., Koustov, A. V., Themens, D. R., and Jayachandran, P. T.: Seasonal and diurnal variations of PolarDARN F region echo occurrence in the polar cap and their causes, *J. Geophys. Res.-Space*, 119, 10426–10439, <https://doi.org/10.1002/2014JA020726>, 2014.
- Gillies, R. G., Hussey, G. C., Sofko, J. G., McWilliams, K. A., Fiori, R. A. D., Ponomarenko, P. V., and St-Maurice, J.-P.: Improvement of SuperDARN velocity measurements by estimating the index of refraction in the scattering region using interferometry, *J. Geophys. Res.*, 114, A07305, <https://doi.org/10.1029/2008JA013967>, 2009.
- Gillies, R. G., Hussey, G. C., Sofko, J. G., Wright, D. M., and Davies, J. A.: A comparison of EISCAT and SuperDARN F-region measurements with consideration of the refractive index in the scattering volume, *J. Geophys. Res.*, 115, A06319, <https://doi.org/10.1029/2009JA014694>, 2010.
- Gillies, R. G., van Eyken, A., Spanswick, E., Nicolls, M. J., Kelly, J., Greffen, M., Knudsen, D., Connors, M., Schutser, M., Valentic, T., Malone, M., Buonocore, J., St-Maurice, J.-P., and Donovan, E.: First observations from the RISR-C incoherent scatter radar, *Radio Sci.*, 51, 1645–1659, <https://doi.org/10.1002/2016RS006062>, 2016.
- Gillies, R. G., Perry, G. W., Koustov, A. V., Varney, R. H., Reimer, A. S., Spanswick, E., St-Maurice, J.-P., and Donovan, E.: Large-scale comparison of polar cap ionospheric velocities, measured by RISR-C, RISR-N, and SuperDARN, *Radio Sci.*, 53, 624–639, <https://doi.org/10.1029/2017RS006435>, 2018.
- Greenwald, R. A., Baker, K. B., Dudeney, J. R., Pinnock, M., Jones, T. B., Thomas, E. C., Villain, J.-P., Cerisier, J.-C., Senior, C., Hanuise, C., Hunsaker, R. D., Sofko, J. G., Koehler, J. A., Nielsen, E., Pellinen, R., Walker, A. D. M., Sato, N., and Yamagishi, H.: DARN/SuperDARN: A global view of the dynamics of high-latitude convection, *Space Sci. Rev.*, 71, 761–796, <https://doi.org/10.1007/BF00751350>, 1995.
- Heinselman, C. J. and Nicolls, M. J.: A Bayesian approach to electric field and E-region neutral wind estimation with the Poker Flat Advanced Modular Incoherent Scatter Radar, *Radio Sci.*, 43, RS5013, <https://doi.org/10.1029/2007RS003805>, 2008.
- Koustov, A. V., St-Maurice, J.-P., Sofko, G. J., Andre, D., McDougall, J. W., Hairston, M. R., Fiori, R. A., and Kadochnikov, G. E.: Three-way validation of the Rankin Inlet PolarDARN radar velocity measurements, *Radio Sci.*, 44, RS4003, <https://doi.org/10.1029/2008RS004045>, 2009.
- Koustov, A. V., Lavoie, D. B., and Varney, R. H.: On the consistency of the SuperDARN radar velocity and $E \times B$ plasma drift, *Radio Sci.*, 51, 1792–1805, <https://doi.org/10.1002/2016RS006134>, 2016.
- Milan, S. E., Davies, J. A., and Lester, M.: Coherent HF radar backscatter characteristics associated with auroral forms identified by incoherent radar techniques: A comparison of CUTLASS and EISCAT observations, *J. Geophys. Res.*, 104, 22591–22604, <https://doi.org/10.1029/1999JA900277>, 1999.
- Mori, D., Koustov, A. V., Jayachandran, P. T., and Nishitani, N.: Resolute Bay CADI ionosonde drifts, PolarDARN HF velocities and SuperDARN cross polar cap potential, *Radio Sci.*, 47, RS3003, <https://doi.org/10.1029/2011RS004947>, 2012.
- Ponomarenko, P. V., Waters, C. L., and Menk, F. W.: Factors determining spectral width of HF echoes from high latitudes, *Ann. Geophys.*, 25, 675–687, <https://doi.org/10.5194/angeo-25-675-2007>, 2007.
- Ponomarenko, P. V., St-Maurice, J.-P., Waters, C. L., Gillies, R. G., and Koustov, A. V.: Refractive index effects on the scatter volume location and Doppler velocity estimates of ionospheric HF backscatter echoes, *Ann. Geophys.*, 27, 4207–4219, <https://doi.org/10.5194/angeo-27-4207-2009>, 2009.
- Ruohoniemi, J. M. and Baker, K. B.: Large-scale imaging of high-latitude convection with Super Dual Auroral Radar Network HF radar observations, *J. Geophys. Res.*, 103, 20797–20811, <https://doi.org/10.1029/98JA01288>, 1998.
- Ruohoniemi, J. M., Greenwald, R. A., Baker, K. B., Villain, J.-P., and McCready, M. A.: Drift motions of small-scale irregularities in the high-latitude F region: An experimental comparison with plasma drift motions, *J. Geophys. Res.*, 92, 4553–4564, <https://doi.org/10.1029/JA092iA05p04553>, 1987.
- Ruohoniemi, J. M. and Greenwald, R. A.: Statistical patterns of high-latitude convection obtained from Goose Bay HF radar observations, *J. Geophys. Res.*, 101, 21743–21763, <https://doi.org/10.1029/96JA01584>, 1996.

- Shepherd, S. G. and Ruohoniemi, J. M.: Electrostatic potential patterns in the high-latitude ionosphere constrained by SuperDARN measurements, *J. Geophys. Res.*, 105, 23005–23014, <https://doi.org/10.1029/2000JA000171>, 2000.
- Themens, D. R., Jayachandran, P. T., Galkin, I., and Hall, C.: The Empirical Canadian High Arctic Ionospheric Model (E-CHAIM): *NmF2* and *hmF2*, *J. Geophys. Res.-Space*, 122, 9015–9031, <https://doi.org/10.1002/2017JA024398>, 2017.
- Thomas, E. G. and Shepherd, S. G.: Statistical patterns of ionospheric convection derived from mid-latitude, high-latitude, and polar SuperDARN HF radar observations, *J. Geophys. Res.: Space Physics*, 123, 3196–3216 <https://doi.org/10.1002/2018JA025280>, 2018.
- Uspensky, M. V., Koustov, A. V., Williams, P. J. S., Huuskonen, A., Kangas, J., and Nielsen, E.: Effect of unresolved electrojet microstructure on measurements of irregularity drift velocity in auroral radar backscatter, *Adv. Space Res.*, 9, 119–122, [https://doi.org/10.1016/0273-1177\(89\)90349-9](https://doi.org/10.1016/0273-1177(89)90349-9), 1989.
- Xu, L., Koustov, A. V., Thayer, J., and McCready, M. A.: SuperDARN convection and Sondrestrom plasma drift, *Ann. Geophys.*, 19, 749–759, <https://doi.org/10.5194/angeo-19-749-2001>, 2001.