

Recent changes in measured wind in the NE Atlantic and variability of correlation with NAO

P. A. Pirazzoli¹, A. Tomasin², and A. Ullmann³

¹CNRS – Laboratoire de Géographie Physique, 1 Place Aristide Briand, 92195 Meudon cedex, France

²ISMAR-CNR, Venezia and Dipartimento di Matematica Applicata, Università “Ca’ Foscari”, Venezia, Italy

³Centre de Recherches de Climatologie, Université de Bourgogne, BP 138, 21004 Dijon cedex, France

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Abstract. The paper deals with wind measurements, recorded since the 1950s, at twelve meteorological stations along a transect near the westernmost European border, between 64° and 44° N. Extreme wind speed tends to decrease sharply near the northern boundary (at Reykjavick), near the middle of the study area (at Shannon and Valentia) and near the southern boundary (at Brest and Cap Ferret), to increase at Thorshavn, with less significant trends at the other stations. Average wind speeds confirm the above tendencies, with an additional increasing speed at Lerwick, Kirkwall, Malin Head, Belle-Ile and Cap Ferret.

To compare changes in wind activity, the data have been subdivided into three periods: until 1975, 1976–1992 and 1993–2008. Frequencies have been computed also for the “winter” (October to March) period, per quadrants, and for occurrences exceeding the speed of 15 m s^{-1} .

At Reykjavick a recent increase in the frequency of strong winds has occurred from various directions. Between 62° N (Thorshavn) and 59° N (Kirkwall) strong wind has been increasing since 1975. Minor changes can be observed at Stornoway, whereas at Malin Head the greatest increase for southerlies and westerlies is observed during the 1976–1992 period. At Belmullet, the frequency of strong southerlies has almost doubled since 1992, while at Shannon and Valentia it remains quite low. Finally at Brest and Belle-Ile, westerlies are predominant among winds $> 15 \text{ m s}^{-1}$.

Important changes in time and latitude appear in the correlation with the NAO (North Atlantic Oscillation) index. The highest correlation coefficients, calculated with monthly or seasonal means between the early 1950s and 1975, are observed from between 58° N (Stornoway) and Iceland, whereas low positive coefficients are reported more

south. During the period 1976–1992, when increasing NAO index is predominant, positive correlation improves southwards as far as 54° (Belmullet) with some improvement also at Shannon and Valentia, while it remains low or even negative near the French Atlantic coast. Finally in the 1993–2008 period, correlation improves for all the stations south of 54° N (Belmullet), while it weakens more north.

Keywords. Meteorology and atmospheric dynamics (Climatology; General circulation)

1 Introduction

The majority of storms that affect the northeastern Atlantic area originate in the midlatitude westerly wind belt and are generically referred to as extratropical storms (Lozano et al., 2004). Their recent evolution is often discussed, but a general agreement about recent changes in storm and cyclone activity in the extratropical northeast Atlantic is still missing.

The North Atlantic Oscillation (NAO) is often considered as the most robust climate teleconnection in the NE Atlantic (Hurrell and Deser, 2009). According to Hurrell et al. (2003), it swings from one NAO phase to the other to produce large changes in the mean wind speed and direction over the Atlantic and the intensity and number of storms. Surface winds are strongest during the cold season. Nevertheless, month-to-month and even year-to-year changes in the phase and amplitude of the NAO remain largely unpredictable.

Reanalysis results often show a noticeable increase in the number or in the intensity of cyclone events after 1970 (Lambert, 1996; Schmith et al., 1998; Paciorek et al., 2002). Some studies make a distinction between the high-latitude areas, where a significant increasing trend would be observed, and a gradual decrease over the eastern coasts of the Atlantic (Geng and Sugi, 2001; McCabe et al., 2001). Wang et al. (2006)



Correspondence to: P. A. Pirazzoli
(pirazzol@cnrs-bellevue.fr)

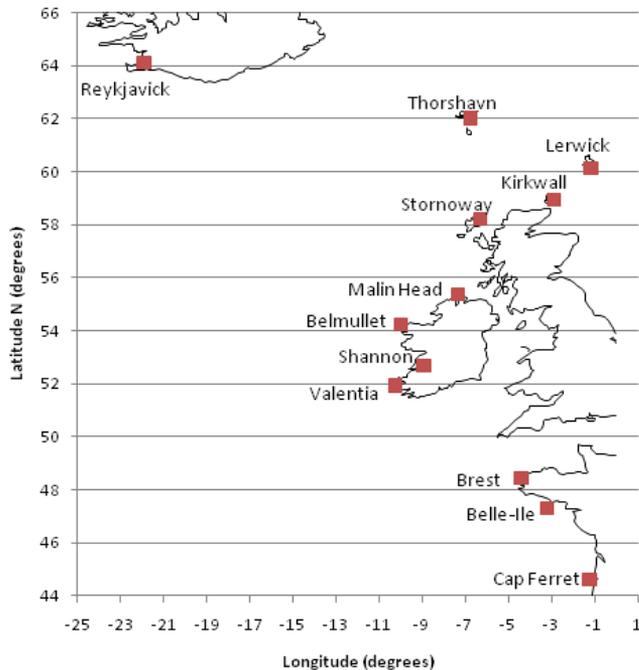


Fig. 1. Location map.

associate such changes in winter with a shifting of about 181 km northward of the mean position of the storm track. According to Gulev et al. (2001), Atlantic cyclone frequency demonstrates a high correlation with NAO.

On the other hand, station measurements of pressure change between October and March would show in Iceland a tendency towards less extreme severe events after 1980, while a tendency towards larger magnitude events was observed in the UK regions in recent decades (Alexander and Tett, 2005). According to Alexandersson et al. (2000), an increasing trend in NW Europe storminess – derived from station pressure measurements available since the late 19th century – started in the 1960s, but shows a clear tendency of a reduction in storminess after a maximum around 1990.

In an attempt to clarify recent climate changes in the eastern North Atlantic area, we collected and analysed statistically hourly or three-hourly wind measurements recorded at twelve meteorological stations close to the western border of Europe, i.e. directly exposed to Atlantic storms, between Iceland and France. In this paper, preliminary results and correlation with the NAO index of boreal cold season are presented. The data and methods used are specified in Sect. 2. Results (trends for wind speed, directions of wind occurrences and the temporal variability of strong wind frequency) are summarized in Sect. 3. Finally the difficulty of explaining the changes observed, or even of relating them in detail to changes in the NAO index, are discussed in Sect. 4.

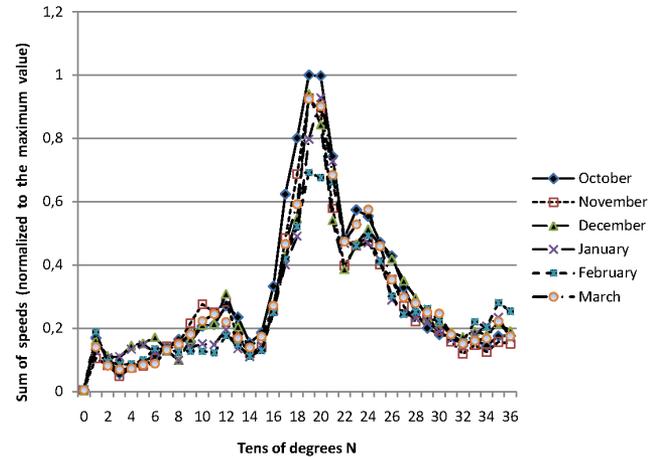


Fig. 2. Angular distribution of wind at Stornoway between October and March.

2 Data and methods

This work is based on the analysis of hourly-type records from twelve meteorological stations: Reykjavik in Iceland, Thorshavn in the Faeroes (Denmark), Lerwick in Zetland (UK), Kirkwall in Orkney (UK), Stornoway in the Outer Hebrides (UK), Malin Head, Belmullet, Shannon and Valentia in Ireland, and Brest, Belle-Ile and Cap Ferret in France (Fig. 1). Wind data have been generally provided by the Icelandic Meteorological Office, the Danish Meteorological Institute, the Met Office, the Irish Meteorological Service and the Météo France. In hourly-type measurements, wind directions are recorded clockwise from the geographical north and correspond to the average direction during the ten minutes preceding the report time. Wind speed for a specific hour corresponds to the average speed during the same 10 min period to which the direction is referred. In this work, all speeds have been reduced to the m s^{-1} unit. The records available usually start in the 1950s; some series have been completed for the period 1982–1997 by using the CD “International Surface Weather Observations” produced by the US Department of Commerce. For the most recent period, the records have been updated to the end of March 2009 by downloading from the Russia’s Weather Service (<http://meteo.infospace.ru/wcarch/html/>) (accessed April 2009). The frequency of measurements available is most often hourly for Met Office records, hourly or three-hourly for Irish records (with the possible absence of measurements between 15H and 0H UTC), three-hourly for Danish and French records. The data reported by the Russia’s weather server are only three-hourly or six-hourly.

The analysis is based on observations on several wind time series from different sources, without any assurance that they have already been homogenized. Possible changes over time in station location, anemometer elevation or anemometer exposure, may cause spurious temporal trends. Caution is,

Table 1. Location of the meteorological stations considered, sources of the data and periods of observation.

Number	Station			Available record length		
	Name	Lat. N	Long. W	National authorities	NOAA	Russia's weather server
4030	Reykjavik	64°08'	21°54'	1950–1981, 1998–2000	1982–1997	2000–2009
6011	Thorshavn	62°01'	6°46'	1953–1981, 1998–2000	1982–1997	2000–2009
3005	Lerwick	60°08'	1°11'	1957–1982, 1988–2000		2000–2009
3017	Kirkwall	58°57'	2°54'	1957–1963, 1965–1966, 1971–1972, 1976, 1981, 1998–2000	1982–1997	2000–2009
3026	Stornoway	58°13'	6°19'	1957–1982, 1988–2000	1983–1987	2000–2009
3980	Malin Head	55°22'	7°20'	1955–1981, 1998–2000	1982–1997	2000–2009
3976	Belmullet	54°14'	10°00'	1956–2000		2000–2009
3962	Shannon	52°42'	8°55'	1950–1981, 1998–2000	1982–1997	2000–2009
3953	Valentia	51°56'	10°15'	1950–2000		2000–09
7110	Brest	48°27'	4°42'	1949–2004		2005–2009
7207	Belle-Ile	47°18'	3°17'	1949–2004		2005–2009
7500	Cap Ferret	44°63'	1°25'	1949–2004		

therefore, necessary for the preliminary nature of the results of this study. Due to the inhomogeneous temporal distribution, 99th percentiles have been preferred to maximum values for the estimation of annual extremes. However, an inhomogeneity may have been introduced also in the 99th percentiles by using hourly or three-hourly data prior to 2000, or three or six-hourly data from 2000 on (with the shift towards the Russian data source). While mean values may not be affected by changes in the sampling frequency, higher values (like the 99th percentiles) might be slightly modified. The location of the meteorological stations, sources of data and periods of observations have been specified in Table 1. In addition to the usual representations, a quantity more related to dynamics was computed: given a direction of interest, from each observed wind vector a contribution was taken by projection on it, and added if positive.

The Gibraltar-Iceland NAO index, originally produced by Jones et al. (1997) and updated by Osborn (2006), is used. Monthly data until the year 2000 are available via ftp at the url: <ftp://ftp.cru.uea.ac.uk/data>. 2000-to-present monthly data are available at the url: http://www.cru.uea.ac.uk/~timo/projpages/nao_update.htm.

Statistical values were computed for full years as well as for “winter” (boreal cold season: ONDJFM) periods. The period ONDJFM was preferred to the DJFM period, often used for “winter” by Hurrell (1996) or by Osborn (2004), because in the area considered dominant monthly wind peaks are often recorded in October and sometimes also in November (Fig. 2, Table 2). In this paper, “Winter”, e.g. for the year 2008, means the period from October 2008 up to and including March 2009. With the same yearly convention, the cold season NAO indexes used are summarized in Fig. 3.

To identify changes in tendency or evolution in correlation coefficients, three periods have been considered: a) from

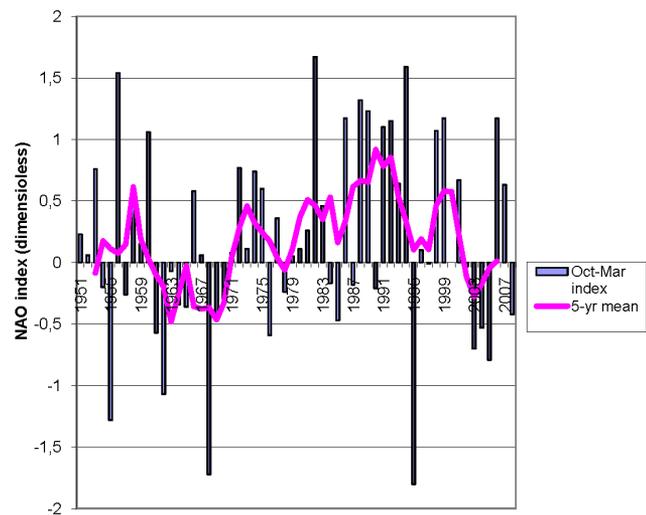


Fig. 3. Boreal cold season (October to March average) NAO index from 1950–1951 to 2008–2009.

the beginning of the record to 1975; b) from 1976 to 1992; c) from 1993 to 2008. The 1975 limit has been chosen because a number of studies have suggested that a change in certain climatic trends may have occurred around the mid-1970s in the Northern Hemisphere (e.g., Folland et al., 2001; Pirazzoli and Tomasin, 2003); the year 1992 has the advantage to cut the 1976–2008 period into two similar parts.

Linear trends have been determined by regressions with slope uncertainties and the statistical confidence.

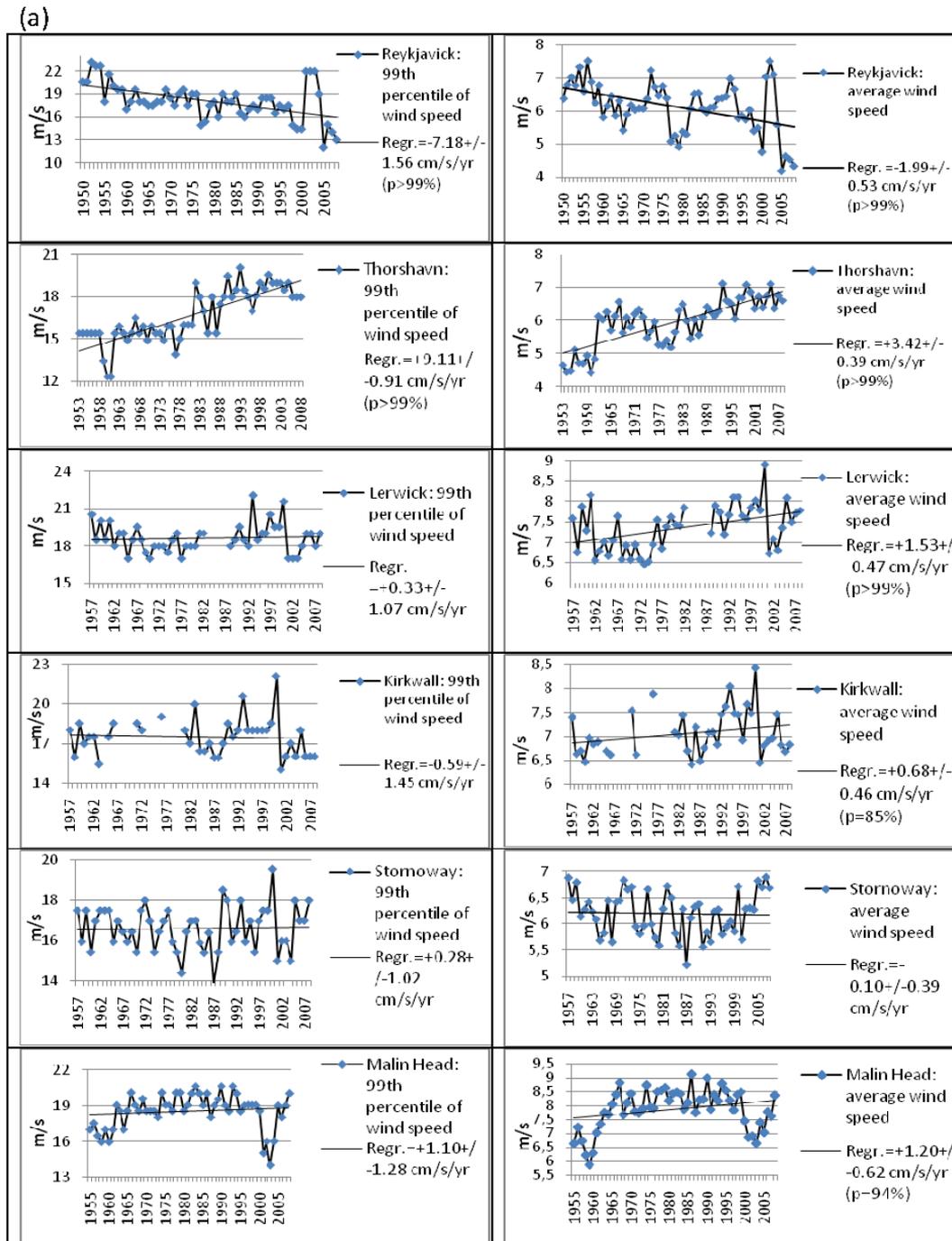


Fig. 4a. Recent changes in extreme (99th percentile) and average wind speed (all directions): (a) at Reykjavick, Thorshavn, Lerwick, Kirkwall, Stornoway and Malin Head; (b) at Belmullet, Shannon, Valentia, Brest, Belle-Ile and Cap Ferret.

3 Results

3.1 Overall trends for wind speed

Recent changes in extreme (99th percentile) and average wind speeds are reported for all directions in Fig. 4. Ex-

treme wind speed tends to decrease sharply near the northern boundary of the transect considered (at Reykjavick), near the middle of the study area (at Shannon and Valentia) and near the southern boundary (at Brest and Cap Ferret), to increase rapidly at Thorshavn, with less significant trends at the other stations. Average wind speeds confirm the above tendencies,

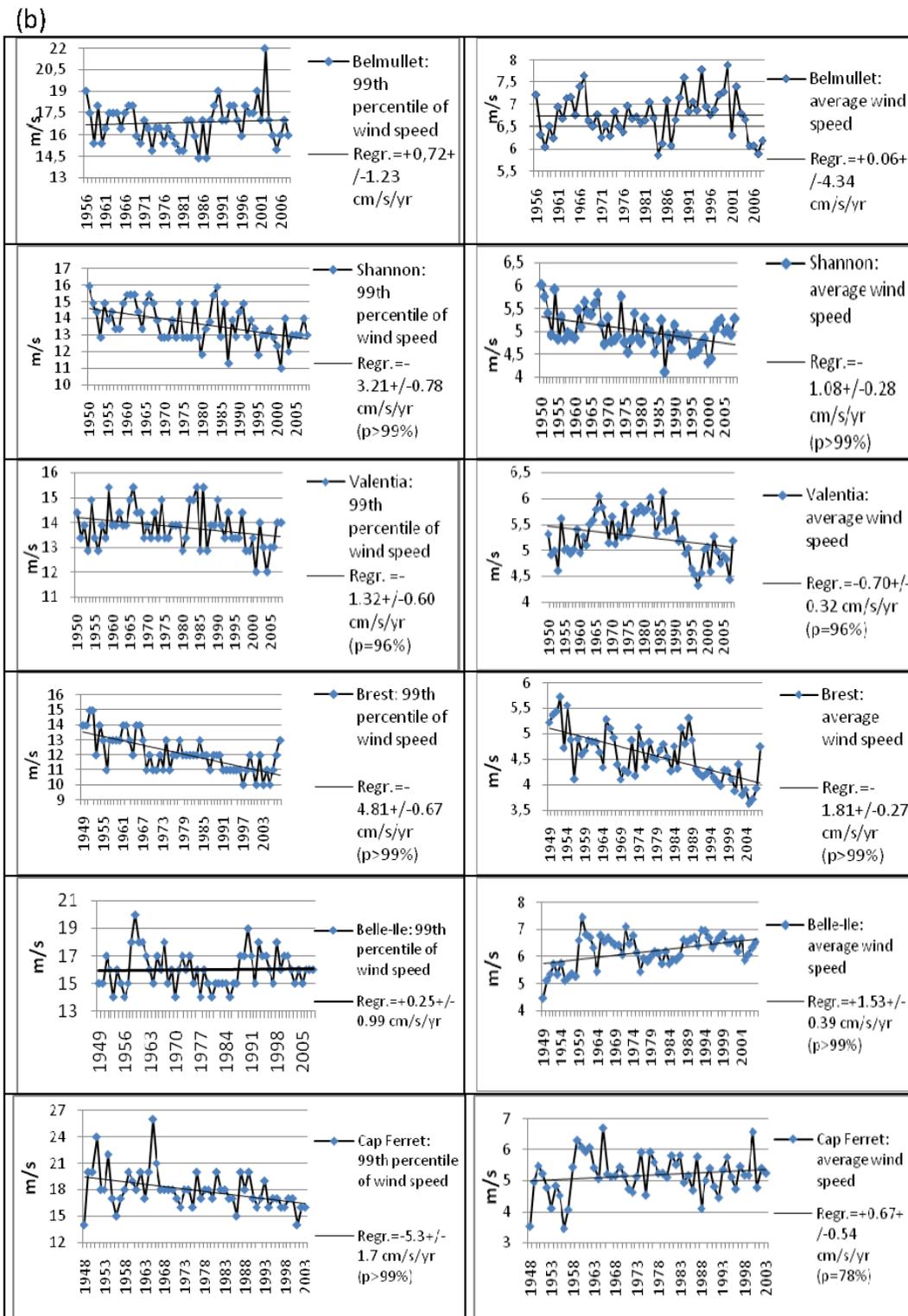


Fig. 4b. Continued.

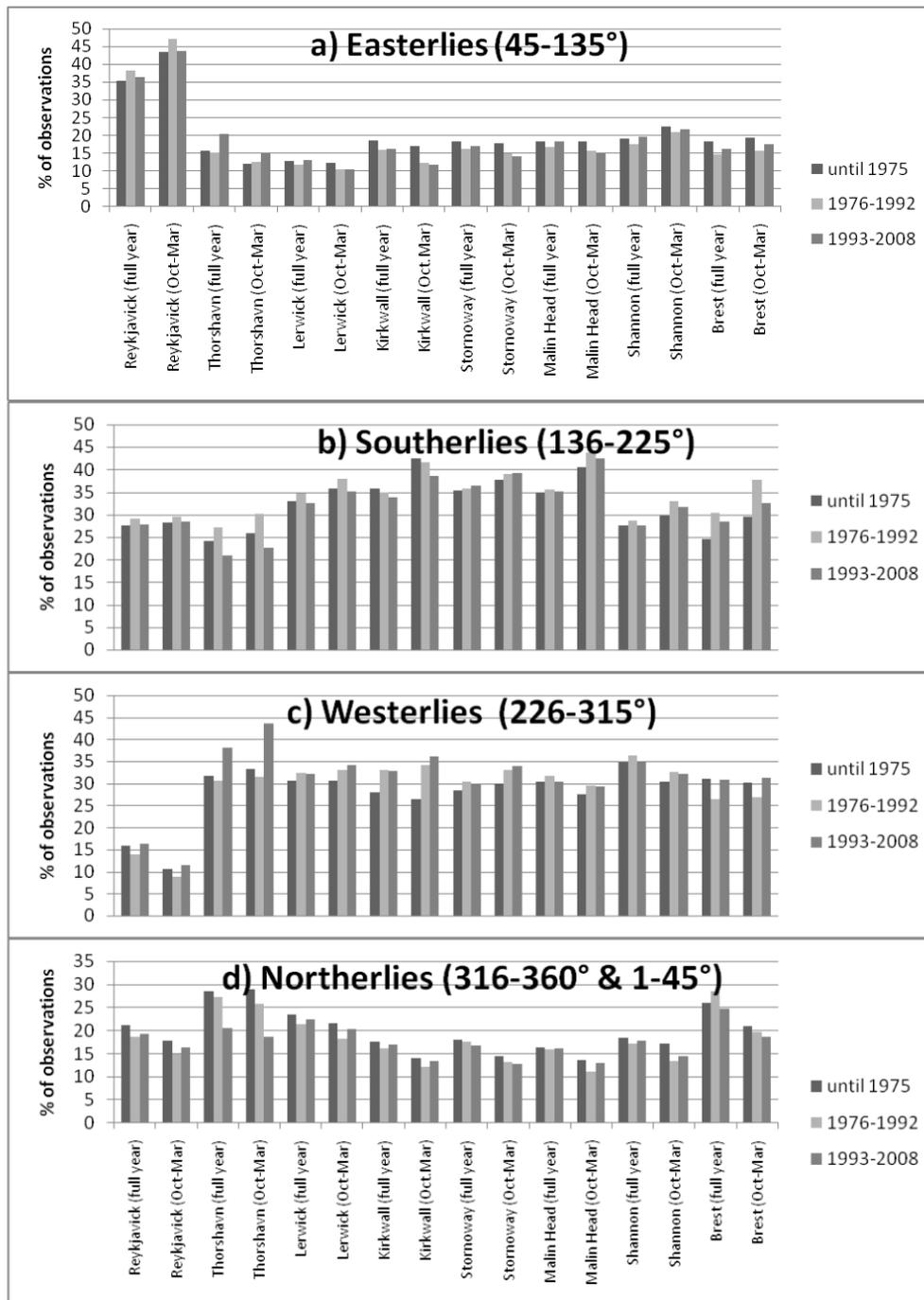


Fig. 5. Frequency of wind measurement in different periods (full year and October to March) at Reykjavick, Thorshavn, Lerwick, Kirkwall, Stornoway, Malin Head, Shannon and Brest. (a) easterlies, (b) southerlies, (c) westerlies, (d) northerlies.

with additional increasing speeds also at Lerwick, Kirkwall, Malin Head, Belle-Ile and Cap Ferret.

3.2 Direction of wind occurrences

In Fig. 5, the percentage frequency from each quadrant is given for the periods until 1975, 1976–1992 and 1993–2008,

for the full year, and for the October–March periods. Similar frequencies for Belmullet, Valentia and Belle-Ile were summarized in a previous publication (Pirazzoli, 2005, Fig. 6).

The wind distribution at Reykjavick is clearly different from that of the other stations considered, with much more frequent easterlies and less frequent westerlies, especially for the winter periods; easterlies generally decrease slightly

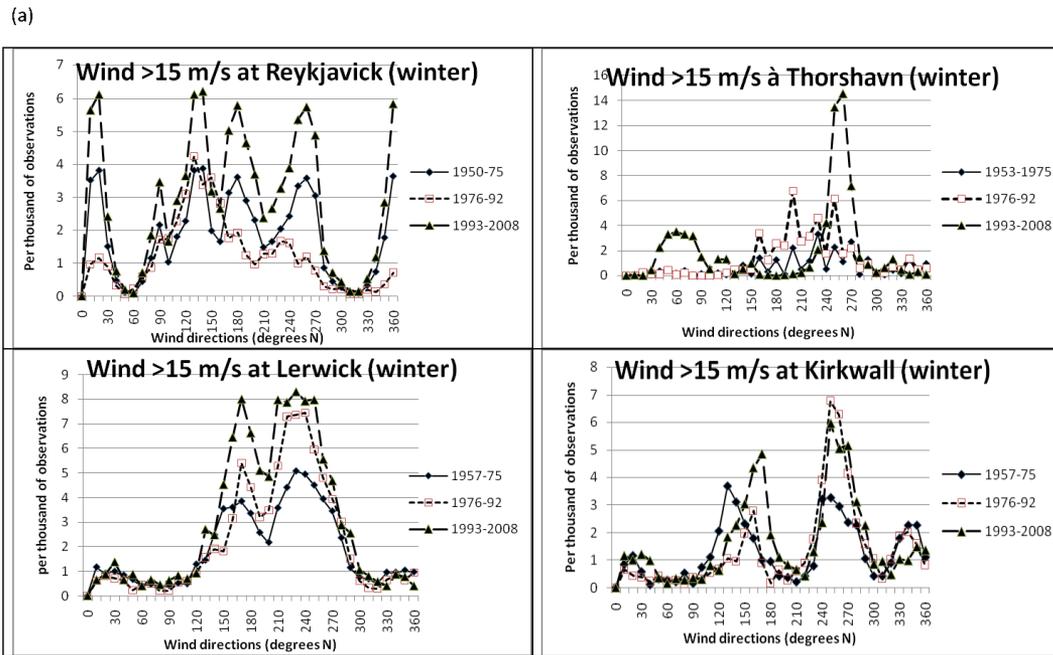


Fig. 6a. Distribution of wind directions for October to March measurements of wind speed $> 15 \text{ m s}^{-1}$ for the periods until 1975, 1976–1992 and 1993–2008: (a) Reykjavick, Thorshavn, Lerwick, Kirkwall; (b) Stornoway, Malin Head, Belmullet and Shannon; (c) Valentia, Brest, Belle-Ile and Cap Ferret.

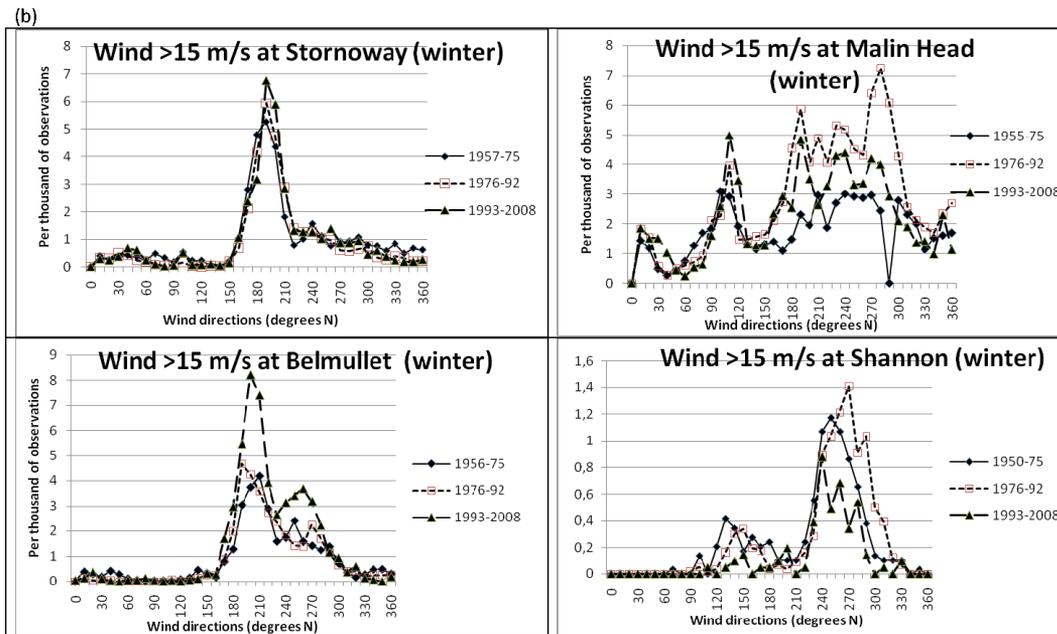


Fig. 6b. Continued.

during winter and after 1975 in the other stations, except at Shannon, where they show a small increase in winter. Similar winter increases for easterlies have been observed at Belmullet and Valentia (Pirazzoli, 2005).

Southerlies predominate between Lerwick and Malin Head, especially in winter and generally show an increase between 1976 and 1992. Westerlies represent about 30% of the observations at most stations (except Reykjavick) and tend to increase since 1993 between Thorshavn and Stornoway

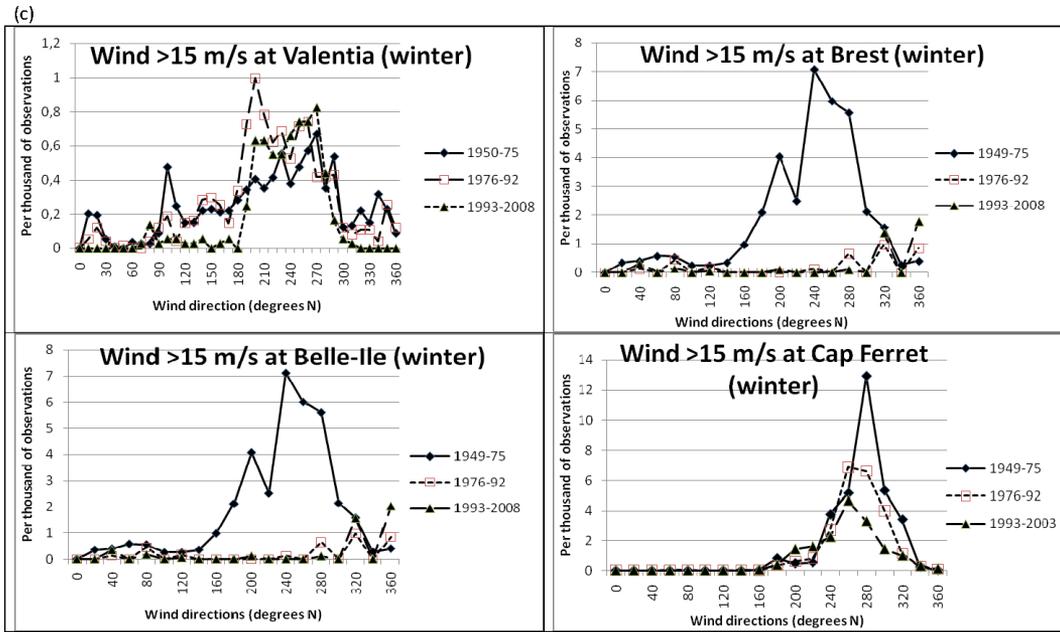


Fig. 6c. Continued.

Table 2. Months between October and March, at each station considered, exhibiting peaks in the angle of wind origin (using sum of speeds).

Station	Uppermost peak	Second peak	Third peak
Reykjavick	O	M	D
Thorshavn	N	M	D
Lerwick	D	M	O
Kirkwall	N	O	D
Stornoway	O	D	J
Malin Head	M	O	J
Belmullet	O	M	J
Shannon	O	M	D
Valentia	O	J	M
Brest	O	D	J
Belle-Ile	O	M	D

(especially in winter). Northerlies tend to decrease with time in most stations, particularly in winter.

3.3 Temporal variability of strong wind frequency

In order to clarify the recent evolution of strong wind occurrences, the frequency of directions for winds exceeding 15 m s^{-1} during the cold season has been reported in Fig. 6 for the periods before 1976, 1976–1992 and 1993–2008. It appears that at Reykjavick (Fig. 6a) the frequency of strong wind has been decreasing between 1976 and 1992, then strongly increasing from various directions after 1993.

At Thorshavn, strong wind has been increasing from all directions after 1975, with a remarkable recent peak of 14‰ of occurrences for westerlies during the most recent period. At Lerwick the increase after 1975 is general for all strong southerlies and westerlies, with a frequency of occurrences that is almost double of that observed before 1976. Also at Kirkwall strong westerlies have increased substantially after 1975, though the highest peak has been recorded during the 1976–1992 period.

Minor changes can be observed at Stornoway (Fig. 6b), whereas at Malin Head the greatest increase for southerlies and westerlies has occurred during the 1976–1992 period. At Belmullet, the frequency of strong southerlies has almost doubled since 1992, while at Shannon and Valentia the frequency of wind speed exceeding 15 m s^{-1} remains low (less than 1.5‰ of the total measurements).

More south, an impressive drop in the frequency of strong westerly wind occurrences has taken place after 1975 along the Atlantic coast of France (Fig. 6c). More detailed information on wind tendencies at Belmullet, Valentia and Belle-Ile has been reported by Pirazzoli (2005).

4 Correlation with the NAO index

Correlation with the boreal cold season NAO index has been computed at all stations for the periods until 1975, 1976–1992 and 1993–2008, for several quantiles of wind speed and also for the positive projection upon certain dominant wind directions deduced from angular distribution of energy flux or from certain peaks of Fig. 6. Among the quantiles of wind

Table 3. Correlation coefficients between wind speed (A: average from all directions; B: the sum of positive projections upon certain dominant wind directions deduced from Fig. 6; C: the sum of positive projections upon the dominant wind direction deduced from angular distribution of energy flux), and NAO index during the cold season (October to March) along a transect of meteorological stations in the NE Atlantic.

Station	Average wind speed	Periods considered		
		Until 1975	1976–1992	1993–2008
Reykjavick	A:	0.59	0.55	0.14
	C: 90°	−0.145	0.21	0.16
Thorshavn	A:	0.02	0.78	0.72
	B: 250°	0.39	0.86	0.44
	C: 260°	0.41	0.85	0.43
Lerwick	A:	0.70	0.81	0.47
	B: 240°	0.38	0.54	0.44
	B: 170°	0.10	−0.04	0.11
	C: 250°	0.36	0.56	0.47
Kirkwall	A:	0.43	0.69	0.28
	B = C: 260°	−0.02	0.82	0.35
	B: 180°	0.04	0.58	0.11
Stornoway	A:	0.42	0.70	0.50
	B = C: 200°	0.23	0.84	0.38
Malin Head	A:	0.01	0.44	0.48
	B = C: 190°	0.57	0.60	0.07
	B: 280°	0.33	0.72	0.41
	B: 110°	−0.46	0.00	−0.40
Belmullet	A:	0.21	0.69	0.72
	B: 210°	0.89	0.75	0.42
	C: 260°	0.62	0.89	0.51
Shannon	A:	0.05	0.27	0.50
	B: 250°	0.46	0.66	0.41
	C: 260°	0.22	0.65	0.40
Valentia	A:	0.20	0.35	0.42
	C: 180°	0.54	0.665	0.305
Ouessant	A:	−0.08	0.47	
Brest	A:	0.16	−0.16	0.56
	C: 210°	0.44	−0.08	0.37
Belle-Ile	A:	−0.05	0.18	0.51
	B = C: 260°	0.28	−0.01	0.45
Cap Ferret	A:	−0.25	0.27	0.00 ^a

^a incomplete record limited to the year 2003.

speed, the most significant for correlation with NAO is indeed the average wind speed for all directions. Results are summarized in Table 3.

At Reykjavick, where no wind direction is dominant in a stable manner (Fig. 6a), a sharp decrease in correlation with NAO has occurred since 1993, corresponding to the increase in frequency of strong wind from various directions. The peak of energy flux from 90° seems unrelated to NAO variability.

At Thorshavn, a marked increase in correlation with the winter NAO index appears after 1975, for the average wind speed from all directions, as well as for projections upon 250° and 260°, that reaches a maximum of over 0.86 during the period 1976–1992. The peak of strong westerlies since 1993 (Fig. 6a) was accompanied by a slight decrease in correlation with NAO.

At Lerwick, a gradual increase in strong southerlies and westerlies is evident since the beginning of the records

(Fig. 6a). Southerlies seem unrelated to NAO (Table 3), whereas westerlies maintain a relatively high level of correlation.

At Kirkwall, correlation is at best during the period 1976–1992. It is difficult, however, to ascribe this result to the strong westerlies (that reach a peak during the same period) because the same winds occurred also (though less frequently) during the preceding period, when projections upon the dominant direction of 260° were uncorrelated with NAO. A similar difficulty exists for projections upon 180° , with a correlation coefficient of 0.58 between 1976–1992 and an absence of significant correlation before 1976 and since 1993, that can be hardly ascribed to clear changes in strong wind frequency for southerlies.

No major changes appear at Stornoway, where the level of correlation with NAO remains relatively high and significant during the various periods considered. Positive correlation is found for all wind directions, with even some improvement, especially between 1976 and 1992, when projections around 200° are considered.

At the Irish stations, strong wind remains limited at Malin Head until 1975 (Fig. 6b) and there is no correlation between the average wind speed and the NAO index (Table 3). Positive correlation is found for wind projection upon 190° (that becomes however insignificant after 1992), as well as around 280° , while a clearly negative correlation exists for projections around 110° before 1976 and after 1992. At Belmullet, southerly and westerly strong wind peaks appear after 1992 (Fig. 6b). Correlation with NAO is positive for all periods, especially for projections of winds around 210° . At Shannon, a gradual increase in correlation can be observed throughout the periods considered for the average wind speed. Westerlies show a strong increase after 1975, followed by a slight decrease after 1992 (Fig. 6b), that may explain the small changes in correlation observed (Table 3). Finally at Valentia correlation with NAO remains limited (between 0.2 and 0.42 for the average wind speed, but some improvement appear for projection upon 180°).

Along the French Atlantic coast, no clear relationship appears between the drop in the frequency of strong winds after 1975 (Fig. 6c) and the absence of correlation with NAO during the period 1976–1992, or with the improved correlation after 1993. At Cap Ferret, where our record is limited to 2003, the average wind speed seems anticorrelated to NAO.

In short, it is possible to cut the area considered into two different parts:

1. in the northern part (Northern Atlantic stations and probably also in the North Sea area), changes in correlation between wind speed and NAO are usually consistent for most winds and also for the dominant ones. When the NAO becomes more positive (negative), the correlation with wind speed increases (decreases), according to storm track variations;

2. in the southern part (French Atlantic stations), when the NAO becomes more positive (negative), the correlation with wind speed decreases (increases). Then, for the long-term trends in wind speed, it is more difficult to find a clear explanation. Nevertheless, the discrepancies between stations may be associated with the direction of local dominant winds at each station.

For example, at Reykjavick, easterlies are dominant. This is consistent with the fact that the main storm tracks pass within the $60\text{--}65^\circ$ N belt, thus, south of Reykjavick. At Thorshavn, westerlies are dominant for a similar reason (this station is just below the main storm tracks).

5 Conclusions

This work considers the “fire front” for storms in the NE area of the Atlantic Ocean, down through the coasts of Western Europe, from 1950 to 2008, using a large collection of wind data. It is shown that there is a great difference in the evolution of recent storm frequency between Iceland and the northernmost European islands located a few degrees eastward. In the former, prevailing winds (especially in winter time) are easterlies, whereas in the latter westerlies or southerlies predominate; in the former, extreme and average wind speed decrease, whereas they increase sharply in Thorshavn and average wind speed increases moderately also at Lerwick and Kirkwall.

Connection with the values of NAO is unavoidable: the evolution of local winds is compared to the one of this large scale parameter (a lot more robust, being related to the atmospheric pressure) and something will be satisfactory, something will not. Over the period 1950–2008, considering the boreal cold season, the NAO pattern shows an important decadal variability and appears to be intense, with its changes in polarity. Observing the three periods of the present study, the central one is impressive with its positive excess (Fig. 3). In this case, the main storm track is displaced more and more northwards. This could be consistent with the decrease of wind speed at Reykjavick, where the speed of the easterlies should decrease. Similarly, Thorshavn increase of westerlies speed appears fully consistent with the positive polarity of the NAO.

This contrasts with the sharp decrease in extreme and average wind speed along the west coast of Ireland, that seems to extend more south to most Atlantic French coast (Pirazzoli et al., 2008) and even to the Mediterranean (Pirazzoli and Tomasin, 2003).

Some studies stress the observed correlation with NAO of climate changes over the NE Atlantic or even in western Europe. The present study, based on wind measurements along a transect on the western border of Europe, has shown that even if some recent changes in wind activity indeed exist, they are not always in agreement with NAO. A great temporal and spatial variability can be observed even for stations

that are relatively close to each other and could, therefore, be expected to be affected by similar storm tracks. Of course, local geographical characteristics could play an important role in the discrepancies between close stations.

In short, there is not a uniform tendency, characterised by a general northward shifting of the storm tracks, as claimed by Wang et al. (2006), but a more variable evolution, suggesting rather the occurrence of mutable individual storm trajectories.

With some generalizing, if the average wind speed is considered, it can be said that correlation coefficients above 0.4 for the period between the early 1950s and 1975 are observed at Reykjavick and between 60° (Lerwick) and 58° N (Stornoway), whereas little positive coefficients are reported more south. During the period 1976–1992, characterised by a clear increase of the NAO index (Fig. 3), positive correlation improves southwards as far as 54° (Belmullet), with some improvement also at Shannon and Valentia, while it remains low or even negative near the French Atlantic coast. Finally in the most recent 1993–2008 period, correlation improves at all stations south of 55° N (Malin Head), while it weakens more north (Reykjavick, Kirkwall).

If projections upon dominant wind directions are considered, the variability is even greater, with a general increase in correlation between 64° and 52° during the period 1976–1992 in relation to the preceding period, while a decrease predominates south of Valentia. Lastly, since 1993, a decrease in correlation between wind activity and NAO is general, except south of 52°, where a clear increase is observed at Brest and Belle-Ile.

Though it can be confirmed that station measurements of “winter” pressure increase in Iceland (as reported in the literature) correspond to an overall tendency towards less extreme wind events at Reykjavick (Fig. 4a), nevertheless, the frequency of wind $> 15 \text{ m s}^{-1}$ from several directions has been increasing from several directions during the 1993–2008 period (Fig. 5a). Since wind measurements could be quite inhomogeneous, and pressure measurements are usually more reliable, the interpretation of our results should be more focused on pressure analysis: when there is a consistency between wind and pressure, it means that the results are robust. To clarify this point a new study of changes in measured air pressure at the same stations of the present study and for the same periods is in preparation by the same authors.

Finally this paper shows: (i) the occurrence of average changes in meteorological condition between different latitudinal areas, and (ii) that discrepancies may exist between close stations within each of these areas, probably linked with some small variations in storm tracks (especially in the northern part, that is more sensitive to these small variations).

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