

Imaging observations of nighttime mid-latitude F-region field-aligned irregularities by an MU radar ultra-multi-channel system

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Abstract. Mid-latitude F-region field-aligned irregularities (FAIs) were studied by using the middle-and-upper atmosphere (MU) radar ultra-multi-channel system with the radar imaging technique. On 12 June 2006, F-region FAI echoes with a period of about one hour were observed intermittently. These echoes were found to be embedded in medium-scale traveling ionospheric disturbances (MSTIDs) observed as variations of total electron content (TEC). The echoes drifting away from (toward) the radar were observed in the depletion (enhancement) phase of the MSTID. The Doppler velocity of the echoes is consistent with the range rates in the range-time-intensity (RTI) maps. Fine scale structures with a spatial scale of 10 km or less were found by the radar imaging analysis. Those structures with positive Doppler velocities (moving away from the radar) appeared to drift north-(up-) westward, and those with negative Doppler velocities south- (down-) eastward approximately along the wavefronts of the MSTID. FAIs with positive Doppler velocities filling TEC depletion regions were observed.

Keywords. Ionosphere (Ionospheric irregularities; Mid-latitude ionosphere; Plasma waves and instabilities)

1 Introduction

In the mid-latitude ionosphere, several kinds of plasma irregularities are known to exist in both the E- and F-regions. Nighttime F-region 3-m-scale field-aligned plasma irregularities (FAIs) associated with the mid-latitude spread-F were first reported by Fukao et al. (1988). Compared with those in the E-region, such as quasi-periodic (QP) echoes associ-

ated with the sporadic-E (E_s) layer (Yamamoto et al., 1991), they have not been studied well. Only a few studies in Japan (Fukao et al., 1988, 1991; Saito et al., 1998, 2002) and in Puerto Rico (Swartz et al., 2000) using radars have reported the mid-latitude F-region FAIs. This is partly because the echoes associated with the mid-latitude F-region FAIs are relatively weak and a powerful radar is required to observe them.

Several characteristics have been revealed by the studies cited above. Nighttime mid-latitude F-region FAIs associated with medium-scale traveling ionospheric disturbances (MSTIDs) with typical wavelengths of 100–500 km have been observed (Saito et al., 1998, 2002). However, F-region FAIs are less frequently observed than MSTIDs. Accordingly, the mechanism of generating 3-m-scale FAIs in the F-region has not been understood well.

Perkins (1973) has proposed that the ionosphere can be unstable and the wave structure excited by the instability, called the Perkins instability, preferentially propagates southwestward or northeastward (see also Zhou and Mathews, 2006). Nighttime mid-latitude MSTIDs are often explained by the Perkins instability. However, it is also known that the linear growth rate of the instability is too small to explain the MSTID phenomenon. The Perkins instability cannot explain how 3-m-scale FAIs are generated, because they are supposed to be secondary waves generated after non-linear evolution of primary waves.

Kelley et al. (1991) proposed that interchange-type instability could be excited by the electron density gradient associated with MSTIDs and the neutral wind. In this case, a certain slope of MSTIDs with $\nabla n_e \cdot \mathbf{u} < 0$, where n_e and \mathbf{u} are the electron density and the neutral wind velocity, respectively, would be unstable, and the other slope would be stable. Using the middle-and-upper (MU) atmosphere radar (Fukao et

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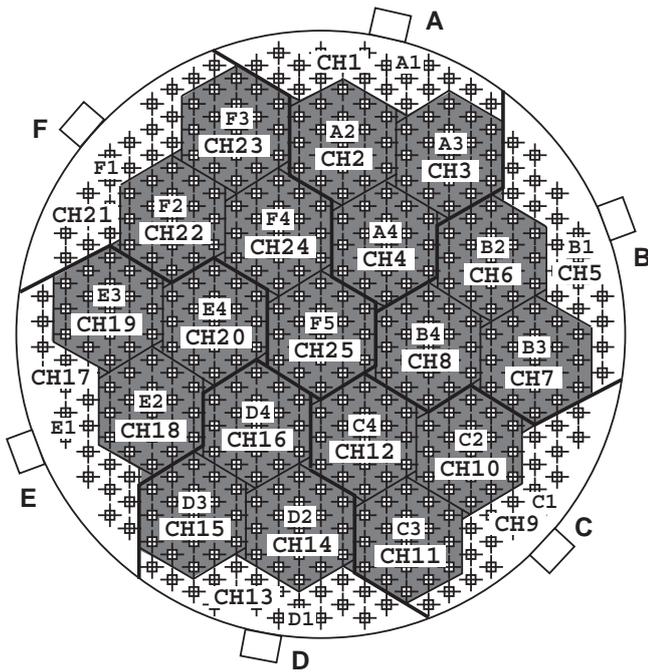


Fig. 1. Layout of the MU radar antenna sub-arrays. The 19 shaded sub-arrays were used to take data independently for imaging analysis. All 25 sub-arrays were used for transmission.

al., 1985a, b) multi-beam experiment, Saito et al. (2002) reported that echoes tend to appear at the northern slope of MSTIDs. However, it was difficult to see the precise relationship between the MSTIDs and FAIs, nor fine scale structures of FAIs, as their observations were conducted without interferometry.

Zhou et al. (2005) showed by numerical simulation that the non-linear Perkins instability would have a larger growth rate and may generate 10-km-scale irregularities. They also showed that the interchange-type instability can generate small-scale irregularities.

The radar imaging technique is now widely used for studying fine structures of ionospheric irregularities from equatorial to high latitudes (Hysell, 1996; Hysell and Burcham, 2000; Hysell and Chau, 2002; Hysell et al., 2002, 2004; Bahcivan et al., 2005; Saito et al., 2006, 2007). With this technique, structures smaller than the radar beam width can be resolved. Using the MU radar ultra-multi-channel system, we conducted multi-beam radar-imaging experiment to study the structures of mid-latitude F-region FAIs in detail.

2 Experiment

All 25 antenna sub-arrays of the MU radar ultra-multi-channel system were used for transmission, and 19 sub-arrays with a regular hexagonal shape recorded returned signals independently (Fig. 1). The observations were con-

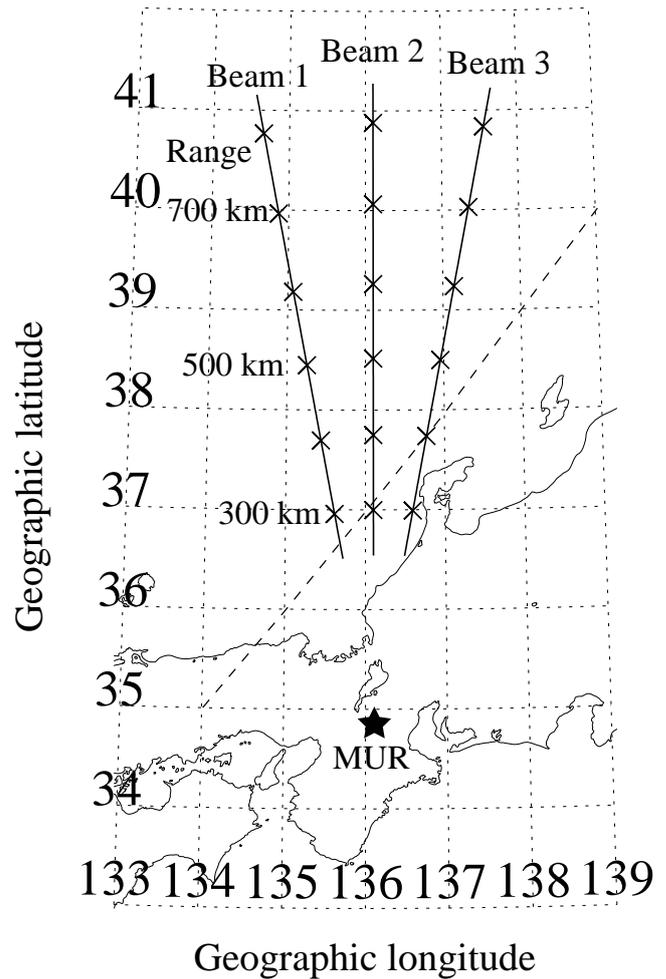
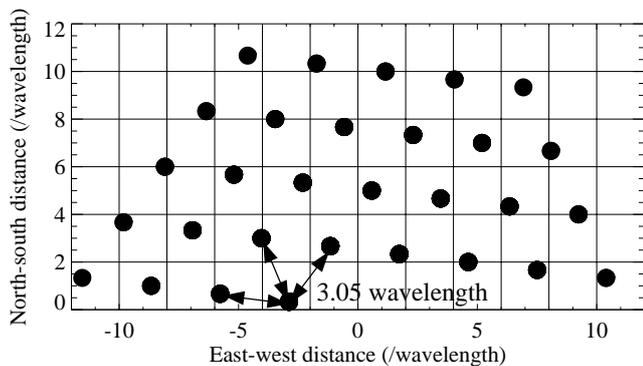


Fig. 2. Geometry of the observation. Three thick lines show the center of the radar beam with crosses marked at every 100 km range.

ducted from 20:00 to 04:00 JST (UT+9 h) on 12–15 June and 11–13 July 2006. E- and F-regions were observed alternately. Parameters of the MU radar observations of the F-region are summarized in Table 1. Figure 2 shows the geometry of the experiment. The beams used for the observations were named beams 1, 2, and 3 from west to east, respectively. Due to the high aspect angle sensitivity of FAI echoes, echoes can be observed only when the radar k-vector is perpendicular to the geomagnetic field line. The perpendicularity condition in the -3 dB beam width (4.5°) was achieved at altitudes 208–336 km, 224–347 km, and 161–295 km for beams 1, 2, and 3, respectively. These altitudes correspond to ranges 350–619 km, 384–648 km, and 274–547 km, respectively. Although the MU radar can switch beam directions on a pulse-to-pulse basis, beam directions in this experiment were switched after completion of a measurement for each beam in order to keep their Doppler spectrum widths wide enough. One experiment cycle took 91 s, and nominal dwell time of each F-region beam was about 17 s.

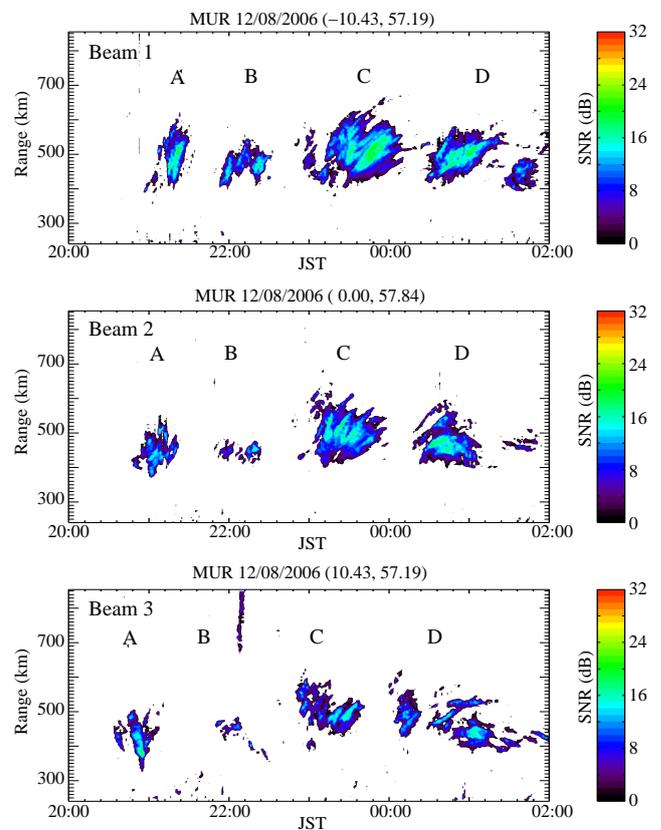
Table 1. Parameters of the MU radar observations.

Number of beams	3
Beam direction (Azimuth, Zenith)	(-10.43° , 57.19°), (0.00° , 57.84°), (10.43° , 57.19°)
Sensitive altitude range	208–336 km, 224–347 km, 161–295 km
Pulse code	11-bit Barker code
Sub-pulse length	$32 \mu\text{s}$ (4.8 km)
Inter pulse period	8 ms
Doppler velocity bandwidth	$\pm 201 \text{ m s}^{-1}$
Number of incoherent integration	5
Peak power	1 MW

**Fig. 3.** Non-redundant baselines achieved by the 19 antenna sub-arrays and used for imaging analysis.

3 Imaging analysis

The analysis procedures of radar echoes are basically the same as used for the E-region measurements by Saito et al. (2006). Complex Doppler spectra were calculated by fast Fourier transform (FFT). The two-dimensional visibility function was estimated for 30 non-redundant baselines achieved by the 19 antenna sub-arrays (Fig. 3). The coherence functions were divided into four segments in the Doppler velocity domain. Images of echoes were reconstructed by the maximum entropy method (Hysell, 1996) for three of the four segments of Doppler velocities, toward the radar (-151 to -50 m s^{-1}), around zero (-50 to $+50 \text{ m s}^{-1}$), and away ($+50$ to $+151 \text{ m s}^{-1}$) from the radar. The fastest segment (-151 to -201 m s^{-1} and $+151$ to $+201 \text{ m s}^{-1}$), which had more chance to be frequency aliased than the other segments, was not used, because the F-region echo spectra are broad with relatively small Doppler shifts that are lower than the ion acoustic velocity (about 300 to 400 m s^{-1}). In the present study, the positive and negative Doppler velocities refer to the motion away from and toward the radar, respectively.

**Fig. 4.** Echo power observed on 12 June 2006 in the RTI format (JST=UT+9 h). Upper, middle, and lower panel show the data from beams 1, 2, and 3, respectively.

4 Results

Figure 4 shows the echo strength observed by the three beams on 12 July 2006 in the range-time-intensity (RTI) format. Echoes were observed at the expected ranges in all beams. Echoes with signal-to-noise ratio (SNR) as strong as 15 dB were intermittently observed around 21:00, 22:00, 23:00, and 01:00 JST (UT+9 h) with duration of 0.5 to 1 h as marked as A to D, respectively, in the figure. Most of echoes were observed with positive range rates, i.e. the range of the

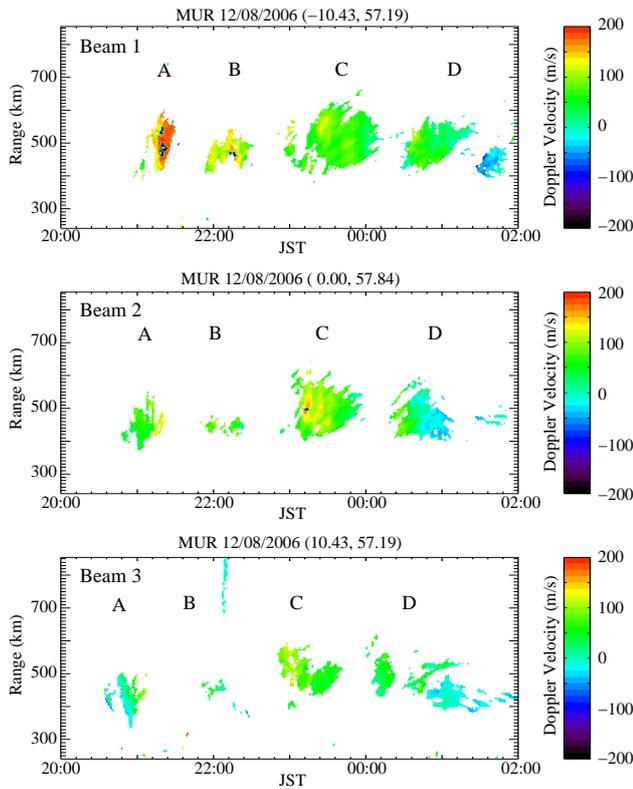


Fig. 5. Same as Fig. 4, except the Doppler velocity is shown.

echo pattern increased with time. Echo cluster A in beam 3 and echo cluster D in all beams have both positive and negative range rates. The first half of echo cluster D has positive range rates, and the latter half of the echoes has negative range rates.

Figure 5 shows the Doppler velocities observed by the three beams over the same period. Doppler velocities of echoes with SNR larger than 0 dB are plotted. The Doppler velocities were mostly positive, and about 50 to 150 m s^{-1} . However, some echoes have negative Doppler velocities of about -50 m s^{-1} . There are several echoes with very high negative Doppler velocities as fast as -200 m s^{-1} in the echo clusters A and B in beam 1 and in the echo cluster C in beam 2. However, spectral analysis of these echoes shows that they are most likely due to frequency aliasing by very fast, positive Doppler velocity exceeding the Nyquist limit (202 m s^{-1}). Echoes with positive range rates have positive Doppler velocities (moving away from the radar) and those with negative range rates had negative ones (moving toward the radar). The observed Doppler velocities were consistent with the range rates, as reported in previous works (e.g. Fukao et al., 1991).

Figure 6 shows an example of a sequence of radar images, when a strong echo cluster (marked as C in Fig. 4) was observed. Images for each one of the Doppler velocity segments (-150 to -50 m s^{-1} , -50 to $+50 \text{ m s}^{-1}$, and $+50$ to

$+150 \text{ m s}^{-1}$) were merged by assigning weights proportional to echo power and image intensity. The colors blue, green, or red, was assigned to each Doppler velocity segment: toward, zero, or away, respectively. Each color represents a range from 0 to 3 dB with six color steps to show weak echoes more clearly. The images were projected on a horizontal plane at an altitude of 300 km. Echo cluster C contains a few echoes with Doppler velocities higher than 150 m s^{-1} . These echo spectra are broad, and a significant amount of echo power in the spectral ranges that were used for imaging analysis (-150 to $+150 \text{ m s}^{-1}$). Thus, the images of such echoes with very high Doppler velocities were also successfully reconstructed. The images shown in Fig. 6 are dominated by a red color, indicating that the Doppler velocity was higher than 50 m s^{-1} . The Doppler echoing regions consists of substructures with scale sizes less than 10 km. As indicated by white circles in each panel, these substructures appeared to drift north- (up-) westward. This is consistent with Swartz et al. (2000)'s finding of 1 to 100 km scale layering in the mid-latitude F-region irregularities. After 00:40 JST on 13 June 2006, echoes with negative range rates were observed in beam 2 in the latter half of the echo cluster D in Fig. 4. Figure 7 shows a sequence of radar images of the echoes from the latter half of the cluster D in beam 2. All echoes in this echo cluster had Doppler velocities between -150 and $+150 \text{ m s}^{-1}$. The images were dominated by a green color, indicating that the Doppler velocity was mostly negative (toward) and slower than -50 m s^{-1} . Echoing regions also consist of patches with scale sizes less than 10 km. These echoing regions appeared to drift south (down) eastward. Similar images were also observed in the echo clusters A and B.

From 20:00 JST on 12 July 2006 to 02:00 JST on 13 July 2006, in the region where MU radar echoes were observed, strong medium-scale traveling ionospheric disturbances (MSTIDs) were observed in the total electron contents (TECs) measured by the GEONET GPS receiver network. MSTIDs were observed as TEC perturbation which was derived every 30 s as deviation from one hour moving average of TECs. In converting slant TECs to vertical ones, the contribution of electrons to the slant TECs is assumed to be entirely from a virtual thin layer at an altitude of 300 km. This is the same method as that developed by Saito et al. (1998). Figure 8 shows the temporal variations of TEC perturbation along the line passing through the region observed by the MU radar which is drawn in Fig. 2 as an oblique dashed line. Oblique stripes from top left to bottom right indicate the occurrence of MSTIDs propagating southwestward. The amplitude was about $\pm 1 \text{ TECU}$ (10^{16} m^{-2}). The echoes with positive (negative) range rates were observed in the negative (positive) phase of MSTIDs. Figure 9 shows the relationship between FAI images and MSTIDs. The FAI images are projected on a horizontal plane at an altitude of 300 km. All these echoes have positive Doppler velocity (moving away from the radar) as indicated by reddish colors. It can be seen that the FAIs filled the TEC depletion region,

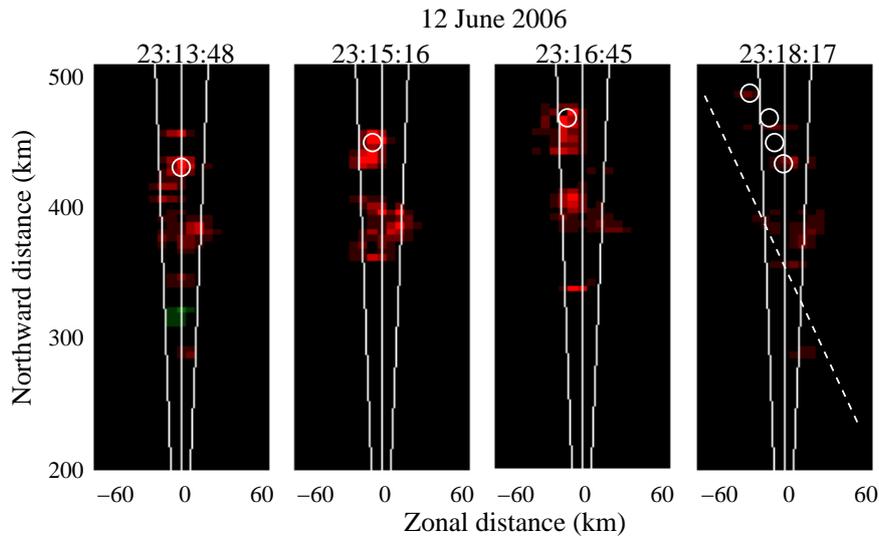


Fig. 6. Sequential images of echoes from beam 2 cluster C from 23:13:48 to 23:18:17 JST on 12 June 2006. The images are projected on a horizontal plane at an altitude of 300 km along the geomagnetic field line. White circles in the rightmost panel show the approximate positions of echo patches in these four images as also indicated by circles in other panels. The dashed line in the rightmost panel indicates the wavefront direction of the MSTID (at an arbitrary phase).

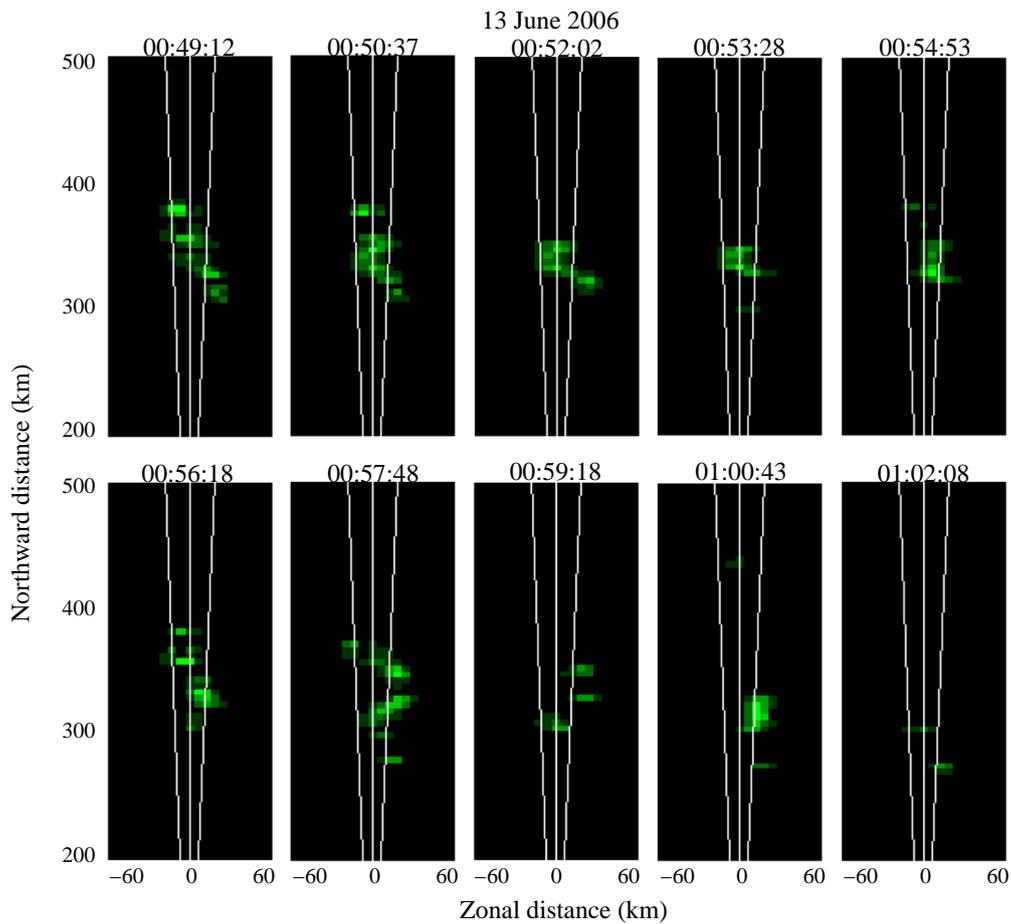


Fig. 7. Same as Fig. 6, except echoes from beam 2 cluster D from 00:49:12 to 01:02:08 JST on 13 June 2006 are shown.

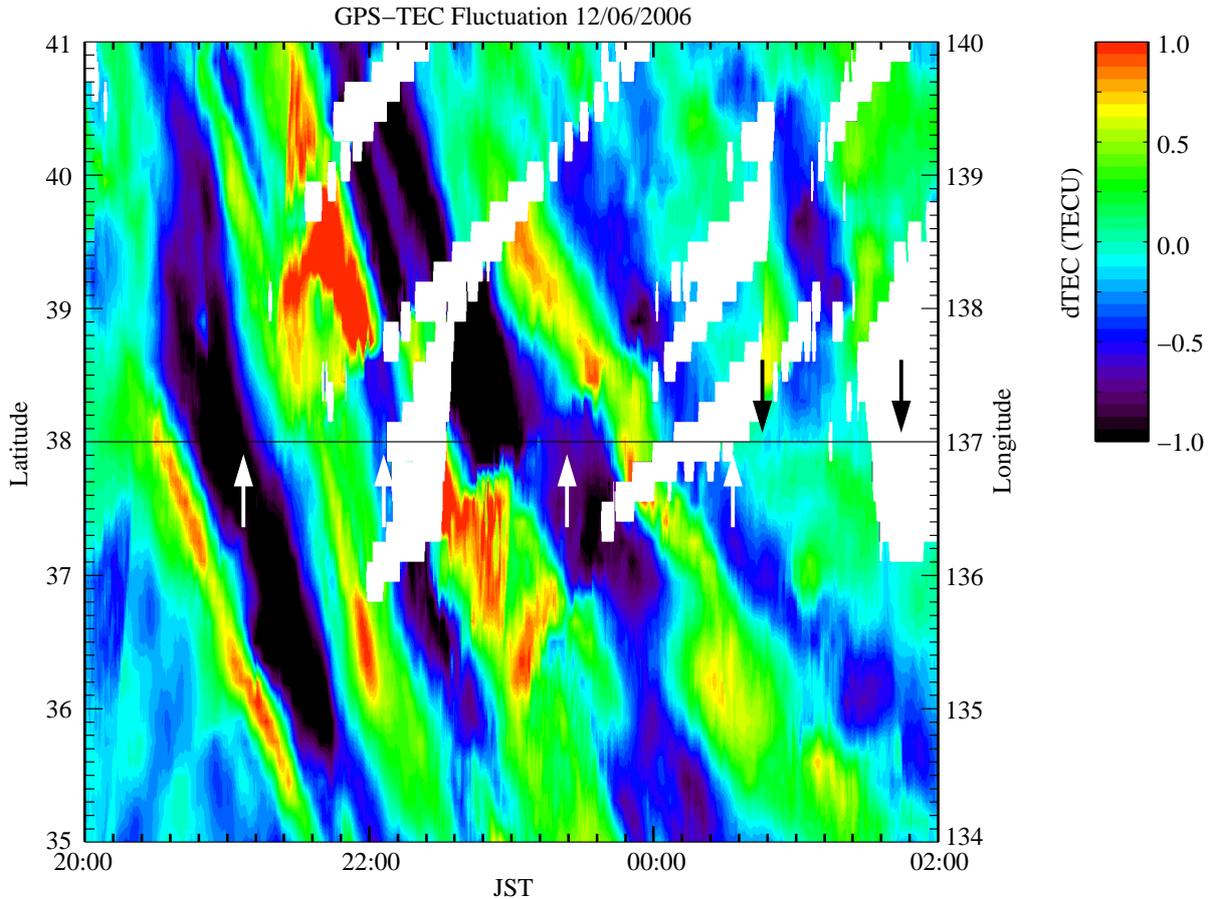


Fig. 8. TEC perturbation along the line from (35° N, 134° E) to (41° N, 140° E) (dashed line in Fig. 2 approximately perpendicular to the MSTID wavefront). Up and down arrows indicate the times when FAIs with positive and negative range rates, respectively, appeared in beam 2.

and moved southwestward as a whole clearly together with the MSTID.

5 Discussion and summary

Our observations by the MU radar ultra-multi-channel system show that substructures (~ 10 km-scale) of FAI echoes drifted approximately along the MSTID wave front, while echoing regions as a whole drifted southwestward together with the MSTID. These 10-km-scale FAI substructures could be related to the kilometer-scale layered structure reported by Mathews et al. (2001). The drift of these echo patches may be explained by the electric field associated with the MSTID. Figure 10 illustrates the motion of the irregularity patches and its relation to the MSTID. At night in the Northern Hemisphere, the neutral wind velocity in the F-region is generally southeastward. The wind, \mathbf{U} , generates the Pedersen current, $\mathbf{J} = \sigma_P \mathbf{U} \times \mathbf{B}$, where σ_P is the Pedersen conductivity, as shown in Fig. 10. According to Saito et al. (2002), in the TEC depletion region ($dN_e < 0$), the polarization electric field, E_p ,

would be northeastward and lift up the ionosphere. On the other hand, in the TEC enhancement region ($dN_e > 0$), the polarization electric field, E_p , would be southwestward. By the E_p associated with the MSTID and the geomagnetic field, $\mathbf{E}_p \times \mathbf{B}$ drift would be north- (up-) westward (south- (down-) eastward) in the TEC depletion (enhancement) region (see also Otsuka et al., 2007). Using satellite observations and a 630 nm airglow imager, Shiokawa et al. (2003) found that the plasma drift was along the MSTID wavefront and was north-westward (southeastward) in the 630 nm airglow depletion (enhancement) region. This corresponds to the ionosphere that is uplifted (pushed down), and hence the TEC depletion (enhancement) region, and is consistent with our finding of the small-scale echo patches drift. The small-scale patches seem to move with the background plasma. This agrees with the numerical simulation results by Zhou et al. (2006). For quantitative verification if the FAI patches really move with background plasma, neutral wind measurements are necessary. The Fabry-Perot interferometer of Nagoya University at the MU radar site was unfortunately not in operation on 12

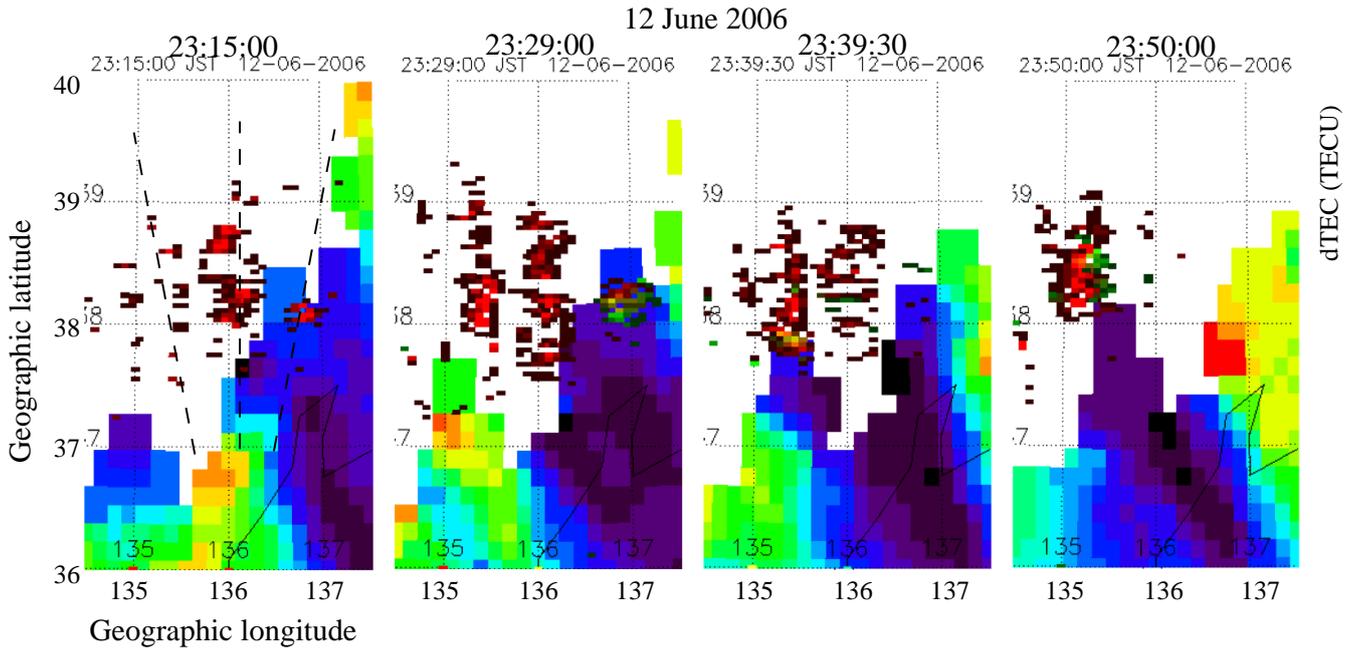


Fig. 9. Sequential images of echoes from beam 1, 2, and 3 together with TEC perturbation observed from 23:15 to 23:50 JST on 12 July 2006. The images are projected on a horizontal plane at an altitude of 300 km along the geomagnetic field line. Dashed lines in the images show the center of the radar beams.

June 2006 due to maintenance. In our future studies, simultaneous observations of FAI images and the neutral wind are planned.

Recently, Saito et al. (2007) reported that MSTIDs are closely related to quasi-periodic (QP) echoes in the E-region as evidence of the E- and F-region coupling. They have also pointed out different scale sizes in the E- and F-regions: 100 to 500 km for MSTIDs and several kilometers for QP echoes. Existing theories regarding E- and F-region coupling, which assume the same scale sizes in both regions, would have difficulty in explaining the difference in scale sizes. In this experiment, we found smaller scale (~10 km) structures in the F-region FAIs associated with MSTIDs. In fact, on 12 June 2006, E-region echoes, that could be classified as QP echoes, were observed. Although MSTIDs do not always accompany FAIs, our results show possibility that small-scale structures exist in the F-region with a scale size comparable to that of the E-region.

We showed that the FAIs were found filling the TEC depletion region. Using an MU radar multi-beam experiment without interferometry, Saito et al. (2002) reported that echoes tend to appear at the northern slope of MSTIDs. However, our experiment with better spatial resolution showed that these echoes with positive range rate were distributed over the TEC depletion region of the MSTID irrespective of TEC gradients. This situation is quite similar to the case of plasma bubbles. The western wall of a plasma bubble is more unstable with the interchange instability due to the eastward thermospheric neutral wind. However, 3-m-scale FAIs fill entire

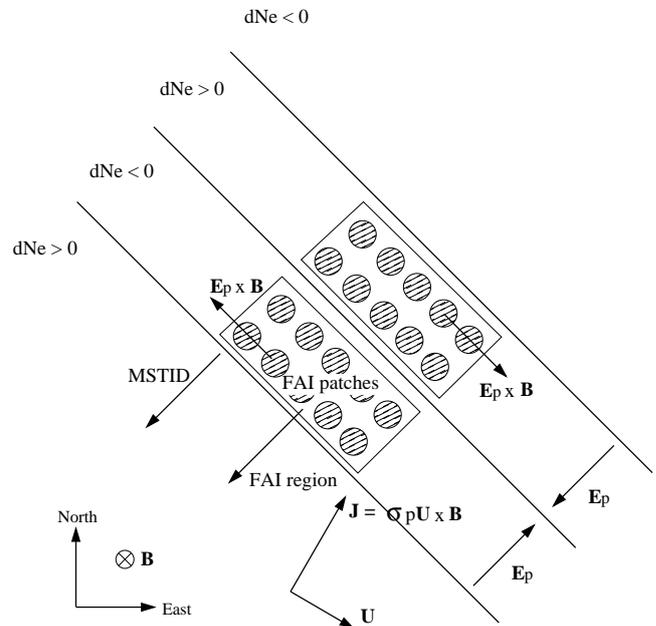


Fig. 10. Schematic illustration of motion of FAI patches with MSTIDs.

plasma bubbles from the western to the eastern walls (Otsuka et al., 2004). In the case of MSTIDs, even though FAIs could be generated at a certain phase of MSTIDs, they could soon develop into other regions. To test this hypothesis, we need to trace the development of FAIs. Multi-beam imaging

observations can distinguish whether they are generated there or generated far away and drifted into the radar field-of-view. Unfortunately, no FAIs that were observed in our experiment were generated in the radar field-of-view. Even the echo cluster D, which may look generated in the radar field-of-view if we look at the RTI map of only one beam, was in fact generated outside the radar field-of-view. It is necessary to conduct more experiments of this kind to enhance the chance to encounter “fresh” F-region FAIs.

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