

An interpretation of ion composition diurnal variation deduced from EISCAT observations

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Abstract. Physical interpretation is made of the O⁺/Ne diurnal variations in summer, revealed by Litvine et al. (1998) from the EISCAT observations. It is shown that the observed anti-correlation between the Z_{50} parameter, corresponding to the transition region between 50% of molecular and atomic ions, and the width D_z of the transition, defined as the altitude width between 10% and 90% of the O⁺/Ne ratio, can be reproduced in model calculations and the result of different recombination laws (quadratic in the lower and linear in the upper ionosphere) as well as diurnal variations in the photo-ionization rates.

Key words. Ionosphere (ion chemistry and composition; modeling and forecasting)

1 Introduction

Relative ion composition (O⁺/Ne ratio) deduced from EISCAT observations demonstrates diurnal variations as was depicted by Lathuillère and Pibaret (1992) and Litvine et al. (1998). The Z_{50} parameter, corresponding to the transition region between 50% of molecular and atomic ions, and the width D_z of the transition, defined as the altitude width between 10% and 90% of the O⁺/Ne ratio, correlate in their diurnal variations during winter and equinoctial periods with Z_{50} and D_z being the smallest around noon hours and the largest around midnight. Such variations are typical of winter and equinoctial periods and they were revealed in the earlier ion composition model by Lathuillère and Pibaret (1992). On the contrary, Z_{50} and D_z anti-correlate in their diurnal variations for the summer period both at solar maximum and minimum, as it was shown for the first time by Litvine et al. (1998). This peculiarity of summer O⁺/Ne variations needs physical interpretation. The aim of this paper is to explain these summer Z_{50} and D_z diurnal variations using model

calculations of ion composition in the F-region, along with EISCAT observations for summer sunlit ionosphere.

2 Observations

The summer geomagnetically quiet period of 02/03 July 1990, with $A_p = 6/7$ and $F_{10.7} = 276.6/253.8$, was chosen for our analysis. The electric fields observed with EISCAT were small (2–5 mV/m) during the whole period in question. Observed NmF2 and hmF2 variations are shown in Fig. 1. F2-layer maximum electron concentration diurnal variation is observed to be rather small, with NmF2 maximizing soon after midnight (02 UT), while maximum hmF2 are observed around midnight. This is the so-called effect of diurnal anomaly related to the direct solar photo-ionization of the F2-region during nighttime hours and large hmF2 (Fig. 1, bottom) which results from the equatorward thermospheric wind during this part of the day. Decreased recombination efficiency due to large hmF2, along with direct (although small) photo-ionization of the F2-region, result in nighttime NmF2 values larger than the daytime ones.

The EISCAT CP-1 program provides range profiles of N_e , T_e , T_i and V_i every 5 min, with the antenna beam directed along the local geomagnetic field line. They were used to calculate median profiles over 1.5–2 hours of observations (17–25 values at each height) for the chosen periods. These median vertical profiles were then smoothed by a polynomial (up to the 5th degree) fitting before being used in calculations.

3 Model calculations

The sunlit auroral ionospheric F-region above the EISCAT facility in geomagnetically quiet conditions when electric fields are small is controlled by local processes typical for the midlatitude ionosphere (Farmer et al., 1984; Lathuillère and Brekke, 1985). Therefore, a model which takes into account the photo-ionization of neutral atmospheric species by

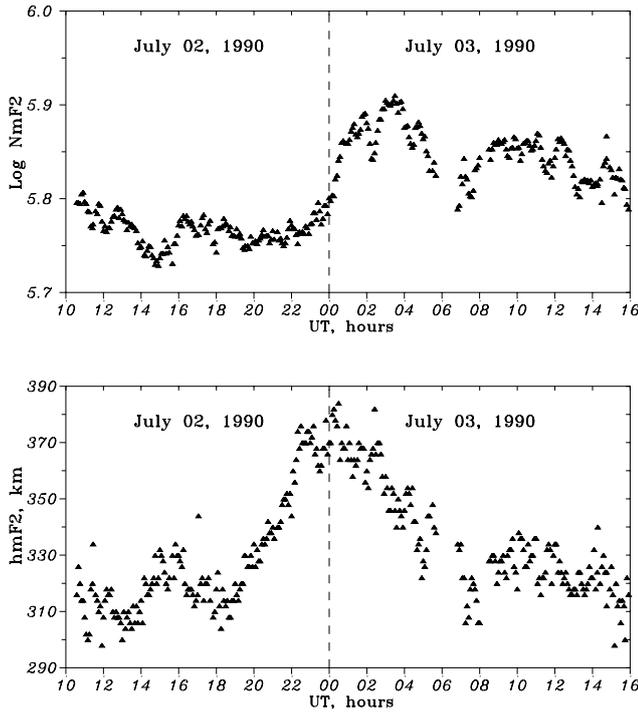


Fig. 1. Observed with EISCAT diurnal variations of NmF2 and hmF2 on 02/03 July 1990.

solar EUV, a set of main chemical processes, vertical plasma transport due to diffusion, thermospheric winds and electric fields may be used for such an analysis. The model used earlier for EISCAT and Millstone Hill incoherent scatter (IS) data analysis (Mikhailov and Foster, 1997; Mikhailov and Förster, 1997; Mikhailov and Schlegel, 1997, 1998) includes: transport process for $O^+(^4S)$ and photo-chemical processes only for $O^+(^2D)$, $O^+(^2P)$, $O_2^+(X^2\Pi)$, N^+ , N_2^+ and NO^+ ions in the 120–620 km height range. A two-component model of the solar EUV from Nusinov (1992) is used to calculate the photo-ionization rates in 35-wavelength intervals (100–1050 Å). The photo-ionization and photo-absorption cross-sections are obtained from Torr et al. (1979) and Richards and Torr (1988). The only difference from the previous version of the model is in the $O^+ + N_2$ reaction rate constant. Recent flowing afterglow laboratory measurements by Hierl et al. (1997) are included in the model. These measurements were made at $T_n = T_i = T_v$ (where T_n is neutral, T_i is ion temperature and T_v is vibrational temperature of the excited N_2) in a wide temperature range and take into account the effects of vibrationally excited N_2^* . This may be important for summer high solar activity conditions (e.g. Pavlov, 1986; Ennis et al., 1995; Pavlov and Buonsanto, 1997; Pavlov et al., 1999, and references therein) considered in the present study. A comparison of different $O^+ + N_2$ reaction rate constants using EISCAT observations has shown that the Hierl et al. (1997) rate coefficient for this reaction may be recommended for aeronomic calculations (Mikhailov and Schlegel, 2000).

Table 1. Dependence of calculations on the upper boundary height specification for daytime conditions (13 UT). Vertical plasma drift W is given at 300 km. Factors for neutral concentrations are given with respect to MSIS-83 model values with T_{ex} , S and T_{120} shown in the table.

Height of the upper boundary, km	400	450	500	550	600
T_{ex} , K	1480	1475	1489	1479	1455
S , km^{-1}	0.0158	0.0159	0.0157	0.0156	0.157
T_{120} , K	389	390	383	383	386
Fac[O]	0.98	0.98	0.97	0.96	0.96
Fac[O ₂]	0.93	0.99	0.98	0.96	0.99
Fac[N ₂]	1.00	1.01	0.98	0.98	0.98
W , $m s^{-1}$	−9.0	−9.2	−8.4	−7.9	−7.6

Vertical plasma drift W used in the continuity equation for O^+ ions is obtained from the observed parameters as a difference between measured total vertical plasma velocity and diffusion velocity for O^+ ions. This is a standard approach usually used to find meridional thermospheric winds from IS observations (e. g., Buonsanto and Wittasse, 1999 and references therein). The expression (19.59) from Banks and Kockarts (1973) is used for W specification

$$W = V_z + \frac{k}{m_i \sum v_{ij}} \sin^2 I \left\{ T_i \frac{\ln N_i}{dh} + T_e \frac{d \ln N_e}{dh} + \frac{g m_i}{k} + \frac{d(T_e + T_i)}{dh} \right\} \quad (1)$$

where $V_z = V_i \sin I$, v_{ij} are diffusion collision frequencies for O^+ related to momentum transfer collision frequencies ν^* by the expression (see Eq. 19.13 in Banks and Kockarts, 1973) $v_{ij} = m_j / (m_i + m_j) \nu_{ij}^*$, where i applies to O^+ ions and j applies to other neutral or ionized gas species; all other symbols are standard. Collisions of O^+ ions with neutral O, O₂, N₂ and NO⁺, O₂⁺, N₂⁺, N⁺ ions were taken into account. All O^+ ion collision frequencies were taken from Banks and Kockarts (1973). Ion concentrations used in Eq. (1) are known at each iteration of fitting calculated $Ne(h)$ to the experimental one. Calculations are made with the upper boundary conditions specified at different heights in the 400–600 km height range. Normally the results are rather insensitive to the choice of the upper boundary height (see Table 1) and any calculation can be used as a solution. But the scatter of the measured V_i may increase with height (as observations show) and the reliability of the calculated median V_i decreases at high altitudes for such cases. This may result in some dependence of the solutions on the upper boundary height choice. Therefore, we usually specify the upper boundary at 500 km.

Observed (median and smoothed) $T_e(h)$ and $T_i(h)$ profiles are used in the calculations. No correction was ap-

Table 2. Calculated thermospheric parameters in comparison with MSIS-83 values (second line). Concentrations and meridional winds are given at 300 km.

Periods	T_{ex} , K	T_{120} , K	S , km^{-1}	$\log [\text{O}]$	$\log [\text{O}_2]$	$\log [\text{N}_2]$	V_{nx} , m s^{-1}
03 July 90	1411	407	0.019	8.615	7.326	8.614	+108.0
02 UT	1406	403	0.018	8.627	7.186	8.621	
03 July 90	1419	415	0.018	8.707	7.304	8.669	+32.3
04 UT	1422	415	0.017	8.669	7.216	8.627	
03 July 90	1480	422	0.017	8.768	7.338	8.744	-19.3
07 UT	1458	425	0.017	8.769	7.292	8.690	
03 July 90	1500	415	0.016	8.835	7.250	8.734	-8.7
10 UT	1508	412	0.017	8.810	7.304	8.730	
02 July 90	1489	383	0.016	8.714	7.136	8.595	-38.0
13 UT	1521	398	0.016	8.774	7.231	8.682	
02 July 90	1517	414	0.014	8.767	7.152	8.667	-15.6
16 UT	1556	410	0.015	8.728	7.239	8.667	
02 July 90	1507	400	0.015	8.645	7.249	8.571	+47.3
19 UT	1524	418	0.016	8.707	7.273	8.698	
02 July 90	1457	393	0.016	8.539	7.230	8.528	+145.7
21 UT	1454	410	0.017	8.682	7.248	8.692	

plied to these profiles, as the calculated O^+/Ne ratio usually is close to the standard EISCAT model for quiet time conditions (Fig. 4). Such a correction makes sense only for storm conditions when deviations from the model are essential (Mikhailov and Schlegel, 1997, 1998; Mikhailov and Foster, 1997; Mikhailov and Förster, 1999). In our case, the maximal deviation from the standard model is about 15% at 200 km (Fig. 4). This should result in a 15% correction for T_i and 13% for T_e at the 200 km height (Waldeufel, 1971) and this is not very important for the calculated aeronomic parameters, as our previous analysis has shown.

Using standard multi-regressional methods, we fit the calculated $N_e(h)$ profile to the observed one and find by this method: T_{ex} , shape parameter S for the $T_n(h)$ profile, T_{120} and factors for the MSIS-83 $[\text{O}]$, $[\text{O}_2]$, $[\text{N}_2]$ concentrations, as well as for the total EUV flux from the Nusinov (1992) model. The method by Mikhailov and Schlegel (1997, 2000) is still under development and various versions of this method exist. In one of them used in present study, exospheric temperature T_{ex} is included to the list of unknown parameters. This way of searching for T_{ex} is different from the earlier approach used in Mikhailov and Schlegel (1997), Mikhailov and Foster (1997), Mikhailov and Förster, (1997), and it turned out to be more straightforward and efficient, as well as more general, as it uses the most reliable parameter, $N_e(h)$ observed with the IS method, while $T_e(h)$ and $T_i(h)$

profiles depend on the ion composition model applied during the incoherent scatter data analysis (e.g. Lathuillère et al., 1983; Alcaydé et al., 1996). There is also a problem with the specification of the frictional term in the equation of ion energy conservation when electric fields are strong enough. Moreover, for strong convection electric fields, the ion velocity distribution is no longer Maxwellian (St.-Maurice and Schunk, 1979; Hubert and Kinzelin, 1992) and this basic assumption in the data analysis is not valid. There are also doubts if the energy conservation equation for O^+ ions in the F2-region normally used to find T_n takes into account all necessary processes (Oliver, 1997).

Although we are not dealing with the heights below 150 km, it was found that the method works better if one included the MSIS T_{120} value; thus, T_{120} was formally added to the list of searched parameters. But it should be stressed that this is just a technical step and the extension of T_n , $[\text{O}]$, $[\text{O}_2]$, $[\text{N}_2]$ down to 120 km height is just an extrapolation, as we do not fit any $N_e(h)$ profile below 160 km height.

The stationary form of the continuity equations is used in our method and daytime sunlit ionosphere is assumed. Therefore, in principle, only periods of relative stability in NmF2 and hmF2 variations around noon hours may be used for the analysis. But the analyzed period of 02/03 July 1990 shows very small NmF2 variations for most of the day (Fig. 1); therefore, the method was applied to more than just

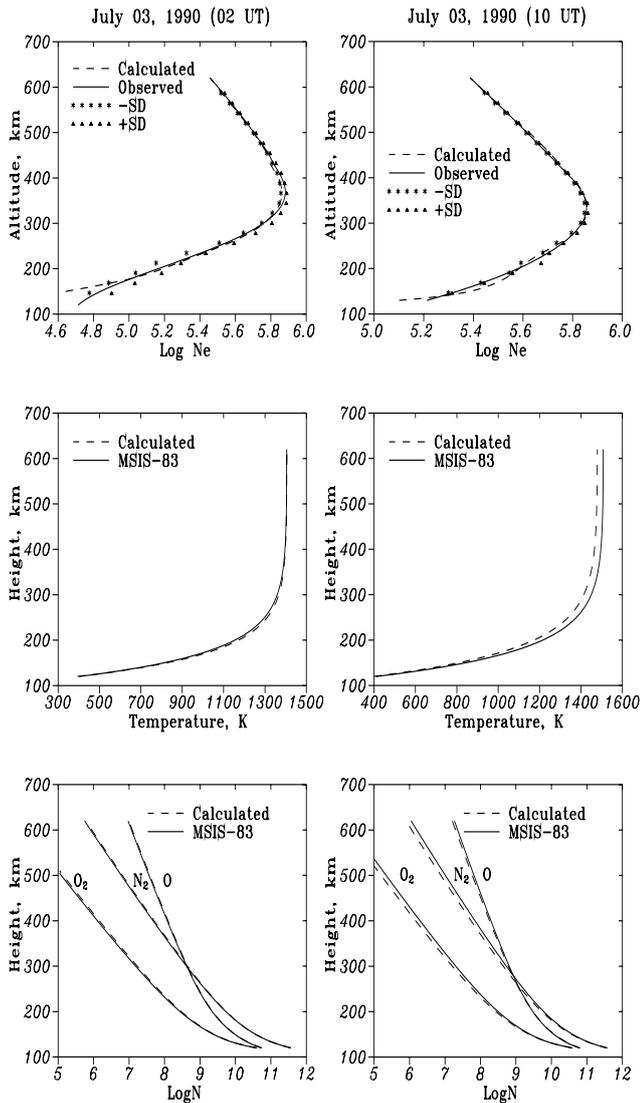


Fig. 2. Observed (together with a band of standard deviations) and calculated $N_e(h)$ profiles for two UT moments (top panel). Calculated along with MSIS-83 T_n , O, O_2 , and N_2 height profiles (middle and bottom panels).

the hours around noon and acceptable results were obtained (see Table 2). Only the hours around midnight, when the solar zenith angle was close to 90° , could not be developed properly. Two examples of $N_e(h)$ fitting and the calculated $T_n(h)$, [O], [O_2], and [N_2] height profiles, along with the MSIS-83 thermospheric model, are shown in Fig. 2.

Dependence of the calculations on the height of the upper boundary condition specification (where observed Ne values are used) is shown in Table 1 for daytime (13 UT) conditions ($LT=UT+1.3$). In general, the effect of the upper boundary height change is seen as small, but the large heights of the upper boundary (550 and 600 km) give W at 300 km, which differ from the other values. We suppose that this is the effect of decreasing the vertical velocity determination accuracy at high altitudes.

The analyzed periods and calculated thermospheric pa-

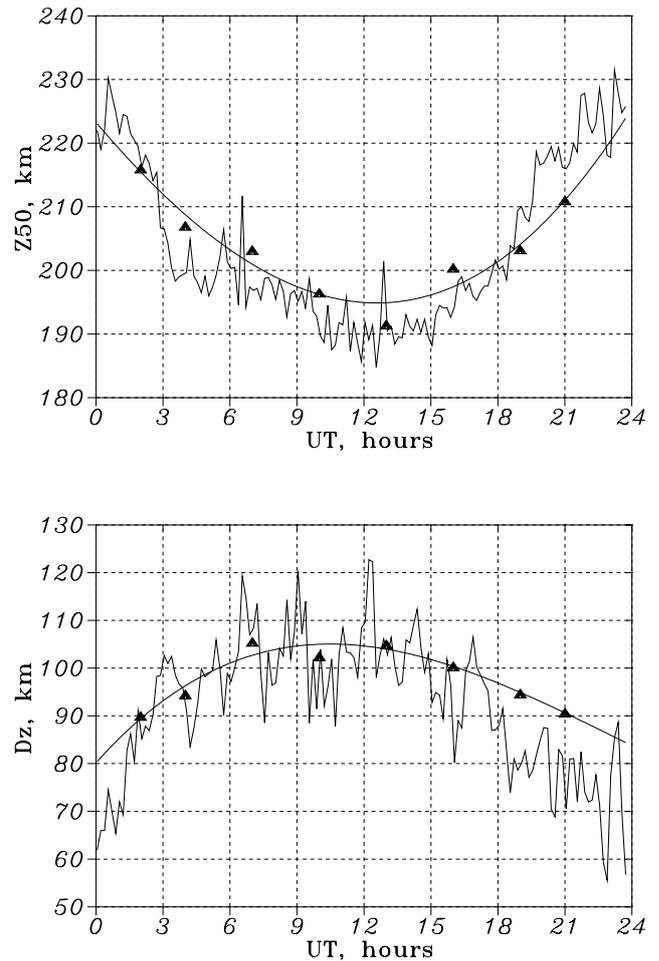


Fig. 3. Calculated (triangles) and experimental from Litvine et al. (1998) Z_{50} and D_z diurnal variations. A smooth curve is the least squares approximation of the calculated (triangles) values.

rameters, in comparison with the MSIS-83 model predictions, are given in Table 2. The calculated thermospheric parameters for this quiet day are seen to be close to the MSIS-83 model values (see also Fig. 2). Average absolute differences are: less than 1% for T_{ex} , 1.8% for T_{120} , 12% for [O], 18% for [O_2] and 16.5% for [N_2]. This is a normal result for quiet time periods provided by this method (e.g. Mikhailov and Schlegel, 1997; Mikhailov and Förster, 1999). The largest deviations from the MSIS-83 take place in the evening (21 UT) when the F2-layer obviously was non-stationary; abrupt changes of $hmF2$ are seen during these periods (Fig. 1). The calculated from W meridional thermospheric wind, V_{nx} shows usual diurnal variation being strong equatorward during nighttime and poleward during daytime hours (e.g. Lathuillère and Lilensten, 1997; Buonsanto and Witasse, 1999 and references therein).

The calculated Z_{50} and D_z diurnal variations; along with the Litvine et al. (1998) experimental values for summer high solar activity, are given in Fig. 3. Our calculations demonstrate the anti-correlation between Z_{50} and D_z variations, in accordance with the empirical model by Litvine et al. (1998).

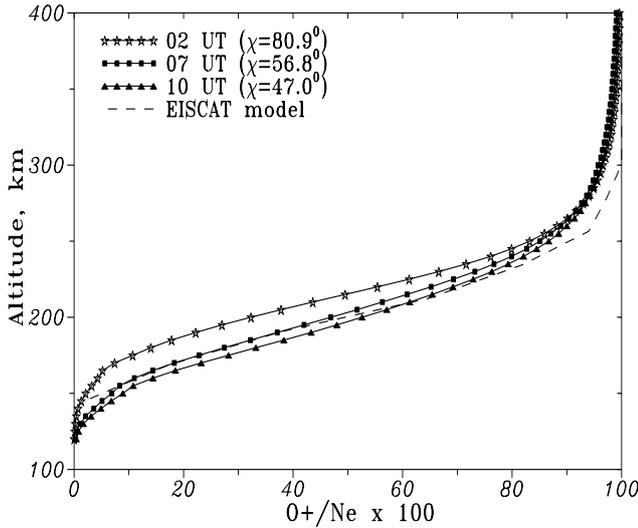


Fig. 4. Calculated height profiles of the O^+/Ne ratio at three solar zenith angles. The standard EISCAT ion composition model is shown for a comparison (dashes).

The transition level between atomic and molecular ions is the lowest (191 km) during daytime hours (13 UT) and the highest (216 km) during nighttime. The width of the 10–90% layer varies in an opposite way, maximizing (105 km) during daytime hours and minimizing (90 km) in the nighttime. On the other hand, the experimental values from Litvine et al. (1998) are seen to be shifted a little with respect to the calculated curves. This may be explained by the chosen period for our analysis (see later). Let us analyze the physical reasons for such diurnal variations of Z_{50} and D_z .

4 Interpretation

Calculated height profiles of the O^+/Ne ratio are shown in Fig. 4 for three solar zenith angles χ corresponding to 02, 07 and 10 UT. This ratio demonstrates systematic changes with the solar zenith angle. Any level of constant O^+/Ne ratio in the 10–90% range descends with the decreasing of the solar zenith angle. Figure 5 shows these variations for the three levels (10, 50, and 90%) discussed in the paper. The height variation of the 10% level is larger (about 35 km) compared to the 90% level variation (about 15 km). Therefore, the decrease in D_z is mainly due to the uplifts of the 10% level when we pass from daytime to nighttime hours. These results obtained from our calculations also follow from an analysis of a scheme of photochemical processes in the daytime ionosphere. For the analyzed period of very high solar activity and daytime hours, the photochemical equilibrium is valid for the main ions up to 200 km and even higher. For instance, the estimated characteristic times for O^+ ions, with respect to recombination $1/\beta = 1.5 \cdot 10^2$ s to diffusion H^2/D , is about $1.0 \cdot 10^5$ s, and to vertical transport H/W , is about $2.0 \cdot 10^4$ s at 200 km. A comparison of $N_e(h)$ profiles calcu-

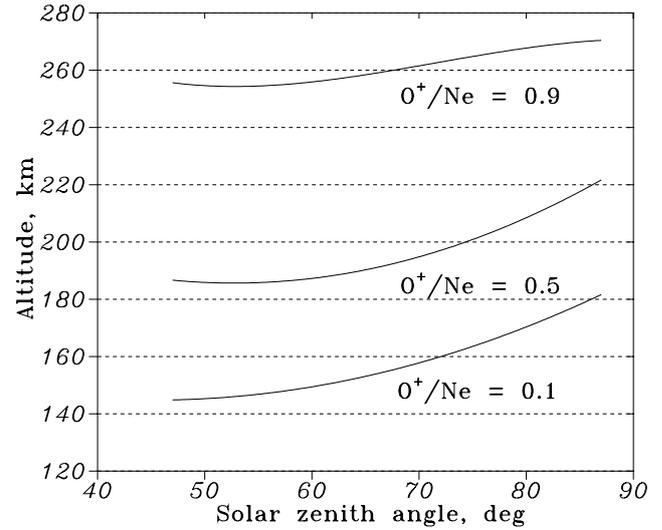


Fig. 5. Calculated variations of the O^+/Ne ratio at three levels (10, 50 and 90%) with solar zenith angle. Note the steepness of the variation is decreased at high altitudes (90% level).

lated with a complete set of processes and in a photochemical equilibrium is shown in Fig. 6 for daytime (13 UT) conditions. The observed $N_e(h)$ profile is shown as well in order to demonstrate the quality of model $N_e(h)$ fitting. Photochemical equilibrium is valid up to 200 km. In accordance with the mechanism of the F2-layer formation, additional plasma is transferred from the topside to the area just below the F2 maximum, where electron density turns out to be a little higher than what photochemical equilibrium provides. Therefore, a qualitative analysis of the ion composition variations can be done in the framework of the photochemical equilibrium.

If one leaves out only the main processes which control the ionosphere formation in the 140–220 km height range, then ion concentrations may be written as follows (e.g. Ivanov-Kholodny and Nikoljisky, 1969):

$$\begin{aligned}
 [O^+] &= \frac{q(O^+)}{\gamma_1[N_2] + \gamma_2[O_2]} \\
 [N_2^+] &= \frac{q(N_2^+)}{\gamma_3[O]} \\
 [O_2^+] &= \frac{q(O_2^+) + \gamma_2[O_2][O^+]}{\alpha_2 n_e} \\
 [NO^+] &= \frac{\gamma_1[N_2][O^+] + \gamma_3[O][N_2^+]}{\alpha_1 n_e} \\
 n_e &= [O^+] + [O_2^+] + [NO^+]
 \end{aligned}$$

Equilibrium concentration of N_2^+ ions is negligible compared to the main ions (e.g. Goldberg and Blumle, 1970). From these expressions, a quadratic equation may be written for the O^+/n_e ratio

$$\left(\frac{n_e}{O^+}\right)^2 - \frac{n_e}{O^+} - A = 0, \quad (2)$$

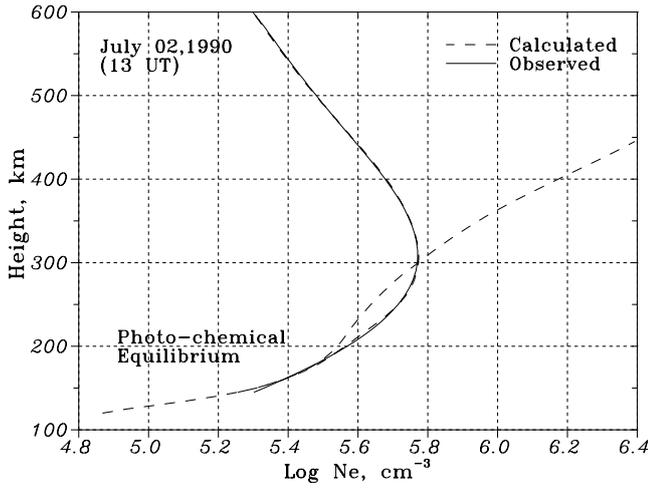


Fig. 6. Calculated $N_e(h)$ profiles with a complete set of processes and in a photo-chemical equilibrium for daytime conditions. The observed profile is shown as well.

where

$$A = \frac{\beta^2}{q(\text{O}^+)^2} \left[\frac{q(\text{O}_2^+)}{\alpha_2} + \frac{q(\text{N}_2^+)}{\alpha_1} \right] + \frac{\beta}{q(\text{O}^+)} \left[\frac{\gamma_2[\text{O}_2]}{\alpha_2} + \frac{\gamma_1[\text{N}_2]}{\alpha_1} \right] \quad (3)$$

and $\beta = \gamma_1[\text{N}_2] + \gamma_2[\text{O}_2]$

A positive solution of (2) for O^+/n_e is

$$\frac{[\text{O}^+]}{n_e} = \left(\frac{1}{2} + \sqrt{\frac{1}{4} + A} \right)^{-1} \quad (4)$$

The case of $A \ll 1/4$ corresponds to large heights, where $[\text{O}^+]/n_e \approx 1$. The other case of $A \gg 1/4$ corresponds to lower heights, where $[\text{O}^+]/n_e \ll 1$ and molecular ions dominate.

Height variations of the two terms in Eq. (3), in comparison with $1/4$, are given in Fig. 7 for the daytime hours of 02 July 1990. The condition $A \gg 1/4$ is valid up to 200 km height. The other result is that the two terms in Eq. (3) show similar height variations which are close to the height variation of A , at least up to 250 km; therefore, any of the terms in Eq. (3) can be used for further analysis.

Let us analyze the solution (4) with respect to the $[\text{O}^+]/n_e$ dependence on the solar zenith angle and height. For the sake of simplicity, we may assume that the isothermal neutral atmosphere consists of atomic oxygen and molecules, M (O_2 and N_2), distributed in accordance with the barometric law: $[\text{O}] = [\text{O}]_0 \exp(-z/H)$ and $[M] = [M]_0 \exp(-2z/H)$, where $H = kT_n/mg$ is the atomic oxygen scale height. The photoionization rates may be written as $q(\text{O}^+) = j_0[\text{O}] \exp(-a\text{Ch}\chi)$ and $q(M^+) = j_M[M] \exp(-a\text{Ch}\chi)$, where j is the ionization efficiency depending on the incident solar EUV flux and ionization cross-sections, $\text{Ch}\chi$ is

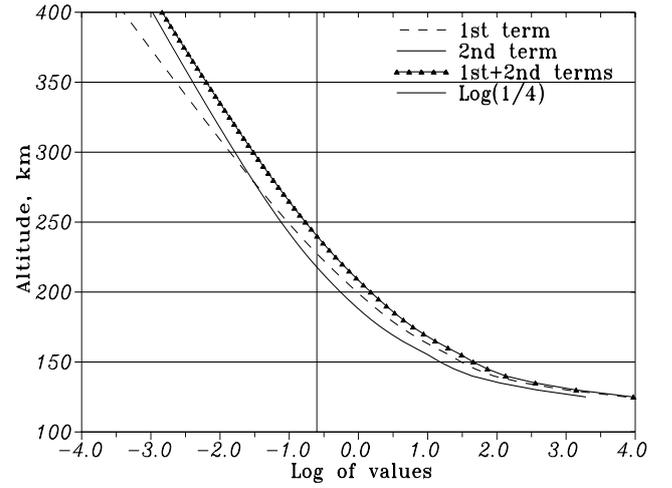


Fig. 7. Comparison of the two terms from the expression (3). Vertical line corresponding to $\log_{10}(1/4)$ (see (4)) divides the ionosphere by the lower and the upper parts where molecular and atomic ions dominate.

the Chapman function for the solar zenith angle χ , and a includes the column density of neutrals multiplied by absorption cross-sections, with a being the same for the atomic and molecular species. The linear loss coefficient β may be written as $\beta = \gamma[M]$. In this case, it may be shown that

$$A \cong \left[C_1 \exp\left(\frac{-4z}{H}\right) + C_2 \exp\left(\frac{-3z}{H}\right) \right] \exp(a\text{Ch}\chi) \quad (5)$$

where C_1 and C_2 are coefficients which are independent on height, z and the solar zenith angle, χ . As described above, any term in the expression for A can be used for the analysis below 200 km, as both demonstrate height variations similar to the height variation of A (see Fig. 7). Therefore, let us analyze the first term in Eq. (5). For lower altitude case ($A > 1/4$), we obtain from Eqs. (4) and (5):

$$[\text{O}^+]/n_e \approx C \exp\left(\frac{2z}{H}\right) \exp\left(-\frac{1}{2}a\text{Ch}\chi\right) \quad (6)$$

where $C = 1/\sqrt{C_1}$ is a coefficient independent on z and χ . Therefore, in the height range where $A > 1/4$, the $[\text{O}^+]/n_e$ ratio increases with height at any solar zenith angle and decreases with solar zenith angle at any height. The steepness of the $[\text{O}^+]/n_e$ dependence on $\text{Ch}\chi$ should decrease as we move to higher altitudes, where A becomes $< 1/4$. Using Eq. (6), it is possible to write an explicit expression for the height z of a given $R = [\text{O}^+]/n_e$ ratio

$$z = \frac{H}{2} \ln(R/C) + \frac{Ha}{4} \text{Ch}\chi \quad (7)$$

Using these analytical expressions, it is possible to interpret the results of our calculations obtained with an ionospheric model, which takes into account the whole set of processes.

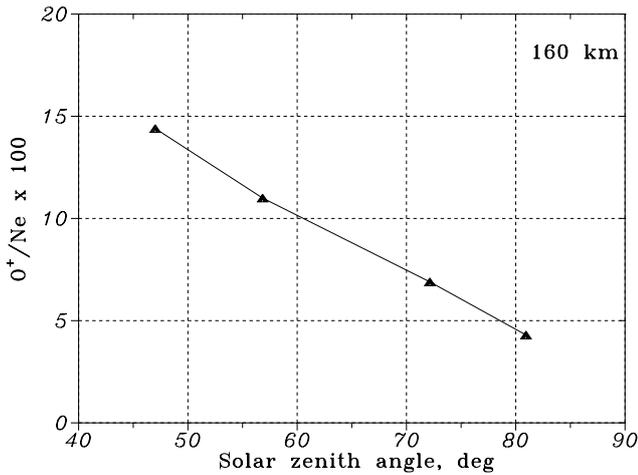


Fig. 8. Calculated dependence of the O^+/Ne ratio on the solar zenith angle at a fixed (160 km) height.

5 The dependence of Z_{50} on solar zenith angle

The $[O^+]/n_e$ ratio always increases with height at any solar zenith angle (Fig. 4), as it follows from Eq. (6). This is due to different laws of plasma recombination: a quadratic one at lower heights, where molecular ions dominate, and a linear one at high altitudes, where atomic oxygen ions dominate. On the other hand, the $[O^+]/n_e$ ratio decreases at a given height with solar zenith angle (Fig. 8), as it follows also from Eq. (6). Therefore, we have to shift to higher altitudes to keep a given $[O^+]/n_e$ ratio when the solar zenith angle increases. This dependence is clearly seen from Eq. (7) as well. This explains the observed diurnal variation of the Z_{50} parameter. From a physical point of view, the increase of the Z_{50} height during the night hours is just due to the shift of the ionization production maximum to higher altitudes, when $Ch\chi$ increases. Figure 9 gives height profiles for the production rates of O^+ , O_2^+ , N_2^+ ions, as well as the total production rate along with ion concentrations for the main ions for 13 and 19 UT. The effect of the sharp decrease in the rate of ion production, and a corresponding decrease in ion concentrations at lower heights is clearly seen. There is an overall shift of the experimental Z_{50} values with respect to the calculations (Fig. 3, top). This may be explained by the choice of the 02/03 July 1990 period for our analysis. Summer conditions at high solar activity used by Litvine et al. (1998) are presented primarily in 1992 when the average $F_{10.7}$ was around 120 for summer months, while for the analyzed period, the three monthly mean $F_{10.7}$ was 190 and the daily $F_{10.7}$ was 276. The higher neutral temperature on the day analyzed explains the shift between the experimental and calculated Z_{50} values in Fig. 3. The experimental Z_{50} values in Litvine et al. (1998, their Fig. 3, left hand, top) are also shifted to lower heights for low solar activity compared to solar maximum, and this may also be related to different neutral temperatures (see expression 7).

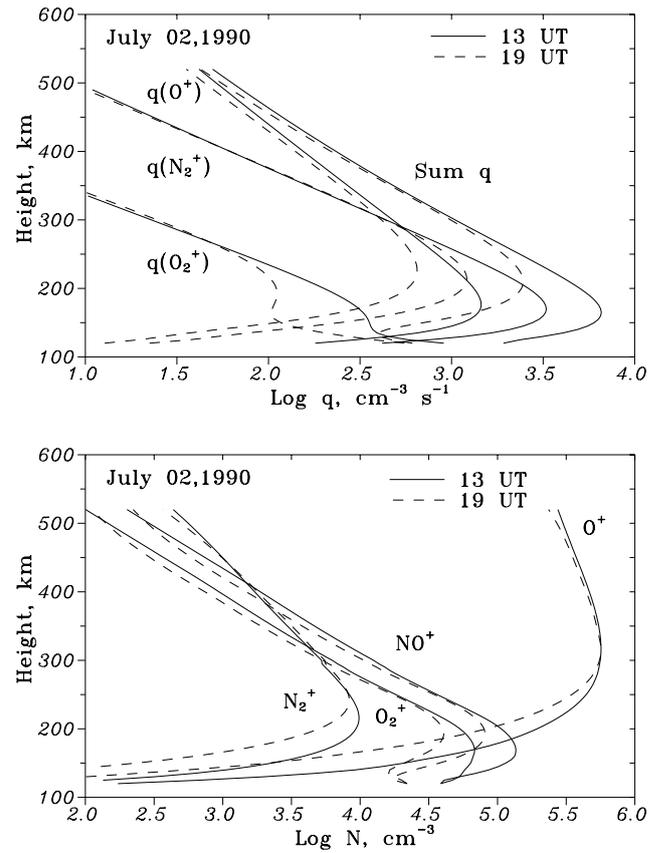


Fig. 9. Calculated height profiles of photo-ionization rates and corresponding ion concentrations for daytime (13 UT) and evening (19 UT) hours. Note the strong decrease in the photo-ionization rates at lower heights in the evening.

6 The dependence of D_z on solar zenith angle

As it follows from the results of model calculations (Fig. 5), the steepness of the $[O^+]/n_e$ dependence on the solar zenith angle is different at high and low altitudes. This follows from Eq. (4) when A becomes $< 1/4$ (high altitudes) and no dependence on the solar zenith angle for the $[O^+]/n_e$ ratio is expected. Therefore, the observed decrease of D_z (Fig. 3) is mainly due to the uplift of the 10% level when we pass to large solar zenith angles in the evening. Again, this is due to the variations of ionization production function for large solar zenith angles (Fig. 9). The calculated D_z values are seen to be larger than the experimental ones during nighttime hours (Fig. 3, bottom). This may be due to the chosen period (02/03 July) for our analysis, when the F2-layer, in fact, was sunlit during the nighttime hours, while the average experimental values include nighttime conditions with $\chi > 90^\circ$. Direct photo-ionization, as it takes place on 02/03 July, produces sufficient ion concentrations at the 10% O^+/Ne level, broadening the F2-layer and increasing D_z .

7 Conclusions

The anticorrelation between Z_{50} and D_z parameters in their diurnal variations, revealed for the first time by Litvine et al. (1998), may be explained using the present day understanding of the physical processes in the ionospheric F1 and lower F2-regions. Model calculations reproduce such diurnal variations of Z_{50} and D_z parameters, which are explained by different variations of ion composition at different ionospheric heights. The main reason for the observed Z_{50} and D_z variation is due to different ion composition (molecular at low and atomic at high altitudes), as well as to diurnal variations of the photo-ionization rates.

During low solar activity, variations of the Z_{50} and D_z parameters are practically the same as those at solar maximum (Litvine et al., 1998), but shifted a little bit to lower heights for the Z_{50} parameter and with less amplitude for the D_z variation. This may be explained by lower neutral temperatures during low solar activity.

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